

A peer-reviewed version of this preprint was published in PeerJ on 4 December 2014.

[View the peer-reviewed version](https://doi.org/10.7717/peerj.685) (peerj.com/articles/685), which is the preferred citable publication unless you specifically need to cite this preprint.

Bothwell LD, Selmants PC, Giardina CP, Litton CM. 2014. Leaf litter decomposition rates increase with rising mean annual temperature in Hawaiian tropical montane wet forests. PeerJ 2:e685
<https://doi.org/10.7717/peerj.685>

Leaf litter decomposition rates increase with rising mean annual temperature in Hawaiian tropical montane wet forests

Decomposing litter in forest ecosystems supplies nutrients to plants, carbon to heterotrophic soil microorganisms and is a large source of CO₂ to the atmosphere. Despite its essential role in carbon and nutrient cycling, the temperature sensitivity of leaf litter decay in tropical forest ecosystems remains poorly resolved, especially in tropical montane wet forests where the warming trend may be amplified compared to tropical wet forests at lower elevations. We quantified leaf litter decomposition rates along a highly constrained 5.2 °C mean annual temperature (MAT) gradient in tropical montane wet forests on the Island of Hawaii. Dominant vegetation, substrate type and age, soil moisture, and disturbance history are all nearly constant across this gradient, allowing us to isolate the effect of rising MAT on leaf litter decomposition and nutrient release. Leaf litter decomposition rates were a positive linear function of MAT, causing the residence time of leaf litter on the forest floor to decline by ~31 days for each 1 °C increase in MAT. Our estimate of the Q_{10} temperature coefficient for leaf litter decomposition was 2.17, within the commonly reported range for heterotrophic organic matter decomposition (1.5 - 2.5) across a broad range of ecosystems. The percentage of leaf litter nitrogen (N) remaining after six months declined linearly with increasing MAT from ~ 88% of initial N at the coolest site to ~74% at the warmest site. The lack of net N immobilization during all three litter collection periods at all MAT plots indicates that N was not limiting to leaf litter decomposition, regardless of temperature. These results suggest that leaf litter decay in tropical montane wet forests may be more sensitive to rising MAT than in tropical lowland wet forests, and that increased rates of N release from decomposing litter could delay or prevent progressive N limitation to net primary productivity with climate warming.

1 **Title:** Leaf litter decomposition rates increase with rising mean annual temperature in Hawaiian
2 tropical montane wet forests.

3 Lori Bothwell¹, Paul C. Selmants^{2*}, Christian P. Giardina³ and Creighton M. Litton²

4 ¹Natural Sciences Division, University of Hawaii at Hilo, Hilo, HI 96720 USA

5 ²Department of Natural Resources and Environmental Management, University of Hawaii at
6 Manoa, Honolulu, HI 96822 USA

7 ³Institute of Pacific Islands Forestry, Pacific Southwest Research Station, USDA Forest Service,
8 Hilo, HI 96720 USA

9 *Corresponding author, email: selmants@hawaii.edu

10 **Introduction:**

11 Litter decomposition is a fundamental biogeochemical process influencing rates of carbon
12 and nutrient cycling in forest ecosystems (Perry et al. 2008). Global syntheses indicate that
13 temperature is a primary factor controlling litter decay rates (Aerts 1997, Gholz et al. 2000, Adair
14 et al. 2008), but these datasets are dominated by temperate forest ecosystems. The factors
15 controlling litter decomposition in tropical wet forest ecosystems are less well studied (Cusack et
16 al. 2009, Wieder et al. 2009) and data from tropical montane wet forests are particularly sparse
17 (Malhi et al. 2010). The paucity of litter decomposition data from tropical montane wet forests
18 represents a significant gap in knowledge given recent evidence that litter decomposition in
19 montane tropical wet forests may be more sensitive to rising temperature than in lowland tropical
20 wet forests (Waring 2012) and that warming in the tropics may be occurring faster at higher
21 elevations (Giambelluca et al. 2008).

22 One of the more tractable approaches to estimate the temperature sensitivity of litter decay in
23 forests is through the use of elevation gradients, which can be powerful tools to examine climatic
24 controls on ecosystem functioning (Malhi et al. 2010). Observational studies along elevation
25 transects substitute space for time by examining litter decay rates in forests across a range of
26 temperature environments, and so have the advantage of representing the long-term, integrated
27 response of decomposition to changing mean annual temperature (MAT). Elevation gradients are
28 seldom a perfect proxy for climate warming, however, because other factors that influence
29 ecosystem processes may also vary with elevation, including plant species composition,
30 precipitation and soil moisture, geologic substrate, and soil chemical and physical properties.
31 These potentially confounding factors can complicate efforts to isolate the influence of
32 temperature on ecosystem functioning along elevation gradients (Wood et al. 2012).

