Cardonatto MC, Melchor RN. 2018. Large mammal burrows in late Miocene calcic paleosols from central Argentina: paleoenvironment, taphonomy and producers. PeerJ 6:e4787
https://doi.org/10.7717/peerj.4787

# Large mammal burrows in late Miocene calcic paleosols from central Argentina: palaeoenvironment, taphonomy and producers

María Cristina Cardonatto[1], Ricardo Néstor Melchor[1,2]

*[1]Geology Department, National University of La Pampa, Av. Uruguay 151, 6300 Santa Rosa, La Pampa, Argentina. Email: mccardonatto@exactas.unlpam.edu.ar*

*[2]Instituto de Ciencias de La Tierra y Ambientales de La Pampa (Universidad Nacional de La Pampa-Consejo Nacional de Investigaciones Científicas y Técnicas), Av. Uruguay 151, 6300 Santa Rosa, La Pampa, Argentina.*

**Corresponding autor:**

María Cristina Cardonatto[1]

PeerJ Preprints | https://doi.org/10.7287/peerj.preprints.26632v1 | rec: 7 Mar 2018, publ: 7 Mar 2018

## Abstract

Large cylindrical structures with a dominantly laminated fill occurring in the loess-paleosol sequence of the late Miocene Cerro Azul Formation (central Argentina) are interpreted as mammal burrows. A total of 115 burrow fills from three localities were measured. The fills typically are shallowly dipping, subcylindrical, unbranched structures with a rounded end and lacking enlargements. The horizontal diameter of the burrows range between 0.15 and 1.50 m, with most of the burrows in the interval of 0.39 to 0.98 m. Geometric morphometric analysis of transverse cross-sections support the distinction of subcircular and elliptical (horizontally flattened) fills. Burrows fills are typically laminated in the lower part and massive in the upper part. The laminated intervals reflect pulses of aqueous currents entering the abandoned burrow during moderate rains, whereas massive intervals reflect *en masse* input of sediment-water mixtures during heavy rains that produced sheet floods. About 1% of the burrow fills contained fragmentary, disarticulated and weathered mammal bones that were introduced in the open burrow by currents along with other sedimentary particles. Analysis of the tetrapod burrow fossil record suggests that Miocene burrows, including those studied herein; reflect a remarkable increase in the average size of the fossorial fauna. We conclude that large late Miocene mammals dug burrows essentially as a shelter against environmental extremes and also to escape predation. Several mammal groups acquired fossorial habits in response to cold and seasonally dry climatic conditions that prevailed during the late Miocene in southern South America. The ample range of horizontal diameters of the studied burrow fills can be attributed to a variety of producers, including dasypodids, the notoungulate *Paedotherium minor*, Glyptodontidae and *Proscelidodon* sp. The simple architecture of the burrows suggests that the producers essentially foraged aboveground.

Keywords: tetrapod burrow; laminated burrow fill; Xenarthra; taphonomy

## Introduction

Fossil vertebrate burrows are relatively common biogenic structures and the oldest convincing evidence of tetrapod burrows are Early Permian (Asselian-Artinskian) lysophorid amphibian burrows from Kansas, USA (Hembree, Martin & Hasiotis, 2004). Most pre-Cretaceous tetrapod burrows have been attributed to therapsids, in part on the basis of the finding of articulated skeletons in a few Late Permian-Early Triassic burrows (Smith 1987; Groenewald, Welman & MacEachern, 2001; Damiani et al., 2003; Modesto & Botha-Brink 2010). A common architecture for Permian to Jurassic tetrapod burrows is a shallowly inclined ramp with a rounded and not enlarged end, of reduced horizontal diameter (Fig. 1), with discrete scratch marks, always showing a horizontally flattened elliptical cross-section, and commonly with a bilobed bottom (e.g., Damiani et al., 2003; Sidor et al., 2008; Riese, Hasiotis & Odier, 2011; Liu & Li 2013; Melchor & Loope 2016; Krummeck & Bordy, 2017). Most of Jurassic tetrapod burrows occur in eolian sequences including the oldest subhorizontal burrow systems that have been assigned to primitive mammals from the Early Jurassic Navajo Sandstone of USA (Riese, Hasiotis & Odier, 2011) (Fig. 1). In contrast, there is a dearth of reports of Cretaceous tetrapod burrows that can be related with the equable climates that existed for most of this period. The more unusual occurrence for the Late Cretaceous is the unique dinosaur burrow containing an adult and two juvenile remains of ornithopods, further suggesting denning behavior and parental care (Varricchio, Martin & Katsura, 2007).

Most of the Cenozoic tetrapod burrows have been attributed to mammals, mainly to Rodentia and Xenarthra (e.g., Voorhies, 1975; Martin & Bennett 1977; Benton, 1988; Gobetz & Martin 2006), whereas a few examples are related to Carnivora (e.g., Hunt, Xiang-Xu & Kaufman, 1983; Hembree & Hasiotis 2008). The record of Paleogene tetrapod burrows is meager and can also be linked to dominantly benign climate conditions (Fig. 1). The Miocene record of tetrapod burrows is more varied and abundant, with a diversification of the architectural patterns and behavioral strategies that, commonly, appeared under stressed volcaniclastic and eolian environments. The early Miocene volcaniclastic floodplains of

Nebraska, USA, witnessed the appearance of 1) the celebrated vertical helical burrows ending in a ramp and terminal chamber named *Daimonelix* (e.g., Barbour, 1892; Martin & Bennett, 1977), as well as 2) smaller, complex subhorizontal rodent burrow systems with terminal chambers and subcircular cross-section (Gobetz & Martin, 2006), and 3) the first carnivore den (Hunt, Xiang-Xu & Kaufman, 1983; Hunt 1990). Also in the early Miocene, the coastal dunes of Germany preserved the oldest fossil food cache (Gee et al., 2003). The main tetrapod burrowing innovation during the middle Miocene is represented by cylindrical, subhorizontal, unbranched tunnels with a meniscate backfill interpreted as foraging tunnels of small Dasypodidae from southern South America (Melchor et al., 2012; 2016). Pliocene burrows are small (horizontal diameter less than 250 mm) and those from the Atlantic coast of the Buenos Aires province, Argentina have received a considerable attention, because they are common and a significant amount of the burrows contains bone remains (e.g., Genise, 1989; Quintana, 1992; Fernández, Vassallo & Zárate, 2000; Elissamburu, Dondas & De Santis, 2011). These are subcircular burrows assigned to rodents and notoungulates. The Pleistocene megafauna of South America is also reflected in the burrow trace fossil record in the form of huge tunnels (up to 2 m wide), with horizontally flattened elliptical cross-sections from Argentina and Brazil (e.g., Quintana, 1992; Vizcaíno et al., 2001; Buchmann, Pereira Lopes & Caron, 2009; Genise & Farina 2012; Frank et al., 2012; 2015). These burrows were attributed to Dasypodidae and Pampatheriidae (the small ones) and to ground sloths (intermediate and large ones). It has been suggested that the acquisition or generalization of burrowing behavior by large Pleistocene herbivorous mammals may reflect the arrival of large carnivorans after the Great American Biotic Interchange, just before the Pliocene-Pleistocene boundary (Soibelzon et al., 2009).

A trend towards larger burrows (as suggested by average Dh) is evident from the early Miocene to the late Pleistocene (Fig. 1). Pre-Miocene burrows attain a maximum average horizontal diameter of 420 mm (the examples from the Navajo Sandstone described by Loope, 2006), however, most are in the range of 100-200 mm. From the Miocene on, burrows with an average horizontal diameter in excess of 500 mm are recorded (Fig. 1), including those studied herein. The Neogene record also includes more common and smaller burrows (Dh~< 200 mm) that are mostly attributed to rodents and small Dasypodidae (Fig. 1).

In this context, the tetrapod burrows from the late Miocene Cerro Azul Formation are the largest pre-Pleistocene tetrapod burrows and can help to understand the reasons for early acquisition of burrowing habits in large Cenozoic mammals from South America. These structures have been partially and briefly described (Genise et al., 2013), but a detailed description is missing. The purposes of this work are: 1) to infer the likely producers of these late Miocene large burrows and 2) to interpret the taphonomy of the burrows and its paleoecological and paleoenvironmental meaning.

## Material & methods

The studied localities with late Miocene burrow fills are distributed in a latitudinal belt of approximately 25 km (Fig. 2): Salinas Grandes de Hidalgo (SG, 37º 12’ 55”S, 63º 35’ 25”W; 100 masl); Laguna Chillhué (LC, 37°19'15.13"S; 64°14'31.52"W; 145 masl); and Laguna La Paraguaya near Carhué city (LLP, 37° 5'53.57"S; 62°47'34.98"W; 101 m.a.s.l.). The first two localities belong to La Pampa province, whereas the remaining is located in the adjacent Buenos Aires province. The burrow fills appear in outcrop as transverse to oblique, longitudinal, tangential and plan view exposures.

Another locality of the Cerro Azul Formation, a roadcut in national road 154 (R154, 37º 49’ 28.5”S, 64º 4’ 8.9”W), has been previously described as having vertebrate burrows (De Elorriaga & Visconti, 2001). However, this locality is not considered herein because the burrow fills probably postdate significantly the deposition of the unit. The reasons for this inference are: 1) burrow fill diameters are considerably larger than those described herein (Dv up to 2.25 m) and more consistent with burrows attributed to a large Pleistocene megafauna (e.g., Vizcaíno et al., 2001; Frank et al., 2012); and 2) the burrow fill cut the carbonate nodules of the capping calcrete. Although absolute dating of the calcrete is not available, it has been suggested that the calcretization process postdate significantly the deposition of the Cerro Azul Formation (Vogt, Carballo & Calmels, 1999; Melchor & Casadío 2000; Visconti et al., 2010). The main argument for this inference is that the calcrete is developed in sedimentary sequences ranging in age from late Miocene to Pleistocene.

Standard measurements in burrow fills were horizontal diameter (Dh), vertical diameter (Dv), preserved length, mean azimuth of burrow fill and inclination of internal laminae. The mean azimuth was measured using a compass and considering the burrow fill boundary and dominant plunge of laminated fill. When changes of dip direction or inclination were observed in a single burrow fill they were recorded separately. Burrow fill diameter (especially Dh) was measured orthogonal to the main axis of the structure. The horizontal diameter was obtained from almost all types of exposures (except for longitudinal ones), whereas the vertical diameter was mostly an apparent value, except for the rare transverse sections, where it can be considered the true vertical diameter. At each location, GPS coordinates were recorded; the burrow fills were photographed and sketched.

The burrow fills exposed in transverse section (n=24) were used for a bidimensional geometric morphometric analysis. From field photographs, the outline of the burrow boundary was sketched in Corel Draw™. These images were used to build a file with the TPSutil software. For each image a total of 16 landmarks were digitised using the TPSdig2 software, arranged with an equidistant pattern on the burrow outline (Fig. 3). Landmarks were aligned using the Procrustes superposition method (McLeod, 2009) and the principal component analyses (PCA) using MorphoJ software. Results are presented by eigenvalue diagrams and PC scores, along with transformation grids.

From scaled sketches of transverse sections of burrow fills, the cross-sectional area of the burrow fill was estimated using ImageJ software. This cross-sectional area was used to estimate the body mass of the producer using the allometric relations by Wu et al. (2015).

At each locality, a detailed sedimentary log was measured and samples from host rock and burrow fill were obtained for petrographic analysis. At selected locations also the burrow fill was logged using standard sedimentologic techniques. Mammal bones found *in situ* within burrow fills were recorded and collected for preparation and detailed taxonomic assignation. Fossil bones found outside burrow fills were not recorded nor collected. Fossil remains from La Pampa province collected during this study are housed at the Paleontological Collection of the Facultad de Ciencias Exactas y Naturales, Universidad Nacional de La Pampa, Santa Rosa city, La Pampa, Argentina, under the acronym GHUNLPam. Fossil material collected in burrows from Laguna La Paraguaya locality (Buenos Aires province) are housed at the Museo Histórico Regional de Guaminí “Coronel Marcelino E. Freyre” under the acronym MHG-P. Field work was approved by the Dirección Provincial de Museos y Preservación Patrimonial, under the Project "Vertebrados del Mioceno tardío-Plioceno en el área de las lagunas encadenadas del oeste de la provincia de Buenos Aires. Aportes a la bioestratigrafía del Cenozoico tardío de la Región Pampeana", permit nr. 2015-3-P-156-2.

## Geological setting

The Cerro Azul Formation outcrops are located in the north-western, central, and eastern part of La Pampa province and adjacent western Buenos Aires province, Argentina (Linares, Llambías & Latorre, 1980; Folguera & Zárate 2009; Visconti et al., 2010). The unit is characterized by a monotonous succession of loess containing moderately developed paleosols (Fig. 4) that has been assigned to the late Miocene (Huayquerian Land Mammal age) essentially on the basis of its mammal remains (Montalvo & Casadío, 1988; Verzi, Montalvo & Vucetich, 1999; Verzi, Montalvo & Visconti, 2003). In particular, the formation is considered as representing the interval between 10 and 5.7 Ma (Cione et al., 2000; Verzi, Montalvo & Deschamps, 2008). The maximum exposed thickness in outcrop is 54 m, although the unit reaches about 180 m in the subsurface (Visconti et al., 2010). The formation is essentially composed of structureless, light brown (5YR 6/4), pale reddish brown (10R 5/4) or grayish orange pink (5YR 7/2), sandy siltstones and fine-grained sandstones, showing moderate selection and common carbonate cementation.

## Results

### *Sedimentology of the burrow fill bearing sections*

Macroscopic features of the studied sections are very similar to those of the classical localities of the Cerro Azul Formation outcrops (e.g., Visconti et al., 2010; Genise et al., 2013; Cardonatto et al., 2016), especially those of the Salinas Grandes de Hidalgo and Laguna La Paraguaya. Paleosol profiles are typically composed of two horizons (Fig. 4). The upper horizon is a clayey siltstone that is distinguished by the presence of subangular blocky, granular or prismatic peds and a darker color (5 YR 6/4) than the underlying horizon. Carbonates are rare except for the local occurrence of carbonate rhizoliths. This upper horizon can be compared with a Bt horizon and its thickness averages 0.60 m (range = 0.35-1.00 m, n = 5), which is very close to the average for the formation (Cardonatto et al., 2016).

The lower horizon is characterized by lighter colored (5 YR 7/2), siltstone to fine- grained sandstone with pervasive carbonate cementation, both in the form of nodules and rhizoliths. Rhizoliths are small to medium sized and commonly 1 -30 mm in diameter. The lower horizon can be up to more than 5 m thick and contain relicts of primary sedimentary structures, like tangential cross-bedding or horizontal bedding, as seen in the Salinas Grandes de Hidalgo section (Fig. 4). Mud intraclasts are common throughout. The remains of primary sedimentary structures and carbonate cementation suggest comparison with a Bk or Bk/C horizon. The trace fossils described in this paper occur in both horizons.

The section at Laguna Chillhué also contains similar paleosols (Fig. 4) and differs from the remaining localities by the presence of a 2 m thick, dark-red, laminated mudstone interval in the lower part of the section (Montalvo et al., 1995). The presence of a mudstone interval in the lower section of the Cerro Azul Formation has been questioned by Lorenzo, Mehl & Zárate (2013), who supposed a younger age for this mudstone interval on the basis of geomorphological inferences. However, at this location the laminated mudstone interval is overlain, through a normal sedimentary contact, by lithologies typical of the Cerro Azul Formation containing mammal remaims of Huayquerian (late Miocene) age, with no evidence of reworking. Vertebrate burrow fills were not observed in the laminated mudstone interval.

### ***Description of large burrow fills***

A total of 115 burrow fills were measured from three localities: 53 from Salinas Grandes de Hidalgo, 59 from Laguna La Paraguaya, and 3 from Laguna Chillhué (see Supplemental Table S1). The studied burrow fills are distinguished on the basis of the presence of a thinly laminated siltstone to mudstone interval that contrasts with the structureless host rock. When part of the fill is massive and similar to the host rock, the distinction of the burrow boundary is obscured. Sedimentary rocks composing the burrow fills exhibit an induration, composition and cementation similar to the hosting rocks, suggesting that they are of late Miocene age. At Laguna La Paraguaya locality, the preferential carbonate cementation of the burrow fills resulted in tridimentional exposures (Fig. 5A). At this locality the density of burrow fills is locally high and may be difficult to find unburrowed intervals. Cross-cutting relationships between two or more burrows are common (Figs. 5B, 5C).

#### *Size and plan view pattern*

Observed horizontal diameter ranges from 0.15 to 1.50 m (n= 112) and the frequency distribution suggests a roughly normal distribution where three subpopulations can be distinguished (Fig. 6). The small subpopulation has a Dh from 0.15 to 0.34 m (8 %), the intermediate subpopulation has a Dh from 0.39 to 0.98 m (84 %), and the large subpopulation exhibits a Dh from 1.05 to 1.50 m (8 %).

