

A peer-reviewed version of this preprint was published in PeerJ on 12 May 2015.

[View the peer-reviewed version](https://doi.org/10.7717/peerj.935) (peerj.com/articles/935), which is the preferred citable publication unless you specifically need to cite this preprint.

Blakeway D, Hamblin MG. 2015. Self-generated morphology in lagoon reefs. PeerJ 3:e935 <https://doi.org/10.7717/peerj.935>

Self-generated morphology in lagoon reefs

The shapes and forms of coral reefs are generally attributed to external controls such as substrate topography or hydrodynamic influences. Little is known about inherent reef morphology in the absence of external control. Here we use reef growth simulations, based on observations in the cellular reefs of Western Australia's Houtman Abrolhos Islands, to show that reef morphology is fundamentally determined by the mechanical behaviour of the reef-building organisms themselves—specifically their tendency to either remain in place or to collapse. A spectrum of realistic reef forms can be generated by varying a single parameter representing this characteristic. Reef-building organisms that tend to remain in place, such as massive and encrusting corals or coralline algae, produce nodular reefs, whereas those that tend to collapse, such as branching *Acropora*, produce cellular reefs. Inherent reef growth forms are best expressed in sheltered lagoons. The purest forms arise where a single type of reef builder prevails, as in the cellular reefs of the Abrolhos. In these cases reef morphology can be considered a phenotype of the predominant reef building organism.

Self-generated morphology in lagoon reefs

David Blakeway^{1,2} and Michael G. Hamblin³

¹School of Earth and Environment, University of Western Australia, Perth, Western Australia

²Fathom 5 Marine Research, Lathlain, Western Australia

³School of Mechanical and Chemical Engineering, University of Western Australia, Perth, Western Australia

Corresponding author: David Blakeway, Fathom 5 Marine Research, 17 Staines Street, Lathlain, Western Australia, 6100, Australia, +61 439 134 618, d_blakeway@yahoo.com.au

Introduction

The three-dimensional form of a coral reef influences, and is influenced by, virtually all processes occurring within and around it (Roberts et al. 1975; Hopley, Smithers & Parnell 2007). Understanding the morphological development of reefs can therefore help provide a context for interpreting present-day reef processes and predicting their future.

Reef morphology is determined by different influences at different scales, from the inherent forms of reef-building organisms at the small scale to the configuration of continental shelves at the large scale. Over intermediate scales (metres to kilometres) reefs exhibit a great diversity of forms, reflecting the multitude of interacting processes affecting them. But within the diversity is a subset of globally recurring forms, indicating that there are some universal rules governing reef morphology (Hatcher 1997; Blanchon 2011; Purkis and Schlager 2014). Traditionally, these rules have been envisaged as external controls operating at or above the scale of the morphological features, for example substrate topography (MacNeill 1954, Purdy 1974; Choi & Ginsburg 1982) or the wave field (Munk & Sargent 1954; Roberts 1974). While these factors undoubtedly influence the morphology of many reefs, they raise an interesting question: what would reefs look like in the absence of such external influences? This question brings the focus down to the reef-building organisms themselves. Because these organisms cumulatively become the reef, there is significant potential for behaviour and events at their scale to be expressed in reef morphology at the intermediate scale. Such ‘emergence’ of self-organised patterns from small scale processes is ubiquitous in physical and biological systems (Nicolis and Prigogine 1977; Camazine et al. 2001). While it is recognised that coral reefs are likely to exhibit this trait (Mistr and Bercovici 2003; Rietkirk and van de Koppel 2008; Blanchon 2011; Schlager and Purkis 2014), it has not been demonstrated.

Lagoons are the most likely setting for inherent reef growth forms to arise, as they are generally flat-floored and sheltered. Several characteristic lagoon reef forms are repeated worldwide, ranging from simple mound-like patch reefs to complex cellular¹ reefs (Stoddart 1969; Hopley 1982; Blanchon 2011; Fig. 1). The development of patch reefs can be readily visualised in terms of expansion from a nucleus, and this mode of growth has been

¹ Cellular reefs (terminology after Hoskin 1963) are also widely known as reticulate reefs, from the latin *reticulum*: a network or net-like structure. However reticulate is also applied to another characteristic lagoon reef form: anastomosing reefs. Therefore we consider cellular and anastomosing as subdivisions of reticulate. As in most natural classifications, boundaries are transitional.

43 demonstrated repeatedly, from various nuclei including topographic highs in underlying
44 limestones (Halley et al 1977; Mazzullo et al. 1992), sedimentary structures (Perry et al.
45 2012; Novak et al. 2013), or early-colonising corals (Jones 1977; Crame 1981). Cellular reef
46 morphology, in contrast, is not an intuitive growth form. Cellular reefs distinctly resemble
47 negative landforms, particularly karst terrains (terrestrial erosion landforms created in
48 limestone through dissolution by rainwater). Based on this resemblance, and the recognition
49 that the foundations of most reefs have been exposed to at least 100,000 years of weathering
50 during Interglacial periods, cellular reef morphology has long been interpreted as an
51 inheritance from underlying karst (Fairbridge 1948; Purdy 1974; Guilcher 1988; Searle 1994;
52 Macintyre, Precht & Aronson 2000; Purkis et al. 2010; Kan et al. 2014). However there has
53 always been an opposing view attributing cellular morphology to reef growth (Dakin 1919;
54 Rossfelder 1990; Collins et al. 1993; Wyrwoll et al. 2006; Barott et al. 2010; Blanchon 2011;
55 Schlager and Purkis 2014). The growth alternative is gradually gaining acceptance, having
56 been confirmed for the cellular reefs of Mataiva Atoll in French Polynesia (GIE Raro Moana
57 1985; Rossfelder 1990) and the Houtman Abrolhos Islands in Western Australia (Collins et al
58 1993; Collins, Zhu & Wyrwoll 1996; 1998; Wyrwoll 2006). Seismic surveys and coring at
59 Mataiva showed the Holocene reef to be 10-20 m thick and demonstrated that, although the
60 reef is underlain by a karst Tertiary limestone, the karst features are relatively small-scale and
61 were infilled before submergence, such that the present reef morphology is independent of
62 the substrate (Rossfelder 1990). Similarly, seismic and coring in the Abrolhos lagoons
63 recorded a Holocene reef thickness of 40 m over a flat Last Interglacial grainstone, again
64 demonstrating independence from the substrate (Collins et al. 1993; Collins, Zhu & Wyrwoll
65 1996; 1998).

66
67 While this work has removed all doubt that the cellular reefs of Mataiva and the Abrolhos
68 have grown into their present configuration, it has not provided a generally accepted growth
69 mechanism. Three alternative mechanisms have been proposed. The first, developed
70 independently by GIE Raro Moana (1985) at Mataiva and Barott et al. (2010) at Millenium
71 Atoll, is the colonisation of the lagoon floor by networks of massive corals, which are
72 subsequently colonised by other corals and grow upward to the surface. The second,
73 proposed by Wyrwoll et al. (2006) for the Abrolhos, is growth to sea level of isolated
74 branching *Acropora* pinnacle, reticulate and stellate (star-shaped) patch reefs, which
75 subsequently extend and coalesce to surround enclosed depressions. The third, proposed by
76 Blanchon (2011), is a self-limitation mechanism based on negative feedback between reef

growth and water circulation—reef growth reduces water circulation which reduces reef growth, such that the cellular depressions become self-reinforcing as the surrounding reefs grow.

The network mechanism (GIE Raro Moana 1985; Barott et al. 2010) has an inherent problem explaining how massive corals might colonise the lagoon floor in a network. While it can't be discounted it seems unlikely. Coalescence of pinnacle, reticulate and stellate patch reefs (Wyrwoll 2006) seems a better mechanism but doesn't explain why patch reefs composed of a uniform coral type developing on a uniform surface would develop non-uniform shapes. Blanchon's (2011) self-limiting concept is arguably a much better proposition as nothing is predefined; structure emerges from the interaction between reef builders and their environment. However it is a loosely defined concept with no details of the interaction and no specific mechanism initiating the cellular morphology.

Here we examine the process of cellular reef growth using computer simulation based on field observations in the Abrolhos cellular reefs. These reefs provide an ideal case study for cellular reef development due to their flat pre-Holocene substrate, known accretion history and very pure reef-building community—cores through the Abrolhos cellular reefs consist almost entirely of branching *Acropora*, with a few tabular *Acropora* appearing as the reefs approached sea level (Collins et al. 1993; Collins, Zhu & Wyrwoll 1996; 1998).

Survey methods

An apparent sequence of cellular reef development is evident in the Abrolhos lagoons, progressing from circular patch reefs to coalescing stellate reefs surrounding semi-enclosed depressions to a reef platform surrounding enclosed depressions (Wyrwoll et al. 2006; Fig. 1; Fig. 2). Under the assumption that these are sequential stages of reef development, surveys of the circular-stellate-enclosed sequence represent surveys through time. Space-for-time substitution (Darwin 1842; Davis 1899) may therefore be applied to describe the evolution of the Abrolhos cellular reefs and, potentially, to reveal their formative mechanism. We examined replicate sites of each stage in a 15 km² cellular reef complex known as 'the Maze' in the Easter Group of the Abrolhos (Fig. 2; 28°41'S, 113°49'E). We surveyed fifteen sites in detail and many more in brief visits, including some in the Pelsaert Group to the south and the Wallabi Group to the north. At each of the fifteen sites we established four transects oriented to the cardinal directions, running from the deepest to shallowest habitat. We

quantified substrate composition by point counting video records of each transect, and constructed topographic profiles by recording tide-corrected depth every metre along each transect. We used twenty five categories in point count analysis and condensed these into seven categories for graphical representation: tabular *Acropora*, branching *Acropora*, other scleractinia, soft coral, macroalgae, dead coral, and sediment/rubble.

Survey results

Underwater observations show that the different reef stages are all connected in a continuous reef blanket with a distinctive 'egg carton' form. The relationship between the reef shapes can be envisaged by imagining sequentially deeper horizontal slices through a solid egg carton. The initial slices contact the peaks, producing circular reefs. Deeper slices reach the ridges between adjacent peaks, producing stellate reefs. Subsequent slices produce a platform surrounding enclosed depressions and eventually a solid platform. Below we describe the sequence in the three idealised stages: circular, stellate and enclosed. However it should be noted that the sequence is a continuum and that sites within each stage often have features of earlier and/or later stages. Fig. 3 shows representative transect profiles from each stage and Fig. 4 is a schematic block diagram incorporating the main features of the idealised reef stages.