33 We know of only two studies that have specifically examined the effect of temperature on leaf
34 litter decomposition along elevation transects in the tropics (Scowcroft et al. 2008, Salinas et al.
35 2011), both of which indicate that temperature is a primary factor controlling leaf litter
36 decomposition rates. However, the two studies vary widely in their estimates of the apparent Q_{10}
37 temperature coefficient, the proportional change in litter decay rate due to a 10 °C increase in
38 temperature, potentially because of the confounding effects of precipitation and soil moisture,
39 which also exert a strong control on litter decay (Schuur 2001). For example, precipitation
40 differences with elevation in Salinas et al. (2011) resulted in a three-fold variation in mean annual
41 soil moisture across sites of different elevation. Relatively few litter decomposition studies and
42 the wide range of results from those studies highlight the need for more research aimed at
43 isolating the influence of temperature on leaf litter decay in the carbon dense and highly
44 productive tropical montane wet forest biome.

45 Here we present results of a leaf litter decomposition experiment across a well-constrained
46 5.2 °C MAT gradient consisting of nine permanent plots in native-dominated tropical montane
47 wet forests spanning 800 m elevation along the eastern slope of Mauna Kea Volcano on the Island
48 of Hawaii. Many potentially confounding factors remain constant across this MAT gradient, such
49 as dominant canopy tree species, disturbance history, soil water content, geological substrate and
50 soil type (Litton et al. 2011, Iwashita et al. 2013, Selmants et al. 2014). We examined decay of a
51 common substrate (senescent *Metrosideros polymorpha* [Myrtaceae] leaves from a mid-elevation
52 plot) across the MAT gradient over a six-month time period to address two main research
53 questions: (i) Does rising MAT increase rates of leaf litter decomposition in tropical montane wet
54 forests when other environmental factors are held constant?; and (ii) does variation in MAT alter
55 the rate of nitrogen (N) release from decomposing leaf litter in tropical montane wet forests?

56 **Materials and methods:**

57 *Study sites*

58 We tested the effect of rising MAT on leaf litter decomposition rates by using a highly
59 constrained MAT gradient along the eastern flank of Mauna Kea volcano on the Island of Hawaii.
60 This MAT gradient consists of nine permanent 20 x 20 m plots in native-dominated tropical
61 montane wet forests spanning 800 – 1600 m in elevation, which corresponds to a 5.2 °C MAT
62 gradient (13.8 - 18.2 °C). All nine MAT plots are similar in factors other than mean annual
63 temperature that can affect ecosystem processes, including vegetation, disturbance history, soils,
64 parent material, and soil water balance (Litton et al. 2011, Iwashita et al. 2013, Selmants et al.
65 2014). Specifically, all nine MAT plots are in moderately aggrading mature forests with a canopy
66 dominated by *M. polymorpha* and a mid-story dominated by three species of tree fern (*Cibotium*
67 spp.). Soils are all well-drained Acrudoxic Hydruclands according to USDA Soil Classification
68 System, and are all derived from ~20,000 year-old volcanic ash deposits on top of a Pleistocene-
69 age lava flow dominated by the minerals hawaiiite and mugarite (Wolfe and Morris 1996, Litton
70 et al. 2011). Rainfall is not constant across the MAT gradient, but tends to decrease with
71 elevation (Table 1). However, potential evapotranspiration also decreases with elevation (Table
72 1; Giambelluca et al. 2014), which results in near constant mean annual soil water content across
73 the gradient (Litton et al. 2011). Solar radiation is also nearly constant across the MAT gradient
74 (Table 1), and all MAT plots are below the trade wind inversion layer (Cao et al. 2007). The
75 seven lower elevation MAT plots are in the Hawaii Experimental Tropical Forest and were
76 accessed by permission from the USDA Forests Service and the State of Hawaii Department of
77 Land and Natural Resources. The two highest elevation MAT plots are in the Hakalau Forest
78 National Wildlife Refuge and were accessed by permission from the US Fish and Wildlife
79 Service.