In plan view exposure, which is found only at SG and LLP localities (n=78), a number of morphologies can be distinguished (Fig. 7). 1) The more common are straight to slightly curved burrow fills (89 % of cases), which exhibits a Dh= 0.15-1.15 m, showing an uniform inclination of internal laminae (ranging from ≈ 0° to 27°), the maximum height difference between the proximal and distal portion of a burrow fill is 0.6 m, and the maximum preserved length is 5.18

m (Figs. 7A, 7B). Some burrow fills in this category display a shallowing of internal laminae toward more distal positions (i.e., from 27° to 8°). 2) A sinuous pattern is herein used to refer burrow fills that comprise two opposite curves in plan view, which was recorded in 5% of the cases (Figs. 7C, 7D). The horizontal diameter of sinuous burrow fills ranges from 0.42 to 0.80 m, dip of internal laminae is subhorizontal to slightly inclined (up to 8°), and the maximum observed length is 8 m. 3) The third plan view pattern is a C-shaped closed curve observed in 6% of the burrow fills, with an horizontal diameter ranging from 0.44 to 0.72 m (Figs. 7E, 7F), which commonly appears as a ramp with a height difference of up to 0.55 m, the inclination of internal laminae can be uniform (from 3° to 12°) or show a shallowing toward the distal position (from 14° to subhorizontal).

In a few cases, the distal portion of burrow fill showed an enlargement of up to 23% of the Dh, commonly having a subhorizontal lamination (Figs. 7G, 7H). Other burrow fills exhibit a rounded end with no enlargement that can be accompanied by an upward bending of mudstone laminae at the burrow boundary.

### *Cross-sectional shape and body mass*

The analysis of the well defined cross-sectional shape of burrows fills (n= 24 from all localities) suggest a distinction between elliptical (with the major axis subhorizontal) and subcircular cross-sections. Elliptical cross-sections are more common (n=18) and the corresponding Dh ranges from 0.39 to 1.50 m (belonging to the intermediate and large subpopulations, Fig. 6), with an average Dv/Dh ratio of 0.55. The burrows fills with elliptical cross-section include a few cases (n=4) with a flat bottom and convex top. The subcircular cross-sectional shape (n=6) is represented in the intermediate subpopulation with a Dh ranging from 0.39 to 0.56 m, and an average Dv/Dh ratio of 0.88.

Morphometric analyses suggest that 90.13 % of the variation is explained by the first two principal components (Fig. 8B), and deformation grids range from elliptical (score= -0.12) to subcircular (score= 0.17) (Fig. 8A).

Body mass estimates of the producers of the burrow on the basis of the cross-sectional area (using the method by Wu et al., 2015) suggest that there are two ranges (Table S1). Most of the estimates (n= 18) belongs to the intermediate subpopulation with a range from 37 to 439 kg, whereas the remaining estimates comes from the large subpopulation (n= 7) with a range 708 to 1623 kg. Burrows with subcircular cross-section from the intermediate subpopulation, are linked with a producer having body mass from 92.84 to 186.0 kg.

### *Orientation*

Readings of plunge azimuths of burrow fills from all localities are variable but most values are located in the northeast to southeast quadrants (i.e., between N20° and N140°) (Figs. 9A, 9C). The average dip angle of all measured burrow fills with respect to the paleohorizontal is 7.25° and ranges from nearly 0 to 27° (Fig. 9B). Most orientation data comes from the intermediate subpopulation (Dh= 0.39-0.98 m) and especially from LLP locality.

### *Composition of burrow fills*

The vertebrate burrows are easily spotted in the field because of its laminated structure that contrasts with the massive aspect of the hosting rock. The filling is composed of an alternation of laminated claystone and siltstone with massive fine-grained sandstone and siltstone having claystone intraclasts. Laminated intervals are a few millimeters to about 50 mm thick, whereas massive intervals tend to be thicker. Most of the burrow fills display a laminated interval in the lowermost part of the fill, whereas the upper part is massive, especially in the

Salinas Grandes de Hidalgo locality (Figs. 10A, 10B, 10C). A few burrows display a poorly defined lamination to massive structure (Fig. 10D). Claystone and siltstone at the bottom of the structure typically onlap on the burrow wall, which is a good criterion to distinguish burrow fills that are mostly eroded. Lamination tends to be horizontal but successive laminated packages resting at low angles were also identified. Individual laminae are normally graded (typically siltstone grading to claystone), and locally disrupted giving a brecciated aspect. Both synsedimentary faulting and deformation were identified (Fig. 10F). Some rare but distinctive feature of the fillings the pseudomeniscate structure, which was identified in two cases (one from Salinas Grandes de Hidalgo and the other from Laguna La Paraguaya), This structure is composed of massive siltstone or fine-grained sandstone arranged in adjacent crescent-shaped bodies with the convex margin pointing downslope that span the full width of the fill (Fig. 10E). Horizontal width of individual pseudomenicate bodies taken parallel to the burrow axis is 120 mm.

### *Associated ichnofossils*

Only the ichnofossils found within or very close to the burrow fills are considered. We found within the fills abundant rhizoliths and rare smaller vertebrate burrows, vertebrate footprints and ?*Rosellichnus* isp. In the paleosol adjacent to the fills we found *Taenidium barretti*, *Attaichnus kuenzelli* and *Coprinisphaera* isp.

Calcareous rhizoliths, including rhizocretions and root casts, are abundant and were recorded in all studied localities. Commonly, rhizoliths are arranged in vertical, oblique and subhorizontal position (Fig. 10C). Rhizoliths are submillimeter to 45 mm thick and the maximum preserved length is 500 mm. Rhizoliths occurring inside burrow fills are similar in shape and cementation to those of the hosting rock.

A single small burrow fill (Fig. 11A) that cuts the laminated interval of a larger vertebrate burrow fill was identified at Salinas Grandes de Hidalgo (# 638). The 43 mm in diameter structure displays a subcircular outline and a poorly laminated siltstone fill.

A partially eroded burrow fill from Laguna La Paraguaya (# 708) exposed an internal bedding plane of the filling showing elliptical footprints with a noticeable marginal rim (Fig. 11C). At least four footprints are recognized that probably compose a 0.7 m long quadrupedal trackway, with a width of 316 mm. If these footprints do compose a trackway a pace angulation of 112° and a stride length of 600 mm can be inferred. Measurements on individual footprints indicate that average footprint length is 144 mm, average footprint width is 93 mm, and the marginal rim is of uniform thickness and about 50 mm wide.

A group of five subcircular rings in the upper part of a burrow fill (# 659A from SG) is tentatively identified as a cluster of bee cells and compared with the ichnogenus *Rosellichnus* (Fig. 11D). The presence of the ichnogenus at this locality, although at other section, was already documented by Cardonatto et al. (2016).

Adjacent to the burrow fills at Salinas Grandes de Hidalgo, several specimens of *Attaichnus kuenzelii* were identifed, in some cases very close, but no cross-cutting relationship were seen. A few specimens of *Taenidium barretti* also occur at this locality, in the form of subcylindrical burrows, 12 mm wide and 80 mm long with an average meniscus thickness of 2.2 mm. At Laguna La Paraguaya we also found two specimens of cemented and compact spherical chambers (diameter 18.7-23.0 mm) with a large emergence hole (10-11 mm) assigned to *Coprinisphaera* isp. (Fig. 11B).

### *Bone remains found within burrow fills*

Mammal bones within the burrow fills are scarce (only 1% of burrows contained fossil remains) and usually appear disarticulated and poorly preserved, in some cases with signs of abrasion (Fig. 12A). The fossil remains from the SG locality are *Proscelidodon* sp. and Glyptodontidae indet., whereas the rest of the fossil material was found at LLP locality, including: *Paedotherium minor* (two specimens), *Doellotatus* sp., *Eosclerocalyptus* sp., Mesotheriinae indet., Gyptodontidae indet. (three specimens), and undeterminate mammals (two specimens). For details about the taxonomy and illustrations of mammal remains, see Supplemental Article S2 and Supplemental Figure S1. The only articulated remains are glyptodont osteoderms found at Laguna La Paraguaya (# 702) that are assigned to *Eosclerocalyptus* sp. (Fig. 12B), and remains of a carapace with several articulated osteoderms from the same locality (# 670) assigned to Glyptodontidae indet. *Proscelidodon* sp. remains (including a hemimandible with teeth and postcranial elements) appeared disarticulated but associated within a single burrow fill. The fossil remains display different degree of weathering and corrosion, as well as bioestratinomic fractures.

## Discussion

### *Producers*

The studied fossil burrows are unbranched and display a significant variation in the horizontal diameter, which ranges from 0.15 m to 1.5 m (Fig. 6). The simple, ramp type morphology of the studied burrows suggests that the animals foraged aboveground (e.g., Reichman & Smith, 1990). In order to infer the likely producers of the fossil burrows there are several constraints that need to be considered: 1) the faunal remains found inside the burrow fills; 2) the fossorial mammals that were recorded in the Cerro Azul Formation, especially those from the studied localities; 3) the size of burrows, as expressed by the Dh; and 4) the overall architecture and cross-section of burrows (including the Dv/Dh ratio) and the extrapolated body mass of its digger. The surface ornamentation of burrows is commonly a very useful clue to the producer (e.g., Seilacher, 2007), however, it is not preserved in the studied cases.

### *Faunal remains found in burrow fillings*

In general, bone remains found inside a burrow can be considered as belonging to its producer or occupant only if they are articulated, nearly complete, commonly are found in a terminal portion and fit in the size (cross-section) of the burrow (e.g., Smith, 1987; Groenewald, Welman & MacEachern, 2001; Damiani et al., 2003). The remains found inside the studied burrows do not fulfill any of these criteria. In most cases, these bone remains have been passively introduced and it is uncertain if they belong to the producers. The remains are essentially fragmentary, disarticulated, with evidence for abrasion and weathering (Fig. 12A, S1); suggesting that they spent some time in the surface and then were introduced into the burrows by currents along with other sedimentary particles. The fragmentary and disarticulated state of *Doellotatus* sp. and one of the specimens of *Paedotherium minor* and the considerably small size of the animals (body mass about 1-2 kg, Table 1) in comparison with the containing burrows; further suggest that these remains were introduced by currents. In the case of *Proscelidodon* sp., the bones are disarticulated but associated, which suggest that they can belong to a single specimen, and the partial horizontal diameter of the burrow match the size of this ground sloth. The only articulated remains are fragments of the dorsal carapace of Glyptodontidae that occur in burrows large enough to hug these animals (Dh= 0.78 to 1.50 m) (Table 1). In consequence, the unique remains that can belong to the producer of the burrows are *Proscelidodon* sp. and those of Glyptodontidae.

*Fossorial mammals of the Cerro Azul Formation and size of burrows*

The fossorial mammals recorded in the Cerro Azul Formation include xenarthrans, notoungulates and rodents (e.g., Goin, Montalvo & Visconti, 2000; Cerdeño & Montalvo 2001; Urrutia, Montalvo & Scillato-Yané, 2008). Among the Xenarthra, the Glyptodontidae, Dasypodidae and Mylodontidae display fossorial adaptations. The same is true for Mesotheridae and Hegetotheriidae (Notoungulata); and Caviidae, Octodontidae, and Chinchillidae (Rodentia). We will discuss the potential producers for each size class of the burrow fills (Table 1) as expressed by the horizontal diameter and cross-sectional area of the burrows.

For the small subpopulation (Dh=0.15-0.34 m, 8 % of cases), with a body mass ranging from 1 to 13 kg, the likely candidates are the notoungulate *Paedotherium minor*, the dasypodids *Doellotatus*, *Chorobates, Proeuphractus*, and *Chasicotatus*; and the rodent *Lagostomus.*

*Paedotherium* (Hegetotheriidae) is a medium-sized rodent-like ungulate native to South America. This taxon is very common in the Cerro Azul Formation, both in La Pampa and Buenos Aires provinces (e.g., Montalvo, Tomassini & Sostillo, 2016). Articulated remains of this genus have been found within Pliocene fossil burrows (about 0.16-0.22 m wide) from the Atlantic coast of Buenos Aires province (e.g., Genise, 1989; Scognamillo, 1993; Elissamburu, Dondas & De Santis, 2011) and a morphofunctional analysis of its postcranial skeleton suggest a digging capacity (Elissamburu 2004).

The Dasypodidae show a neotropical geographic distribution and were important components of the late Miocene-Pliocene South American fauna (Scillato-Yané, 1982; Ortiz Jaureguizar, 1998). Dasypodids exhibit fossorial habits and were abundant during the late Miocene in the Pampean region of Argentina, suggesting preference for open environments and well drained soils (Scillato-Yané et al., 2013). Most dasypodids recorded in the Cerro Azul Formation were small- to medium-sized, with body mass in the range 1-10 kg for *Doellotatus*, *Chasicotatus*, *Proeuphractus* and *Chorobates* (Table 1). In particular, the holotype of *Chasicotatus ameghinoi* is a nearly complete carapace about 150 mm wide (Scillato-Yané, Kromptic & Esteban, 2010), which match the lower size range of the small subpopulation. Modern dasypodid burrows are usually simple ramps lacking significant enlargements (e.g., González, Soutullo & Altuna, 2001, Abba, Udrizar & Vizcaíno, 2005), which is similar to the architecture of the fossil burrows.

In the same localities of *Paedotherium*-bearing burrows from the Atlantic coast of the Buenos Aires province, there are also burrows containing articulated remains of *Lagostomus* that partially overlap in diameter with those containing *Paedotherium* remains (Genise, 1989; Elissamburu, Dondas & De Santis, 2011). The extant *Lagostomus maximus* (plains vizcacha) is well known for its digging adaptations and for living in communal burrow systems (e.g., Jackson, Branchi & Villarreal, 1996). Plains vizcacha burrow systems show an average entrance horizontal diameter of 0.26 m and a range of 0.17-0.37 m (Llanos & Crespo, 1952), which matches the range of the small subpopulation. However, extant *L. maximus* burrow systems have several entrance ramps that typically converge into a central chamber or a much more complex architecture (e.g., Llanos & Crespo, 1952; Rafuse et al., 2017), which contrast with the simple ramp type morphology of the fossil burrows. The 43 mm in diameter subcircular burrow identified in the fill of a larger burrow at Salinas Grandes de Hidalgo (# 638) is probably related to a caviomorph rodent (Caviidae, Octodontidae or Chinchillidae).

For the dominant intermediate subpopulation (Dh= 0.39-0.94 m, 83% of measured burrows), with an estimated body mass ranging from 37 to 438 kg, the likely candidates are the Mesotheriinae (Mesotheriidae, Notoungulata); *Eosclerocalyptus*, *Coscinocercus*, and *Aspidocalyptus* (Xenarthra, Glyptodontidae); *Macrochorobates* and *Macroeuphractus* (Xenarthra, Dasypodidae); and *Proscelidodon* (Xenarthra, Mylodontidae). The fossil remains found in this size range that are likely candidates are those of Glyptodontidae and *Proscelidodon* sp. (Table 1). There are two Mesotheriinae species recognized for the late Miocene of central Argentina: *Pseudotypotherium subinsigne* and *Typotheriopsis silveyrai* (Cerdeño & Montalvo, 2001). These species exhibited a small to medium size (20.88 to 60.13 kg after Croft, Flynn & Wyss, 2004) (Table 1). The Mesotheriidae shows modifications in the appendicular skeleton that suggest a scracth-digging habit and fossorial adaptations and are envisaged as having used its hypsodont teeth to cut roots and break the substrate, to aid digging with claws (Shockey et al., 2007).

Kraglievich (1934) and Quintana (1992) suggested that glyptodonts were not functionally suited for digging. However, a geometric morphometric study of the limb bones of five glyptodont species of Miocene and Pleistocene age and comparison with extant armadillos led Vizcaíno et al. (2011) to conclude that were generalized diggers, as modern Dasypodini and Euphractini. Generalized diggers are species that dig short burrows for protection or in search of food and that feed on the surface or just below it by making ‘food probes’ (Abba, Udrizar & Vizcaíno, 2005). In order to asses if glyptodonts were likely producers of the fossil burrows we compared the width of the dorsal carapace and the dorsal carapace height / width ratio with comparable values of the fossil burrows. Dorsal carapace width of Miocene-Pliocene glyptodonts range between 0.40 and 0.77 m (Perea, 2005; Vizcaíno et al., 2011; Zurita et al., 2011), well in the range of horizontal diameter of the fossil burrows. Information on the ratio between carapace height and width for Miocene-Pliocene glyptodonts is incomplete, and similar data for Pleistocene South American glyptodonts (Duarte, 1997; Zurita et al., 2010) average 0.87 (range= 0.78-0.91; n=4). In our case study, glyptodonts are considered good candidates for constructing the subcircular burrows of the intermediate subpopulation, which are 0.39-0.56 m wide and display an average

Dv/Dh ratio of 0.88. Regarding the large dasypodids *Macrochorobates* and *Macroeuphractus*, the available body mass estimates suggest a range of 10 to 100 kg (Vizcaíno & Fariña, 1999) and little is known about their paleoecology.

