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Nitrogen addition increases the contents of glomalin-related soil protein and soil organic carbon but retains aggregate stability in a *Pinus tabulaeformis* forest

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Background: Glomalin-related soil protein (GRSP) and soil organic carbon (SOC) contribute to the formation and stability of soil aggregates, but the mechanism by which global atmospheric nitrogen (N) deposition changes soil aggregate stability when it alters the distribution of GRSP and SOC in different aggregate fractions remains unknown.

Methods: We used a gradient N addition (0–9 g N⁻² y⁻¹) in *Pinus tabulaeformis* forest for 2 years in northeast China and then examined the changes in SOC contents, total GRSP (T-GRSP), and easily extractable GRSP (EE-GRSP) contents in three soil aggregate fractions (macro-aggregate: >250 μm, micro-aggregate: 250–53 μm, and clay-silt aggregate: <53 μm) and their relationship with aggregate stability. **Results:** (1) The soil was dominated by macro-aggregates. Short term N addition had no significant effect on mean weight diameter (MWD) and geometric mean diameter (GMD). (2) GRSP varied among aggregate fractions, and N addition had variable effects on the distribution of GRSP in aggregate fractions. The EE-GRSP content in the macro-aggregates increased initially and then decreased with increasing N addition levels, having a peak value of 0.480 mg/g at 6 g N⁻² y⁻¹. The micro-aggregates had the lowest EE-GRSP content (0.148 mg/g) at 6 g N⁻² y⁻¹. Furthermore, the T-GRSP content significantly increased in the aggregate fractions with the N addition levels. (3) The macro-aggregate had the highest SOC content, followed by the micro-aggregate and the clay-silt aggregate had the lowest SOC content. N addition significantly increased the SOC content in all the aggregate fractions. (4) GRSP and SOC contents were not significantly correlated with MWD. **Conclusion:** The distributions of GRSP and SOC varied with aggregate fractions. GRSP and SOC contents increased by N addition, but this increase did not enhance aggregate stability in short term, and the improvement of stability might depend on binding agents and incubation time.

Author Cover Page**Nitrogen addition increases the contents of glomalin-related soil protein and soil organic carbon but retains aggregate stability in a *Pinus tabulaeformis* forest****Lipeng Sun¹, Guoliang Wang¹, Hang Jing¹, Guobin Liu^{1*}**

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2 **carbon but retains aggregate stability in a *Pinus tabulaeformis* forest**

3
4 **Abstract**

5 **Background:** Glomalin-related soil protein (GRSP) and soil organic carbon (SOC)
6 contribute to the formation and stability of soil aggregates, but the mechanism by which global
7 atmospheric nitrogen (N) deposition changes soil aggregate stability when it alters the
8 distribution of GRSP and SOC in different aggregate fractions remains unknown. **Methods:** We
9 used a gradient N addition (0–9 g N⁻² y⁻¹) in *Pinus tabulaeformis* forest for 2 years in northeast
10 China and then examined the changes in SOC contents, total GRSP (T-GRSP), and easily
11 extractable GRSP (EE-GRSP) contents in three soil aggregate fractions (macro-aggregate: >250
12 μm, micro-aggregate: 250–53 μm, and clay–silt aggregate: <53 μm) and their relationship with
13 aggregate stability. **Results:** (1) The soil was dominated by macro-aggregates. Short term N
14 addition had no significant effect on mean weight diameter (MWD) and geometric mean
15 diameter (GMD). (2) GRSP varied among aggregate fractions, and N addition had variable
16 effects on the distribution of GRSP in aggregate fractions. The EE-GRSP content in the macro-
17 aggregates increased initially and then decreased with increasing N addition levels, having a
18 peak value of 0.480 mg/g at 6 g N⁻² y⁻¹. The micro-aggregates had the lowest EE-GRSP content
19 (0.148 mg/g) at 6 g N⁻² y⁻¹. Furthermore, the T-GRSP content significantly increased in the
20 aggregate fractions with the N addition levels. (3) The macro-aggregate had the highest SOC
21 content, followed by the micro-aggregate and the clay–silt aggregate had the lowest SOC content.
22 N addition significantly increased the SOC content in all the aggregate fractions. (4) GRSP and
23 SOC contents were not significantly correlated with MWD. **Conclusion:** The distributions of
24 GRSP and SOC varied with aggregate fractions. GRSP and SOC contents increased by N
25 addition, but this increase did not enhance aggregate stability in short term, and the improvement
26 of stability might depend on binding agents and incubation time.

27 **Keywords:** Aggregate; Glomalin-related soil protein; Soil organic carbon; Nitrogen deposition;

28 Mean weight diameter

29 **Introduction**

30 Human activities have led to a rapid increase in global nitrogen (N) deposition, which is
31 expected to increase by 70% in 2050 than in 2000 (Galloway et al. 2004). the increasing N
32 deposition has some negative effects on soil quality. for example, N deposition can affect soil
33 aggregate stability by changing plant growth rates(Bai et al. 2010), soil microbial activities
34 (Keeler, Hobbie and Kellogg 2009), and soil chemical properties in forest communities (Yin et al.
35 2016). However, the effects of available N in soil on aggregate stability remains controversial
36 (Liu et al. 2015, Tripathi, Kushwaha and Singh 2008, Zhong et al. 2017), especially with respect
37 to the mechanisms by which available N in soil influences aggregate stability when it changes
38 glomalin-related soil protein (GRSP) and soil organic carbon (SOC) in aggregate fractions.

39 GRSP can bind soil particles to form aggregates, and the stability of these aggregate may
40 increase with increasing GRSP content (Wright and Upadhyaya 1998). Several studies showed
41 that soil available N has variable effects on GRSP content in bulk soil. For instance, some
42 researchers found that N addition significantly increased GRSP content in bulk soil (Zhang et al.
43 2015, Garcia et al. 2008), but Wuest found that N addition had no significant effects on GRSP
44 content in bulk soil (Wuest et al. 2005). These inconsistent results may be related to the soil
45 initial N content (Kathleenk, Katiem and Michellec 2007), as well as N addition level and
46 duration (Garcia et al. 2008, Treseder 2004). For example, N addition could increase the growth
47 of arbuscular mycorrhizal fungi (AMF) in the soil with low initial N content (Treseder and Allen
48 2002), thus it increases GRSP content in bulk soil because glomalin is mainly produced by AMF
49 (Kathleenk et al. 2007). Meanwhile, N addition may inhibit the growth of AMF in soil with high
50 initial N content and may thus decrease GRSP content in bulk soil (Treseder 2004). However, a
51 recent study showed that N addition reduced plant fine root length and AMF infection rate in soil
52 with low N content (Wang et al. 2017), indicating that N addition can decrease AMF and GRSP
53 content even in such soil (Kathleenk et al. 2007). Therefore, exploring the relationship between
54 GRSP and available N content in soil with low N content is necessary.