80 *Experimental design*

81 We examined the effect of MAT on leaf litter decomposition by following the decay of
82 senescent *M. polymorpha* leaves in each of the nine MAT plots for a six month period (June to
83 December, 2012). Rainfall is fairly evenly distributed throughout the year on windward slopes in
84 the Hawaiian Islands (Giambelluca et al. 2013) and seasonal changes in solar zenith angle,
85 although much less pronounced than in higher latitude temperate regions, lead to slightly warmer
86 air temperatures in summer months than in winter (Giambelluca et al. 2014).

87 We used intact senescent *M. polymorpha* leaves sorted from oven-dried litterfall samples
88 collected monthly in elevated trays at 1116 m (16.1 °C; Table 1) from July 2010 through March
89 2012. These leaves were composited and well mixed for construction of individual litterbags. We
90 used litter collected from one species at a single elevation as a standard substrate to isolate the
91 effect of MAT on litter decay by minimizing variation in litter nutrient concentration. Each
92 litterbag was 10 x15 cm, constructed of fiberglass screen with 1.5 mm mesh size and contained ~
93 2.5g of well-mixed senescent *M. polymorpha* leaves. We deployed 15 litterbags per MAT plot in
94 June of 2012. We placed a cluster of three litterbags tethered with nylon fishing line to a central
95 stake in five 5x5 m subplots within each MAT plot, resulting in five replicate litterbags for each
96 collection period in each MAT plot. All litterbags were placed on the soil surface (top of the O
97 horizon). We collected litterbags at one, three, and six months after initial placement. After each
98 collection, litter was oven-dried at 70 °C, weighed and finely ground in a ball mill for chemical
99 analysis. We did not correct for ash-free dry mass, but, because they were placed on top of the
100 undisturbed O horizon, there was no evidence of soil accumulation on either litterbag surfaces or
101 litter itself. Total N of initial, undecomposed leaf litter and of partially decomposed litter from the
102 three collection periods was determined by combustion using a Costech Elemental Analyzer at
103 the University of Hawaii at Hilo Analytical Laboratory (Costech Analytical Technologies,
104 Valencia, CA USA).

105 *Data analysis*

106 We estimated the decomposition rate of *M. polymorpha* leaf litter for each of the five replicate
107 litterbag clusters within each MAT plot by fitting a single-pool negative exponential model to the
108 litter mass data (Olson 2007) using non-linear regression with initial mass fixed at the measured
109 value (Adair et al. 2010):

$$110 \quad X(t) = e^{-kt} \quad (1)$$

111 where $X(t)$ is the proportional mass remaining at time t (in days) and k is the decomposition rate.
112 The proportional initial mass, $X(0)$, is equal to 1 by definition and so is not estimated as a model
113 parameter (Adair et al. 2010). The mean residence time of leaf litter on the forest floor was
114 calculated as $1/k$ for each set of litterbags. We calculated the Q_{10} temperature coefficient for litter
115 decomposition rates across the MAT gradient as:

$$116 \quad Q_{10} = (R_2/R_1)^{10/(T_2-T_1)} \quad (2)$$

117 Where Q_{10} is the proportional change in k due to a 10° C increase in MAT, R_1 and R_2 are
118 regression-derived estimates of k at the coolest and warmest MAT plots, and T_1 and T_2 are MAT
119 values of the coolest and warmest MAT plots (13° C and 18.2° C, respectively). The proportion
120 of initial N remaining at the end of the six-month incubation period was calculated by dividing
121 the mass of N in litter collected after six months by the mass of N in the initial undecomposed
122 litter (Schuur 2001).

123 We used ordinary least-squares linear regression to determine whether leaf litter
124 decomposition rate (k), leaf litter residence time ($1/k$) and the percentage of leaf litter N
125 remaining after six months varied significantly as a function of MAT. The plot ($n = 9$) was the
126 smallest experimental unit to which the treatment (MAT) was applied, and so we use within-plot
127 means of the five replicate litterbag clusters as the response variable for all linear regression
128 analyses. We also used regression analysis at the plot level to determine whether annual rainfall

129 and soil water content were significant predictors of leaf litter decay rates across the MAT
130 gradient. We calculated 95% confidence intervals of the percentage of initial N remaining in leaf
131 litter after one, three and six months to examine net N mineralization and net N immobilization at
132 each stage of decomposition across the MAT gradient. For all statistical tests, we set $\alpha = 0.05$ and
133 confirmed that the assumptions of normality and homoscedasticity were met. All statistical
134 analyses were performed in R version 3.0.2 (R Core Team, 2013).