Among the mylodontids, the Scelidotherinae, endemic to South America (McDonald, 1987;Taglioretti et al., 2014); are only represented for the Huayquerian – Chapadmalian SALMAs (late Miocene to early Pliocene) by *Proscelidodon*, a ground sloth related to open environments with grasslands, under temperate and warm climate (Miño Boilini et al., 2011; Pujos et al., 2012; McDonald & Perea 2002). A digging habit was inferred for *Proscelidodon* after a morphofunctional study of a Montehermosian (latest Miocene-early Pliocene) forelimb (Aramayo, 1988). Body mass estimates are only available for Pleistocene scelidotherines (Table 1), and range from 584 to 1057 kg (De Esteban-Trivigno, Mendoza & De Renzi, 2008; Bargo et al., 2000; Fariña, Vizcaíno & Bargo, 1998). These would be maximum estimates for late Miocene scelidotherines because the primitive Mylodontidae were smaller and there seems to be a trend toward progressively larger sizes in the Pleistocene (e.g., McDonald & Perea 2002). Large Pliocene-Pleistocene fossil burrows near Mar del Plata city (Buenos Aires province) have been attributed to mylodonts on the basis of the finding of bone remains inside the fill (Frenguelli 1955) and using the surface ornamentation of the burrows (Zárate et al., 1998; Dondas, Isla & Carballido, 2009).

For the large subpopulation, with a Dh ranging from 1.05 and 1.50 m (9 % of cases) and an extrapolated body mass of 700-1600 kg, the more likely producer is *Proscelidodon* sp. and, secondarily, the Glyptodontidae.

To summarize, the studied fossil burrows can be attributed to several producers, according to their horizontal diameter. The more likely producers of the studied fossil burrows are: 1) for the small subpopulation, the smaller dasypodids (*Doellotatus*, *Chasicotatus*, *Proeuphractus* and *Chorobates*) on the basis of body mass, the fossorial habit and architecture of modern dasypodid burrows; and, secondarily *Paedotherium minor*. 2) For the intermediate and large subpopulations, the Glyptodontidae and Mylodontidae (*Proscelidodon* sp.) are good candidates as these were the largest representatives of the late Miocene burrowing fauna of the Cerro Azul Formation. The Glyptodontidae were generalized diggers, like modern dasypodids, and exhibited a carapace fitting especially the subcircular burrows. *Proscelidodon* sp. is also a likely candidate of the elliptical and larger burrows. For the intermediate subpopulation, probably the large dasipodids (*Macrochorobates* and *Macroeuphractus*) and Mesotheriinae should be considered.

### *Taphonomy of burrow fills*

The horizontally laminated and massive fill of the burrows suggest that the material entered the excavation without the participation of its producer. The burrows were abandoned and received sediments both in successive small pulses and large catastrophic ones. Although we cannot discard some secondary input of dust by wind, most of the filling of the burrows is related to water transport as suggested by the well laminated and direct graded siltstone to mudstone laminae (Figs. 10A-C, 10F). Laminated intervals are linked to successive pulses of sediment-laden water that eventually ponded in the terminal tracts of the burrows. This is in agreement with the theoretical interpretation by Imbellone, Teruggi & Mormeneo (1990) of similar Quaternary burrows and experimental results by Woodruff & Varricchio (2011). Experiments by Woodruff & Varricchio (2011) indicate that well laminated fills were obtained by adding small amounts of sediment-water mixtures entering at a low angle (5°) into the burrow. In contrast, *en masse* pouring of the sediment-water mixture at high angle (30°) produced a massive and poorly sorted sediment fill, whereas *en masse* pouring at a low angle (5°) produced thicker graded beds

(Woodruff & Varricchio 2011). *En masse* filling experiments also produced “arcuate structures” (Woodruff & Varricchio 2011) that are very similar to the pseudomeniscate structures described herein. Both features are comparable to “arcuate surface ridges” produced in experimental debris flows that reflect the pulsatory nature of experimental and natural debris flows (Major 1997). The experiments by Woodruff & Varricchio (2011) lend further support to the interpretation of the massive intervals as result of catastrophic input of large volume of unsorted sediment. As the burrows are related to an essentially flat landscape and no fluvial channels deposits were observed in any of the studied localities, the sediment pulses should be related with rains. We speculate that one or more laminae may result from moderate to heavy rains. In contrast, massive intervals can be related to heavy rains producing sheet flooding, which can generate hyperconcentrated flows (e.g., Smith & Lowe 1991) capable of transporting enough material to fill, at least, the terminal portion of a burrow in a single event. High-energy sheet floods can also saturate burrow walls and produce roof collapse.

Our studies also support the generalization that well laminated burrow fills will not contain remains of the producer and that massive fills of the whole burrow or most of the lower part have a greater chance of containing remains of the tetrapod that dug the burrow, as proposed by Scognamillo (1993) and Groenewald, Welman & MacEachern (2001). For the laminated burrow fills, the most likely scenario is that the burrow was vacated or, if the animal died inside, it may result scavenged and/or weathered, which produces incomplete and disarticulated remains. In the case of a massive fill, both live entombing (Scognamillo, 1993; Groenewald, Welman & MacEachern, 2001) and fast burial after death (Woodruff & Varricchio 2011) are required to produce an nearly articulated and complete skeleton. Massive fills in the upper half of the burrow will not preserve remains of its producer.

The episodic nature of the filling processes is evidenced by the laminated fill and further supported by the presence of footprints in the surface of some laminae and the cluster of bee cells (?*Rosellichnus* isp.) found inside the fill. These trace fossils suggest that partially filled burrows with a surface communication were explored or reoccupied by other tetrapods and used by bees to nest. Alternatively, the bee cells may be constructed after the complete filling of the burrow in the soil profile. Among the presumed producers of burrows of intermediate size, the outline and size of the footprints match those of Pleistocene glyptodonts, but are quite different from those of ground sloths (compare Aramayo et al., 2015). Disruption of laminae composing the fill of the burrows is explained by drying and cracking of mud after a flood event, whereas synsedimentary faulting can be related to trampling by tetrapods.

*Attaichnus kuenzelli* occur profusely in the SG locality, in some cases very close, but never were cut by a large mammal burrow. These relationships suggest that the producers of the burrows were apparently not foraging on *A. kuenzelli*, considered a nest chamber of leaf-cutting ants (Genise et al., 2013).

### *Paleoecological and paleoenvironmental meaning*

Detailed inferences about the paleoecological and paleoenvironmental meaning of the studied large mammal burrows can be gained through sedimentological study of the hosting rocks, analysis of orientation of burrows fills and considering the associated trace fossils. This information, along with the potential producers will help to understand the reason for acquisition of burrowing habits in large late Miocene mammals.

### Sedimentology

Thick, monotonous, massive continental successions of siltstone showing moderate to good sorting with associated paleosols, as those described for the Cerro Azul Formation, are typical of loess deposits, whose dominantly eolian origin is well established (e.g., Johnson, 1989; Pye, 1995). The presence of pedogenic calcite is indicative of well-drained soil profiles in sub-humid, semi-arid, and arid climates with low rainfall (less than 800 mm/yr) and high evapotranspiration (see review in Sheldon & Tabor 2009). Previous estimation of mean annual precipitation for the development of the paleosols of the Cerro Azul Formation is 449 ± 147 mm (Cardonatto et al., 2016). Paleosols showing a Bt horizon and blocky or prismatic peds can be compared with mollisols (Cardonatto et al., 2016). Some paleoenvironmental constraints can also be derived from the composition of the mammal fauna, and the stable isotopic composition of enamel teeth. Vertebrate remains of the Cerro Azul Formation, mainly notoungulates and rodents, suggest that these sediments were deposited in open landscapes like steppes or herbaceous plains (Montalvo et al., 2008). Carbon isotope composition from late Miocene herbivorous enamel teeth from Salinas Grandes de Hidalgo and nearby localities indicates a dominance of C3 plants in lowland areas (MacFadden, Cerling & Prado, 1996), which are favoured in climates with cool growing season (Ehleringer, Cerling & Helliker, 1997)

### Orientation of burrow fills

Comparison with orientation data from modern Dasypodidae burrows can help to interpret the orientation pattern of fossil burrow fills. As xenarthrans are imperfect homeotherms, their body temperatures do change with the environment (e.g., McNab, 1980; 1985). It has been suggested that the burrow entrance orientation of armadillos avoid prevailing winds and both uniform and preferential orientation has been documented (e.g., McDonough & Loughry, 2008). The cases of no preferential orientation are related to the invasive armadillo *Dasypus novemcinctus* from southern USA (Texas, Alabama, Oklahoma) and Belize (Clark, 1951; Zimmerman, 1990; Platt et al., 2004; Sawyer et al., 2012). All these cases are mostly related to forested areas. Studies documenting a preferred orientation of Dasypodidae burrows are from Argentina, Uruguay and Brazil, involving open environments and several species (Crespo, 1944; Carter & Encarnaçao, 1983; González, Soutullo & Altuna, 2001; Abba, Udrizar & Vizcaíno, 2005; Ceresoli & Fernandez Duque, 2012). The pioneer study by Crespo (1944) included three localities from western Argentina, ranging from 27° 37′ S to 34° 13′ S and including annual precipitation ranging from less than 200 mm to 500 mm. The vegetation ranges from low bushes, to shrubland and psammophilous grassland with sparse trees. These localities belong to the Monte and Espinal biogeographic provinces (e.g., Roig, Roig-Juñent & Corbalán, 2009) and the included armadillo species are: *Chaetophractus vellerosus*, *C. villosus* and *Zaedyus pichy*. A compilation of the entrance orientation data from the three localities of Crespo (1944) suggests a dominant entrance orientation toward the west (Fig. 9D). This distribution is remarkably similar to the fossil burrows if we assume that entrance orientation was at 180° of dipping azimuth (Fig. 9C). Dominant surface wind patterns in northern Argentina are humid and sometimes hot winds from the east and north (e.g., Barros et al., 2015), whereas cold winds are from the south. In consequence, the orientation pattern described by Crespo (1944) from open environments of the semiarid region of Argentina can be interpreted as preferential orientation of entrances avoiding dominant hot and cold winds. Similar patterns of armadillo burrow entrance orientation avoiding prevailing winds were documented by Carter & Encarnaçao (1983) in Minas Gerais, Brazil; González, Soutullo & Altuna (2001) in Uruguay (Fig. 9E); Abba, Udrizar & Vizcaíno (2005) in Buenos Aires province of Argentina (Fig. 9F); and Ceresoli & Fernández Duque (2012) in

Formosa province, northern Argentina. Alternative explanations for this preferential orientation are that, as the armadillos seek food following an odour in the wind, they tend to approach a site from downwind and dig in the lee side (Carter & Encarnaçao, 1983) and to maximize sun exposure during cold winters (Ceresoli & Fernández Duque, 2012). In particular, the most adequate example to evaluate the orientation of the fossil burrows is the data from dasypodid burrows by Crespo (1944), which were collected in open semiarid settings similar to those of the late Miocene of central Argentina. In consequence, it is possible to propose that the late Miocene wind pattern of central Argentina was similar to the present one with hot winds from the east and north and cold winds from the south.

### *Associated trace fossils*

The trace fossil assemblage of the Cerro Azul Formation is of low diversity and abundance and dominated by insect trace fossils (*Celliforma*, *Rosellichnus*, *Fictovichnus*, *Rebuffoichnus* and *Teisseirei*), and was compared with the *Celliforma* ichnofacies (Cardonatto et al., 2016). The *Celliforma* ichnofacies is typical of well-drained calcareous paleosols developed under low vegetation coverage (Genise et al., 2010, 2016). The reduced size of associated rhizoliths suggests that the vegetation was dominated by scrubs with minor participation of herbaceous plants.

The local occurrence of cemented *Coprinisphaera* at LLP and additional occurrences of fossil dung-beetle brood balls (*Quirogaichnus coniunctus* Laza, 2006) from the formation in a nearby locality (Laza, 2006) is indicative of the presence of the *Coprinisphaera* ichnofacies, suggesting herbaceous communities and wetter climatic conditions (Genise et al., 2016) for the easternmost locations of the formation.

### *Burrowing habits in large late Miocene mammals*

Mammals burrows are typically constructed as shelters from environmental extremes and predators, and also for food storage, foraging and reproduction (e.g., Reichman & Smith ,1990; Kinlaw, 1999). From these common uses of burrows, protection from environmental extremes and predators are more likely for the studied fossil burrows and no evidence supporting the remaining functions is available. Top predators during deposition of the Cerro Azul Formation are the Phorusrhacidae (Cenizo, Tambussi & Montalvo, 2012; Vezzosi, 2012) that occupied the role of large carnivorans, as well as the Sparassodonta (Goin, Montalvo & Visconti, 2000). However, the main factor controlling the occurrence of large mammal burrows during the late Miocene (Fig. 1) is herein related to environmental changes. It has been suggested that different mammal groups acquired fossorial habits during the Cenozoic as a response to the expansion of open, savanna-like environments under cold, dry and seasonal climates (Nevo, 1979; 1995; 2011). During the late Miocene (the Huayquerian SALMA), southern South America experienced a global cooling as response to the increase in the Antarctic ice sheet (Zachos et al., 2001) and the uplift of the Andes (e.g., Amidon et al., 2017), which favored cold and seasonally dry climatic conditions. This regional framework is confirmed by the inferences on the sedimentology, faunal remains and invertebrate ichnology of the studied succession. This evidence suggests open environments, with well-drained soils and low vegetation coverage, and a semiarid and seasonal climate. Considering that the more likely candidates for the largest burrows are xenarthrans (Glyptodontidae and Mylodontidae), which are imperfect homeotherms (e.g., McNab, 1980; 1985), the necessity and convenience for excavating an underground refuge is clear. In addition to escape from predation, these animals used burrows to avoid extremely cold or warm climatic conditions in order to conserve energy and water, and to breed because of the particular physiology of xenarthrans (Vizcaíno et al., 2001).

## Conclusions

Subcylindrical structures with a laminated fill occurring in a late Miocene loess-paleosol sequence from central Argentina are interpreted as the burrow fills of fossorial mammals. The burrows occur associated with calcareous paleosols developed under a semiarid climate in a flat landscape. A detailed analysis of more than one hundred structures permitted to conclude that:

1. The more common geometry is a shallow ramp with a rounded end, lacking bifurcations.

2. The ample variation of the horizontal diameter of the burrow fills, along with cross-sectional shape and inferred body masses suggest that several producers were involved.

3. The smaller burrows (Dh= 0.15-0.34 m, 8 % of cases, body mass ranging from 1 to 13 kg) were produced by small dasypodids and, secondarily, by the notoungulate *Paedotherium minor*.

3. For the dominant burrow fills exhibiting an intermediate (Dh= 0.39-0.94 m, 83% of measured burrows, producer body mass of 37 to 438 kg), and large horizontal diameter (Dh= 1.05-1.50 m, 9 % of measured burrows, producer body mass of 700-1600 kg), the Glyptodontidae and Mylodontidae (*Proscelidodon* sp.) are the best candidates. The Glyptodontidae are related to the subcircular burrows of intermediate size and *Proscelidodon* sp. would be the producer of the elliptical and largest burrows.

4. About 10% of the burrow fills contained fragmentary, disarticulated, abraded and weathered bone remains that were introduced into the burrows by aqueous currents and do not belong to the excavator of the burrow.

5. After abandonment, the burrows received sediments from the surface that progressively filled the structure. The filling process was produced dominantly by several pulses of sediment laden currents related to moderate rains (well laminated fill) and *en masse* input of dense sediment-laden currents related to heavy rains producing sheet flooding (massive fill). During filling, the abandoned burrows were visited or reoccupied by other tetrapods and used by bees to nest,

6. It is not expected that well laminated burrow fills will not contain remains of the producer, whereas massive fills have a greater chance of containing remains of the tetrapod that dug the burrow.
7. The main factor explaining the appearance of large mammal burrows in the late Miocene are environmental changes, including open environments with low vegetation and semiarid and seasonal climate.

## Acknowledgements

We are indebted to Fátima Mendoza Belmontes for help during field work, Claudia Montalvo for guidance with the taxonomy of vertebrate remains and for appropriate comments on an earlier version of the manuscript; Marcelo Zárate for noticing about the burrow fills of the LLP locality; and María F. Vera Candioti for his help in morphometric analysis.

## References

Abba AM, Udrizar DE, Vizcaíno, SF. 2005. Distribution and use of burrows and tunnels of *Chaetophractus villosus* (Mammalia, Xenarthra) in the eastern Argentinean pampas. Acta Theriologica, 50: 115-124.

Amidon WH, Fisher GB, Burbank DW, Ciccioli PL, Alonso RN, Gorin AL, Silverhart PH, Kylander-Clark ARC, Christoffersen MS. 2017. Mio-Pliocene aridity in the south-central Andes associated with Southern Hemisphere cold periods. Proceedings of the National Academy of Sciences, 114: 6474-6479.

Aramayo SA. 1988. Nuevos restos de *Proscelidodon* sp. (Edentata, Mylodontidae) del Yacimiento Monte Hermoso (Plioceno inferior a medio) provincia de Buenos Aires, Argentina. Estudio morfológico funcional. Actas Segundas Jornadas Geológicas Bonaerenses: 99-107.