Circular reefs

Circular reefs occur around the margin of the Maze (e.g. sites K and M; Fig. 2). The reef at these sites has a sinusoidal profile, curving up over dome-shaped reef tops then descending into bowl-shaped depressions (Fig. 3). The wavelength and amplitude of the profile varies within a range of approximately 40-100 m and 15-30 m respectively. Adjacent circular reef tops are linked by saddle-shaped ridges, and the topography is covered by a blanket of branching *Acropora*, primarily *A. formosa* and *A. abrolhosensis*. Tabular *A. spicifera* is abundant at site M on the exposed northern margin of the Maze but absent from the more sheltered site K to the south. Live *Acropora* cover is 60-100% on the reef tops and ridges at both sites, decreasing to approximately 30% within the site M depressions and 0% within the more enclosed and restricted site K depressions. Dead *Acropora* branches at depth are occupied by macroalgae at both sites, predominantly *Sargassum* spp. at site M and *Lobophora variegata* at site K. Beyond the outermost circular reefs, the reef surface slopes down to a flat sandy seafloor at 35 m without breaking into isolated patches. Corals on the

outer slopes are predominantly tabular and branching *Acropora* to 25 m depth. Below 25 m the community is more diverse, consisting predominantly of foliose agariciids and pectiniids.

Three islands of storm-deposited *Acropora* rubble line the eastern margin of the Maze (Collins, Zhao & Freeman 2006), and a series of submerged east-west trending linear reef banks occur on the northern margin of the Maze. Two banks can be seen in Fig. 2 and two deeper banks lie beyond them. The bank crests are 2 to 15 m deep, sloping downward at approximately 30° to U-shaped troughs at 20 to 30m. The outermost bank reaches the seafloor at 35m. Coral cover and zonation on the banks is equivalent to that of the circular reefs.

Stellate reefs

In the stellate stage the circular reef tops and shallow ridges reach sea level, producing a discontinuous network of flat-topped, coralline algal-cemented and macroalgal-covered stellate platforms (sites B, E, F, J, L, N). Water circulation within the intervening depressions is further reduced and the water column is often stratified and stagnant. Live coral is consequently restricted to shallow depths, often less than 15 m in the more enclosed depressions. Shallow subtidal habitat remains dominated by live branching *Acropora* (Fig. 5) and occasional foliose *Montipora*. Dead *Acropora* branches at depth are colonised by *Nephthea* soft corals and *Lobophora variegata*, and fine sediment accumulates in the bases of the depressions. A distinct shallow coral community begins to appear in the stellate stage, consisting of diverse massive and encrusting corals, predominantly faviids, pectiniids, *Montipora* and *Astreopora*. Several apparent developmental stages of this community are present, initiating as a discontinuous cover of small colonies on dead *Acropora* branches (Fig. 6A) and culminating in vertical or overhanging walls descending from the surface to as deep as ten metres, but typically between two and eight metres (Fig. 6B).

Enclosed depressions

In the enclosed stage (sites C, D, I, O) the previously-established patterns strengthen: most depressions are rimmed by vertical walls, the water column is often stratified and turbid, live coral cover declines rapidly with depth, and most depressions have a deep sediment fill. Late-stage enclosed depressions (A, G, H) gradually fill with sediment to the level of the surrounding reef flat. As they fill, the fringe of live *Acropora* beneath the vertical walls

migrates upward, eventually overgrowing the walls and encroaching over the depression floors.

Based on the descriptions above, the Abrolhos cellular reefs appear to exhibit a straightforward developmental sequence. However they provide little insight into the origin of cellular morphology because the cellular form is already present at the circular reef stage. Given the higher coral cover on the reef tops and ridges than the depressions, subsequent growth will inevitably lead to enclosed depressions from the circular reef stage. Several processes appear to be suppressing live coral cover at depth within the circular reef depressions, including reduced water circulation and macroalgal colonisation. But without being able to see further back into the Holocene it is impossible to determine whether any of the current processes could have initiated the cellular morphology or whether they are consequential.

Model

In the absence of conclusive indicators of cellular origin in the real reefs, we used a cellular automaton to model their development (Blakeway 2000; see Data S1 for code). A cellular automaton (Ulam 1962; von Neumann 1963) is an array of identically-programmed interacting cells. This structure is well-suited to modeling reef development because each cell in the array can be considered to represent a square metre of seafloor, and reefs can develop on the seafloor as three-dimensional stacks of cubic ‘corals’. Using this approach we first describe a basic patch reef system in which growth is essentially random and unconstrained but for sea level, and subsequently introduce a parameter representing branching *Acropora*.

Each run of the model is initialised by defining the dimensions of the array, the seafloor depth, and the number of coral ‘recruits’. We used a consistent configuration for most model runs: an array of 160 x 160 cells (representing 160 x 160 m, or 2.56 hectares, of seafloor), a uniform depth of 30 m, and 64 randomly-spaced coral recruits, occupying 0.25% of the array (variations to this configuration are described later in supporting figures). We chose 30 m as the default depth, rather than the 40 m of the Maze, because cellular reefs elsewhere appear to be thinner than the Maze; those at Mataiva for example are only 10-20 m thick (Rossfelder 1990). The 160 x 160 m array is embedded within a larger array of 250 x 250 m, giving the reefs a 45 metre margin for lateral growth.

Corals can grow upward and outward in increments of one metre per iteration, with vertical growth halted at sea level. Each cell in the array is assigned two growth probabilities in each iteration: a vertical probability representing the likelihood of the cell growing upward itself and a neighbour probability representing the likelihood of the cell being overgrown by a neighbouring coral (Fig. 7). The vertical probability of vacant seafloor cells is zero and the vertical probability of coral-filled cells is random. The neighbour probability of each cell is the product of a random number between zero and one and a 'neighbour value' that depends on the state of the eight surrounding cells. Cells with no shallower neighbours are assigned a neighbour value of zero; otherwise the cell's neighbour value rises incrementally for each shallower neighbour. If a cell becomes surrounded by shallower neighbours it is guaranteed to be overgrown. Otherwise growth is determined by comparing the cell's vertical and neighbour probabilities against two random numbers between zero and one. If either or both probabilities exceed their respective random numbers, the cell grows by one metre when the array is updated prior to the next iteration.

Patch reefs created with the random model appear circular in plan view and steeply conical in oblique view (Fig. 8A and B). They maintain their conical form as they enlarge, and become flat-topped at sea level. In the multiple patch reef system, each individual patch develops similarly, eventually coalescing and fusing with neighbouring patches (Fig. 8C and D). While these reefs appear realistic they bear little resemblance to cellular reefs. In fact their shapes are the inverse of cellular reefs; the model reefs appear nodular and subcircular in plan view whereas cellular reefs are concave and stellate, surrounding subcircular depressions. However, the random model is generic and does not intentionally represent any particular coral type.

In considering modifications to represent branching *Acropora*, one aspect of the random reefs is immediately apparent—coral spires projecting up to four metres above neighbouring cells (Fig. 8B). Such projections cannot occur on real branching *Acropora* reefs because, due to their 'brittle tree' morphology, colonies that grow to project more than a metre or two above their surroundings will inevitably collapse (Maragos 1972; Bak 1976; Fig. 9). This is not necessarily a disadvantage. Because broken fragments can survive and grow to form new colonies, fragmentation is recognised as an inherent and significant mode of reproduction and short range dispersal in branching *Acropora* (Gilmore and Hall 1976; Tunnicliffe 1981; Bothwell 1982; Highsmith 1982). Collapse is represented in a branching coral model by

imposing a maximum height differential between neighbours (hereafter termed collapse limit) of two metres, such that corals growing to project two metres above any neighbouring cell are prevented from growing upward until the deeper cell grows. Although they cannot grow upward, projecting corals contribute to the growth probability of neighbouring cells in two ways: first, they ‘support’ neighbouring corals, ensuring they are unrestricted by the collapse limit, and second they may ‘collapse into’ deeper neighbouring cells (i.e. they increase the neighbour probability of those cells). Although this representation of collapse involves no subtraction of height from the collapsing colony, it remains valid because it is equivalent to the projecting colony growing a metre then collapsing back a metre during the iteration.

Reefs created with the branching coral model closely resemble the Abrolhos cellular reefs (Fig. 10). The model reproduces the characteristic egg carton form of the real reefs and all its corollaries including circular reefs, stellate reefs with radiating ridges, scalloped platform margins, bowl-shaped depressions and the presence of multiple small depressions within larger multi-lobed depressions. The 45° reef slopes of the model reefs are steeper than the mean of the real *Acropora* slopes (37° ± SD 6°) but within their recorded range.

Fig. 11 shows sequential stages in the development of the branching coral reefs, demonstrating the emergence of their egg carton form. The key process is the formation of ridges between adjacent patch reefs. This process begins when the patch reefs meet, whereupon the valleys between them grow rapidly upward to become saddle-shaped ridges (Fig. 11C and D). The depressions surrounded by the reefs and ridges are initially irregular in outline but are progressively smoothed to subcircular shapes as the surrounding reef grows. Eventually the depressions become completely enclosed within the reef platform and infilled by coral (Fig. 11E).

Progressively relaxing the collapse limit produces intermediate forms between the branching and random models. A three metre collapse limit creates reefs with weakly developed subcircular depressions (Fig. 12A and B) and a four metre collapse limit creates reefs with very few depressions (Fig. 12C and D). Collapse limits of more than four metres produce nodular reefs equivalent to those of the random model.

We examined the morphological effects of several variations to the model, including altering depth (Fig. S1), altering colonisation density (Fig. S2), and incorporating sea level rise and

depth-dependent accretion (Fig. S3). With respect to the cellular morphology, the most influential modification was reducing water depth. Reducing depth reduces reef thickness, which constrains the morphological expression of the collapse limit such that the appearance of the branching coral reefs transforms from cellular to nodular as depth decreases. In the extreme case of reefs grown in only one or two metres water depth, where the collapse limit has no effect, all variants of the model produce identical nodular reefs.

Discussion

Model

The resemblance in shape and form between the model branching coral reefs and the Abrolhos cellular reefs suggests that the model adequately represents reality. This interpretation is supported by the model's simplicity: it has only one rule—collapse if too steep—which is intuitively reasonable and supported by field observations. Model reef morphology is hyper-sensitive to that rule, running through a nodular to cellular spectrum as permissible steepness is reduced and collapse becomes more frequent.