135 **Results:**

136 Leaf litter decomposition rate (k) ranged from $1.67 \times 10^{-3} \text{ d}^{-1}$ to $4.03 \times 10^{-3} \text{ d}^{-1}$ across the MAT
137 gradient and k was a positive linear function of MAT ($R^2 = 0.65$, $p < 0.01$). Decomposition rates
138 increased by $0.21 \times 10^{-3} \text{ d}^{-1}$ for each 1° C increase in MAT. Consequently, leaf litter residence
139 time declined by 31 days for each 1° C increase in MAT (Fig. 1). The estimated Q_{10} for leaf litter
140 decomposition was 2.17.

141 Initial N concentration of mixed *M. polymorpha* litter was 8.5 mg g^{-1} . After six months of
142 decomposition, the percentage of N remaining in decomposing *M. polymorpha* leaves declined
143 significantly as a function of increasing MAT (Fig. 2), from $\sim 88\%$ of initial N at the coolest site
144 to $\sim 74\%$ of initial N at the warmest site, a decline of approximately two percentage points for
145 each 1° C increase in MAT. Nitrogen remaining in leaf litter was never significantly larger than
146 100% at any stage of decomposition within any of the nine MAT plots (Fig. 3), indicating there
147 was no net N immobilization in decaying leaf litter. Neither annual rainfall nor soil water content
148 during the six month experiment period (Table 1) were significant predictors of leaf litter decay
149 rates ($R^2 = 0.013$, $p = 0.34$ for annual rainfall; $R^2 = 0.03$, $p = 0.67$ for soil water content).

150 **Discussion:**

151 Quantifying the temperature sensitivity of leaf litter decomposition and nutrient release is
152 critical to understanding how forest ecosystem processes will respond to climate change. We used
153 a highly constrained MAT gradient to demonstrate that both mass loss and N release during leaf
154 litter decay increase linearly in response to rising MAT in Hawaiian tropical montane wet forests.
155 Globally, there is evidence that leaf litter decay is slowed by cool temperatures in tropical
156 montane forests (Waring 2012) and that climate warming in the tropics is occurring faster at
157 higher elevations (Bradley et al. 2004, Giambelluca et al. 2008). When combined with results
158 presented here, this evidence suggests leaf litter decomposition rates in tropical montane wet
159 forests may increase substantially with climate warming in the coming decades.

160 The rates of leaf litter decomposition reported here ($1.67 \times 10^{-3} \text{ d}^{-1}$ to $4.03 \times 10^{-3} \text{ d}^{-1}$ across the
161 MAT gradient) are consistent with other studies in montane wet forests in Hawaii (Hobbie and
162 Vitousek 2000, Schuur 2001, Scowcroft et al. 2008) and well within the two orders of magnitude
163 range of $3 \times 10^{-4} \text{ d}^{-1}$ to $3 \times 10^{-2} \text{ d}^{-1}$ reported for tropical wet forests globally (Gholz et al. 2000,
164 Cusack et al. 2009, Waring 2012). There are a number of factors aside from temperature that can
165 affect leaf litter decomposition rates in tropical montane wet forests, most notably leaching from
166 precipitation (Wieder et al. 2009), soil oxygen availability related to soil water content (Schuur
167 2001) and leaf litter chemistry (Wieder et al. 2009, Salinas et al. 2011). By decomposing a
168 common substrate across a highly constrained MAT gradient, we were able to isolate the effect of
169 temperature by largely controlling for the effects of precipitation, soil moisture, initial litter
170 chemistry and other potential confounding factors, demonstrating a strong linear increase in leaf
171 litter decomposition rate with rising MAT. This increase in leaf litter decay rate is in line with,
172 and likely contributes to, the substantial increase in soil-surface CO_2 efflux across this MAT
173 gradient (Litton et al. 2011). Notably, soil organic carbon storage remains constant across the
174 MAT gradient despite increased rates of litter decay and soil-surface CO_2 efflux with rising MAT
175 (Selmants et al. 2014, Giardina et al. 2014).