Barbour EH. 1892. Notice of new gigantic fossils. Science N.S. 19: 99-100.

Bargo MS, Vizcaíno SF, Archuby FM, Blanco, RE. 2000. Limb bone proportions, strength and digging in some Lujanian (Late Pleistocene-Early Holocene) mylodontid ground sloths (Mammalia, Xenarthra). Journal of Vertebrate Paleontology, 20: 601-610.

Barros VR, Boninsegna JA, Camilloni IA, Chidiak M, Magrín GO, Rusticucci M. 2015. Climate change in Argentina: trends, projections, impacts and adaptation. Wiley Interdisciplinary Reviews: Climate Change, 6: 151-169.

Benton MJ. 1988. Burrowing by vertebrates. Nature, 331: 17-18.

Bown TM. 1982. Ichnofossils and rhizoliths of the nearshore fluvial Jebel Qatrani Formation (Oligocene), Fayum Province, Egypt. Palaeogeography, Palaeoclimatology, Palaeoecology, 40: 255-309.

Bown TM, Kraus MJ. 1983. Ichnofossils of the alluvial Willwood Formation (lower Eocene), Bighorn Basin, northwest Wyoming, U.S.A. Palaeogeography, Palaeoclimatology, Palaeoecology, 43: 95-128.

Buchmann FS, Pereira Lopez R, Caron F. 2009. Icnofósseis (paleotocas e crotovinas) atribuídos a mamíferos extintos no sudeste e sul do Brasil. Revista brasileira paleontological, 12: 247-256.

Cabrera A. 1939. Sobre vertebrados fósiles del Plioceno de Adolfo Alsina. Rev. Mus. de La Plata (Nueva Serie), 2: 3-35.

Cardonatto MC, Sostillo R, Visconti G, Melchor RN. 2016. The *Celliforma* ichnofacies in calcareous paleosols: An example from the late Miocene Cerro Azul Formation, La Pampa, Argentina. Palaeogeography, Palaeoclimatology, Palaeoecology, 443: 203-215.

Carter TS, Encarnaçao CD. 1983. Characteristics and Use of Burrows by Four Species of Armadillos in Brazil. Journal of Mammalogy, 64: 103-108.

Cenizo MM, Tambussi CP, Montalvo CI. 2012. Late Miocene continental birds from the Cerro Azul Formation in the Pampean region (central-southern Argentina). Alcheringa, 36: 47-68.

Cerdeño E, Montalvo CI. 2001. Los Mesotheriinae (Mesotheriidae, Notoungulata) del Mioceno superior de La Pampa, Argentina. Revista Española de Paleontología, 16: 63-75.

Ceresoli N, Fernandez-Duque E. 2012. Size and orientation of giant armadillo burrow entrances (*Priodontes maximus*) in western Formosa Province, Argentina. Edentata, 13: 66-68.

Cione AL, Azpelicueta, MM, Bond M, Carlini AA, Casciotta JR, Cozzuol MA, de la Fuente M, Gasparini Z, Goin FJ, Noriega J, Scillato-Yané GJ, Soibelzon L, Tonni EP, Verzi D, Vucetich MG. 2000. Miocene vertebrates from Entre Ríos province, eastern Argentina. En: F.G. Aceñolaza y R. Herbst (Eds.), El Neógeno de Argentina. Serie Correlación Geológica, 14: 191–237.

Clark WK. 1951. Ecological life history of the armadillo in the eastern Edwards Plateau region. American Midland Naturalist, 46: 337–58.

Crespo J. 1944. Contribución al conocimiento de la ecología de algunos dasipódidos (Edentata) Argentinos. Revista Argentina de Zoogeografía, Buenos Aires, 4: 7-39.

Croft DA, Flynn JJ, Wyss, A. 2004. Notoungulata and Litopterna of the early Miocene Chucal Fauna, northern Chile. Fieldiana, Geology, n.s. 50, 1-52.

Damiani R, Modesto S, Yates A, Neveling J. 2003. Earliest evidence of cynodont burrowing. Proceedings of the Royal Society of London B: Biological Sciences, 270: 1747-1751.

De Esteban-Trivigno S, Mendoza M, De Renzi M . 2008. Body mass estimation in Xenarthra: A predictive equation suitable for all quadrupedal terrestrial placentals? Journal of Morphology, 269: 1276-1293.

De Elorriaga EE, Visconti G. 2002. Crotovinas atribuibles a grandes mamíferos del Cenozoico en el sureste de la Provincia de La Pampa. 9° Reunión Argentina de Sedimentología, Resúmenes, Asociación Argentina de Sedimentología, Córdoba: 63.

Delsuc F, Superina M, Tilak M, Douzery EJP, Hassanin A. 2012. Molecular phylogenetics unveils the ancient evolutionary origins of the enigmatic fairy armadillos. Molecular Phylogenetics and Evolution, 62: 673-80

Dondas, A., Isla, F.I. y Carballido, J.L., 2009. Paleocaves exhumed from the Miramar Formation (Ensenadan Stage-age, Pleistocene), Mar del Plata, Argentina. Quaternary International, 210: 44-50.

Duarte RG. 1997. Gliptodontes del Pleistoceno tardío de Agua de las Palomas, Campo del Pucará, Catamarca, Argentina. Variaciones morfológicas del caparazón de *Glyptodon reticulatus* Owen, 1845. Ameghiniana, 34: 345-355.

Ehleringer JR, Cerling TE, Helliker BR. 1997. C4 photosynthesis, atmospheric CO2, and climate. Oecologia, 112: 285–299.

Elissamburu A. 2004. Análisis morfométrico y morfofuncional del esqueleto apendicular de *Paedotherium* (Mammalia, Notoungulata). Ameghiniana, 41: 363-380.

Elissamburu A, Dondas A, De Santis L. 2011. Morfometría de las paleocuevas de la “Fm.” Chapadmalal y su asignación a *Actenomys* (Rodentia), *Paedotherium* (Notoungulata) y otros mamíferos fósiles hospedantes. Mastozoología Neotropical, 18: 227-238.

Fariña RA, Vizcaíno SF, Bargo MS. 1998. Body mass estimations in Lujanian (late Pleistocene–early Holocene of South America) mammal megafauna. Mastozoología Neotropical, 5: 87–108.

Fernández ME, Vassallo AI, Zárate M. 2000. Functional morphology and paleobiology of the Pliocene rodent *Actenomys* (Caviomorpha: Octodontidae): the evolution to a subterranean mode of life. Biological Journal of the Linnean Society, 71:71–90.

Frank HT, Buchmann FS, Gonçalves de Lima L, Fornari M, Caron F. 2012. Cenozoic vertebrate tunnels in southern Brazil. In: Netto RG, Carmona NB, Tognoli, FM (Eds.), Ichnology of Latin America: selected papers, Monografías da Sociedade Brasileira de Paleontología, 2: 141-158.

Frank HT, Althaus CE, Dario EM, Tramontina FR, Adriano RM, Almeida ML, Ferreira GF, Nogueira R, Breier R. 2015. Underground chamber systems excavated by Cenozoic ground sloths in the state of Rio Grande do Sul, Brazil. Revista Brasileira de Paleontologia, 18: 273-284.

Frenguelli, J. 1955. Loess y limos pampeanos. Universidad Nacional de La Plata, Facultad de Ciencias Naturales y Museo, Serie técnica y didáctica, 7:1-88.

Folguera A, Zárate M. 2009. La sedimentación neógena continental en el sector extrandino de Argentina central. Revista de la Asociación Geológica Argentina, 64: 692-712.

Gee CT, Sander PM, Petzelberger BEM. 2003. A Miocene rodent nut cache in coastal dunes of the Lower Rhine Embayment, Germany. Palaeontology, 46: 1133-1149.

Genise J. 1989. Las cuevas de *Actenomys* (Rodentia: Octodontidae) de la Formación Chapadmalal (PIioceno superior) de Mar del Plata y Miramar (Provincia de Buenos Aires). Ameghiniana, 26: 33-42.

Genise JF, Farina JL. 2012. Ants and xenarthrans involved in a Quaternary food web from Argentina as reflected by their fossil nests and palaeocaves. Lethaia, 45: 411-422.

Genise JF, Melchor RN, Bellosi ES, Verde M. 2010. Invertebrate and vertebrate trace fossils from continental carbonates. In: AM Alonso-Zarza, Tanner L, Carbonates in continental settings: facies, environments and processes, Developments in Sedimentology, 61: 319–369.

Genise JF, Melchor RN, Sánchez MV, González MG. 2013. *Attaichnus kuenzelii* revisited: A Miocene record of fungus-growing ants from Argentina. Palaeogeography, Palaeoclimatology, Palaeoecology, 386: 349-363.

Genise JF, Bedatou E, Bellosi E, Sarzetti LC, Sánchez MV, Krause JM. 2016. The Phanerozoic four revolutions and evolution of paleosol ichnofacies. In: LA Buatois, Mángano MG, The trace-fossil record of major evolutionary events. p. 301-370.

Gobetz KE, Martin LD. 2006. Burrows of a gopher-like rodent, possibly *Gregorymys* (Geomyoidea: Geomyidae: Entoptychtinae), from the early Miocene Harrison Formation, Nebraska. Palaeogeography, Palaeoclimatology, Palaeoecology, 237: 305-314.

Goin FJ, Montalvo CI, Visconti G. 2000. Los marsupiales (Mammalia) del Mioceno Superior de la Formación Cerro Azul (Provincia de La Pampa, Argentina). Estudios Geológicos, 56: 101–126.

González EM, Soutullo A, Altuna CA. 2001. The burrow of *Dasypus hybridus* (Cingulata: Dasypodidae). Acta Theriologica, 46: 53-59.

Groenewald GH, Welman J, MacEachern JA. 2001. Vertebrate burrow complexes from the Early Triassic *Cynognathus* Zone (Driekoppen Formation, Beaufort Group) of the Karoo Basin, South Africa. Palaios, 16: 148-160.

Hembree DI, Martin LD, Hasiotis ST. 2004. Amphibian burrows and ephemeral ponds of the Lower Permian Speiser Shale, Kansas: evidence for seasonality in the midcontinent. Palaeogeography, Palaeoclimatology, Palaeoecology, 203: 127-152.

Hembree DI, Hasiotis ST. 2008. Miocene vertebrate and invertebrate burrows defining compound paleosols in the Pawnee Creek Formation, Colorado, U.S.A. Palaeogeography, Palaeoclimatology, Palaeoecology, 270: 349-365.

Hunt RM. 1990. Taphonomy and sedimentology of Arikaree (lower Miocene) fluvial, eolian, and lacustrine paleoenvironments, Nebraska and Wyoming: A paleobiota entombed in fine-grained volcaniclastic rocks. In: Lockley MG, Rice A, Volcanism and fossil biotas. Geological Society of America, Special Paper, 244: 69-111.

Hunt RMJ, Xiang-Xu XUE, Kaufman J. 1983. Miocene burrows of extinct bear dogs: Indication of early denning behavior of large mammalian carnivores. Science, 221: 364-366.

Imbellone, P., Teruggi, M., Mormeneo, L. 1990. Crotovinas en sedimentos cuaternarios del partido de la Plata. In: Zárate, M. (Ed.) Simposio Internacional sobre Loess, pp. 166-172. Mar del Plata.

Jackson JE, Branch LC, Villarreal D. 1996. *Lagostomus maximus*. Mammalian Species, 543: 1-6.

Johnson SY. 1989. Significance of loessite in the Maroon Formation (Middle Pennsylvanian to Lower Permian), Eagle Basin, Northwest Colorado. Journal of Sedimentary Research, 59: 782–791.

Kinlaw A. 1999. A review of burrowing by semi-fossorial vertebrates in arid environments. Journal of Arid Environments, 41: 127–145.

Kraglievich L. 1934. La antigüedad pliocena de las faunas de Monte Hermoso y Chapadmalal, deducidas de su comparación con las que le precedieron y sucedieron. Imprenta "El Siglo Ilustrado", 136 p.

Krummeck WD, Bordy EM. 2017. *Reniformichnus katikatii* (new ichnogenus and ichnospecies): Continental vertebrate burrows from the Lower Triassic, Main Karoo Basin, South Africa. Ichnos, 24: 1-12 (in press)

Laza, JH. 2006. Dung-beetle fossil brood balls: The ichnogenera *Coprinisphaera* Sauer and *Quirogaichnus* (Coprinisphaeridae). Ichnos, 13: 217–235.

Linares E, Llambías E, Latorre C. 1980. Geología de la Provincia de La Pampa, República Argentina y Geocronología de sus rocas metamórficas y eruptivas. Revista de la Asociación Geológica Argentina, 35: 87-146.

Liu J, Li L. 2013. Large Tetrapod Burrows from the Permian Naobaogou Formation of the Daqingshan Area, Nei Mongol, China. Acta Geologica Sinica - English Edition, 87: 1501-1507.

Llanos AC, Crespo JA. 1952. Ecología de la vizcacha (*Lagostomus maximus maximus* Blainv.) en el nordeste de la provincia de Entre Ríos. Revista de Investigaciones Agrícolas, 6:289-378.

Loope DB. 2006. Burrows dug by large vertebrates into rain-moistened Middle Jurassic sand dunes. Journal of Geology. 114: 753-762.

Loope DB. 2008. Life beneath the surfaces of active Jurassic dunes: Burrows from the Entrada Sandstone of south-central Utah. Palaios, 23: 411-419.

Lorenzo FR, Mehl AE, Zárate MA. 2013. Sedimentología y estratigrafía de depósitos del Mioceno tardío de los valles transversales de la provincia de La Pampa, Argentina. Latin American Journal of Sedimentology and Basin Analysis, 20: 67-84.

Lucas SG, Gobetz KE, Odier GP, McCormick T, Egan C. 2006. Tetrapod burrows from the Lower Jurassic Navajo Sandstones, southeastern Utah. In: JD Harris, Lucas SG, Spielmann JA, Lockley MG, Milner ARC, Kirkland JI (Eds), The Triassic-Jurassic Terrestrial Transition New Mexico Museum of Natural History and Science Bulletin, 37: 147-154.

MacFadden BJ, Cerling TE, Prado J. 1996. Cenozoic terrestrial ecosystem evolution in Argentina: Evidence from carbon isotopes of fossil mammal teeth. Palaios, 11: 319–327.

Major, JJ. 1997. Depositional processes in large-scale debris-flow experiments. Journal of Geology, 106: 345-366.

Martin LD, Bennett DK. 1977. The burrows of the Miocene beaver *Palaeocastor*, Western Nebraska, U.S.A. Palaeogeography, Palaeoclimatology, Palaeoecology,. 22: 173-193.

McDonald HG. 1987. A systematic review of the Plio-Pleistocene Scelidotherinae Ground Sloth (Mammalia: Xenarthra: Mylodontidae). PhD Tesis, University of Toronto, Toronto, Canada, 478 p. Unpublished.

McDonald HG, Perea D. 2002. The large Scelidothere *Catonyx tarijensis* (Xenarthra, Mylodontidae) from the Pleistocene of Uruguay. Journal of Vertebrate Paleontology, 22: 677-683.

McDonough CM, Loughry WJ. 2008. Behavioral ecology of armadillos. In: SF Vizcaíno, Loughry WJ (Eds), The Biology of Xenarthra, p. 281-293. Gainesville.

Macleod N .2009. Who is Procrustes and what has he done with my data. Palaeontological Association Newsletter, 70: 21-36.

McNab BK. 1980. Energetics and the limits to a temperate distribution in armadillos. Journal of Mammalogy, 61: 606-27.

McNab BK.1985. Energetics, population biology, and distribution of xenarthrans, living and extinct. In G G Montgomery (Ed), The evolution and ecology of armadillos, sloths and vermilinguas, p. 219-232. Smithsonian Institution Press, Washington, D.C.

Melchor RN, Casadío S. 2000. Descripción Geológica de la Hoja 3766-III “La Reforma” (1:250.000), Provincia de la Pampa. Servicio Geológico Minero Argentino, Boletín, 295: 1-56.

Melchor RN, Bedatou E, de Valais S, Genise JF. 2006. Lithofacies distribution of invertebrate and vertebrate trace-fossil assemblages in an Early Mesozoic ephemeral fluvio-lacustrine system from Argentina: Implications for the *Scoyenia* ichnofacies. Palaeogeography, Palaeoclimatology, Palaeoecology, 239: 253-285.

Melchor RN, Genise JF, Umazano AM, Superina M. 2012. Pink fairy armadillo meniscate burrows and ichnofabrics from Miocene and Holocene interdune deposits of Argentina: Palaeoenvironmental and palaeoecological significance. Palaeogeography, Palaeoclimatology, Palaeoecology, 350-352: 149-170.

Melchor RN, Loope D. 2016. Communal therapsid burrows from interdune facies of the Lower Jurassic Navajo Sandstone, southern Utah, USA. Abstract book Ichnia: 174.