55 Soil is mainly composed of aggregates of various sizes. As far as we know, the mechanism

56 underlying the influence of N addition on GRSP content in aggregates of variable size has not
57 been reported yet. In some studies on farmland ecosystems, manure amendment was found to
58 promote soil nutrient accumulation and microbial activities, considerably increasing GRSP
59 contents in all soil aggregate fractions except micro-aggregates (Xie et al. 2015). In these
60 farmland ecosystems, tillage management methods disrupts soil structure and reduces microbial
61 biomass and extent of enzyme activity, resulting in the significant decrease of GRSP content in
62 all soil aggregate fractions except clay–silt aggregates and small fractions of macro-aggregates
63 (Wright, Green and Cavigelli 2007). Similar to manure amendment and tillage management, N
64 addition may also alter soil nutrient accumulation and microbial activity (Treseder and Allen
65 2002, Treseder 2004). Thus, N addition may change the GRSP content in soil aggregates. We
66 speculated that the effects of N addition on GRSP content in aggregate fractions at forest system
67 might be similar to that of farmland ecosystems with manure amendment and tillage
68 management, as both systems alter soil nutrient accumulation and microbial activity.

69 GRSP can bind soil particles to form aggregates. Thus, changing of GRSP in aggregate
70 fractions can influence aggregate stability directly (Wright and Upadhyaya 1998). A large
71 number of studies have shown that the correlation between GRSP and aggregate stability is
72 significant and positive (Wright et al. 2007, Zhang et al. 2014). However, a recent study on the
73 relationships between GRSP and aggregate stability in citrus rhizosphere found that EE-GRSP in
74 0.25–0.5 mm aggregates was significantly positively correlated with mean weight diameter
75 (MWD), whereas the T-GRSP in 2.0–4.0 mm aggregates and EE-GRSP in 0.5–1.0 mm
76 aggregates were significantly negatively correlated with MWD. However, the GRSP in other size
77 aggregates (1.0–2.0 mm) was insignificantly correlated with MWD (Wu et al. 2013). These
78 results indicated that the relationships between GRSP and aggregate stability may vary with the
79 aggregate sizes. Thus, we hypothesize that low-level N addition may increase the GRSP content
80 in all soil aggregate fractions except micro-aggregates, and high-level N addition may decrease
81 the GRSP content in all soil aggregate fractions except clay–silt aggregates. The relationships
82 between GRSP and aggregate stability will be different according to the aggregate sizes.

83 In most previous studies, N addition significantly increased SOC content in bulk soil (Nave
84 et al. 2009), although some studies indicated that N addition considerably decreased or did not
85 affect SOC content (Zak and Sinsabaugh 2004, Zhong et al. 2017). N addition has variable
86 effects on SOC content in aggregate fractions and bulk soil (KH et al. 2014, Fonte et al. 2009).
87 However, most studies were focused on the effects of N addition on SOC content in bulk soil,
88 and changes in SOC content in aggregate fractions are rarely investigated (Pan et al. 2007). In
89 agriculture ecosystems, N fertilization has a minimal effect on SOC content in aggregate
90 fractions under a conventional maize cropping system (KH et al. 2014). In maize cropping
91 systems utilizing organic residues, N addition significantly increases SOC content in clay–silt
92 aggregates and has no significant effect on the SOC content in other size aggregates (Fonte et al.
93 2009). These inconsistent results may be explained by the following reasons: first, the physical
94 protection on SOC were different in aggregate fractions, thus resulting in different accumulation
95 and decomposition rates of SOC in aggregate fractions (BlancoCanqui and Lal 2004). Second,
96 tillage management affects soil carbon (C) input. In a conventional maize cropping system, N
97 addition has no significant effect on SOC content in aggregate fractions, but it has considerable
98 effect on SOC content in clay–silt aggregates in a maize cropping system utilizing organic
99 residues. The increased SOC content in clay–silt aggregates may be due to the input of organic
100 residues (Fonte et al. 2009). At present, studies on the effect of soil available N on SOC in
101 aggregates are mainly focused on farmland ecosystems, and no study on forestland system has
102 been reported yet (Chen et al. 2017). We speculated that the effects of N addition on SOC
103 content in aggregate fractions at forest system might be similar to that of maize cropping systems
104 utilizing organic residues, as both systems return a large amount of plant biomass to soil.
105 However, this speculation requires experimental validation. Recent studies found that the
106 relationship between SOC and aggregate stability is uncertain. For example, SOC content has a
107 significantly negative relationship with aggregate stability in red soil found in subtropical China
108 (Li et al. 2005). Moreover, the relationships between SOC and aggregate stability were
109 insignificant in different land use tropical Ultisol (Leelamanie and Mapa 2015). SOC content in

110 aggregate fractions had different relationships with aggregate stability in rice–wheat rotation
111 farmlands (Das et al. 2014). For example, SOC content in macro-aggregates is significantly
112 positively correlated with aggregate stability, but SOC content in micro- and clay–silt aggregates
113 significantly negatively correlated with aggregate stability. However, research on the
114 relationships between SOC content in aggregate fractions and aggregate stability under N
115 addition condition remains lacking. In the present study, we proposed that N addition
116 significantly increases SOC content in clay–silt aggregates and minimal affect SOC content in
117 other size aggregate fractions. The relationships between SOC content and aggregate stability are
118 varied among aggregate fractions.

119 Two hypotheses were tested by determining GRSP and SOC contents in aggregate fractions
120 and aggregate stability in a *Pinus tabulaeformis* forestland. This study is expected to determine
121 the underlying mechanism of soil aggregate stability in response to N addition.