176 The Q_{10} temperature coefficient, which describes the rate of change in a biological or
177 chemical process over a 10° C interval, often falls within the range of 1.5 - 2.5 when applied to
178 heterotrophic processes such as organic matter decomposition (Reiners 1968, Kätterer et al. 1998,
179 Gholz et al. 2000, Hyvönen et al. 2005, Zhou et al. 2009). Our Q_{10} estimate for leaf litter
180 decomposition (2.17) is within this range, and is similar to the Q_{10} estimate for soil-surface CO₂
181 efflux (2.26) across the same MAT gradient (Litton et al. 2011), suggesting a consistent response
182 of carbon cycling rates to rising temperature in tropical montane wet forests. In contrast to our
183 results, Salinas et al. (2011) estimated a Q_{10} for leaf litter decay of 3.06 in five tropical forest
184 plots across an elevation gradient spanning 12.8 °C (11.1 – 23.9 °C). Although there is wide
185 variation in estimates of apparent Q_{10} from litter decomposition studies (Gholz et al. 2000, Adair
186 et al. 2008), we suggest the Salinas et al. (2011) estimate of litter decay Q_{10} is ~ 40% higher than
187 ours largely because of confounding factors related to site selection and data analysis technique.
188 First, the lowest temperature site in the Salinas et al. (2011) study was also the driest, with a
189 mean annual soil moisture nearly three-fold lower than the mean of the other four sites,
190 potentially depressing k for this low temperature site. In contrast, mean annual soil moisture is
191 nearly constant across our MAT gradient (Litton et al. 2011). Although there was some variation
192 in soil water content among our MAT plots during the six-month period when leaf litter was
193 decomposing, the driest plots during the experiment period were in the middle of our gradient
194 and the percentage variation in soil moisture was an order of magnitude less than that of Salinas
195 et al. (2011). Second, Salinas et al. (2011) used linear regression of log-transformed mass loss
196 data to estimate decomposition rates, which can result in sizable overestimates of k depending on
197 error structure (Adair et al. 2010). We used non-linear regression of untransformed
198 decomposition data, which consistently yields accurate k estimates (Adair et al. 2010). The
199 potential confounding effect of a three-fold variation in soil moisture combined with the use of

200 log-transformed mass loss data in the Salinas et al. (2011) study suggest that their estimate of
201 apparent Q_{10} may be artificially inflated.

202 We found no evidence of N limitation to leaf litter decomposition across the MAT gradient.
203 At all stages of decomposition, the proportion of initial N remaining in decomposing leaf litter
204 never significantly exceeded 100% at any of the MAT plots. This evidence is consistent with
205 results from a fertilization experiment at a site near the middle of our MAT gradient, in which
206 individual additions of N and phosphorus (P) had no effect on *M. polymorpha* leaf litter
207 decomposition and combined N and P additions had only a weak positive effect (Hobbie and
208 Vitousek 2000). We also found an overall trend of increased rates of N release from decomposing
209 litter with rising MAT. Taken together, this evidence suggests that warming will increase rates of
210 N cycling and availability in these forests, a response consistent with results from warming
211 experiments across a wide range of forest and grassland ecosystems (Rustad et al. 2001). We did
212 not measure leaf litter phosphorus (P) dynamics in this study, so it remains unclear how rates of P
213 release from decomposing leaf litter respond to rising MAT. However, since P does not directly
214 limit *M. polymorpha* leaf litter decomposition at a site near the middle of our gradient (Hobbie
215 and Vitousek 2000), we consider it likely that increasing rates of leaf litter decay with rising MAT
216 will increase rates of P release and availability similar to the trend for N documented here.

217 Results from this leaf litter decomposition experiment across a well-constrained MAT
218 gradient have two potentially countervailing implications for how the carbon balance of tropical
219 montane wet forests will respond to climate warming, at least within the MAT range studied here
220 and where increasing temperature does not drive significant, concomitant changes in leaf litter
221 chemistry or soil water balance. First, our results indicate that warming will increase rates of leaf
222 litter decay in tropical montane wet forests, which explains part of the warming-induced increase
223 in rates of soil-surface CO₂ efflux to the atmosphere (Litton et al. 2011). At the same time, these
224 results suggest that more rapid decomposition with warming should also increase rates of nutrient

225 release from decaying leaf litter, at least in forest ecosystems where decomposition is not
226 currently nutrient limited. The availability of nutrients may strongly regulate whether ecosystem
227 carbon sequestration keeps pace with rising atmospheric CO₂ concentrations (Luo et al. 2004).
228 An increase in rates of nutrient release from decaying leaf litter with climate warming, as
229 suggested by our results, could delay or even prevent the onset of progressive nutrient limitation
230 of ecosystem carbon sequestration.