The nodular reefs produced by the random model appear straightforward and visually 'correct' as growth forms, because the individual patch reefs maintain their forms as they grow and merge. This straightforward morphology is indicative of pure in-situ (in place) growth. Cellular reefs are more complex because the patch reefs transform as they merge, eventually becoming linked by a network of ridges. This transformation results from the high frequency of collapse in the branching coral model. However, it is not simply the frequency of collapse that produces ridges; more important is the distribution of collapse. Because the valleys between merging patch reefs are low points in the reef structure, coral colonies in the valleys are less likely to project above their neighbours than corals elsewhere. Consequently, they are relatively unrestricted by the collapse limit and are therefore more likely to remain in place as they grow, and less likely to collapse, than colonies elsewhere (Fig. 13). The retention of in-situ colonies transforms the V-shaped valleys into U-shaped ridges that grow to sea level, enclosing depressions (Fig. 11C and D). The subcircular shapes of the depressions arise through the same non-uniform distribution of collapse. Colonies in the re-entrant concavities of the early-stage depressions are supported by neighbours and therefore tend to remain in place while those on projecting convexities tend to collapse. Over time this creates smooth rounded shapes, the ultimate smooth shape being a circle.

Abrolhos cellular reefs

The foregoing descriptions of the model branching coral reefs provide two testable predictions regarding real cellular reefs. First, their slopes should have a consistent and relatively low gradient, representing the angle of repose (maximum slope stability angle) of branching *Acropora*. Second, the proportion of collapsed colonies should be lowest in the valleys and ridges and highest on reef slopes. Both predictions are supported in the Abrolhos, where *Acropora* slopes average $37^{\circ} \pm \text{SD } 6^{\circ}$ (Fig. 4) and *Acropora* colonies in valleys and ridges are generally upright (Fig. 5) while those on reef slopes are often overturned (Fig. 9). We conclude that the Abrolhos cellular reefs have developed according to the model and that Fig. 11 closely describes their morphological progression.

One significant difference between the real and model reefs is the reduced accretion rate of the real reefs once they reach sea level. Model reefs first reach sea level in approximately 100 iterations and only require 60 more iterations to fill the platform, whereas the Maze reefs first reached present sea level in approximately 4000 years (Eisenhauer et al. 1993; Collins et al. 1993) but still have not filled the platform 7000 years later. The reduced accretion of the real reefs probably results from two factors not represented in the model. The first is the reduction of *Acropora* cover and vitality at depth as water circulation is restricted by the enclosure of the depressions (Wyrwoll 2006). This is an example of self-limitation through negative feedback between reef growth and water circulation (Blanchon 2011). Self-limitation is therefore a significant influence on the Abrolhos cellular reefs, but operates primarily on their accretion rate not their morphology. The second factor is the colonisation of upper depression slopes by relatively slow-growing massive and encrusting corals. The steep walls created by these corals effectively exclude branching *Acropora*, because any branching *Acropora* that colonise the walls are likely to fall off once they grow too large to be supported by their basal attachment. By ‘engineering’ steep walls (*sensu* Jones et al. 1994), massive and encrusting corals monopolise—for thousands of years—prime shallow subtidal habitat that would otherwise be occupied by fast-growing branching *Acropora*. The combination of reduced water circulation at depth and steep walls in the shallows restricts live *Acropora* to a fraction of their previous distribution, significantly slowing the overall reef accretion rate. Model cellular reefs, in contrast, rapidly fill the platform because ‘live’ *Acropora* occupy all habitats including the depression slopes and floors.

Another significant difference between the real and model reefs is the series of linear reef banks on the northern margin of the Maze. We interpret these as early to mid Holocene wave-deposited structures, resulting either from extratropical storms, cyclones (Scheffers et al. 2012) or tsunami (Scheffers et al. 2008).

Application to other reefs

The morphology of cellular reefs elsewhere appears similar enough to the Abrolhos reefs to suggest they have developed the same way, an inference supported by the abundance of branching *Acropora* in documented examples (Alacran Reef: Hoskin 1963; Solomon Islands: Morton 1974; Cocos-Keeling Atoll: Williams 1994; Pelican Keys: Aronson et al. 1998; Elizabeth Reef: Woodroffe et al. 2004; Pohnpei: Turak and DeVantier 2005; Tun Sakaran: Montagne et al. 2013; Palau-no reference yet; Nagura Bay: Kan et al. 2014). At least two of these examples, the Solomon Islands (Morton 1974) and Nagura Bay (Kan et al. 2014), exhibit vertical walls of massive and encrusting corals above the *Acropora* zone, suggesting they have undergone the late-stage shallow coral community succession observed in the Abrolhos.

We have separated Mataiva Atoll from the list above as *Porites* is abundant there and has been considered responsible for the cellular morphology (GIE Raro Moana 1985; Delesalle 1985). However, branching *Acropora* are also abundant at Mataiva (Delesalle 1985; Rossfelder 1990). We suggest that branching *Acropora* are the primary reef-builders at Mataiva and *Porites* are colonisers of the *Acropora* reef, not framework builders. Another possible exception to the rule of *Acropora* dominance is the 'Type-1' reticulate reef of the Red Sea (Purkis et al. 2010) where *Porites* is also abundant (Bruckner 2011). However, we would not classify all Type-1 Red Sea reefs as cellular because, although deep, the depressions are not always circular, more often resembling the transitional depressions produced by intermediate collapse limits (see Purkis et al.'s Fig. 2). Some Red Sea reefs are distinctly cellular and we predict those to be *Acropora*-dominated (e.g. 27°57'N, 35°13'E). Closer examination of these and other cellular reefs is required to determine whether the predominance of branching *Acropora* is universal, and whether the slope gradients and the distribution of collapsed colonies conform to the Abrolhos example.

We have not classified the previously-mentioned Millenium Atoll with the cellular reefs listed above because the scale of the cellular morphology at Millenium and in many other

Pacific atoll lagoons is up to an order of magnitude larger than the Abrolhos. While the large-scale cellular reefs also seem to consist predominantly of branching *Acropora* (Roy and Smith 1971; Grovhoug and Henderson 1973; Valencia 1977; Barott et al. 2010), we do not believe our model applies directly to them because it cannot produce cells of their dimensions unless it is scaled up massively, to unrealistic depths of at least 100 m. We are currently investigating the large-scale cellular morphology.

The intermediate to nodular shapes produced in the model are common in real reefs (e.g. Fig. 14). The simplest interpretation of these shapes is that they indicate coral types, or mixtures of coral types, progressively less prone to collapse than branching corals. In this interpretation, nodular shapes (i.e. circular in plan view) represent reef builders that rarely collapse, such as massive and encrusting corals or coralline algae. However, the model indicates that intermediate to nodular shapes are not necessarily diagnostic of coral type, because even branching coral patch reefs appear nodular before they merge with adjacent patch reefs, and after they merge in shallow water. This suggests that additional information on water depth, reef slope gradients and reef thickness will be required to reliably infer ecology from morphology in intermediate to nodular reefs. Such three-dimensional data is becoming increasingly available through reef-oriented remote sensing (Goodman et al. 2013 and references therein). In two-dimensional aerial images, however, the only diagnostic morphology is cellular—signifying relatively thick ($\geq 10\text{m}$) reefs constructed by collapse-prone organisms.

The default collapse-prone reef builders on modern reefs are branching *Acropora*. While it seems possible for other branching coral genera or other calcareous branching invertebrates (e.g. *Millepora*) to create cellular reefs, observations worldwide (listed above) suggest it is invariably *Acropora*: *A. cervicornis* in the Caribbean and multiple species in the Indo-Pacific. This is probably because branching *Acropora* have the ultimate strategy for rapid pre-emption of space in lagoon environments. *Acropora* branches not only grow quickly (up to 19 cm.yr^{-1} in the Maze; Blakeway 2000), they regularly develop new branches which themselves branch and rebranch, giving them the potential for exponential expansion (Shinn 1976). Constant growth and branching produces an open three-dimensional structure that allows *Acropora* to rapidly fill lagoons. Our model indicates that, given adequate depth, cellular reefs are the inevitable result. Cellular reefs are essentially a phenotype of the branching

Acropora genome(s), emerging from the innate behaviour of branching *Acropora* colonies just as colony morphology emerges from the innate behaviour of polyps.

If the relationship between cellular reef morphology and branching *Acropora* holds, the distinctive shapes of cellular reefs in remotely-sensed images can potentially be used to identify and map branching *Acropora* habitat. This could be useful in reef health assessments, as the sensitivity of branching *Acropora* to environmental conditions makes them something of a canary in the coral reef coalmine (Loya et al. 2001; Roth and Dehyn 2013). However, care should be taken in attributing cause and effect in apparently degraded *Acropora*-dominated lagoons because natural self-limitation and community succession can drastically reduce *Acropora* cover and vitality in the mid to late stages of reef development. Before assessing anthropogenic impacts, careful work may be required not only to understand natural intrinsic environmental and successional change but also to determine if and how the reef community may have been modified by less predictable natural extrinsic influences (Aronson et al. 1998).

Conclusions

Our research indicates that the fundamental control on reef morphology is the nature of the reef building organisms. This control is best expressed in lagoons, where diminished hydrodynamic and substrate influences allow reefs to grow into their inherent forms. The purest growth forms arise where a single type of reef builder prevails, as in the cellular reefs of the Abrolhos. In these cases reef morphology can be considered a phenotype of the predominant reef building organism.

The extent to which reef building organisms either remain in place or collapse appears to explain the nodular to cellular spectrum of lagoon reef morphology. Many more relationships between reef ecology and morphology remain to be discovered. Most of the recurrent patterns in reef morphology, including the large-scale cellular reefs and anastomosing reefs, are likely to be partly or wholly ecological phenomena (Blanchon 2011; Schlager and Purkis 2014). Investigation of these relationships is benefiting from the increasing availability and resolution of remotely sensed imagery. However, the single most important research technique remains careful underwater observation, preferably without any preconceptions regarding process. Any consistent correlations between reef morphology and underwater

survey data, such as coral type, can be considered potential causal relationships warranting closer examination.

In surveying modern reefs it should be recognised that a reef's current state may not represent its formative Holocene (or earlier) state, particularly if the reef has reached sea level. While seismic and coring can access the history stored within such reefs, both techniques are logistically demanding and expensive. The complementary methods we employed in the Abrolhos, space-for-time substitution and computer simulation, are relatively simple and inexpensive but can provide a comprehensive reconstruction of a reef's history and a sound basis for extrapolating its future development.