122 **Materials and methods**

123 **Study area**

124 The experiment was conducted in the Songyugou watershed in the Loess Plateau region,
125 Shaanxi Province, China (35 390N, 110 060E). The region has a classic mainland monsoon-type
126 climate with a mean annual precipitation of 584.4 mm, mean annual temperature of 9.7 °C, and
127 frost-free period of 180 days. The soil is gray forest soil (Gray Luvisols, FAO soil classification)
128 with a surface soil bulk density of 1.1 g/cm³. The soil total phosphorus (P) content is 1.42 ± 0.38
129 g/kg⁻¹, and the soil pH is 8.6. *P. tabulaeformis* was planted in 1960. The current stand density is
130 1400 – 1800/hm². The average canopy density is 0.7. The average diameter at breast height
131 (DBH) is 10.0 cm. The average height of the tree is 11.2 m. The forest stand volume is 75.5
132 m³/hm². The leaf area index is 6.34. The biomass of tree layer, shrub layer, and herbaceous layer
133 are 112.96, 3.56, and 8.28 t/hm², respectively, The plant diversity index of the community is 0.51
134 (calculated using the Simpson method). Shrubs are mainly *Elaeagnus pungens* Thunb, *Rosa*
135 *xanthina* Lindl, *Spiraea Salicifolia* L, *Lonicera japonica* Thunb, and *Viburnum dilatatum* Thunb.
136 The shrub coverage is 30%. Herbs are mainly *Carex lanceolata* Boott, and the herb coverage is
137 30%–50%.

138 Experimental design

139 In China, the N deposition rate in 2010 was ranged from 0.1 g N⁻² y⁻¹ to 7.43 g N⁻² y⁻¹, with
140 an average of 2.11 g N⁻² y⁻¹, and N deposition is increasing at an annual rate of 0.041 g N⁻² y⁻¹
141 (Lü and Tian 2007). Thus, we designed four levels of N addition treatments (0, 3, 6, and 9 g N
142 m⁻² y⁻¹) in the form of urea (Fumin Agriculture Product Company, Xi'an, China). In each
143 treatment, six 10 m × 10 m plots were established for N addition. A total of 24 pots were
144 subjected to the same climate, stand age, and terrain conditions. In 2014, urea was dissolved in
145 10 L distilled water and evenly sprayed in each plot before the rain in early April, June, August,
146 and October every year. Soil samples were collected in September 2015, and N was added for 2
147 years.

148 Soil sampling

149 In September 2015, undisturbed soil was collected from 0–20 cm soil layer because the
150 short-term N addition directly affected the physical and chemical properties of the surface soil.
151 First, three pits were dug in the central rows of each plot, and undisturbed soil was collected in
152 each pit after removing the surface litter. Second, the undisturbed soil samples were carefully
153 transported back to the laboratory. The undisturbed soils were drilled gently along the natural
154 fissures, and all the soil samples were air-dried after filtering through 8 mm sieve. The three soil
155 samples of each pot were mixed up after air-drying. Finally, all soil samples were stored until
156 physical and chemical analysis.

157 Aggregate distribution

158 Soil aggregates were classified by wet sieving method (Six et al. 2010). Air-dried soil
159 sample (100 g) was soaked with deionized water for 5 min under a 250 µm sieve, and the surface
160 of the water submerged the soil samples. Aggregates were separated by moving the sieve up and
161 down for 3 cm with 50 repetitions/min for 2 min. Aggregates retained in the sieve (>250 µm)
162 were macro-aggregates. The soil and water passed through the 250 µm sieve and then were
163 sieved again as above using a 53 µm sieve for 3 min. Aggregates retained in the sieve (>53 µm)
164 were micro-aggregates, and those that passed through the sieve were clay–silt aggregates. All
165 separated soil and water were collected, oven-dried at 45 °C, and weighed and stored for

166 chemical analysis.

167 **GRSP analyses**

168 GRSP content was determined according to the method of Wright and Upadhyaya (1998)
169 (Wright and Upadhyaya 1998). T-GRSP was extracted from soil sample (1 g) with 8 mL of 50
170 mmol/L citrate solution at pH 7.0 and then autoclaved at 121 °C for 60 min. The supernatant was
171 removed after centrifugation at 10,000 × g for 10 min. This procedure was repeated 5 times on
172 the same sample until the solution was straw-colored. The supernatant was pooled together and
173 stored at 4 °C. EE-GRSP was extracted from soil sample (1 g) with 8 mL of 20 mmol/L citrate
174 solution at pH 7.0 and autoclaved at 121 °C for 30 min. The supernatant was removed after
175 centrifugation at 10,000 × g for 10 min and stored at 4 °C. The T-GRSP or EE-GRSP in the
176 supernatant was determined by Bradford analysis spectrophotometrically at 590 nm light (Wright
177 et al. 1996, Rillig 2004) (Wright and Upadhyaya, 1996; Rillig, 2004) .

178 **SOC analyses**

179 SOC content was measured by dichromate oxidation procedure (Nelson and Sommers
180 1982).

181 **Statistical analyses**

182 Aggregate stability was evaluated by MWD mm and GMD mm. MWD is a comprehensive
183 index for the evaluation of aggregate stability, and GMD is an index of the main grain size
184 distribution of aggregates. The larger value of MWD and GMD indicate higher aggregate
185 stability (Bedini et al. 2009). MWD and GMD were calculated using the following equation:

$$186 \quad \text{MWD} = \sum_{i=1}^n x_i w_i \text{ and } \text{GMD} = \text{EXP} \left[\frac{\sum_{i=1}^n w_i \ln x_i}{\sum_{i=1}^n x_i} \right]$$

187 where x_i is the mean diameter of each class aggregate (mm), w_i is the percentage of each class
188 aggregate (%), and n is the number of aggregate fraction classes.

189 Differences between any two N addition treatments in aggregate percentage, MWD, GMD,
190 GRSP content, and SOC content were tested with one-way ANOVA on SPSS 20.0 statistical
191 software package (SPSS, Inc., USA). Linear regression analyses were used to test the
192 relationships among GRSP, SOC, and MWD.