231 **Acknowledgements:**

232 Thanks to Kupu Hawaii Youth Conservation Corps interns Nathanael Friday and Isaac Ito for
233 assistance with data collection.

234 **References:**

- 235 Adair, E. C., S. E. Hobbie, and R. K. Hobbie. 2010. Single-pool exponential decomposition
236 models: potential pitfalls in their use in ecological studies. *Ecology* 91:1225–1236.
- 237 Adair, E. C., W. J. Parton, S. J. del Grosso, W. L. Silver, M. E. Harmon, S. A. Hall, I. C. Burke,
238 and S. C. Hart. 2008. Simple three-pool model accurately describes patterns of long-term
239 litter decomposition in diverse climates. *Global Change Biology* 14:2636–2660.
- 240 Aerts, R. 1997. Climate, leaf litter chemistry and leaf litter decomposition in terrestrial
241 ecosystems: a triangular relationship. *Oikos*:439–449.
- 242 Bradley, R. S., F. T. Keimig, and H. F. Diaz. 2004. Projected temperature changes along the
243 American cordillera and the planned GCOS network. *Geophysical Research Letters*
244 31:L16210.
- 245 Cao, G., T. W. Giambelluca, D. E. Stevens, and T. A. Schroeder. 2007. Inversion Variability in the
246 Hawaiian Trade Wind Regime. *Journal of Climate* 20:1145–1160.
- 247 Cusack, D. F., W. W. Chou, W. H. Yang, M. E. Harmon, W. L. Silver, The LIDET Team. 2009.
248 Controls on long-term root and leaf litter decomposition in neotropical forests. *Global*

249 Change Biology 15:1339–1355.

250 Gholz, H. L., D. A. Wedin, S. M. Smitherman, M. E. Harmon, and W. J. Parton. 2000. Long-term
251 dynamics of pine and hardwood litter in contrasting environments: toward a global model of
252 decomposition. *Global Change Biology* 6:751–765.

253 Giambelluca, T. W., H. F. Diaz, and M. S. A. Luke. 2008. Secular temperature changes in
254 Hawai‘i. *Geophysical Research Letters* 35:L12702.

255 Giambelluca, T. W., Q. Chen, A. G. Frazier, J. P. Price, Y.-L. Chen, P.-S. Chu, J. K. Eischeid, and
256 D. M. Delporte. 2013. Online Rainfall Atlas of Hawai‘i. *Bulletin of the American
257 Meteorological Society* 94:313–316.

258 Giambelluca, T. W., X. Shuai, M. L. Barnes, R. J. Aliss, R. J. Longman, T. Miura, Q. Chen, A. G.
259 Frazier, R. G. Mudd, L. Cuo, and A. D. Businger. 2014. Evapotranspiration of Hawai‘i. Pages
260 1–178.

261 Giardina, C. P., C. M. Litton, S. E. Crow, and G. P. Asner. 2014. Warming-related increases in
262 soil CO₂ efflux are explained by increased below-ground carbon flux. *Nature Climate
263 Change*.

264 Hobbie, S. E., and P. M. Vitousek. 2000. Nutrient limitation of decomposition in Hawaiian
265 forests. *Ecology* 81:1867–1877.

266 Hyvönen, R., G. I. Ågren, and P. Dalias. 2005. Analysing temperature response of decomposition
267 of organic matter. *Global Change Biology* 11:770–778.

268 Iwashita, D. K., C. M. Litton, and C. P. Giardina. 2013. Coarse woody debris carbon storage
269 across a mean annual temperature gradient in tropical montane wet forest. *Forest Ecology
270 and Management* 291:336–343.

271 Kätterer, T., M. Reichstein, O. Andrén, and A. Lomander. 1998. Temperature dependence of
272 organic matter decomposition: a critical review using literature data analyzed with different
273 models. *Biology and fertility of soils* 27:258–262.