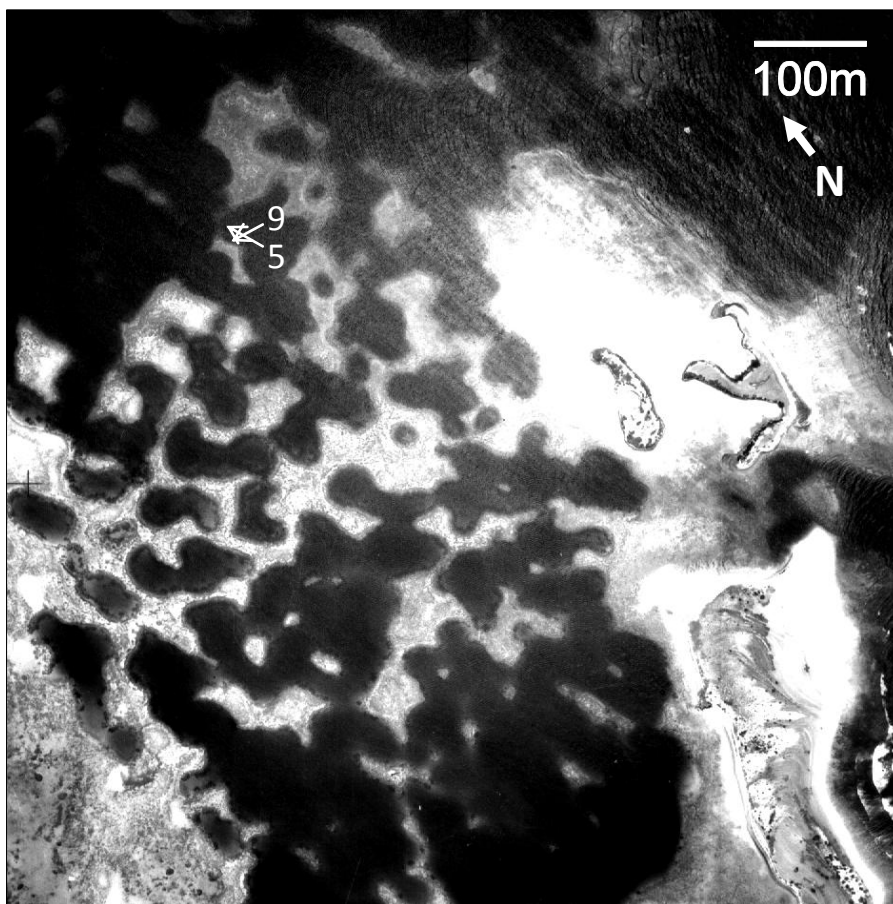


Figure 1. Cellular reefs in the northeastern lagoon of the Pelsaert Group, Houtman Abrolhos Islands, 28°54'S 114°E. Islands of storm-deposited coral rubble line the eastern platform margin. The numbers at top left refer to subsequent figures and the direction of the arrows corresponds to the direction from which the figures (photographs) were taken. (note the arrows clash in the pdf conversion – to be fixed, maybe better as an enlarged inset?).

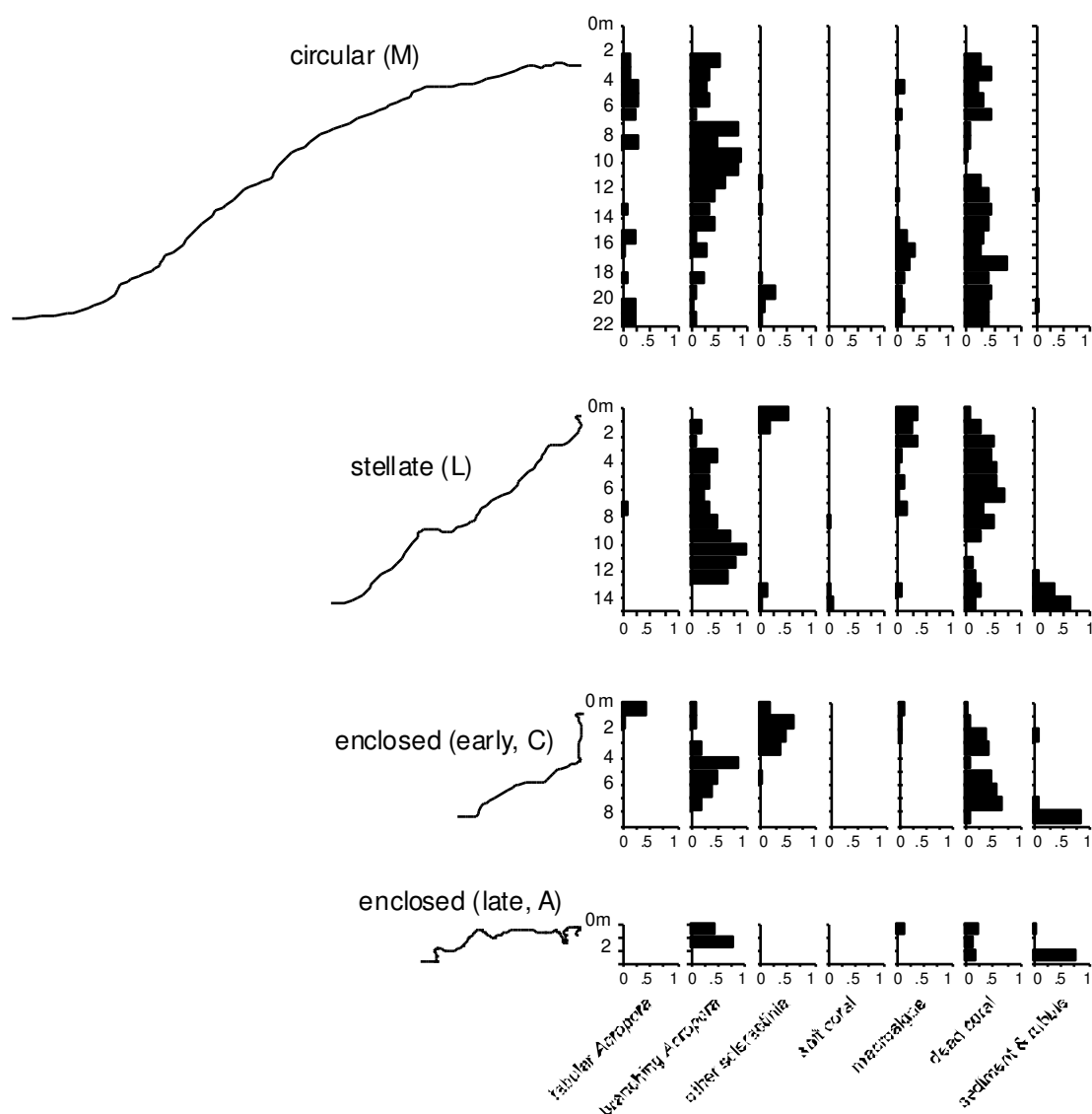


Figure 3. Representative transect profiles and substrate composition from sites of each stage. The x-axis legend (text problem to be fixed) reads tabular *Acropora*, branching *Acropora*, other scleractinia, soft coral, macroalgae, dead coral, sediment & rubble.

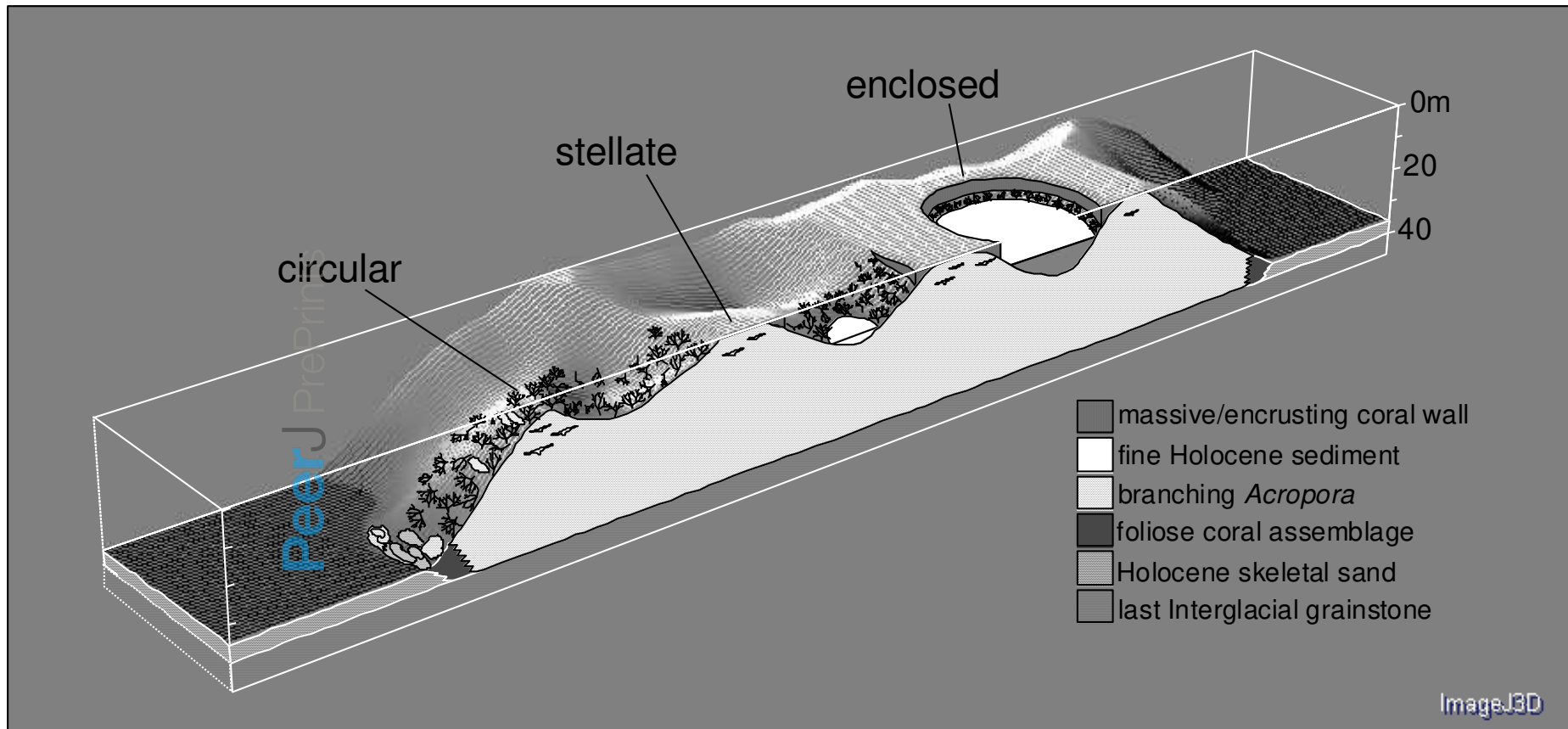


Figure 4. Block diagram illustrating the three idealised reef stages. The cross-section is hypothetical but consistent with survey observations and seismic/core data. (note this figure has come out poorly in pdf, the eps is better)



Figure 5. Dense branching *Acropora* on the ridge extending northeast from the stellate reef shown in Fig. 1.