Miño Boilini AR, Tomassini RL, Oliva C, Manera de Bianco T. 2011. Adiciones al conocimiento de *Proscelidodon* Bordas, 1935 (Mammalia, Xenarthra, Scelidotheriinae). Revista Brasileira de Paleontologia, 14: 269-278.

Montalvo CI, Casadío S. 1988. Presencia del género *Palaeoctodon* (Rodentia, Octodontidae) en el Huayqueriense (Mioceno tardío) de la Provincia de La Pampa. Ameghiniana, 25: 111-114.

Montalvo CI, Visconti G, Púgener L, Cardonatto MC. 1995. Mamíferos de Edad Huayqueriense (Mioceno tardío), Laguna Chillhué, provincia de La Pampa. 4° Jornadas Geológicas y Geofísicas Bonaerenses (Junín), Actas 1: 73-79.

Montalvo CI, Melchor RN, Visconti G, Cerdeño E. 2008. Vertebrate taphonomy in loess-palaeosol deposits: A case study from the late Miocene of central Argentina. Geobios, 41: 133–143.

Montalvo CI, Tomassini RL, Sostillo R. 2016. Leftover prey remains: a new taphonomic mode from the late Miocene (Cerro Azul Formation) in central Argentina. Lethaia, 49: 219-230.

Modesto SP, Botha Brink J. 2010. A burrow cast with *Lystrosaurus* skeletal remains from the Lower Triassic of South Africa. Palaios, 25: 274-281.

Nevo E. 1979. Adaptive convergence and divergence of subterranean mammals. Annual Review of Ecology and Systematics, 10: 269-308.

Nevo E. 1995. Mammalian evolution underground. The ecological-genetic-fenetic interfaces. Acta Theriologica, 3: 9-31.

Nevo E. 2011. Evolution under environmental stress at macro- and microscales. Genome Biology and Evolution, 3: 1039-1052.

Ortiz Jaureguizar, E. 1998. Paleoecología y evolución de la fauna de mamíferos de América del Sur durante la "Edad de las Planicies Australes" (Mioceno superior-Plioceno superior). Estudios Geológicos, 54: 161-169.

Perea D. 2005. *Pseudoplohophorus absolutus* n. sp. (Xenarthra, Glyptodontidae), variabilidad en Sclerocalyptinae y redefinición de una biozona del Mioceno Superior de Uruguay. Ameghiniana, 42: 175-190.

Perea, D. Scillato-Yané, G.J. 1995. *Proeuphractus limpidus* Ameghino, 1886 (Xenarthra, Dasypodidae, Euphractini): Osteología comparada del cráneo y elementos de la coraza asociados (Neógeno del Uruguay). Boletín de la Real Sociedad Española de Historia Natural, 90: 125–130.

Platt SG, Rainwater TR, Brewer SW. 2004. Aspects of the burrowing ecology of nine-banded armadillos in northern Belize. Mammalian Biology, 69: 217-224.

Pujos F, Candela A, Galli CI, Coira BL, Reguero MA, de los Reyes M, Abello MA. 2012. The Scelidotheriine *Proscelidodon* (Xenarthra: Mylodontidae) from the Late Miocene of Maimará (Northwestern Argentina, Jujuy Province). Ameghiniana, 49: 668-674.

Pye K. 1995. The nature, origin and accumulation of loess. Quaternary Science Reviews, 14: 653–667.

Quintana CA. 1992. Estructura interna de una paleocueva, posiblemente de un Dasypodidae (Mammalia, Edentata) del Pleistoceno de Mar del Plata (Provincia de Buenos Aires, Argentina). Ameghiniana, 29: 87-91.

Rafuse DJ, Kaufmann CA, Gutiérrez MA, González ME, Scheifler NA, Álvarez MC, Massigoge A. 2017. Taphonomy of modern communal burrow systems of the Plains vizcacha (*Lagostomus maximus*, Chinchillidae) in the Pampas region of Argentina: implications for the fossil record. Historical Biology:1-17 (in press).

Reichman OJ, Smith SC. 1990. Burrows and burrowing behavior by mammals. Current Mammalogy, 2: 197–244.

Riese DJ, Hasiotis ST, Odier GP. 2011. Synapsid burrows and associated trace fossils in the lower Jurassic Navajo Sandstone, Southeastern Utah, U.S.A., indicates a diverse community living in a wet desert ecosystem. Journal of Sedimentary Research, 81: 299-325.

Roig FA, Roig-Juñent S, Corbalán V. 2009. Biogeography of the Monte Desert. Journal of Arid Environments, 73: 164-172.

Rovereto C. 1914. Los estratos araucanos y sus fósiles. Anales del Museo Nacional de Historia Natural de Buenos Aires, 25: 1-247.

Sawyer CF, Brinkman DC, Walker VD, Covington TD, Stienstraw EA. 2012. The zoogeomorphic characteristics of burrows and burrowing by nine-banded armadillos (*Dasypus novemcinctus*). Geomorphology, 157: 122-130.

Scognamillo DG. 1993. Estructura de las cuevas de *Actenomys* (Rodentia: Octodontidae) de la Aloformación Playa San Carlos, Plioceno tardío (Barranca de los Lobos, Pdo. Gral. Pueyrredón): significado paleoecológico y estratigráfico. Unpublished Licenciate thesis, National University of Mar del Plata, Argentina, 32 pp.

Scillato-Yané GJ. 1982. Los Dasypodidae (Mammalia-Edentata) del Plioceno y Pleistoceno de Argentina. Tesis Doctoral. Facultad de Ciencias Naturales y Museo, Universidad Nacional de La Plata, 159 pp. Inédita.

Scillato-Yané GJ, Krmpotic CM, Esteban GI. 2010. The species of genus *Chasicotatus* Scillato-Yané (Eutatini, Dasypodidae). Revista Mexicana de Ciencias Geológicas, 27: 43-55.

Scillato -Yané G, Góis F, Zurita AE, Carlini A, González Ruiz L, Krmpotic C, Oliva C, Zamorano M. 2013. Los Cingulata (Mammalia, Xenarthra) del “Conglomerado osífero” (Mioceno Tardío) de la Formación Ituzaingó de Entre Ríos, Argentina. Asociación Paleontológica Argentina, Publicación Especial, 14: 118-134.

Seilacher A. 2007. Trace Fossil Analysis. Berlin: Springer, 226 p.

Sheldon ND, Tabor NJ. 2009. Quantitative paleoenvironmental and paleoclimatic reconstruction using paleosols. Earth-Science Reviews, 95: 1-52.

Shockey BJ, Croft DA, Anaya F. 2007. Analysis of function in the absence of extant functional homologues: a case study using mesotheriid notoungulates (Mammalia). Paleobiology, 33: 227-247.

Sidor CA, Miller MF, Isbell JL. 2008. Tetrapod burrows from the Triassic of Antarctica. Journal Vertebrate Paleontology, 28: 277-284.

Smith R. 1987. Helical burrow casts of therapsid origin from the Beaufort Group (Permian) of South Africa. Palaeogeography, Palaeoclimatology, Palaeoecology, 60: 155-170.

Soibelzon LH, Pomi LH, Tonni EP, Rodriguez S, Dondas A. 2009. First report of a South American short-faced bears' den (*Arctotherium angustidens*): palaeobiological and palaeoecological implications. Alcheringa, 33: 211-222.

Taglioretti M, Miño-Boilini AR, Scaglia F, Dondas A. 2014. Presencia de *Proscelidodon patrius* (Xenarthra, Scelidotheriinae) en la Formación Chapadmalal (Plioceno superior), Mar del Plata, Buenos Aires, Argentina: Implicancias bioestratigráficas. Ameghiniana, 51: 420-427.

Torres MR, Borghi CE, Giannoni SM, Pattini, A. 2003. Portal orientation and architecture of burrows in *Tympanoctomys barrerae* (Rodentia, Octodontidae). Journal of Mammalogy, 84: 541-546.

Urrutia JJ, Montalvo CI, Scilato-Yané GJ. 2008. Dasypodidae (Xenarthra, Cingulata) de la Formación Cerro Azul (Mioceno tardío) de la provincia de La Pampa, Argentina. Ameghiniana, 45: 289-302.

Varricchio DJ, Martin AJ, Katsura Y. 2007. First trace and body fossil evidence of a burrowing, denning dinosaur. Proceedings of the Royal Society B: Biological Sciences, 274: 1361–1368.

Verzi DH, Montalvo CI, Vucetich MG. 1999. Afinidades y significado evolutivo de *Neophanomys biplicatus* (Rodentia, Octodontidae) del Mioceno tardío-Plioceno temprano de Argentina. Ameghiniana, 36: 83-90.

Verzi DH, Montalvo CI, Tiranti SI. 2003. Un nuevo *Xenodontomys* (Rodentia, Octodontidae) del Mioceno tardío de La Pampa, Argentina. Patrón evolutivo y bioestratigrafía. Ameghiniana, 40: 229-238.

Verzi DH, Montalvo CI, Deschamps C. 2008. Biostratigraphy and biochronology of the Late Miocene of central Argentina: evidence from rodents and taphonomy. Geobios, 41: 145-155.

Vezzosi RI. 2012. First record of *Procariama simplex* Rovereto, 1914 (Phorusrhacidae, Psilopterinae) in the Cerro Azul Formation (upper Miocene) of La Pampa Province; remarks on its anatomy, palaeogeography and chronological range. Alcheringa, 36: 157-169.

Visconti G. 2007. Sedimentología de la Formación Cerro Azul (Mioceno superior) de la provincia de La Pampa. Tesis doctoral inédita. Biblioteca de la Facultad de Ciencias Exactas y Naturales de la UBA “Luis Federico Leloir”. 203 p.

Visconti G, Melchor RN, Montalvo CI, Umazano AM, de Elorriaga EE. 2010. Análisis litoestratigráfico de la Formación Cerro Azul (Mioceno superior) en la provincia de La Pampa. Revista de la Asociación Geológica Argentina, 67: 257-265.

Vizcaíno SF, Fariña RA. 1999. On the flight capabilities and distribution of the giant Miocene bird *Argentavis magnificens* (Teratornithidae). Lethaia, 32: 271-278.

Vizcaíno SF, Zárate M, Bargo MS, Dondas A. 2001. Pleistocene burrows in the Mar del Plata area (Argentina) and their probable builders. Acta Palaeontologica Polonica, 46 (2): 289-301.

Vizcaíno SF, Blanco RE, Bender JB, Milne N. 2011. Proportions and function of the limbs of glyptodonts. Lethaia, 44: 93-101.

Vogt T, Carballo OC, Calmels AP. 1999. Un esbozo de explicación de la génesis de las toscas de la meseta de La Pampa. 1° Congreso Argentino de Cuaternario y Geomorfología, Actas: 195-205.

Voorhies MR. 1975. Vertebrate burrows. In: RW Frey, The study of trace fossils. p. 325-350. New York.

Woodruff DC, Varricchio DJ. 2011. Experimental modeling of a possible Oryctodromeus cubicularis (Dinosauria) burrow. Palaios, 26: 140-151.

Wu NC, Alton LA, Clemente CJ, Kearney MR, White CR. 2015. Morphology and burrowing energetics of semi-fossorial skinks (*Liopholis* spp.). The Journal of Experimental Biology, 218: 2416-2426.

Zachos J, Pagani M, Sloan L, Thomas E, Billups K. 2001. Trends, Rhythms, and Aberrations in Global Climate 65 Ma to Present. Science, 292: 686-693..

Zárate, M.A., Bargo, M.S., Vizcaíno, S.F., Dondas, A. y Scaglia, O., 1998. Estructuras biogénicas en el Cenozoico tardío de Mar del Plata (Argentina) atribuibles a grandes mamíferos. Revista Asociación Argentina de Sedimentología, 5: 95-103.

Zimmerman JW. 1990. Burrow characteristics of the nine-banded armadillo, *Dasypus novemcinctus*. Southwestern Naturalist, 35: 226-27.

Zurita AE, Soibelzon LH, Soibelzon E, Gasparini GM, Cenizo MM, Arzani H. 2010. Accessory protection structures in *Glyptodon* Owen (Xenarthra, Cingulata, Glyptodontidae). Annales de Paléontologie, 96: 1-11.

Zurita AE, Oliva C, Dondas A, Soibelzon E, Isla FI. 2011. El registro más completo de un Hoplophorini (Xenarthra: Glyptodontidae) para los Pisos/Edades Chapadmalalense-Marplatense (Plioceno tardío-Pleistoceno temprano). Revista del Museo Argentino de Ciencias Naturales,13: 69-75.

**Legend of figures**

Figure 1. Plot of the average horizontal diameter (Dh) of tetrapod burrows against the age of the hosting rock, distinguishing between burrows with remains of the producer, those lacking bone remains and the present study. Note that the age axis contains two gaps in the Cretaceous and Paleocene. Source of information on Supplemental Article S1.

Figure 2. Study area in central Argentina (inset) and localities in La Pampa and Buenos Aires provinces. LC: Laguna Chillhué, SG: Salinas Grandes de Hidalgo, LLP: Laguna La Paraguaya.

Figure 3. Location of landmarks (red points) on burrow fills preserved in cross section.

Figure 4. Detailed sedimentary logs of the measured sections.

Figure 5. Abundance and cross-cutting relationships of burrow fills, from LLP locality. A) General view of the outcrop showing several burrow fills (yellow arrows). B-C) Field view and diagram of cross-cutting relationships between different burrow fills (distinguished in C with different shades of gray).

Figure 6. Histogram showing the frequency distribution of horizontal diameter (Dh) for the studied burrow fills. Three subpopulations can be distinguished.

Figure 7. Burrow fill morphology in plan view. A-B) Field photograph and diagram of slightly curved fill. C-D) Field photograph and diagram of sinuous fill. E-F) Field photograph and diagram of “C” shaped fill. G-H) Field photograph and diagram of burrow fill with slightly enlarged and rounded end. Curved lines in the diagrams represent weathered laminae.

Figure 8. Results of geometric morphometric analysis of burrow fills preserved in cross section. A) Plot of principal components 1 and 2, distinguishing by study locality, and deformation grids for elliptical (PC1 score -0.12) and subcircular (PC1 score 0.17) fills. B) Histogram of variance of principal components.

Figure 9. Orientation of fossil burrow fills compared with modern data from dasypodid burrows. A) Rose diagram showing the dip azimuth of fossil burrow fills. B) Equal area projection of dip azimuth and dip angle of fossil burrow fills. C) Entrance orientation of fossil burrows, assuming that it is located at 180° of measured dip azimuth. The data in A-C is from SG and LLP localities, those from the remaining locality are preserved only in cross-section. D) Entrance orientation of several species of dasypodid burrows from semiarid settings of northern-central Argentina by Crespo (1944). E) Entrance orientation of *Dasypus hibridus* burrows from a grassland setting in Uruguay by González, Soutullo & Altuna (2001). F) Entrance orientation of

*Chaetophractus villosus* burrows from cultivated land in Buenos Aires province, Argentina by Abba, Udrizar & Vizcaíno (2005).

Figure 10. Features of fossil burrow fills. A-B) Detailed sedimentary logs of the fill of selected burrows (see Table S1 for information on these burrow fills). References as for Fig. 4. C) Cross-section of laminated to massive burrow fill # 648 from SG, represented in (B). Elliptical outline of fill indicated by yellow arrows, black arrow point to subvertical rhizolith cutting the laminated fill. D) Cross-section of burrow fill # 714 from LLP. Subcircular outline of poorly laminated to massive burrow fill indicated by yellow arrows. E) Pseudomeniscate structure in burrow fill # 704 from LLP seen in plan view. Yellow arrows point the outline of the burrow fill and black arrows to individual pseudomenisci. F) Detail of laminated burrow fill (# 632 from SG) showing direct grading in siltstone to claystone laminae (yellow arrows), synsedimentary fault (white arrows) and onlap of clay laminae on burrow wall (black arrow).

Figure 11. Ichnofossils associated with burrow fills. A) Small burrow fill (black arrows) cutting the laminated fill of a larger mammal burrow (SG locality). B) *Coprinisphaera* isp. from LLP locality. C) Tetrapod footprints on the top of laminated fill of a burrow from LLP locality. Numbers refers to individual footprints. D) ?*Rosellichnus* isp. (arrowed) inside a burrow fill from SG.

Figure 12. Bone remains inside burrow fills. A) Isolated, weathered and fragmentary glyptodontid osteoderms from a burrow fill of SG locality. B) Partly articulated osteoderms of *Eosclerocalyptus* sp. found inside a burrow fill from LLP locality.

**Legend of table**

Table 1. Relationship between cross-sections and body mass of the putative producers, estimated body mass according to Wu et al. (2015).

PeerJ Preprints | https://doi.org/10.7287/peerj.preprints.26632v1 | CC BY 4.0 Open Access | rec: 7 Mar 2018, publ: 7 Mar 2018

## Table 1(on next page)

Table 1. Body mass estimate of producers.

Relationship between cross-sections and body mass of the putative producers, estimated body mass according to Wu et al. (2015).