193 Results**194 Effects of N addition on soil aggregate distribution and stability**

195 The percentage of aggregate varied with size classes, and macro-aggregates were more
196 dominant than the other size aggregate fractions (Table. 1). Across all N addition treatments, the
197 values of MWD were in the range of 1.38–1.70 mm, and GMD was in the range of 0.86–1.05
198 mm. N addition had no significant effects on aggregate distribution, MWD, and GMD. However,
199 the percentage of macro-aggregate, MWD, and GMD increased first and then decreased with
200 increasing N addition levels, whereas the percentage of micro-aggregate and clay–silt aggregate
201 decreased first and then increased with increasing N addition levels.

202 Effects of N addition on GRSP content

203 Figure 1 shows that the contents of EE-GRSP and T-GRSP were in the range of 0.099–
204 0.551 mg/g and 0.315–4.058 mg/g, respectively. In the control treatment, GRSP content varied
205 among bulk soil and aggregate fractions, and GRSP content increased in the order of macro-
206 aggregate < bulk soil < micro-aggregate < clay–silt aggregate. Across all N addition treatments, EE-
207 GRSP content in the macro-aggregates and bulk soil increased first and then decreased with the
208 increasing N addition levels and had a top value at 6 g N⁻² y⁻¹. Compared with that of control
209 treatment, N addition changed the EE-GRSP content of the macro-, micro-, clay–silt aggregates
210 and bulk soil by –24.8%–222.1%, –30.8%–22.0%, 22.8%–40%, and –8.4%–134.4%. Across all
211 N addition treatments, T-GRSP contents in all size aggregates and bulk soil increased with the
212 increasing N addition levels and had a top value at 9 g N⁻² y⁻¹. Compared with that of control
213 treatment, N addition changed the T-GRSP content in macro-, micro-, clay–silt aggregates, and
214 bulk soil by 412.7%–872.4%, 203.1%–266.9%, –0.3%–14.3%, and 259.3%–468.0%.

215 Effects of N addition on SOC content

216 SOC content ranged from 6.478 mg/g to 21.627 mg/g. Across all aggregate fractions, SOC
217 increased in the following order: clay–silt aggregate < micro-aggregate < macro-aggregate (Fig.
218 2). Compared with that of control treatment, N addition significantly increased SOC content in
219 all size aggregate fractions and bulk soil, and the change ranges of SOC content in the macro-,

220 micro-, clay-silt aggregates, and bulk soil were 71.7%–79.5%, 75.5%–91.5%, 49.6%–53.5%,
221 and 47.6%–74.4%.

222 **Relationships among GRSP, SOC, and MWD**

223 Table 2 shows that the contents of SOC, EE-GRSP, and T-GRSP in aggregate fractions and
224 bulk soil have no significant relationship with MWD (except for the relationship between SOC
225 content in micro-aggregates and MWD).

226 **Discussion**

227 In this study, the soil of *P. tabulaeformis* forest was dominated by macro-aggregates, and
228 the aggregate stability indexes of MWD and GMD were in the range of 1.38–1.70 mm and 0.86–
229 1.05 mm, respectively. N addition for 2 years had no significant effects on the soil aggregate
230 stability indexes, which was consistent with those of the previous studies (Wang et al. 2013,
231 Zhong et al. 2017). For example, Zhong found that N addition for one year insignificantly
232 changed the soil aggregate distribution in subtropical forest. Although N addition had no
233 significant effect on aggregate distribution and stability, the percentage of macro-aggregate and
234 values of MWD and GMD were higher in N addition treatments compared with that in control
235 treatment. Meanwhile, N addition changed the contents of GRSP and SOC in aggregate fractions,
236 which indicated that N addition could have a trend to change the soil aggregate distribution and
237 could increase the aggregate stability with N addition duration.

238 **Effects of N addition on GRSP content**

239 The contents of EE-GRSP and T-GRSP were in the range of 0.099–0.551 mg/g and 0.315–
240 4.058 mg/g, respectively. This finding is consistent with that of the studied on the GRSP contents
241 of different vegetation communities (0.6–5.8 mg/g) (Singh, Singh and Tripathi 2013). In this
242 study, GRSP content varied among aggregate fractions and bulk soil, and N addition had
243 different effects on GRSP distribution in aggregate fractions and bulk soil. The research in
244 farmland system found that tillage management and manure amendment had different effects on
245 GRSP distribution in aggregate fractions (Wright et al. 2007, Xie et al. 2015). Long-term manure
246 amendments redistributed the GRSP in the macro-aggregate fractions. These results indicated
247 that GRSP distribution in aggregate fractions is susceptible to environmental changes. EE-GRSP

248 was produced by AMF most recently, and EE-GRSP was proved to be more sensitive to
249 environmental changes than T-GRSP (Singh et al. 2013). In our study, the content of EE-GRSP
250 in bulk soil increased first and then decreased with the increasing N addition levels and reached
251 the peak (0.403 mg/g) at 6 g N m⁻² y⁻¹. The variation in EE-GRSP content was mainly attributed
252 to the changes in plant growth and soil microbial activities (Xie et al. 2015), which were induced
253 by N addition. In our study, plant growth and soil microbial activities might be increased at 3 g N
254 m⁻² y⁻¹ and inhibited at 9 g N m⁻² y⁻¹. Thus, N addition of 6 g N m⁻² y⁻¹ might be the threshold
255 value between increased and inhibited GRSP production. Xie (Xie et al. 2015) found that the
256 highest EE-GRSP contents were not obtained at the highest amount of organic manure. Our
257 results indicated that low-level N addition could increase the EE-GRSP contents, but high-level
258 N addition inhibited the EE-GRSP contents. The highest EE-GRSP contents were observed at 6
259 g N m⁻² y⁻¹, which indicated a proper N addition level can maximize the EE-GRSP content in
260 this region. These results supported our first hypothesis. Furthermore, N addition had different
261 effects on EE-GRSP content in aggregate fractions. The EE-GRSP content in macro-aggregates
262 increased initially and then decreased with the increasing N addition levels and had a peak value
263 at 6 g N m⁻² y⁻¹. Meanwhile, the EE-GRSP content in micro-aggregates had a low value at 6 g N
264 m⁻² y⁻¹. These results were inconsistent with our first hypothesis. Previous studies showed that
265 organic manure significantly increased the EE-GRSP content in all soil aggregate fractions
266 except micro-aggregates in farmland ecosystems (Xie et al. 2015), and increasing soil depth
267 significantly increased the EE-GRSP content in 0.5–1.0 and 1–2 mm aggregates without other
268 size aggregate fractions in citrus rhizosphere (Wu et al. 2013). These different results indicated
269 that the distribution of EE-GRSP in aggregate fractions varied with environmental conditions,
270 which resulted in the changes of EE-GRSP in bulk soil. In addition, the percentage of macro-
271 aggregates was higher than those of other size aggregate fractions. N addition has the same
272 effects at EE-GRSP content in macro-aggregates with that in bulk soil, which indicated that the
273 EE-GRSP in macro-aggregates was the main factor that determined the EE-GRSP in bulk soil.
274 T-GRSP contained easily and hardly extractable parts. Compared with EE-GRSP, T-GRSP can