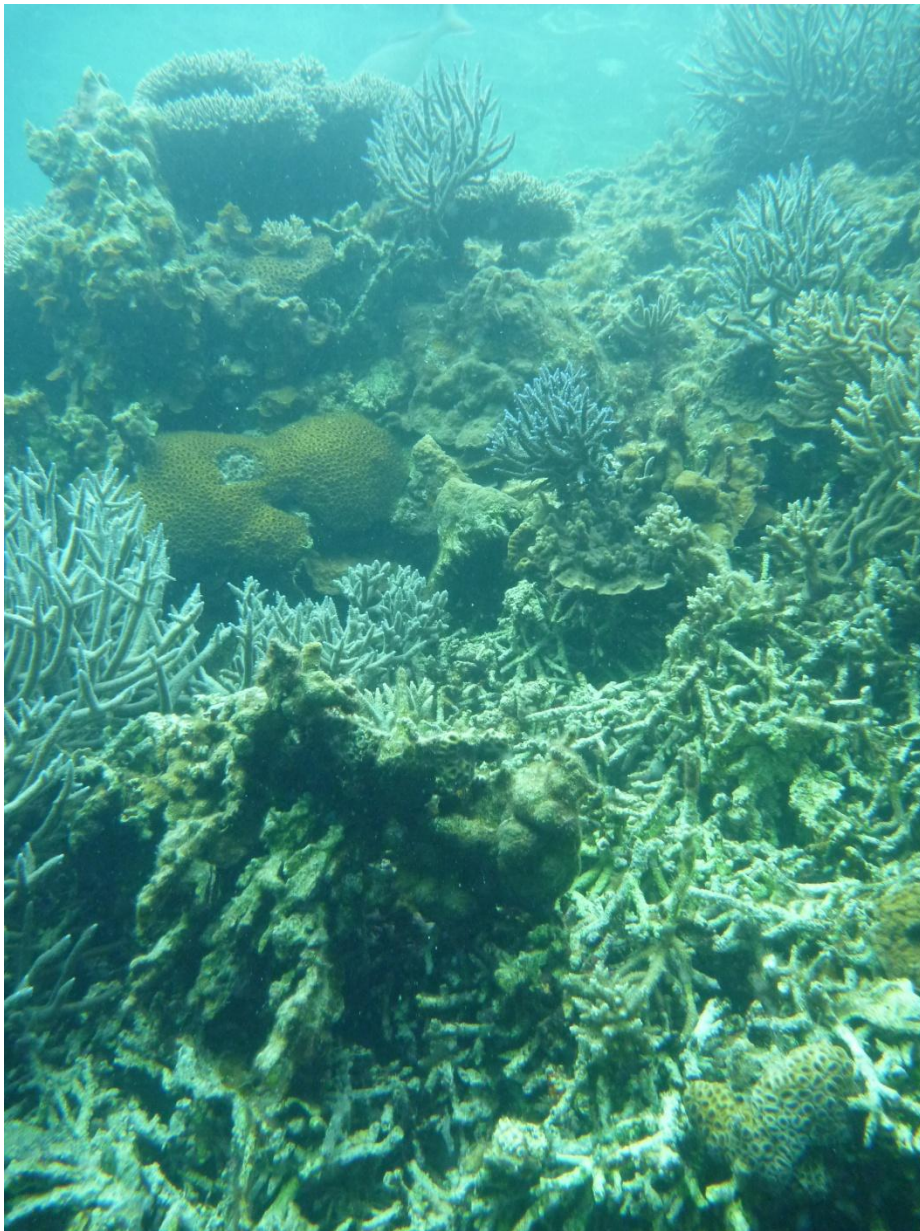


Figure 6A. Inferred early-stage of reef wall development, in which dead *Acropora* branches are colonised by massive and encrusting corals.



Figure 6B. Inferred late-stage of reef wall development, in which massive and encrusting corals have grown to form a vertical wall.

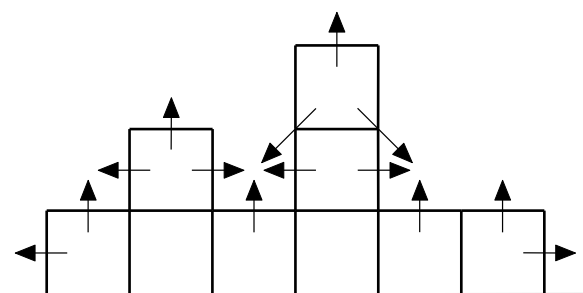


Figure 7. Cross-section through a hypothetical model reef. Upward-pointing arrows indicate vertical growth directions, horizontal and diagonal arrows indicate neighbour growth directions.

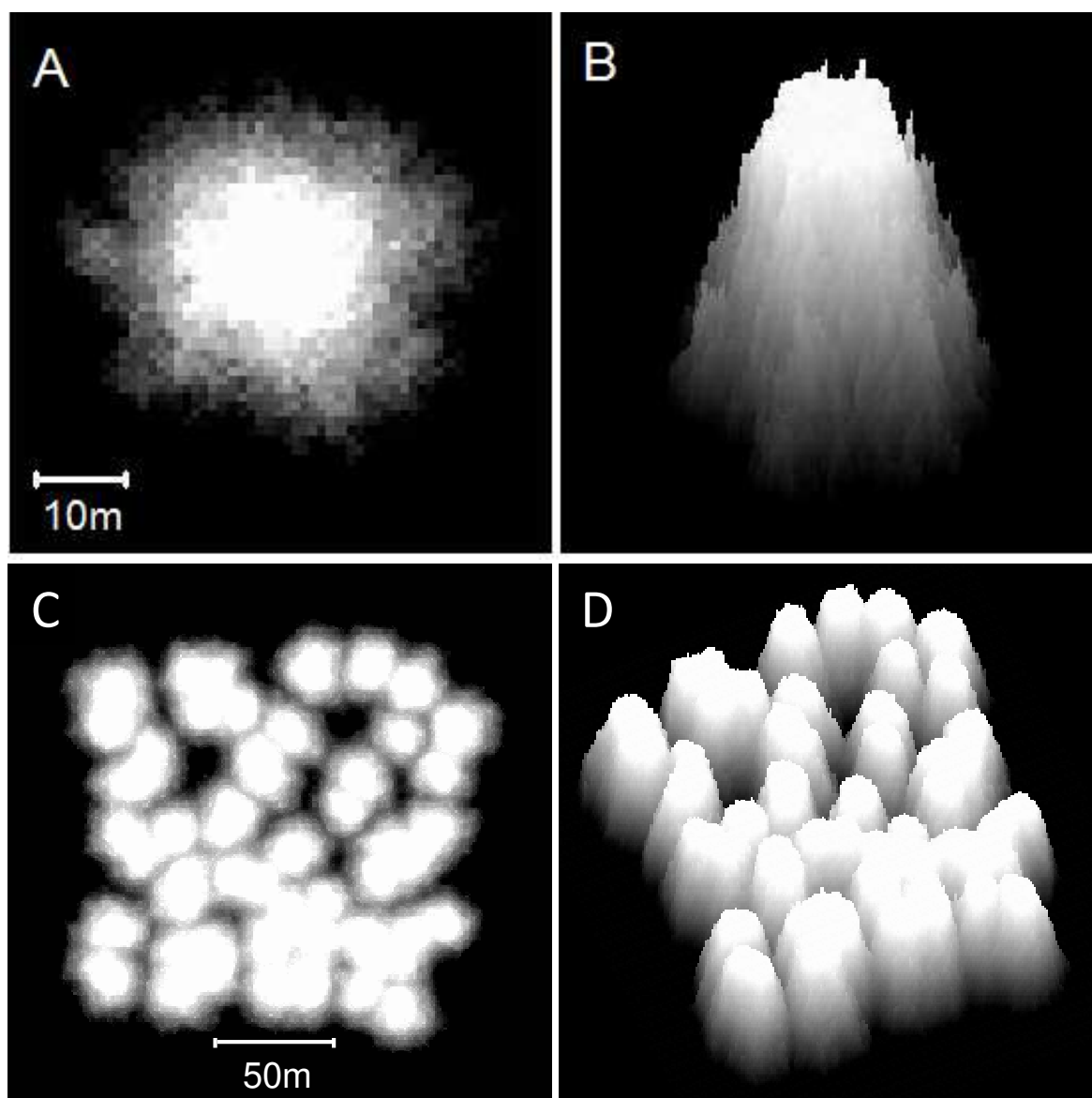


Figure 8. **A)** Two-dimensional plan view of a simulated patch reef after 80 iterations of growth from a single seed coral using the random model. Shading corresponds to depth—the reef top at sea level is white and the surrounding seafloor at 30m depth is black. This patch reef reached sea level in approximately 45 iterations, and by 80 iterations has developed a 15m wide reef flat. **B)** Three-dimensional oblique view of the patch reef in A. **C)** Two-dimensional plan view of multiple patch reefs after 80 iterations of growth from 64 seed corals, occupying 0.25% of the 160 x 160m seed array. Only the uppermost 10m of the reefs is shown, simulating an aerial view with 10m water visibility. **D)** Three-dimensional oblique view of the reefs in C.