PeerJ Preprints | https://doi.org/10.7287/peerj.preprints.26632v1 | CC BY 4.0 Open Access | rec: 7 Mar 2018, publ: 7 Mar 2018

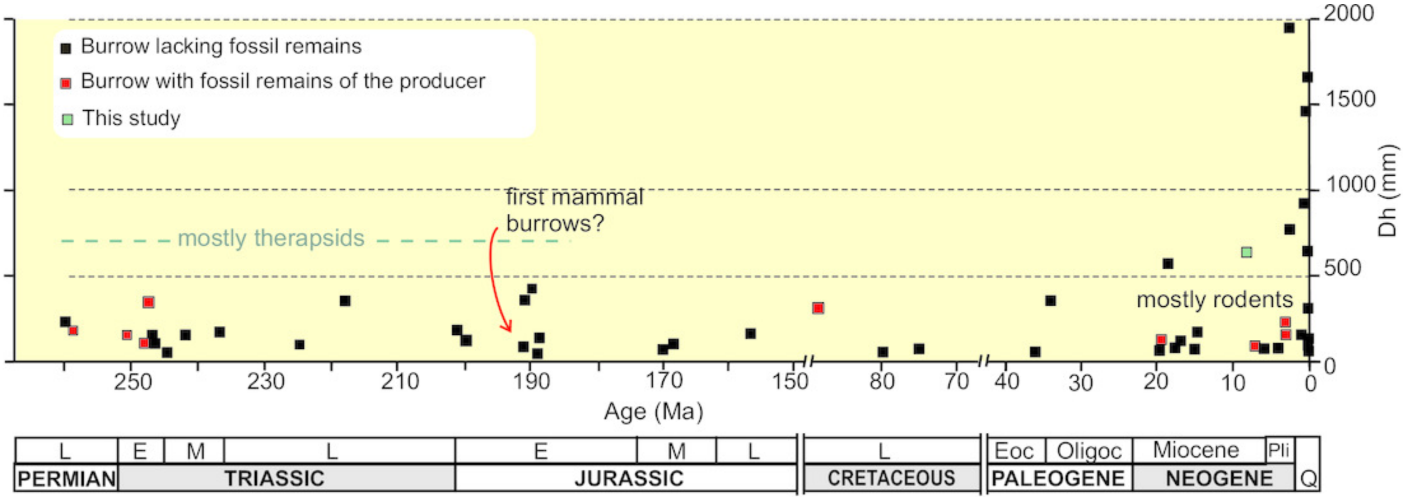
Table 1. Relationship between cross-sections and body mass of the putative producers, estimated body mass according to Wu et al. (2015). Dh: horizontal diameter.

| Range of burrow Dh (m) | Estimaded body mass (kg) | Fossil remains inside burrow fill | Potential burrow producer | Body mass of potential producer (kg) |
|---|---|---|---|---|
| 0.15 – 0.34 | 1 - 13 | *Paedotherium minor* | *Paedotherium* | 1.86 (Elissamburu, 2004) |
| | | | *Doellotatus* | Less than 1 (Vizcaíno & Fariña, 1999) |
| | | | *Chasicotatus* | Less than 1 (Scillato –Yané et al., 2010) |
| | | | *Proeuphractus* | 2–3 (Perea & Scillato–Yané, 1995) |
| | | | *Chorobates* | 1 – 10 (Vizcaíno & Fariña, 1999) |
| | | | *Lagostomus* | 1 – 10 (Vizcaíno & Fariña, 1999); |
| 0.39 – 0.94 | 37 - 439 | Mesotheriinae indet. | Mesotheriinae | 20.88 – 60.13 (Croft, Flynn & Wyss,, 2004) |
| | | *Eosclerocalyptus* sp. | *Eosclerocalyptus* | More than 100 (Vizcaíno & Fariña, 1999) |
| | | Gliptodontidae indet. | *Coscinocercus* | More than 100 (Vizcaíno & Fariña, 1999) |
| | | Gliptodontidae indet. | *Aspidocalyptus* | More than 100 (Vizcaíno & Fariña, 1999) |
| | | | *Macrochorobates* | 10 – 100 (Vizcaíno & Fariña, 1999) |
| | | | *Macroeuphractus* | 10 – 100 (Vizcaíno & Fariña, 1999) |
| | | *Proscelidodon* sp. | *Proscelidodon* | 581.8 (De Esteban-Trivigno, Mendoza & De Renzi,, 2008); 850 (Bargo et al., 2000); 1057 (Fariña, Vizcaíno & Bargo, 1998). Body mass of *S.leptocephalum* |
| | | *Paedotherium minor* | | |
| | | *Doellotatus* sp. | *Doellotatus* | |
| 1.05 – 1.5 | 708 - 1623 | Gliptodontidae indet. | Glyptodontidae | More than 100 (Vizcaíno & Fariña, 1999) |
| | | | *Proscelidodon* | 581.8 (De Esteban-Trivigno, Mendoza & De Renzi,, 2008); 850 (Bargo et al., 2000);1057 (Fariña, Vizcaíno & Bargo, 1998). Body mass of *S. leptocephalum.* |

# Figure 1

Plot of horizontal diameter vs age

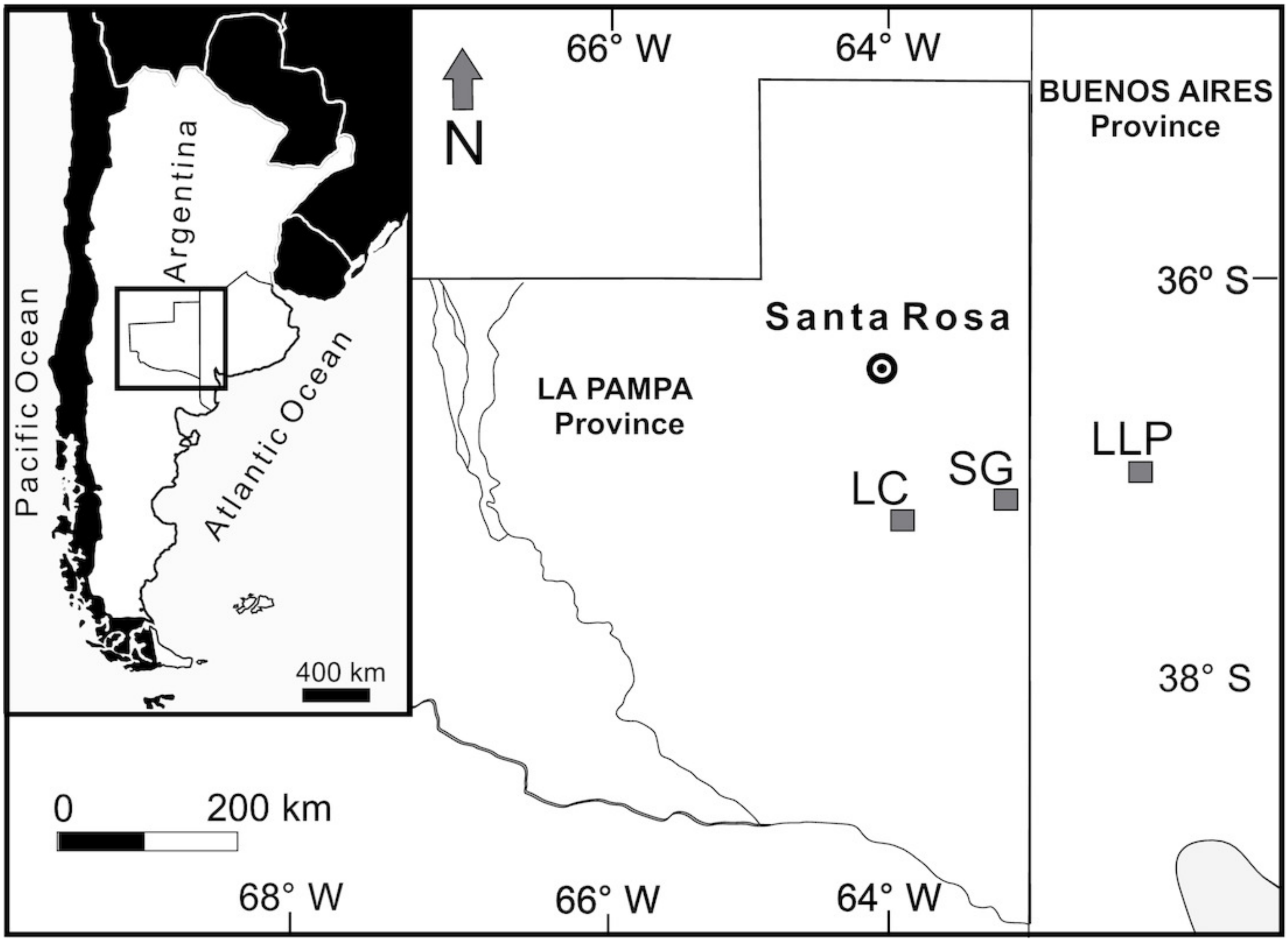
Plot of the average horizontal diameter (Dh) of tetrapod burrows against the age of the hosting rock, distinguishing between burrows with remains of the producer, those lacking bone remains and the present study. Note that the age axis contains two gaps in the Cretaceous and Paleocene. Source of data on Supplemental Information.



# Figure 2

Study area.

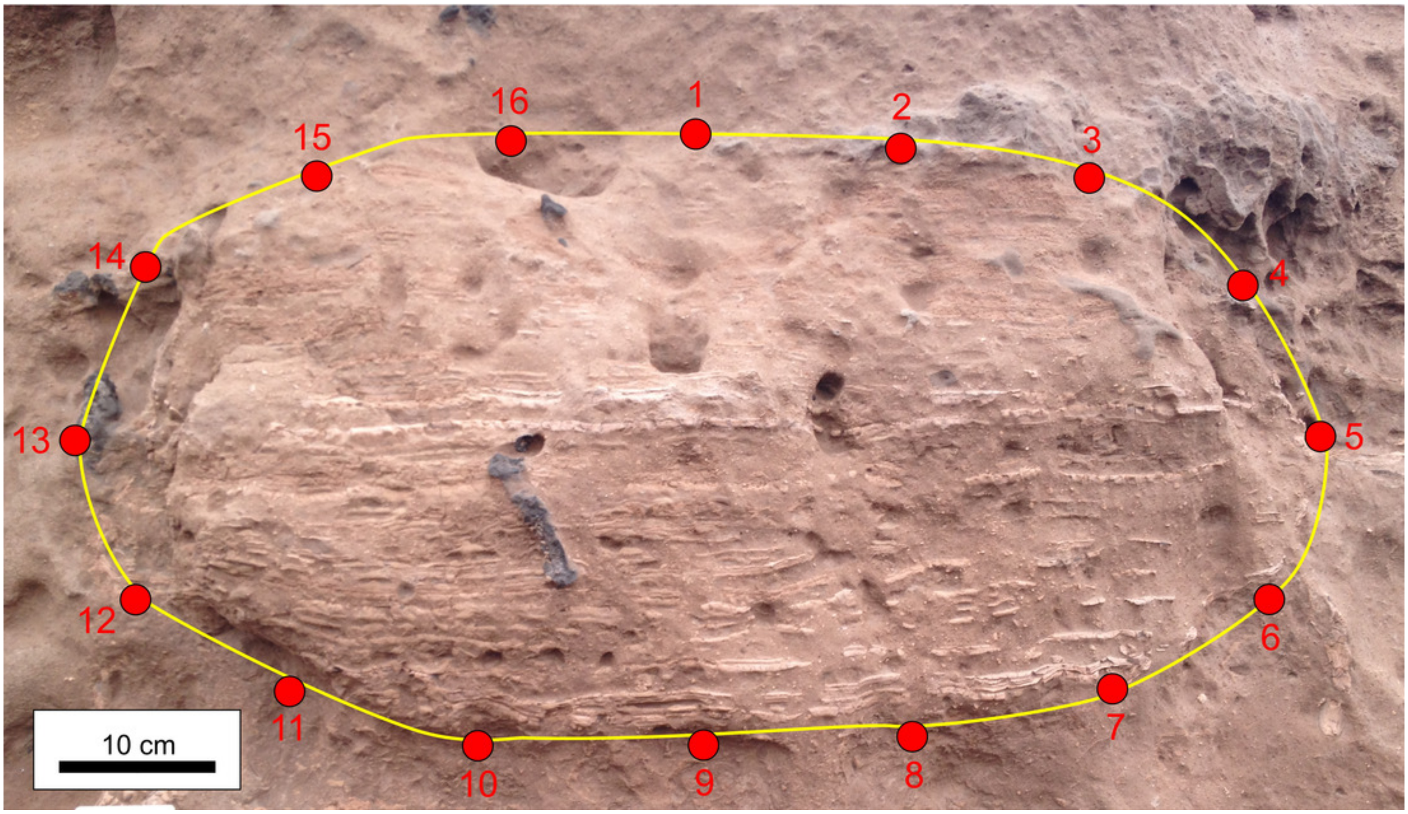
Study area in central Argentina (inset) and localities in La Pampa and Buenos Aires provinces. LC: Laguna Chillhué, SG: Salinas Grandes de Hidalgo, LLP: Laguna La Paraguaya. Image credit: María Cristina Cardonatto.



# Figure 3

Example of location of landmarks.

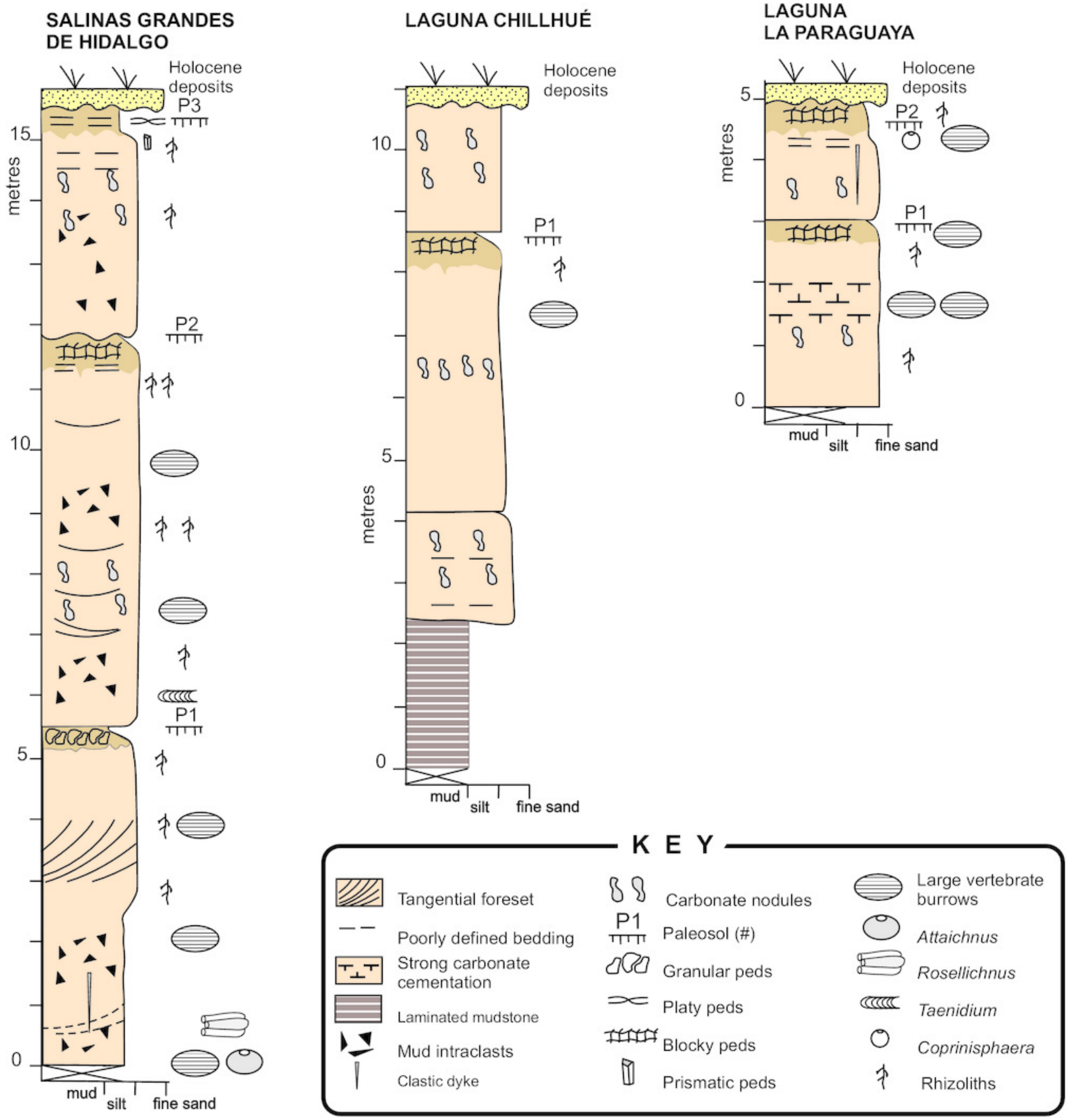
Location of landmarks (red points) on burrow fills preserved in cross section. Photo credit: Ricardo Néstor Melchor.



# Figure 4

Sedimentary logs.