- 274 Litton, C. M., C. P. Giardina, J. K. Albano, M. S. Long, and G. P. Asner. 2011. The magnitude
275 and variability of soil-surface CO₂ efflux increase with mean annual temperature in Hawaiian
276 tropical montane wet forests. *Soil Biology and Biochemistry* 43:2315–2323.
- 277 Luo, Y., B. Su, W. S. Currie, J. S. Dukes, A. C. Finzi, U. Hartwig, B. A. Hungate, R. E.
278 McMurtrie, R. Oren, W. J. Parton, D. E. Pataki, R. Shaw, D. R. Zak, and C. B. Field. 2004.
279 Progressive Nitrogen Limitation of Ecosystem Responses to Rising Atmospheric Carbon
280 Dioxide. *BioScience* 54:731–739.
- 281 Malhi, Y., M. Silman, N. Salinas, M. BUSH, P. Meir, and S. Saatchi. 2010. Introduction:
282 Elevation gradients in the tropics: laboratories for ecosystem ecology and global change
283 research. *Global Change Biology* 16:3171–3175.
- 284 Olson, J. S. 2007. Energy storage and the balance of producers and decomposers in ecological
285 systems. *Ecology* 44:322–331.
- 286 Perry, D. A., R. Oren, and S. C. Hart. 2008. *Forest Ecosystems*. Second edition. Johns Hopkins
287 University Press, Baltimore, MD.
- 288 Reiners, W. A. 1968. Carbon dioxide evolution from the floor of three Minnesota forests. *Ecology*
289 49:471–483.
- 290 Rustad, L. E., J. L. Campbell, G. M. Marion, R. J. Norby, M. J. Mitchell, A. E. Hartley, J. H. C.
291 Cornelissen, J. Gurevitch, and N. A. Not Available. 2001. A meta-analysis of the response of
292 soil respiration, net nitrogen mineralization, and aboveground plant growth to experimental
293 ecosystem warming. *Oecologia* 126:543–562.
- 294 Salinas, N., Y. Malhi, P. Meir, M. Silman, R. Roman Cuesta, J. Huaman, D. Salinas, V. Huaman,
295 A. Gibaja, M. Mamani, and F. Farfan. 2011. The sensitivity of tropical leaf litter
296 decomposition to temperature: results from a large-scale leaf translocation experiment along
297 an elevation gradient in Peruvian forests. *New Phytologist* 189:967–977.
- 298 Schuur, E. A. G. 2001. The effect of water on decomposition dynamics in mesic to wet Hawaiian

299 montane forests. *Ecosystems* 4:259–273.

300 Scowcroft, P., D. R. Turner, and P. M. Vitousek. 2008. Decomposition of *Metrosideros*
301 polymorpha leaf litter along elevational gradients in Hawaii. *Global Change Biology* 6:73–
302 85.

303 Selmants, P. C., C. M. Litton, C. P. Giardina, and G. P. Asner. 2014. Ecosystem carbon storage
304 does not vary with mean annual temperature in Hawaiian tropical montane wet forests.
305 *Global Change Biology* 20:2927–2937.

306 Waring, B. G. 2012. A Meta-analysis of Climatic and Chemical Controls on Leaf Litter Decay
307 Rates in Tropical Forests. *Ecosystems* 15:999–1009.

308 Wieder, W. R., C. C. Cleveland, and A. R. Townsend. 2009. Controls over leaf litter
309 decomposition in wet tropical forests. *Ecology* 90:3333–3341.

310 Wolfe, E. W., and J. Morris. 1996. Geologic Map of the Island of Hawaii. U.S. Geological
311 Survey, Reston, VA.

312 Wood, T. E., M. A. Cavaleri, and S. C. Reed. 2012. Tropical forest carbon balance in a warmer
313 world: a critical review spanning microbial- to ecosystem-scale processes. *Biological*
314 *Reviews* 87:912–927.

315 Zhou, T., P. Shi, D. Hui, and Y. Luo. 2009. Global pattern of temperature sensitivity of soil
316 heterotrophic respiration (Q_{10}) and its implications for carbon-climate feedback. *Journal of*
317 *Geophysical Research* 114:G02016.

Table 1 (on next page)

Environmental characteristics of the nine permanent plots along a 5.2 °C mean annual temperature gradient in tropical montane wet forests on the Island of Hawaii.

^aMean annual air temperature and mean annual soil volumetric water content from Litton et al. (2011).

^bMean monthly soil volumetric water content during the leaf litter decomposition experiment (June to December, 2012).

^cMean annual rainfall estimates from Giambelluca et al. (2013)

^dMean annual potential evapotranspiration and solar radiation estimates from Giambelluca et al. (2014).