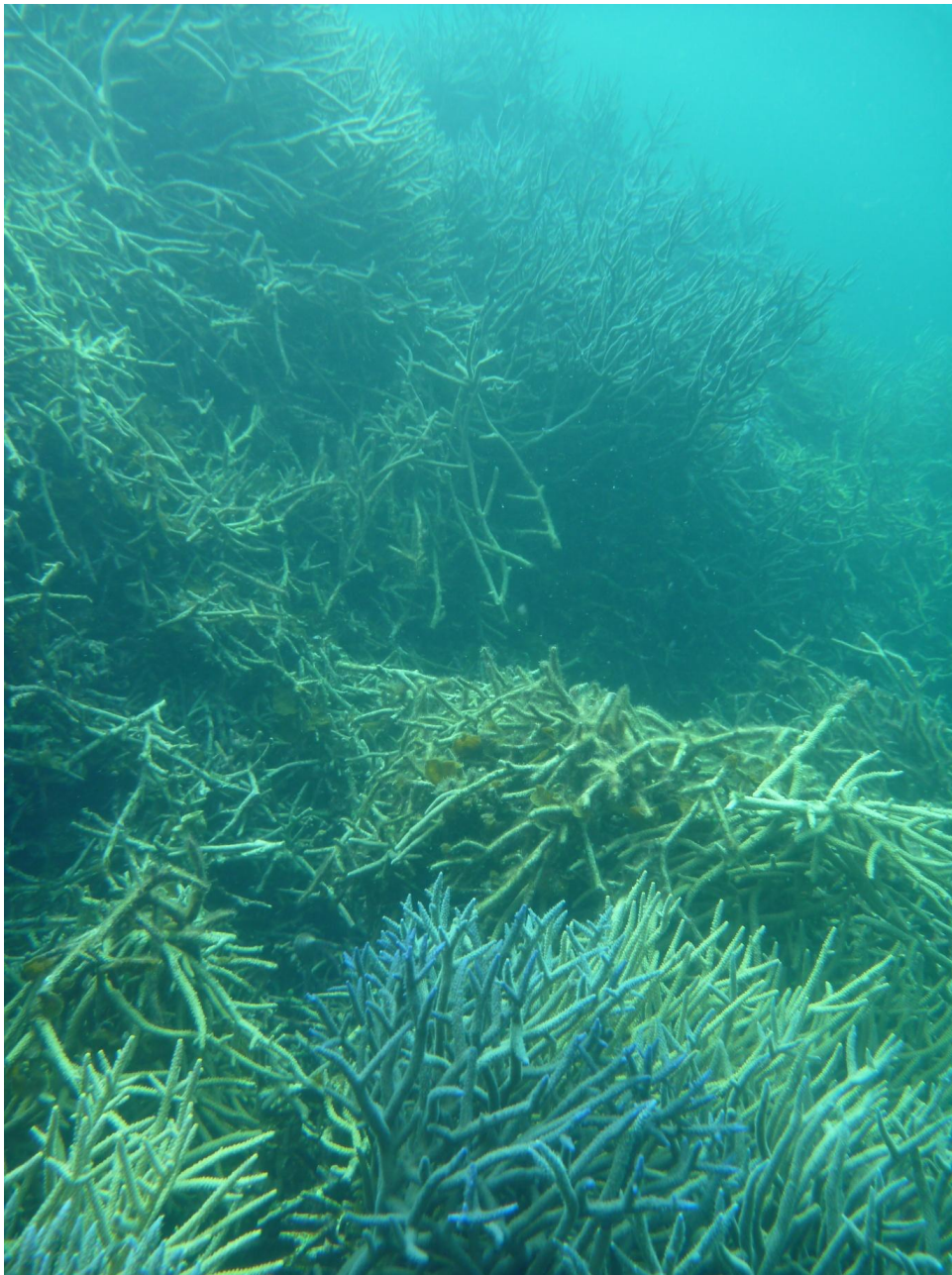


Figure 9. Collapsed branching *Acropora* colonies on the slope of the stellate reef shown in Fig. 1.

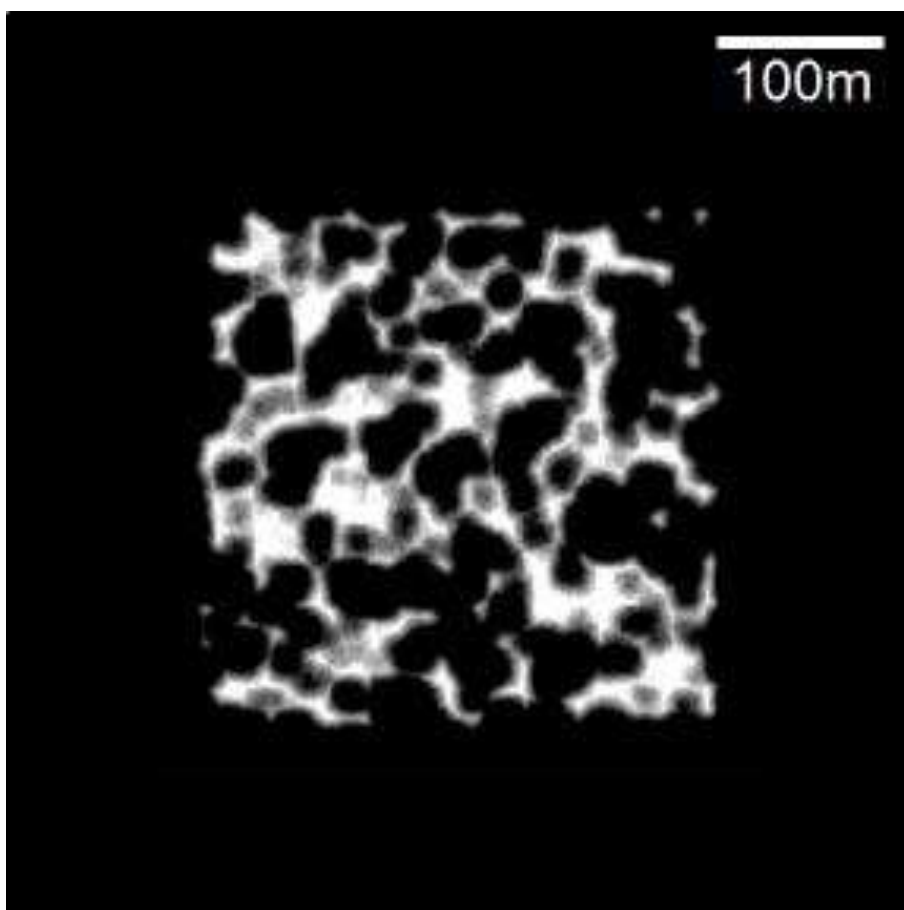


Figure 10. Plan view of a branching coral model reef after 100 growth iterations. Only the uppermost 10m of the reef is shown, simulating an aerial view with 10m water visibility. This reef developed from 225 randomly spaced seed corals, occupying 0.25% of a 300 x 300m seed array.

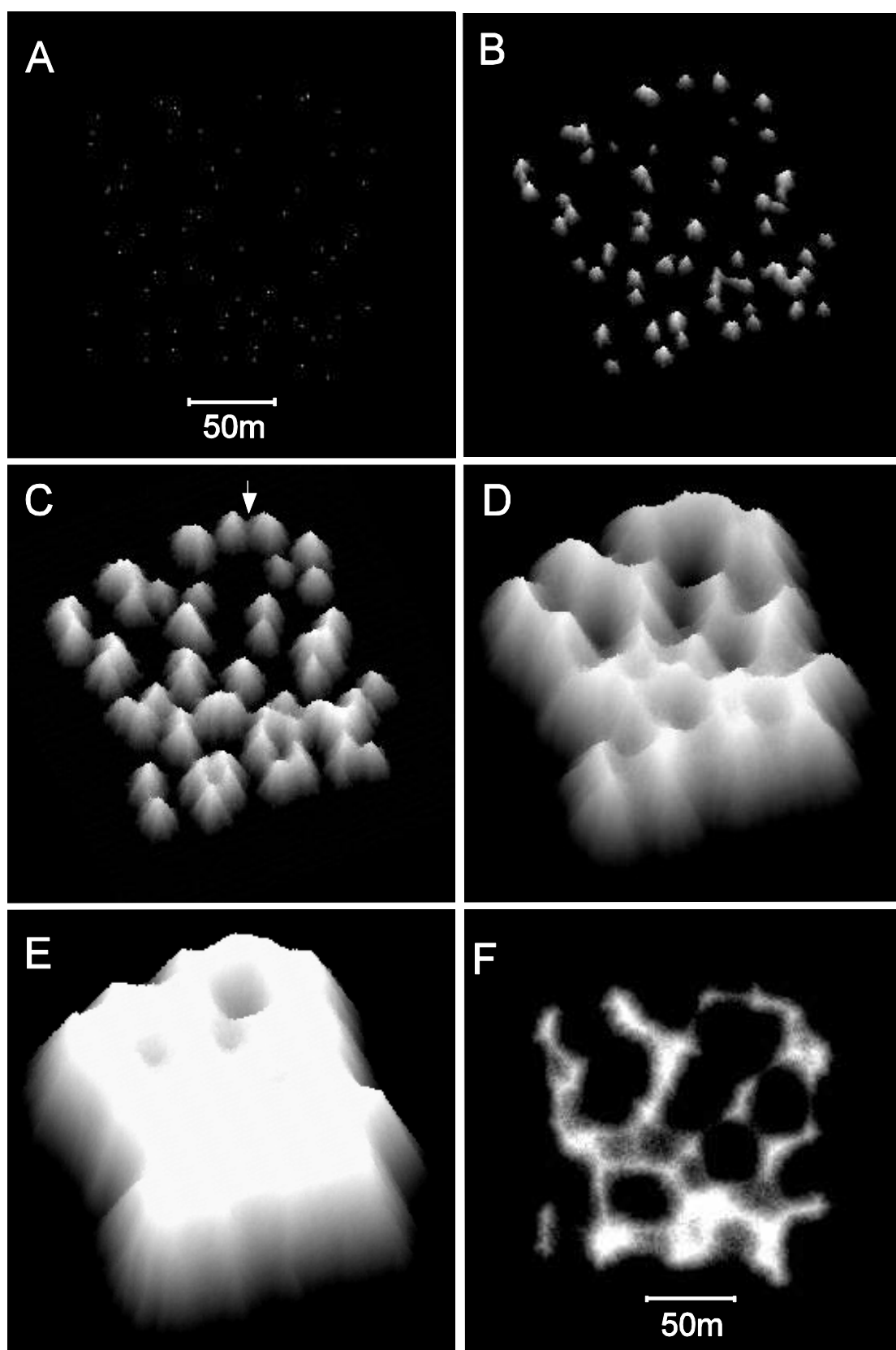


Figure 11. Sequential stages in the development of model branching coral reefs. **A)** The set of 64 randomly spaced seed corals from which the model reefs developed. The seed coral configuration is the same as that used to create the model reefs in Fig. 8C and D. **B)** After 25 iterations the seed corals have developed into cone-shaped patch reefs with 45°

slopes. **C)** After 50 iterations the patch reefs have enlarged and many have merged. When patch reefs meet, the valleys between them grow upward rapidly to become saddle-shaped ridges (arrowed). **D)** After 100 iterations some of the reef tops have reached sea level and the system of ridges has developed to enclose and isolate depressions, producing egg-carton morphology. **E)** After 150 iterations an extensive sea level platform has developed, and most of the depressions have filled. **F)** Plan view showing the uppermost 10m of the reef in D, simulating an aerial view with 10m water visibility.