275 reflect the accumulation of glomalin in the soil (Singh et al. 2013). In our study, N addition
276 significantly increased the T-GRSP content in bulk soil and had a peak value at $9 \text{ g N m}^{-2} \text{ y}^{-1}$,
277 which was different from the changes of EE-GRSP content in bulk soil. A large number of
278 studies found that environmental changes had different effects on T-GRSP and EE-GRSP
279 contents in bulk soil as well (Wright et al. 2007, Xie et al. 2015, Antibus et al. 2006, Zhang et al.
280 2015). A explanation for this result in our study is that high because high-level N addition can
281 quickly relieve N limitation in the soil and promote the microbial activities at the beginning of
282 the experiment, which increase GRSP production in the short term. Given the increasing time,
283 high-level N addition may cause N saturation in the soil and inhibit microbial activities that
284 decrease the GRSP production (Yu et al. 2013). However, the T-GRSP content under high N
285 addition level was higher than that under low N addition level for a period of time due to the
286 cumulative effect. Furthermore, the T-GRSP content in aggregate fractions increased
287 significantly with the N addition, and this result supports our explanation.

288 **Effects of N addition on SOC content**

289 In this study, SOC content ranged from 6.478 and 21.627 mg/g. Across all aggregate
290 fractions, macro-aggregate had the highest SOC content, followed by the micro-aggregate, and
291 those in clay-silt aggregate is the lowest. These results were consistent with that of the previous
292 study (Zhang et al. 2014). However, some studies showed that the highest SOC content was in
293 clay-silt aggregate, and the lowest was in macro-aggregate (Gao et al. 2013, Tang et al. 2016).
294 Different results might be caused by soil type, plant community, and experimental conditions
295 (Wei et al. 2011). N addition for 2 years significantly increased SOC content in bulk soil, which
296 indicated that N addition could quickly increase plant growth in N limitation soil and increase
297 soil C storage (Nave et al. 2009). Interestingly, N addition significantly increased SOC content in
298 all size aggregate fractions. Compared with other size aggregate fractions, macro-aggregate had
299 the highest percentage and SOC content, with the greatest contribution to the soil C storage.
300 Meanwhile, the SOC in clay-silt aggregate had the smallest contribution to the soil C storage.
301 These results were different from our hypothesis. Zhong (Zhong et al. 2017) found that N
302 addition ($0\text{--}300 \text{ kg N ha}^{-1} \text{ y}^{-1}$) for 1 year had no significant influence on SOC in any size

303 aggregate fractions in a subtropical forest. These different results might be related to soil type
304 and initial N content. For example, macro- and micro-aggregate fractions had the highest SOC
305 content among the size aggregate fractions in loamy soils, while clay-silt aggregate fraction had
306 the highest SOC content than those of other size aggregate fractions in sandy soil (Gao et al.
307 2013). Our results indicated that the SOC content varied among aggregate fractions, and N
308 addition had different effects on SOC in aggregate fractions. The effects were not only affected
309 by N addition level but also by soil type and initial N content.

310 **Relationships among GRSP, SOC, and aggregate stability**

311 GRSP and SOC were insignificantly correlated with MWD. This result fails support our
312 hypothesis. A large number of studies found that GRSP and SOC were significantly correlated
313 with MWD because both can bind soil particle for aggregate formation (Rillig 2004, Wright et al.
314 2007, Baldock, Kay and Schnitzer 1987). However, recent studies found that GRSP and/or SOC
315 may have no significant relationship with MWD (Wu et al. 2013, Li et al. 2005, Leelamanie and
316 Mapa 2015) because aggregate stability was not only affected by GRSP and SOC contents but
317 also by soil texture, human disturbance, and other factors (Wei et al. 2011, Tang et al. 2016). In
318 our study, the contents of GRSP and SOC significantly changed, but MWD insignificantly
319 changed with N addition levels which caused insignificant relationships among MWD, GRSP,
320 and SOC. N addition significantly increased the contents of GRSP and SOC and insignificantly
321 increased MWD, indicating that increasing GRSP and SOC contents does not necessarily
322 increase aggregate stability rapidly. Aggregate formation requires not only binding agents but
323 also physicochemical effects, but the underlying mechanism still needs further study.

324 **Conclusions**

325 N addition has different effects on GRSP content in aggregate fractions at *P. tabulaeformis*
326 forestland. The EE-GRSP content in macro-aggregates increased initially and then decreased
327 with the increasing N addition levels and had a peak value at $6 \text{ g N m}^{-2} \text{ y}^{-1}$, indicating a proper N
328 addition level might maximize the GRSP production in this region. The T-GRSP content in
329 aggregate fractions increased significantly with the N addition and had a peak value at 9 g N m^{-2}
330 y^{-1} , because of the cumulative effect. Compare with control treatment, SOC content in all size

331 aggregate fractions significantly increased in N addition treatments, which indicate that N
332 addition could quickly increase plant growth in N limitation soil and increase soil C storage.
333 However, N addition for 2 years has no significant effects on MWD and GMD, and GRSP and
334 SOC are insignificantly correlated with MWD. This study identifies that the underlying
335 mechanism of aggregate formation requires binding agents and reaction time both. Increasing
336 GRSP and SOC contents does not necessarily increase aggregate stability rapidly.