MAT Plot	Elevation (m)	Air temperature (°C) ^a	Rainfall (mm y ⁻¹) ^b	Soil VWC: annual (m ³ m ⁻³) ^a	Soil VWC: experiment (m ³ m ⁻³) ^b	Potential evapotranspiration (mm y ⁻¹) ^c	Solar radiation (W m ⁻² y ⁻¹) ^c
1	800	18.2	4570	0.55	0.67	2298	201.1
2	934	17.3	4292	0.55	0.64	2232	200.9
3	1024	16.7	3975	0.57	0.63	2214	202.4
4	1116	16.1	3734	0.48	0.61	2127	204.9
5	1116	16.1	3433	0.51	0.47	2137	210.1
6	1204	15.5	3181	0.40	0.42	2211	214.5
7	1274	15.1	3101	0.51	0.44	2234	216.2
8	1468	13.8	4119	0.55	0.61	1888	202.6
9	1600	13.0	3282	0.57	0.60	1961	213.1

PeerJ PrePrints

Figure 1(on next page)

Leaf litter residence time across a mean annual temperature gradient on the Island of Hawaii.

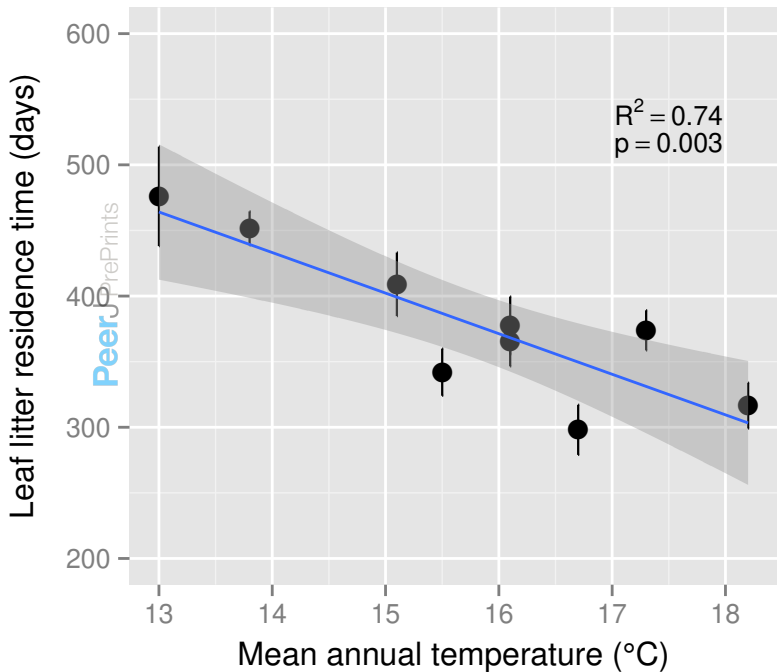


Figure 2 (on next page)

Nitrogen remaining in leaf litter after six months of decomposition across a mean annual temperature gradient on the Island of Hawaii.

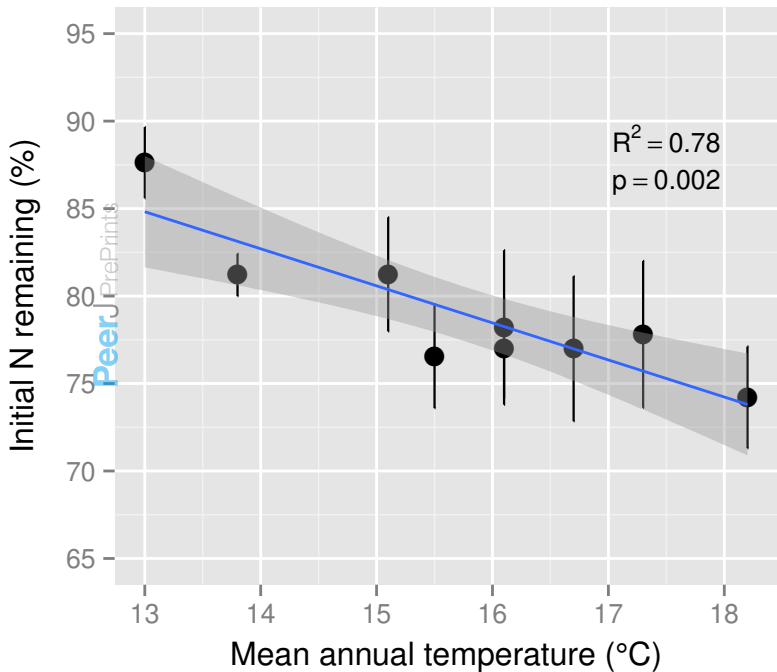


Figure 3(on next page)

Nitrogen remaining in decomposing leaf litter at three stages of decomposition across a mean annual temperature gradient on the Island of Hawaii.