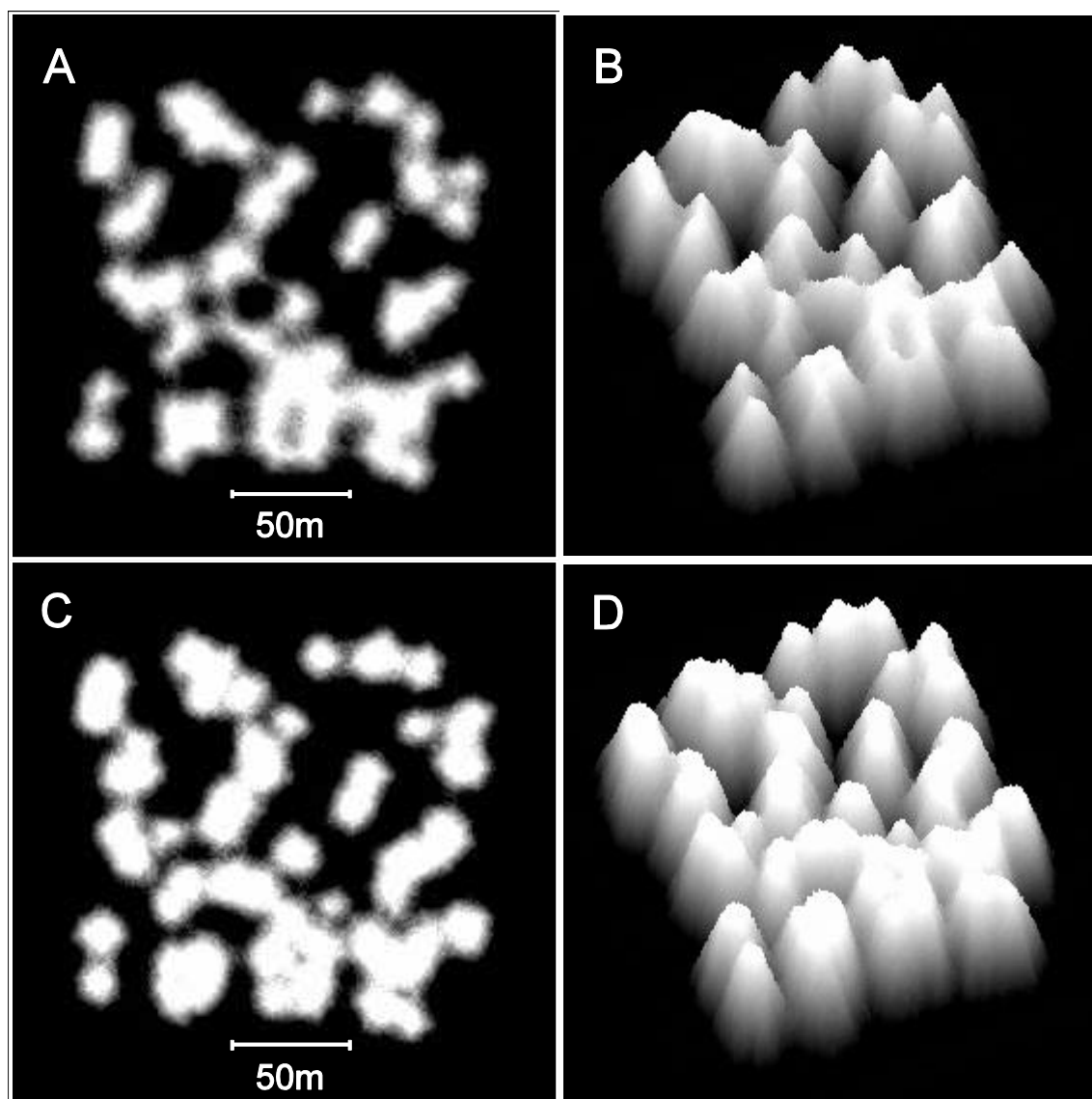


Figure 12. **A)** Plan view, 0-10m depth, of model reef with a 3m collapse limit after 90 growth iterations. **B)** Oblique view of the model reef in A. The reef slopes are approximately 55°. **C)** Plan view, 0-10m depth, of model reef with a 4m collapse limit after 85 growth iterations. **D)** Oblique view of the model reef in C. The reef slopes are approximately 60°.

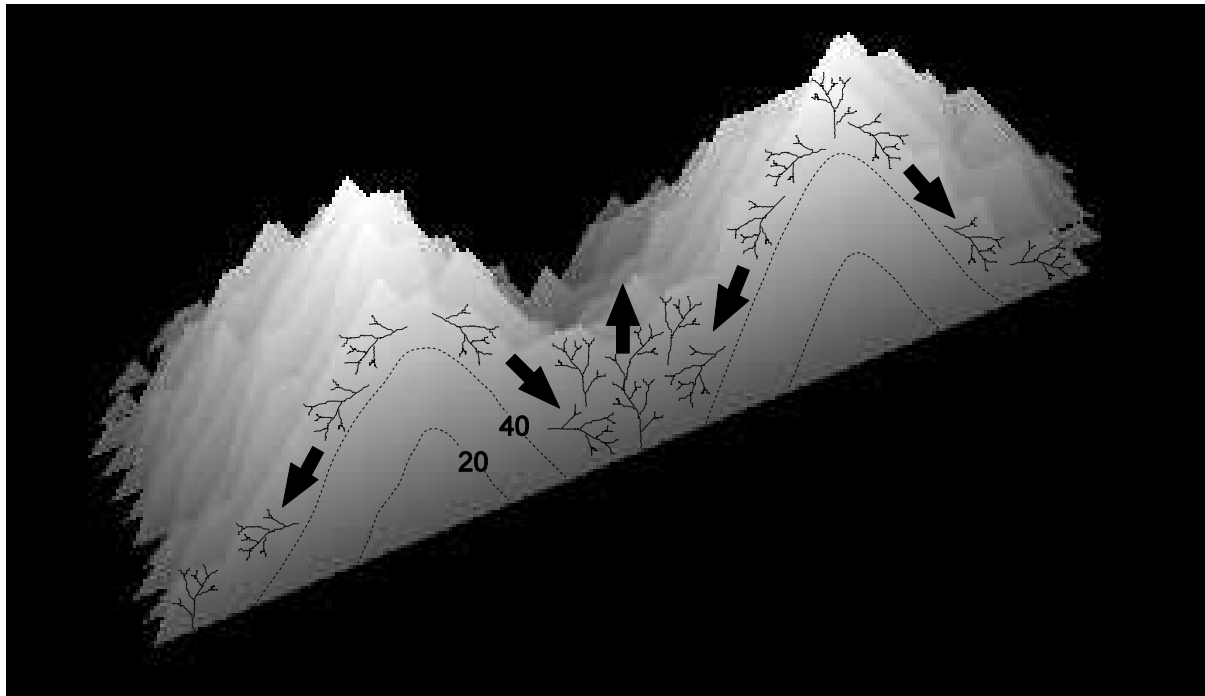


Figure 13. Schematic diagram illustrating the proposed mechanism of ridge formation derived from the branching coral model. The diagram shows a cross-section through two patch reefs after 50 iterations (the patch reefs at the top of Fig. 11C). Isochrons at 20 and 40 iterations show that the patch reefs were initially conical, and that the valley between them has accreted rapidly since the patch reefs merged.



Figure 14. Coalescing nodular patch reefs exposed on a low spring tide at Cockatoo Island, Western Australia. These are ‘true’ nodular reefs, consisting of massive and encrusting corals and having a steep-walled columnar form (D. Blakeway, pers. obs.). Photograph by John MacFadyen.

References

- Aronson, RB, Precht, EF, Macintyre, IG. 1998. Extrinsic control of species replacement on a Holocene reef in Belize: the role of disease. *Coral Reefs* 17: 223-230.
- Bak, RPM. 1976. The growth of coral colonies and the importance of crustose coralline algae and burrowing sponges in relation with carbonate accumulation. *Netherlands Journal of Sea Research* 10(3): 285-337.
- Barott KL, Caselle JE, Dinsdale EA, Friedlander AM, Maragos JE, Obura, D, Rohwer, FL, Sandin, SA, Smith, JE, Zgliczynski, B. 2010. The Lagoon at Caroline/Millennium Atoll, Republic of Kiribati: Natural History of a Nearly Pristine Ecosystem. *PLoS ONE* 5(6): 1-10.
- Blakeway, DR. 2000. Geomorphology and ecology of cellular reefs in the Houtman Abrolhos Islands, Western Australia. D. Phil. Thesis, University of Western Australia, Perth.
- Blanchon, P. 2011. Geomorphic zonation. In Hopley, D, ed. *Encyclopedia of Modern Coral Reefs: Structure, form and process*. Dordrecht: Springer, 469-486.
- Bothwell, AM. (1982) Fragmentation, a means of asexual reproduction and dispersal in the coral genus *Acropora* (Scleractinia: Astrocoeniida: Acroporidae) – A preliminary report. *Proceedings of the 4th International Coral Reef Symposium* Volume 2: 137-144.
- Bruckner, A. (2011) Khaled bin Sultan Living Oceans Foundation Habitat Mapping and Characterization of Coral Reefs of the Saudi Arabian Red Sea: 2006-2009. *Final Report Part II, Ras Qisbah, Al Wajh, Yanbu, Farasan Banks and Farasan Islands*. Arizona: Panoramic Press.
- Camazine, S, Deneubourg, J-L, Franks, NR, Sneyd, J, Theraulaz, G, Bonabeau E. (2001). *Self-organization in biological systems*. Princeton: Princeton University.
- Choi, DR, Ginsburg, RN 1982. Siliciclastic foundations of Quaternary reefs in the southernmost Belize lagoon, British Honduras. *Geological Society of America Bulletin* 93: 116-126.
- Collins, LB, Zhu, ZR, Wyrwoll, K-H, Hatcher, BG, Playford, PE, Chen JH, Eisenhauer, A, Wasserburg, GJ, 1993. Late Quaternary facies characteristics and growth history of a high latitude reef complex: the Abrolhos carbonate platforms, eastern Indian Ocean. *Marine Geology* 110: 203-212.
- Collins, LB, Zhu, ZR, Wyrwoll, K-H, 1996. The structure of the Easter Platform, Houtman Abrolhos reefs: Pleistocene foundations and Holocene reef growth. *Marine Geology* 135: 1-13.
- Collins, LB, Zhu, ZR, Wyrwoll, K-H. 1998. Late Tertiary-Quaternary Geological Evolution of the Houtman Abrolhos Carbonate Platforms, Northern Perth Basin. In Purcell, R, Purcell, P. eds. *The sedimentary basins of Western Australia*, 2. Perth, Western Australia: Petroleum Exploration Society of Australia, 647-663.