337 **References**

- 338 Antibus, R. K., C. Lauber, R. L. Sinsabaugh & D. R. Zak (2006) Responses of Bradford-reactive
339 soil protein to experimental nitrogen addition in three forest communities in northern
340 lower Michigan. *Plant & Soil*, 288, 173-187.
- 341 Bai, Y., W. U. Jianguo, C. M. Clark, S. Naeem, Q. Pan, J. Huang, L. Zhang & G. Xing (2010)
342 Tradeoffs and thresholds in the effects of nitrogen addition on biodiversity and ecosystem
343 functioning: evidence from inner Mongolia Grasslands. *Global Change Biology*, 16, 358-
344 372.
- 345 Baldock, J. A., B. D. Kay & M. Schnitzer (1987) Influence of cropping treatments on the
346 monosaccharide content of the hydrolysates of a soil and its aggregate fractions.
347 *Canadian Journal of Soil Science*, 67, 489-499.
- 348 Bedini, S., E. Pellegrino, L. Avio, S. Pellegrini, P. Bazzoffi, E. Argese & M. Giovannetti (2009)
349 Changes in soil aggregation and glomalin-related soil protein content as affected by the
350 arbuscular mycorrhizal fungal species *Glomus mosseae* and *Glomus intraradices*. *Soil*
351 *Biology & Biochemistry*, 41, 1491-1496.
- 352 BlancoCanqui, H. & R. Lal (2004) Mechanisms of Carbon Sequestration in Soil Aggregates.
353 *Critical Reviews in Plant Sciences*, 23, 481-504.
- 354 Chen, Z. J., S. C. Geng, J. H. Zhang, H. Setälä, Y. Gu, F. Wang, X. Zhang, X. X. Wang & S. J.
355 Han (2017) Addition of nitrogen enhances stability of soil organic matter in a temperate
356 forest. *European Journal of Soil Science*, 68.
- 357 Das, B., D. Chakraborty, V. K. Singh, P. Aggarwal, R. Singh & B. S. Dwivedi (2014) Effect of

- 358 Organic Inputs on Strength and Stability of Soil Aggregates Under Rice-Wheat Rotation.
359 *International Agrophysics*, 28, 163-168.
- 360 Fonte, S. J., E. Yeboah, P. Ofori, G. W. Quansah, B. Vanlauwe & J. Six (2009) Fertilizer and
361 residue quality effects on organic matter stabilization in soil aggregates. *Soil Science*
362 *Society of America Journal*, 73, 961-966.
- 363 Galloway, J. N., F. J. Dentener, D. G. Capone, E. W. Boyer, R. W. Howarth, S. P. Seitzinger, G.
364 P. Asner, C. C. Cleveland, P. A. Green & E. A. Holland (2004) Nitrogen Cycles: Past,
365 Present, and Future. *Biogeochemistry*, 70, 153-226.
- 366 Gao, H., L. Qiu, Y. Zhang, L. Wang, X. Zhang & J. Cheng (2013) Distribution of organic carbon
367 and nitrogen in soil aggregates of aspen (*Populus simonii* Carr.) woodlands in the semi-
368 arid Loess Plateau of China. *Soil Research*, 51, 406.
- 369 Garcia, M. O., T. Ovasapyan, M. Greas & K. K. Treseder (2008) Mycorrhizal dynamics under
370 elevated CO₂ and nitrogen fertilization in a warm temperate forest. *Plant & Soil*, 303,
371 301-310.
- 372 Kathleenk, T., T. Katiem & M. Michellec (2007) Mycorrhizal responses to nitrogen fertilization
373 in boreal ecosystems: potential consequences for soil carbon storage. *Global Change*
374 *Biology*, 13, 78-88.
- 375 Keeler, B. L., S. E. Hobbie & L. E. Kellogg (2009) Effects of Long-Term Nitrogen Addition on
376 Microbial Enzyme Activity in Eight Forested and Grassland Sites: Implications for Litter
377 and Soil Organic Matter Decomposition. *Ecosystems*, 12, 1-15.
- 378 KH, B., B. EM, D. RA, H. KS, J. ES, S. JE & C. MJ (2014) A long-term nitrogen fertilizer
379 gradient has little effect on soil organic matter in a high-intensity maize production
380 system. *Global Change Biology*, 20, 1339.
- 381 Lü, C. & H. Tian (2007) Spatial and temporal patterns of nitrogen deposition in China: Synthesis
382 of observational data. *Journal of Geophysical Research Atmospheres*, 112, 229-238.
- 383 Leelamanie, D. A. L. & R. B. Mapa (2015) Alterations in soil aggregate stability of a tropical
384 Ultisol as mediated by changes in land use. *Biologia*, 70, 1444-1449.

- 385 Li, Z. X., C. F. Cai, Z. H. Shi & T. W. Wang (2005) Aggregate stability and its relationship with
386 some chemical properties of red soils in subtropical China. *Pedosphere*, 15, 129-136.
- 387 Liu, X., G. Yin, J. Wu, J. Chen, Z. He & S. Shi (2015) Effect of organic carbon and total
388 nitrogen distribution in alpine meadow soil aggregates with different nitrogen addition
389 level. *Nongye Gongcheng Xuebao/transactions of the Chinese Society of Agricultural*
390 *Engineering*, 31, 139-147.
- 391 Nave, L. E., E. D. Vance, C. W. Swanston & P. S. Curtis (2009) Impacts of elevated N inputs on
392 north temperate forest soil C storage, C/N, and net N-mineralization. *Geoderma*, 153,
393 231-240.
- 394 Nelson, D. W. & L. E. Sommers (1982) Total carbon, organic carbon and organic matter, in:
395 Methods of Soil Analysis Part 2. Chemical and Microbial Properties.
- 396 Pan, G. X., P. Zhou, L. Q. Li & X. H. Zhang (2007) Core issues and research progresses of soil
397 science of C sequestration. *Acta Pedologica Sinica*.
- 398 Rillig, M. C. (2004) Arbuscular mycorrhizae, glomalin, and soil aggregation. *Canadian Journal*
399 *of Soil Science*, 84, 355-363.
- 400 Singh, P. K., M. Singh & B. N. Tripathi (2013) Glomalin: an arbuscular mycorrhizal fungal soil
401 protein. *Protoplasma*, 250, 663-669.
- 402 Six, J., E. T. Elliott, K. Paustian & J. W. Doran (2010) Aggregation and Soil Organic Matter
403 Accumulation in Cultivated and Native Grassland Soils. *Soil Science Society of America*
404 *Journal*, 62, 1367-1377.
- 405 Tang, F. K., M. Cui, Q. Lu, Y. G. Liu, H. Y. Guo & J. X. Zhou (2016) Effects of vegetation
406 restoration on the aggregate stability and distribution of aggregate-associated organic
407 carbon in a typical karst gorge region. *Solid Earth*, 7, 141-151.
- 408 Treseder, K. K. (2004) A meta-analysis of mycorrhizal responses to nitrogen, phosphorus, and
409 atmospheric CO₂ in field studies. *New Phytologist*, 164, 347-355.
- 410 Treseder, K. K. & M. F. Allen (2002) Direct nitrogen and phosphorus limitation of arbuscular
411 mycorrhizal fungi: a model and field test. *New Phytologist*, 155, 507-515.