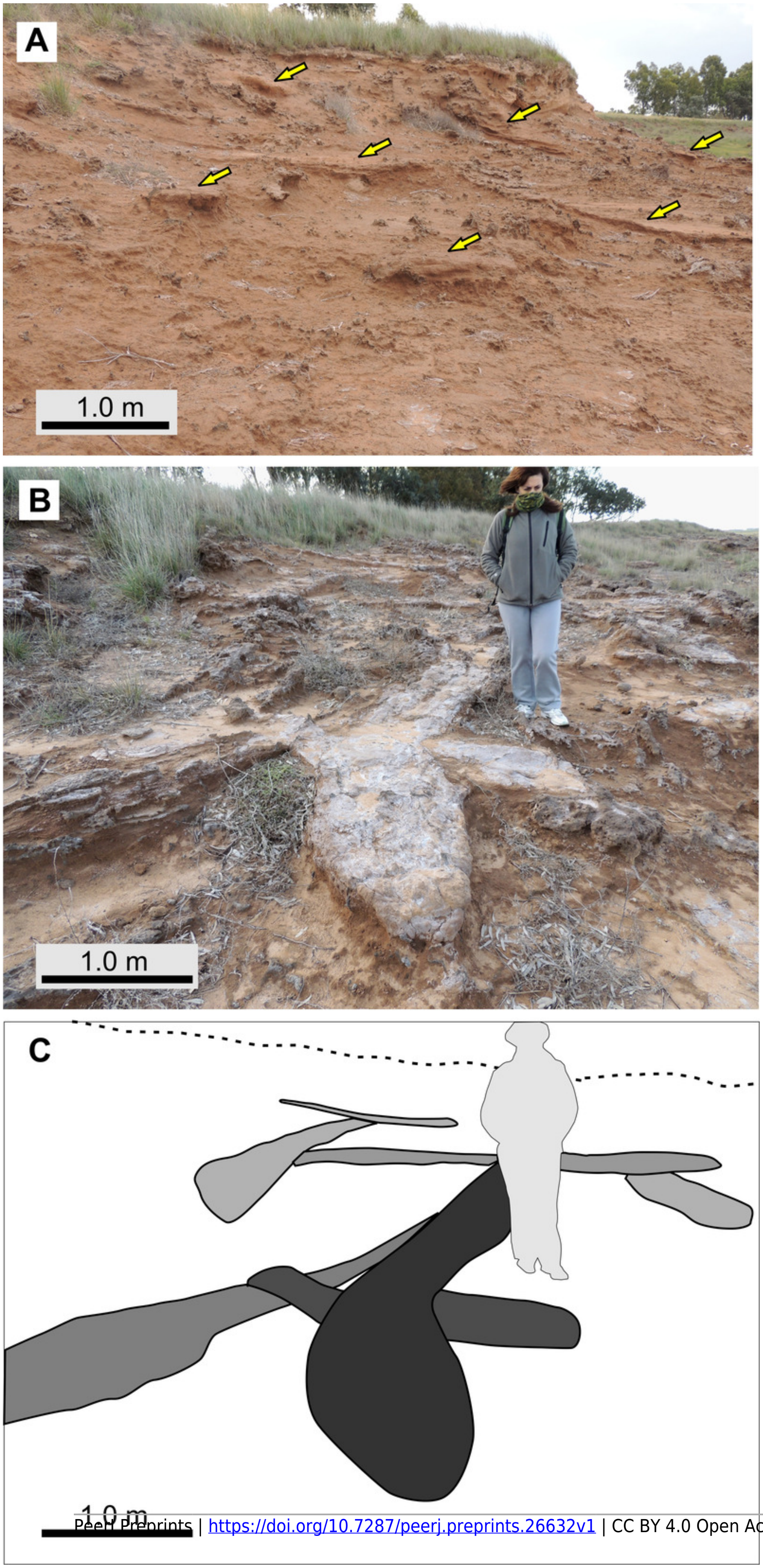
Detailed sedimentary logs of the measured sections from the study localities. Image credit: Ricardo Néstor Melchor.



# Figure 5

Abundance and cross-cutting relationships of burrow fills, from LLP locality.

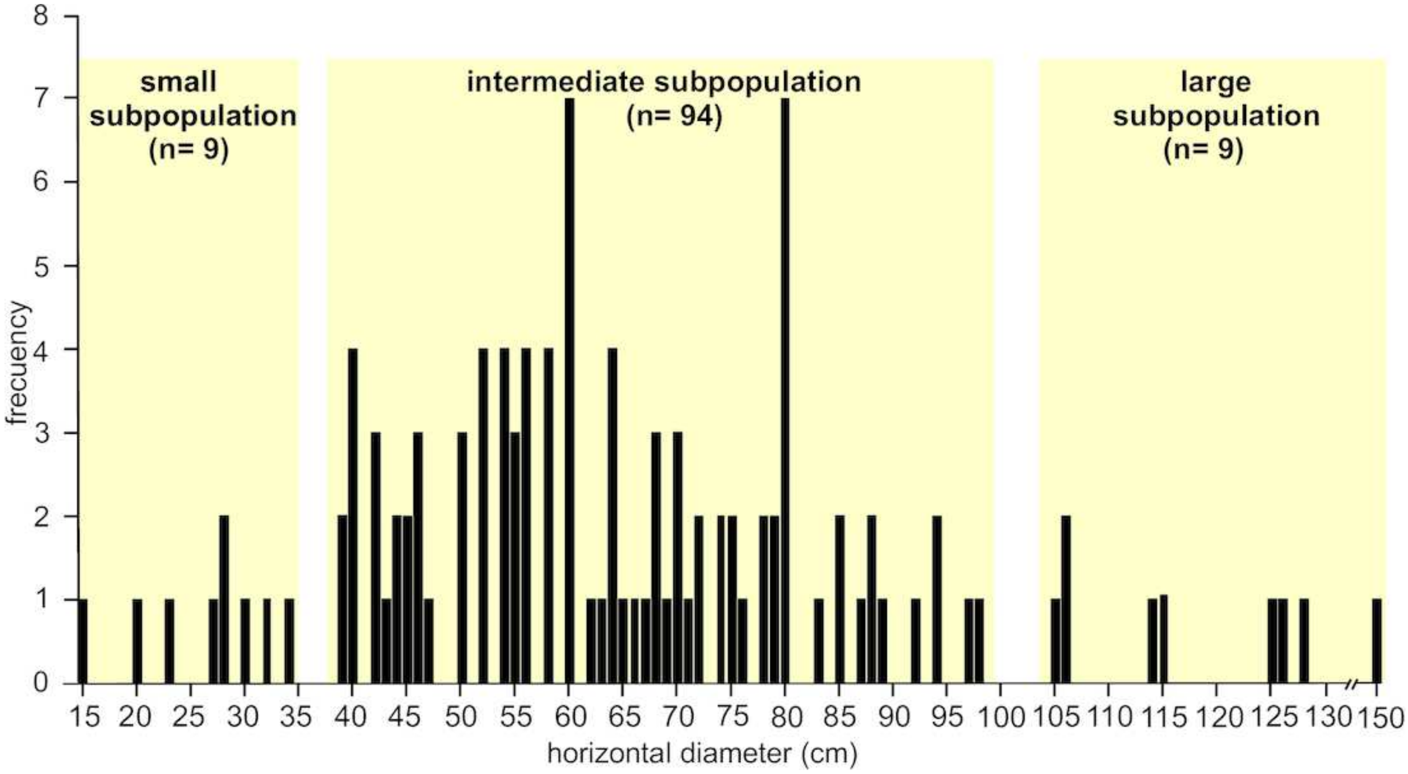
(A) General view of the outcrop showing several burrow fills (yellow arrows). (B-C) Field view and diagram of cross-cutting relationships between different burrow fills (distinguished in C with different shades of gray). Photo and image credit: Ricardo Néstor Melchor.


A
1.0 m
B
1.0 m
C
1.0 m

# Figure 6

Histogram of horizontal diameter.

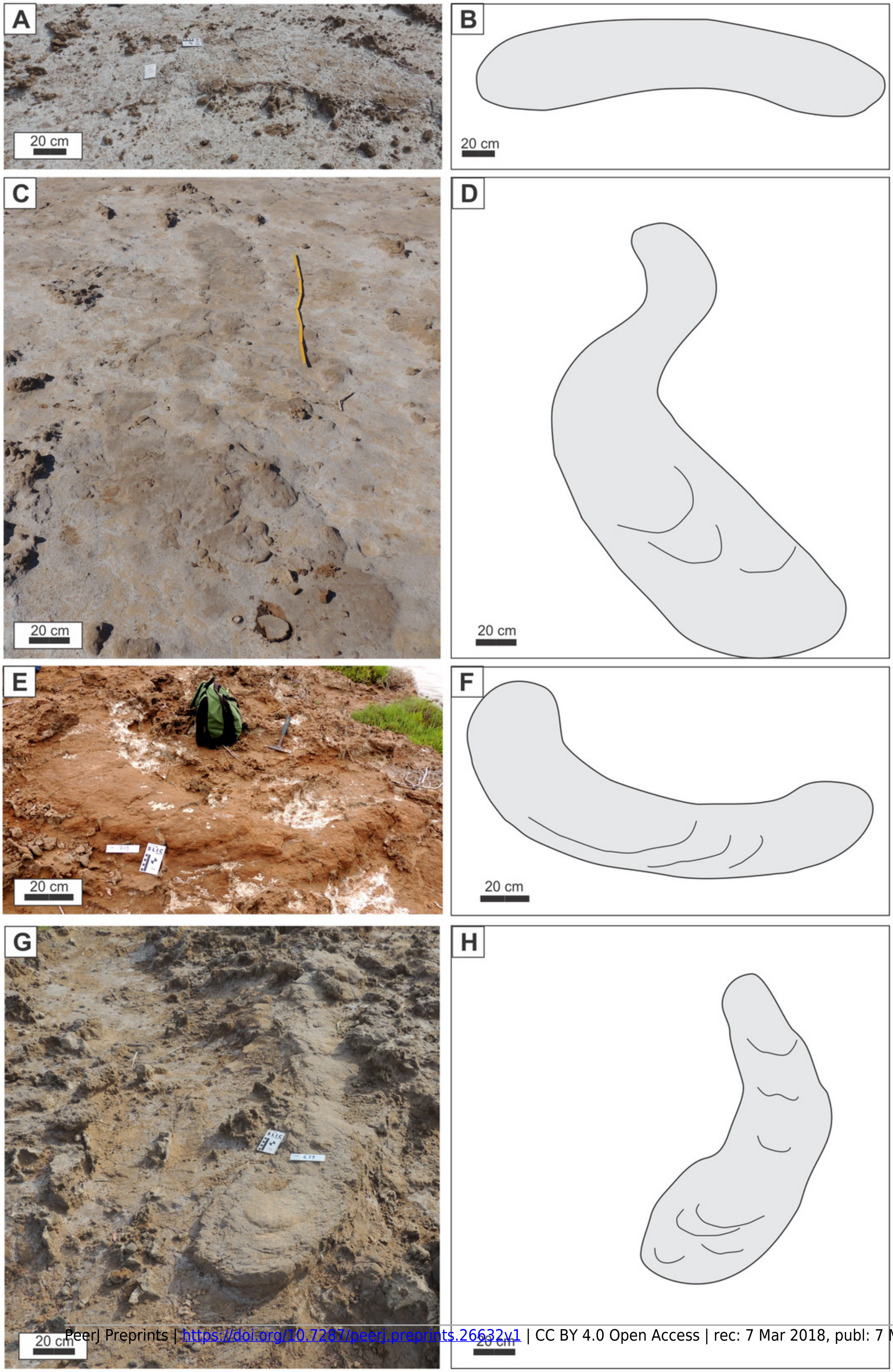
Histogram showing the frequency distribution of horizontal diameter (Dh) for the studied burrow fills. Three subpopulations can be distinguished. Image credit: María Cristina Cardonatto.



# Figure 7

Burrow fill morphology in plan view.

(A-B) Field photograph and diagram of slightly curved fill. (C-D) Field photograph and diagram of sinuous fill. (E-F) Field photograph and diagram of “C” shaped fill. (G-H) Field photograph and diagram of burrow fill with slightly enlarged and rounded end. Curved lines in the diagrams represent weathered laminae. Photo credit: Ricardo Néstor Melchor. Image credit: María Cristina Cardonatto.

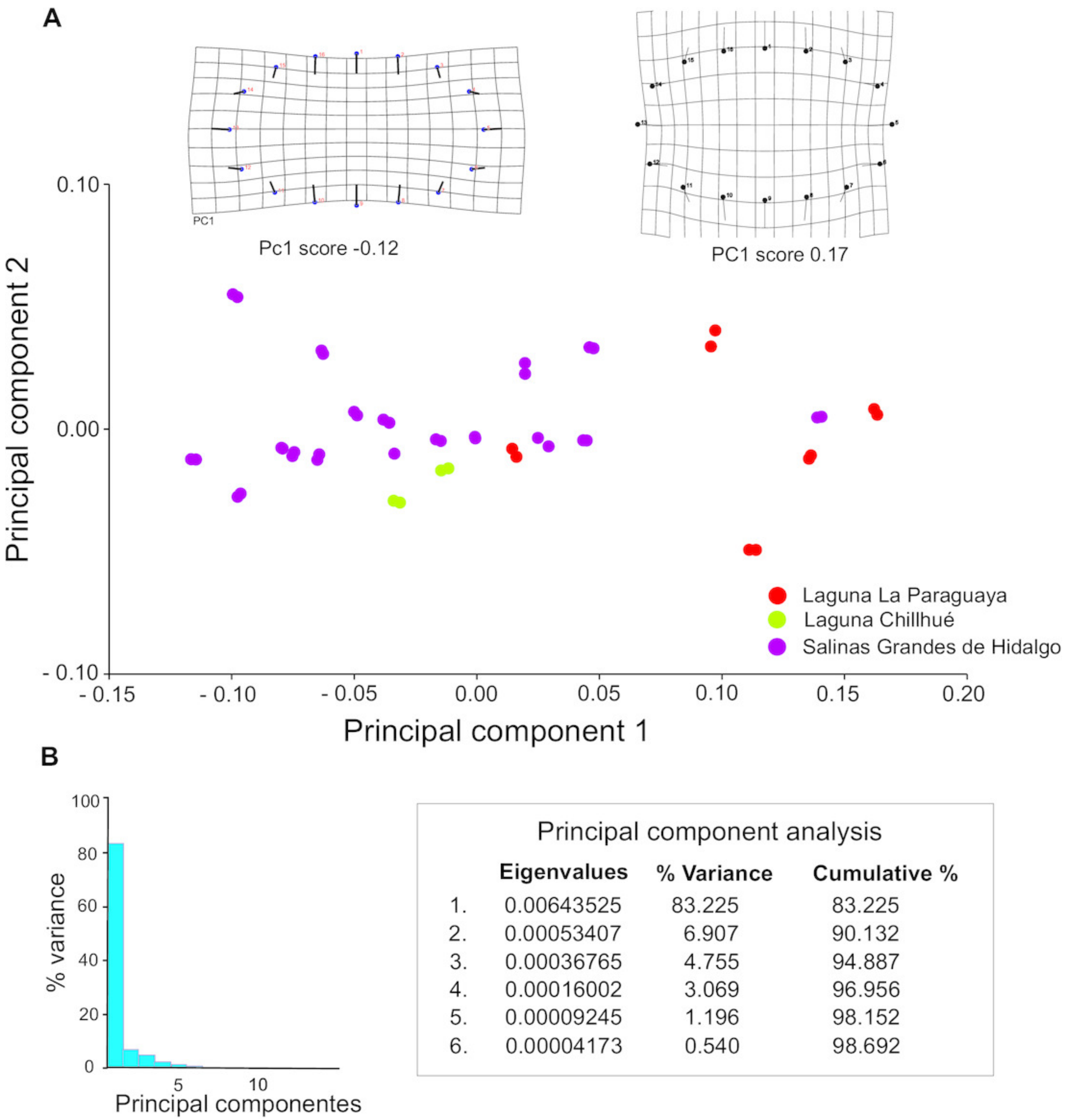

A
20 cm
B
20 cm
C
20 cm
D
20 cm
E
20 cm
F
20 cm
G
20 cm
H
20 cm

# Figure 8

Results of geometric morphometric analysis of burrow fills preserved in cross section.

(A) Plot of principal components 1 and 2, distinguishing by study locality, and deformation grids for elliptical (PC1 score -0.12) and subcircular (PC1 score 0.17) fills. (B) Histogram of variance of principal components. Image credit: María Cristina Cardonatto.