- Collins, LB, Zhao, J, Freeman, H. 2006. A high-precision record of mid-late Holocene sea-level events from emergent coral pavements in the Houtman Abrolhos Islands, southwest Australia. *Quaternary International* 145-146: 78-85.
- Crame, JA, 1981. Ecological stratification in the Pleistocene coral reefs of the Kenya coast. *Palaeontology* 24(3): 609-646.
- Dakin, WJ, 1919. The Percy Sladen Trust Expedition to the Abrolhos Islands (Indian Ocean). *Journal of the Linnean Society of London (Zoology)* 34: 127-180.
- Darwin, CR, 1842. *The Structure and Distribution of Coral Reefs. Being the First Part of the Geology of the Voyage of the 'Beagle.'* London: Smith, Elder and Co.
- Davis, WM, 1899. The geographical cycle. *Geographical Journal* 14: 481-504.
- Done, TJ, DeVantier, LM, Turak, E, Fisk, DA, Wakeford, M, and van Woesik, R (2010) Coral growth on three reefs: development of recovery benchmarks using a space for time approach. *Coral Reefs* 29 (4): 815-833.
- Fairbridge, RW. 1948. Notes on the geomorphology of the Pelsaert Group of the Houtman's Abrolhos Islands. *Journal of the Royal Society of Western Australia* 33: 1-43.
- GIE Raro Moana, 1985. The Phosphates from Mataiva. *Proceedings of the Fifth International Coral Reef Symposium* Volume 1: 317-319.
- Gilmore, MD, Hall, BR. 1976. Life History, Growth Habits, And Constructional Roles of *Acropora Cervicornis* in the Patch Reef Environment. *Journal of Sedimentary Petrology* 46(3): 519-522.
- Goodman, JA, Purkis, SJ, Phinn, SR (eds.). 2013. *Coral reef remote sensing—A guide for mapping, monitoring and management.* New York: Springer.
- Grovhoug, JG, Henderson, RS. (1976) Distribution of Inshore Fishes at Canton Atoll. in SV Smith and RS Henderson, eds. *An Environmental Survey of Canton Atoll Lagoon, 1973*, Naval Undersea Center Technical Report NUC TP 395, San Diego, California: 98-157.
- Guilcher, A. 1988. *Coral reef geomorphology.* Chichester: J. Wiley and Sons.
- Hatcher, BG. 1997. Coral reef ecosystems: how much greater is the whole than the sum of the parts? *Coral Reefs* 16(1): S77-S91.
- Halley, R, Shinn, EA, Hudson, JH, Lidz, B. 1977 Recent and relict topography of BooBee patch reef, Belize: *Proceedings of the Third International Coral Reef Symposium* Volume 1, p. 53-96.
- Highsmith, R.C. 1982. Reproduction by fragmentation in corals. *Marine Ecology Progress Series* 7:207-226.

- Hopley, D. 1982. *The geomorphology of the Great Barrier Reef: Quaternary development of coral reefs*. New York: Wiley-Interscience.
- Hopley, D, Smithers, SG, Parnell, KE. 2007. *The Geomorphology of the Great Barrier Reef: development, diversity and change*. Cambridge: Cambridge University Press.
- Hoskin, CM. 1963. Recent carbonate sedimentation on Alacran Reef, Yucatan, México. *National Academy of Science National Research Council, Foreign Field Research Program, Publication 19*.
- Jones, JA. 1977. Morphology and development of southeastern Florida patch reefs. *Proceedings of Third International Coral Reef Symposium* Volume 2: 231-235.
- Jones, CG, Lawton, JH, Shachak, M. 1996. Organisms as ecosystem engineers. *Oikos* 69: 373-386.
- Loya, Y, Sakai, K, Yamazato, K, Nakano, Y, Sambali, H, van Woesik, R. (2001). Coral bleaching: the winners and the losers. *Ecology Letters* 4(2): 122–131.
- Kan, H, Urata, K, Nagao, M, Hori, N, Fujita, K, Yokoyama, Y. 2014. Submerged karst landforms observed by multibeam bathymetric survey in Nagura Bay, Ishigaki Island, southwestern Japan. *Geomorphology*, 13pp.
- Macintyre, IG, Precht, WF, Aronson, RB. 2000. Origin of the Pelican Cays Ponds, Belize. *Atoll Research Bulletin* 466: 1–12.
- McNeill, FS. 1954. The shape of atolls: an inheritance from subaerial erosion forms. *American Journal of Science* 252(7): 402-427.
- Mazzullo, SJ, Anderson-Underwood, KE, Burke, CD, Bischoff, WD. 1992. Holocene Coral Patch Reef Ecology and Sedimentary Architecture, Northern Belize, Central America. *Palaos* 7(6): 591-601.
- Mistr, S, Bercovici, D. 2003. A theoretical model of pattern formation in coral reefs. *Ecosystems* 6: 61–74.
- Montagne, A, Naim, O, Tourrand, C, Pierson, B, Menier, D. 2013. Status of Coral Reef Communities on Two Carbonate Platforms (Tun Sakaran Marine Park, East Sabah, Malaysia). *Journal of Ecosystems*, Article ID 358183, 15 p.
- Morton, J. 1974. The coral reefs of the British Solomon Islands: a comparative study of their composition and ecology. *Proceedings of the Second International Coral Reef Symposium* Volume 2: 31-53.
- Munk, WH, Sargent, MC. 1954. Adjustment of Bikini Atoll to ocean waves, *U.S. Geological Survey Professional Paper* 260-C: 275–280.
- Nicolis, G, Prigogine, I. 1977. *Self Organization in Non-Equilibrium Systems*. New York: J. Wiley and Sons.

- Novak, V, Santodomingo, N, Rosler, A, Di Marino, E, Braga, JC, Taylor, PD, Johnson, KG, Renema, W. 2013. Environmental reconstruction of a late Burdigalian (Miocene) patch reef in deltaic deposits (East Kalimantan, Indonesia). *Palaeogeography, Palaeoclimatology, Palaeoecology* 374:110–122.
- Perry, CT, Smithers, SG, Gulliver, P, Browne, NK. 2012. Evidence of very rapid reef accretion and reef growth under high turbidity and terrigenous sedimentation. *Geology* 40(8): 719–722.
- Purdy, EG. 1974. Reef configuration: cause and effect. In LaPorte, LF, ed. *Reefs in Time and Space*. Society of Economic Palaeontologists and Mineralogists Special Publications 18: 9–76.
- Purkis, SJ, Rowlands, GP, Riegl, BM, Renaud, PG. 2010. The paradox of tropical karst morphology in the coral reefs of the arid Middle East, *Geology*, 38: 227–230.
- Rietkerk, M, van de Koppel, J. 2008. Regular pattern formation in real ecosystems. *Trends in Ecology and Evolution* 23: 169–175.
- Roberts, HH. 1974. Variability of reefs with regard to changes of wave power around an island. *Proceedings of the Second International Coral Reef Symposium*, Volume 2: 497–512.
- Roberts, HH, Murray, SP, Suhayda, JN. 1975. Physical processes in a fringing reef system. *Journal of Marine Research*, 33: 233–260.
- Rossfelder, A.M. 1990. The submerged phosphate deposit of Mataiva Atoll, French Polynesia. In Burnett, WC and Riggs, SR, eds. *Phosphate deposits of the world. Volume 3. Neogene to Modern Phosphorites*. Cambridge: Cambridge University Press, 195–203.
- Scheffers, SR, Scheffers, A, Kelletat, D, Bryant, EA. 2008. The Holocene paleo-tsunami history of Western Australia. *Earth and Planetary Science Letters* 270 (1–2): 137–146.
- Scheffers, AM, Scheffers, SR, Kelletat, DH, Squire, P, Collins, L, Feng, Y, Zhao, J-X, Joannes-Boyau, R, May, SM, Schellmann, G and Freeman, H. 2012. Coarse clast ridge sequences as suitable archives for past storm events? Case study on the Houtman Abrolhos, Western Australia. *Journal of Quaternary Science* 27(7): 713–724.
- Searle, DE. 1994. Late Quaternary morphology of the Cocos (Keeling) Islands. *Atoll Research Bulletin* 401: 1–13.
- Stoddart, DR. 1969. Ecology and morphology of recent coral reefs. *Biological Reviews* 44(4): 433–498.
- Tunncliffe, V. 1981. Breakage and propagation of the stony coral *Acropora cervicornis*. *Proceedings of the National Academy of Sciences USA* 78(4): 2427–2431
- Ulam, S. 1962. On Some Mathematical Problems Connected with Patterns of Growth and Figures. *American Mathematical Society Proceedings of Symposia in Applied Mathematics* 14: 215–224

- Woodroffe CD, Kennedy DM, Jones BG, Phipps CVG. 2004. Geomorphology and Late Quaternary development of Middleton and Elizabeth Reefs. *Coral Reefs* 23: 249-262
- Hatcher, AI, Hatcher, BG, Wright, GD. 1988. *A preliminary report on the interaction between the major human activities and the marine environment of the Houtman Abrolhos Islands of Western Australia*. Hatcher Research Associates, Perth.
- Maragos, JE. 1972. A study of the ecology of Hawaiian corals. D. Phil. Thesis, University of Hawaii, Honolulu.
- Roth, MS, Deheyn, DD. 2013. Effects of cold stress and heat stress on coral fluorescence in reef-building corals. *Scientific Reports* 3: article 1421.
- Roy KJ, Smith SV. 1971. Sedimentation and coral reef development in turbid water: Fanning Lagoon. *Pacific Science* 25(2): 234-248.
- Turak, E, DeVantier, L. 2005. Reef-building corals and coral communities of Pohnpei, Federated States of Micronesia: Rapid ecological assessment of biodiversity and status. *Final report for the Conservation Society of Pohnpei, September 2005*, 103 p.
- Ulam, S. 1962. On Some Mathematical Problems Connected with Patterns of Growth and Figures. *American Mathematical Society Proceedings of Symposia in Applied Mathematics* 14: 215-224
- Valencia, MJ. 1977. Christmas Island (Pacific Ocean): Reconnaissance geologic observations. *Atoll Research Bulletin* 197: 1-17.
- von Neumann, J. 1963. *Collected Works of John von Neumann*, Taub, A.H, ed., Pergamon Press.
- Wilson BR, Marsh LM. 1979. Coral reef communities at the Houtman Abrolhos, Western Australia, in a zone of bio-geographic overlap. In: *Proceedings of the International Symposium on Marine Biogeography and Evolution in the Southern Hemisphere*. Auckland, New Zealand. Volume 1: 259–278.
- Williams, DG. 1994. Marine habitats of the Cocos (Keeling) Islands. *Atoll Research Bulletin* 406: 1-10.
- Wyrwoll, K-H, Zhu, ZR, Collins, LB, and Hatcher, BG. 2006. Origin of blue hole structures in coral reefs: Houtman Abrolhos, Western Australia: *Journal of Coastal Research* 22: 202–208.