- 412 Tripathi, S. K., C. P. Kushwaha & K. P. Singh (2008) Tropical forest and savanna ecosystems
413 show differential impact of N and P additions on soil organic matter and aggregate
414 structure. *Global Change Biology*, 14, 2572–2581.
- 415 Wang, F., Y. A. Tong, J. S. Zhang, P. C. Gao & J. N. Coffie (2013) Effects of various organic
416 materials on soil aggregate stability and soil microbiological properties on the Loess
417 Plateau of China. *Plant Soil & Environment*, 59, 162-168.
- 418 Wang, G. L., S. Xue, F. Liu & G. B. Liu (2017) Nitrogen addition increases the production and
419 turnover of the lower-order roots but not of the higher-order roots of *Bothriochloa*
420 *ischaemum*. *Plant and Soil*, 423–434.
- 421 Wei, Y. W., Y. R. Su, X. B. Chen, X. Y. He, W. G. Qin & G. F. Wei (2011) [Effects of human
422 disturbance on soil aggregates content and their organic C stability in Karst regions]. *Ying*
423 *Yong Sheng Tai Xue Bao.*, 22, 971-978.
- 424 Wright, F. Sara, Upadhyaya & Abha (1996) EXTRACTION OF AN ABUNDANT AND
425 UNUSUAL PROTEIN FROM SOIL AND COMPARISON WITH HYPHAL PROTEIN
426 OF ARBUSCULAR MYCORRHIZAL FUNGI. *Soil Science*, 161, 575-586.
- 427 Wright, S. F., V. S. Green & M. A. Cavigelli (2007) Glomalin in aggregate size classes from
428 three different farming systems. *Soil & Tillage Research*, 94, 546-549.
- 429 Wright, S. F. & A. Upadhyaya (1998) A survey of soils for aggregate stability and glomalin a
430 glycoprotein of arbuscular mycorrhizal fungi. *Plant Soil* 198: 97-107. *Plant & Soil*, 198.
- 431 Wu, Q. S., X. H. He, M. Q. Cao, Y. N. Zou, W. Shuang & L. Yan (2013) Relationships between
432 glomalin-related soil protein in water-stable aggregate fractions and aggregate stability in
433 citrus rhizosphere. *International Journal of Agriculture & Biology*, 15, 603-606.
- 434 Wuest, S. B., T. C. Caesar-Tonthat, S. F. Wright & J. D. Williams (2005) Organic matter
435 addition, N, and residue burning effects on infiltration, biological, and physical properties
436 of an intensively tilled silt-loam soil. *Soil & Tillage Research*, 84, 154-167.
- 437 Xie, H., J. Li, B. Zhang, L. Wang, J. Wang, H. He & X. Zhang (2015) Long-term manure
438 amendments reduced soil aggregate stability via redistribution of the glomalin-related soil

- 439 protein in macroaggregates. *Scientific Reports*, 5, 14687.
- 440 Yin, J., R. Wang, H. Liu, X. Feng, Z. Xu & Y. Jiang (2016) Nitrogen addition alters elemental
441 stoichiometry within soil aggregates in a temperate steppe. *Solid Earth Discussions*, 7,
442 1565-1575.
- 443 Yu, P. Y., F. Zhu, S. F. Su, Z. Y. Wang & W. D. Yan (2013) [Effects of nitrogen addition on red
444 soil microbes in the *Cinnamomum camphora* plantation]. *Environmental Science*, 34,
445 3231.
- 446 Zak, D. R. & R. L. Sinsabaugh (2004) Nitrogen Deposition Modifies Soil Carbon Storage
447 through Changes in Microbial Enzymatic Activity. *Ecological Applications*, 14, 1172-
448 1177.
- 449 Zhang, J., X. Tang, X. He & J. Liu (2015) Glomalin-related soil protein responses to elevated
450 CO₂ and nitrogen addition in a subtropical forest: Potential consequences for soil carbon
451 accumulation. *Soil Biology & Biochemistry*, 83, 142-149.
- 452 Zhang, X., X. Wu, S. Zhang, Y. Xing, R. Wang & W. Liang (2014) Organic amendment effects
453 on aggregate-associated organic C, microbial biomass C and glomalin in agricultural soils.
454 *Catena*, 123, 188-194.
- 455 Zhong, X. L., J. T. Li, X. J. Li, Y. C. Ye, S. S. Liu, P. D. Hallett, M. R. Ogden & M. Naveed
456 (2017) Physical protection by soil aggregates stabilizes soil organic carbon under
457 simulated N deposition in a subtropical forest of China. *Geoderma*, 285, 323-332.
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Table 1 (on next page)

The soil aggregate characteristic of different N addition treatments

values followed by a different letter indicate significant difference among the treatments.

MWD: Mean weight diameter, GMD □ Geometric mean diameter. ($P < 0.05$ $n = 6$).

Table 1 The soil aggregate characteristic of different N addition treatments

Treatment (g N m ⁻² y ⁻¹)	Aggregate proportion in size class (%)			MWD(mm)	GMD(mm)
	>0.25mm	0.053-0.25mm	<0.053mm		
CK(0)	65.36±7.63a	28.49±6.40a	6.16±1.98a	1.45±0.073a	0.90±0.039a
N3(3)	68.78±7.99a	25.95±7.41a	5.28±0.72a	1.50±0.158a	0.94±0.086a
N6(6)	73.65±5.05a	21.88±4.57a	4.48±0.90a	1.60±0.101a	0.99±0.063a
N9(9)	72.76±7.82a	22.11±6.45a	5.14±1.48a	1.58±0.156a	0.98±0.099a

Note: values followed by a different letter indicate significant difference

among the treatments. MWD: Mean weight diameter, GMD: Geometric mean diameter. ($P < 0.05$ n=6).