PeerJ Preprints | https://doi.org/10.7287/peerj.preprints.26632v1 | CC BY 4.0 Open Access | rec: 7 Mar 2018, publ: 7 Mar 2018



Principal component analysis

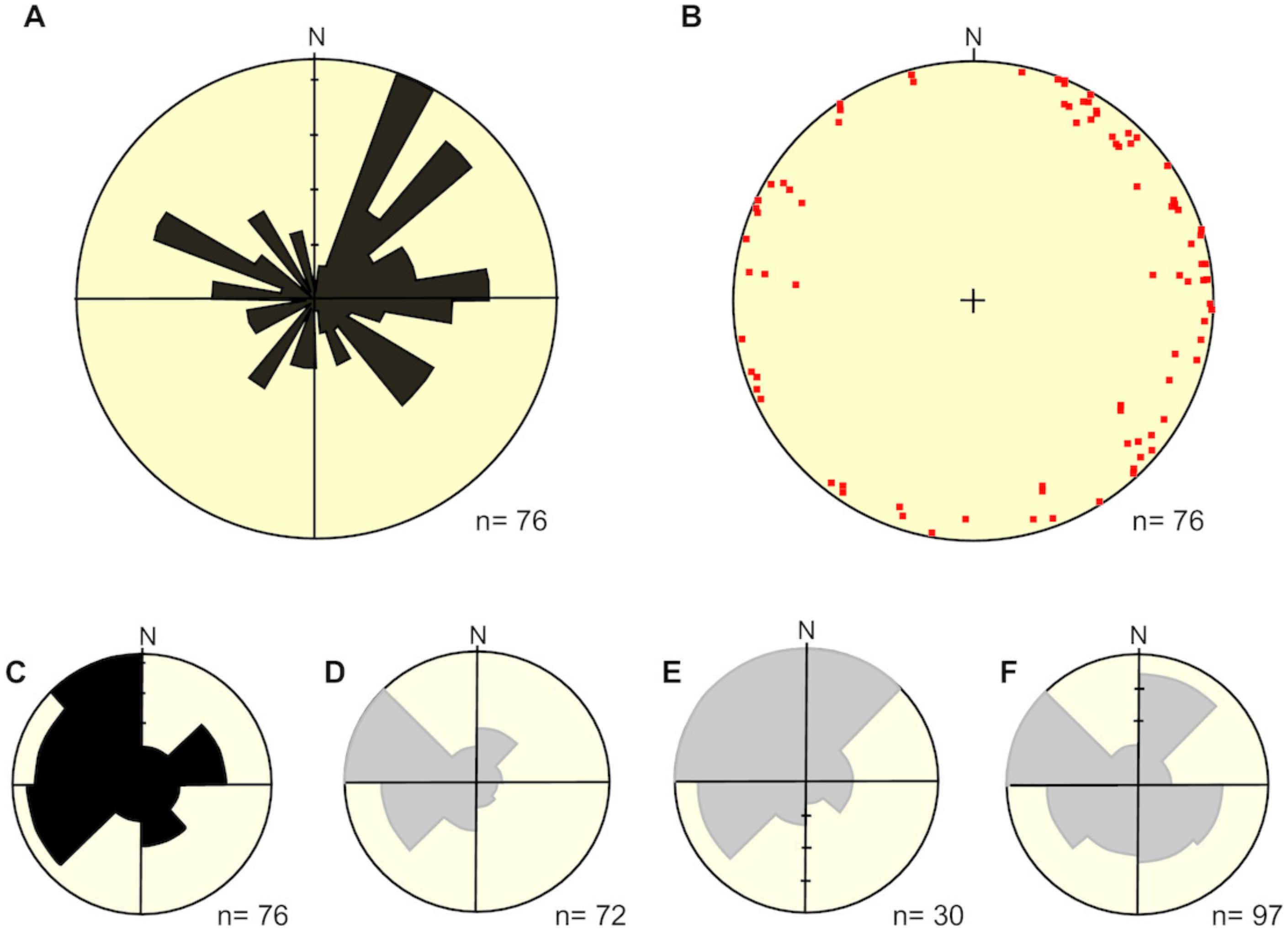
| | Eigenvalues | % Variance | Cumulative % |
|---|---|---|---|
| 1. | 0.00643525 | 83.225 | 83.225 |
| 2. | 0.00053407 | 6.907 | 90.132 |
| 3. | 0.00036765 | 4.755 | 94.887 |
| 4. | 0.00016002 | 3.069 | 96.956 |
| 5. | 0.00009245 | 1.196 | 98.152 |
| 6. | 0.00004173 | 0.540 | 98.692 |

# Figure 9

Orientation of fossil burrow fills compared with modern data from dasypodid burrows.

(A) Rose diagram showing the dip azimuth of fossil burrow fills. (B) Equal area projection of dip azimuth and dip angle of fossil burrow fills. (C) Entrance orientation of fossil burrows, assuming that it is located at 180° of measured dip azimuth. The data in A-C is from SG and LLP localities, those from the remaining locality are preserved only in cross-section. (D) Entrance orientation of several species of dasypodid burrows from semiarid settings of northern-central Argentina build from data by Crespo (1944). (E) Entrance orientation of *Dasypus hibridus* burrows from a grassland setting in Uruguay build from data by González et al. (2001). (F) Entrance orientation of *Chaetophractus villosus* burrows from cultivated land in Buenos Aires province, Argentina build from data by Abba et al. (2005). Image credit: Ricardo Néstor Melchor.

PeerJ Preprints | https://doi.org/10.7287/peerj.preprints.26632v1 | CC BY 4.0 Open Access | rec: 7 Mar 2018, publ: 7 Mar 2018

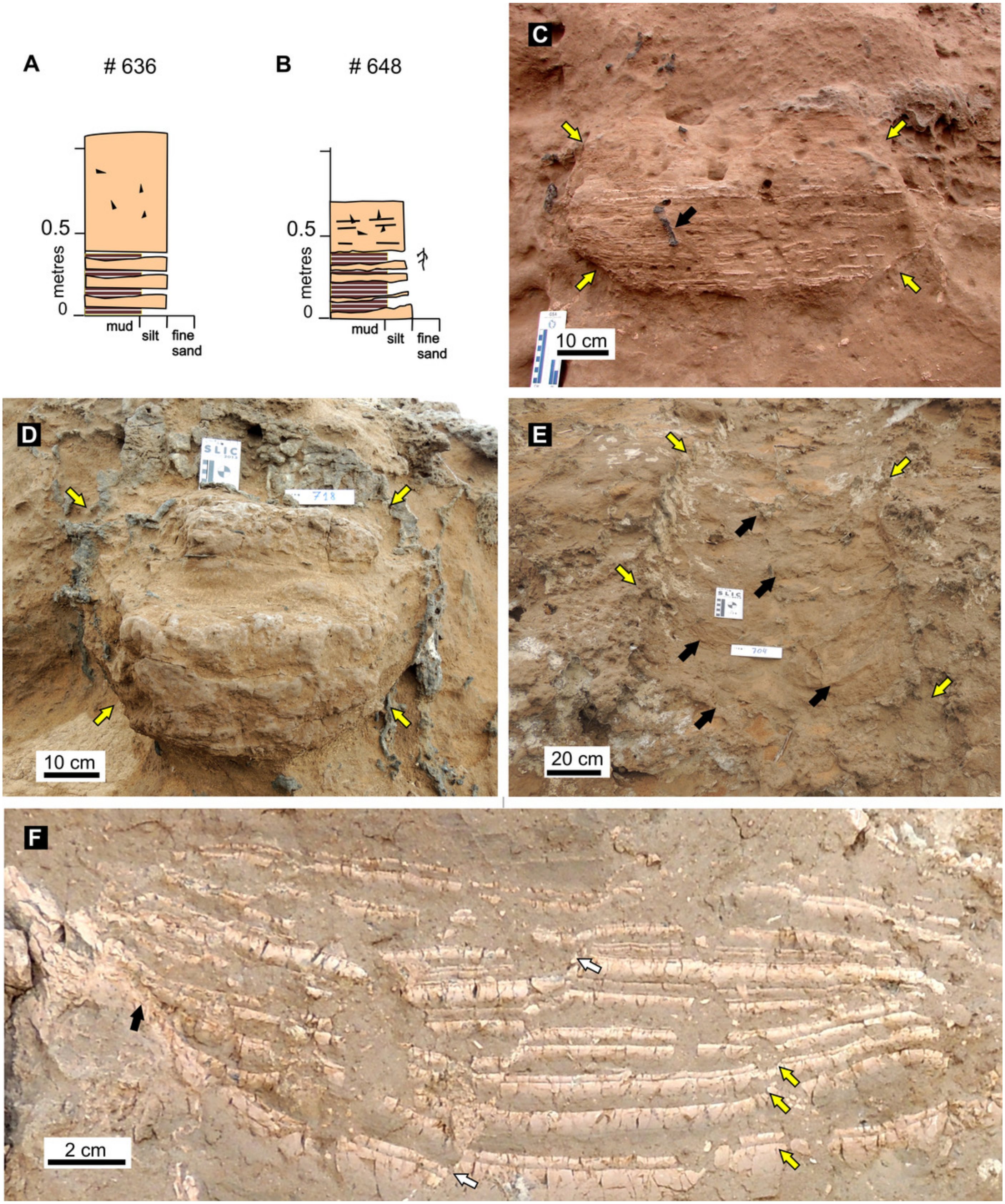


PeerJ Preprints | https://doi.org/10.7287/peerj.preprints.26632v1 | CC BY 4.0 Open Access | rec: 7 Mar 2018, publ: 7 Mar 2018

# Figure 10

Features of fossil burrow fills.

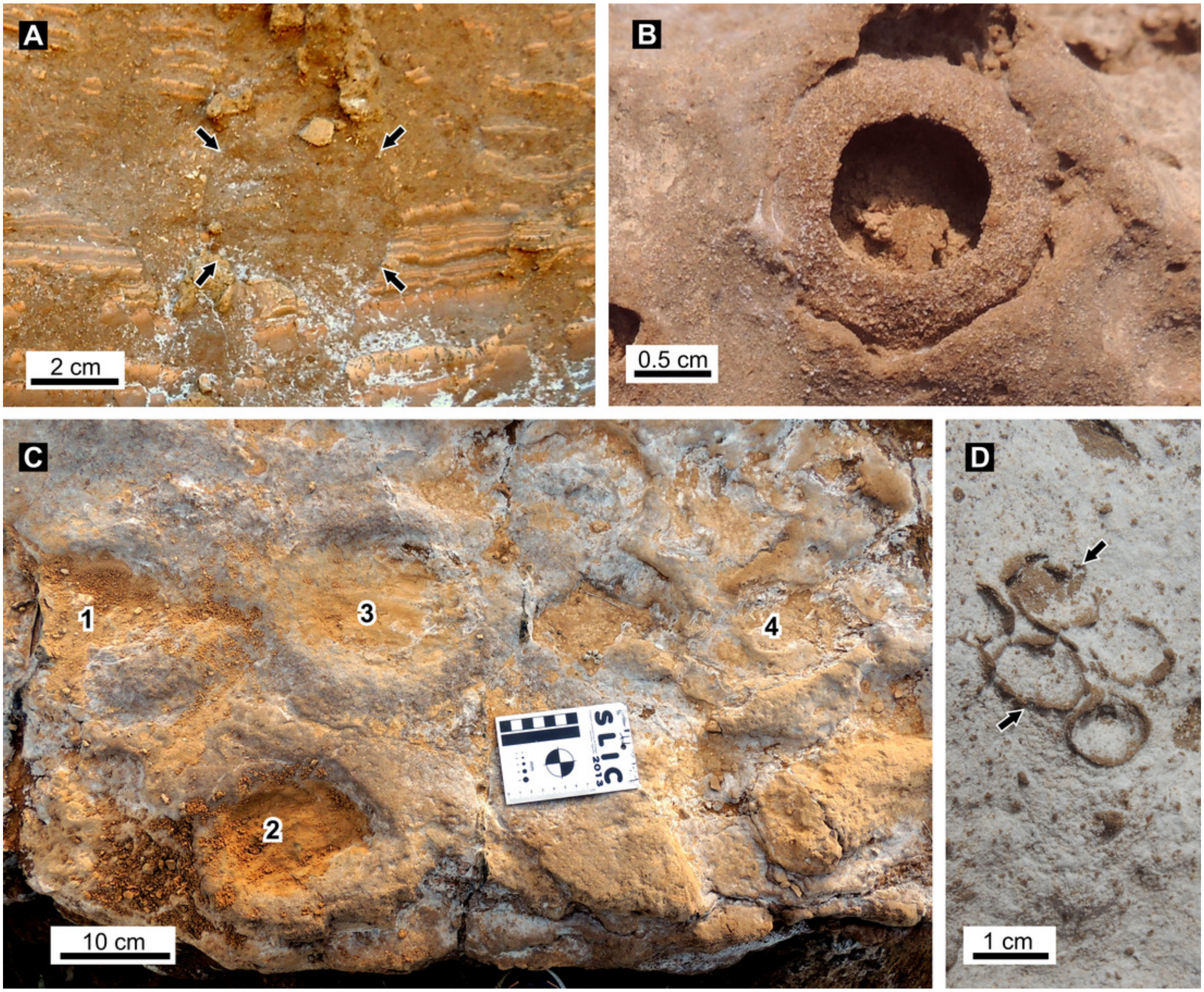
(A-B) Detailed sedimentary logs of the fill of selected burrows (see Table S1 for information on these burrow fills). References as for Fig. 4. (C) Cross-section of laminated to massive burrow fill # 648 from SG, represented in (B). Elliptical outline of fill indicated by yellow arrows, black arrow point to subvertical rhizolith cutting the laminated fill. (D) Cross-section of burrow fill # 714 from LLP. Subcircular outline of poorly laminated to massive burrow fill indicated by yellow arrows. (E) Pseudomeniscate structure in burrow fill # 704 from LLP seen in plan view. Yellow arrows point the outline of the burrow fill and black arrows to individual pseudomenisci. (F) Detail of laminated burrow fill (# 632 from SG) showing direct grading in siltstone to claystone laminae (yellow arrows), synsedimentary fault (white arrows) and onlap of clay laminae on burrow wall (black arrow). Image credit: Ricardo Néstor Melchor.


A
# 636
B
# 648
0.5
metres
0
mud
silt
fine sand
C
10 cm
D
718
10 cm
E
704
20 cm
F
2 cm

# Figure 11

Ichnofossils associated with burrow fills.

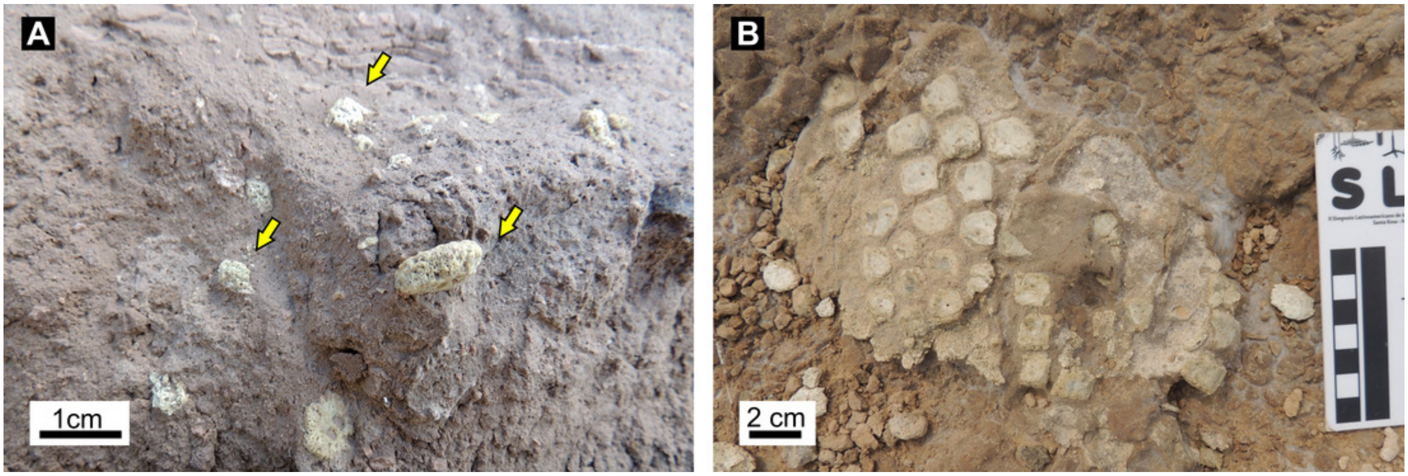
(A) Small burrow fill (black arrows) cutting the laminated fill of a larger mammal burrow from SG locality. (B) *Coprinisphaera* isp. from LLP locality. (C) Tetrapod footprints on the top of laminated fill of a burrow from LLP locality. Numbers refers to individual footprints. (D) ?*Rosellichnus* isp. (arrowed) inside a burrow fill from SG. Photo credit: Ricardo Néstor Melchor.



# Figure 12

Bone remains inside burrow fills.

(A) Isolated, weathered and fragmentary glyptodontid osteoderms from a burrow fill of SG locality. (B) Partly articulated osteoderms of *Eosclerocalyptus* sp. found inside a burrow fill from LLP locality. Photo credit: Ricardo Néstor Melchor.



PeerJ Preprints | https://doi.org/10.7287/peerj.preprints.26632v1 | CC BY 4.0 Open Access | rec: 7 Mar 2018, publ: 7 Mar 2018