10

Table 2 (on next page)

Bivariate correlations (r values) between the mean weight diameter (MWD) and soil organic carbon (SOC), easily extractable and total glomalin-related soil protein (EE-GRSP and T-GRSP).

asterisk indicates a significant correlation, ($P < 0.05$. $n=24$).

1 Table 2 Bivariate correlations (r values) between the mean weight diameter (MWD) and soil organic carbon (SOC), easily
 2 extractable and total glomalin-related soil protein (EE-GRSP and T-GRSP).

3 asterisk 5 significant 8 ($P < 0.05$.		SOC	EE-GRSP	T-GRSP	indicates a correlation, n=24).
9	Bulk soil	0.381	0.176	0.250	
	>0.25mm	0.254	0.156	0.266	
	0.25-0.053mm	0.420*	-0.050	0.318	
	<0.053mm	0.376	0.098	-0.006	

Figure 1(on next page)

Distributions of easily extractable and total glomalin-related soil protein (A: EE-GRSP and B: T-GRSP) concentration in soil aggregates and bulk soil of different N addition treatments.

Values followed by a different lowercase letters indicate significant difference among the N addition treatments, values followed by a different uppercase letters indicate significant difference among the aggregate size. ($P < 0.05$ $n = 6$).

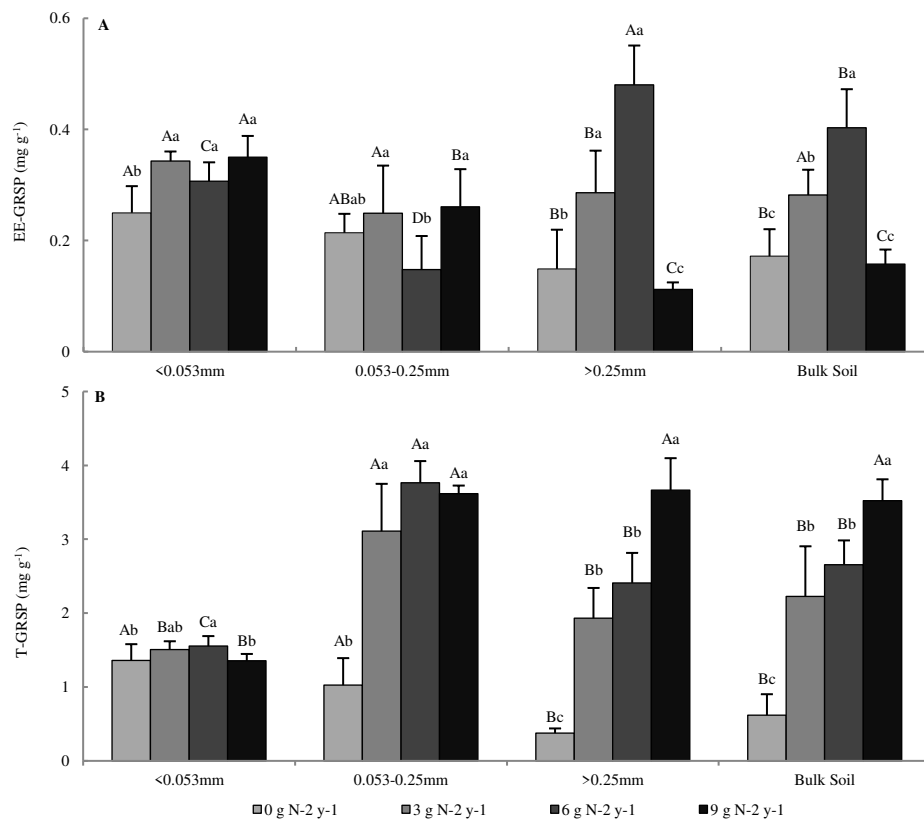


Figure 1 Distributions of easily extractable and total glomalin-related soil protein (A: EE-GRSP and B: T-GRSP) concentration in soil aggregates and bulk soil of different N addition treatments

Note: Values followed by a different lowercase letters indicate significant difference among the N addition treatments, values followed by a different uppercase letters indicate significant difference among the aggregate size. ($P < 0.05$ $n=6$).

Figure 2 (on next page)

Distributions of organic C concentration in soil aggregates and bulk soil of different N addition treatments.

Values followed by a different lowercase letters indicate significant difference among the N addition treatments, values followed by a different uppercase letters indicate significant difference among the aggregate size. ($P < 0.05$ $n=6$).

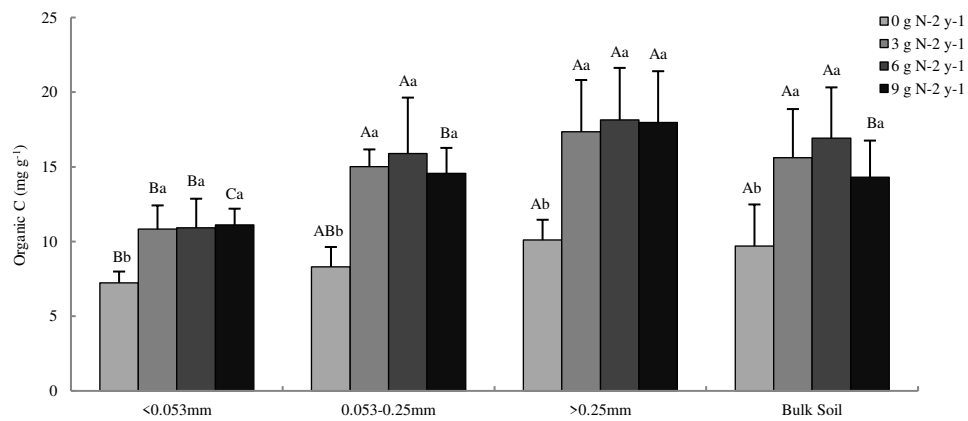


Figure 2 Distributions of organic C concentration in soil aggregates and bulk soil of different N addition treatments

Note: Values followed by a different lowercase letters indicate significant difference among the N addition treatments, values followed by a different uppercase letters indicate significant difference among the aggregate size. ($P < 0.05$ $n=6$).