

# Optimizing alfalfa productivity and persistence versus greenhouse gases fluxes in a continental arid region

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Alfalfa in China is mostly planted in the semi-arid or arid Northwest inland regions due to its ability to take up water from deep in the soil and to fix atmospheric N<sub>2</sub> which reduces N fertilizer application. However, perennial alfalfa may deplete soil water due to uptake and thus aggravate soil desiccation. The objectives of this study were (1) to determine the alfalfa forage yield, soil property [soil temperature (ST), soil water content (SWC), soil organic carbon (SOC) and soil total nitrogen (STN)] and greenhouse gas [GHG: methane (CH<sub>4</sub>), nitrous oxide (N<sub>2</sub>O), and carbon dioxide (CO<sub>2</sub>)] emissions affected by alfalfa stand age and growing season, (2) to investigate the effects of soil property on GHG emissions, and (3) to optimize the alfalfa stand age by integrating the two standard criteria, the forage yield and water use efficiency, and the total GHG efflux (CO<sub>2</sub>-eq). This study was performed in alfalfa fields of different ages (2, 3, 5 and 7 year old) during the growing season (from April to October) in a typical salinized meadow with temperate continental arid climate in the Northwest inland regions, China. Despite its higher total GHG efflux (CO<sub>2</sub>-eq), the greater forage yield and water use efficiency with lower GEI<sub>hay</sub> and high CH<sub>4</sub> uptake in the 5-year alfalfa stand suggested an optimal alfalfa stand age of 5 years. Results show that ST, SOC and RBM alone had positive effects (except RBM had no significant effect on CH<sub>4</sub> effluxes), but SWC and STN alone had negative effects on GHG fluxes. Furthermore, results demonstrate that in arid regions SWC superseded ST, SOC, STN and RBM as a key factor regulating GHG fluxes, and soil water stress may have led to a net uptake of CH<sub>4</sub> by soils and a reduction of N<sub>2</sub>O and CO<sub>2</sub> effluxes from alfalfa fields. Our study has provided insights into the determination of alfalfa stand age and the understanding of mechanisms regulating GHG fluxes in alfalfa fields in the continental arid regions. This knowledge is essential to decide the alfalfa retention time by considering the hay yield, water use efficiency as well as GHG emission.

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17

18 **ABSTRACT**

19 Alfalfa in China is mostly planted in the semi-arid or arid Northwest inland regions due to its ability to  
20 take up water from deep in the soil and to fix atmospheric N<sub>2</sub> which reduces N fertilizer application.

21 However, perennial alfalfa may deplete soil water due to uptake and thus aggravate soil desiccation. The  
22 objectives of this study were (1) to determine the alfalfa forage yield, soil property [soil temperature (ST),  
23 soil water content (SWC), soil organic carbon (SOC) and soil total nitrogen (STN)] and greenhouse gas

24 [GHG: methane (CH<sub>4</sub>), nitrous oxide (N<sub>2</sub>O), and carbon dioxide (CO<sub>2</sub>)] emissions affected by alfalfa  
25 stand age and growing season, (2) to investigate the effects of soil property on GHG emissions, and (3) to  
26 optimize the alfalfa stand age by integrating the two standard criteria, the forage yield and water use  
27 efficiency, and the total GHG efflux (CO<sub>2</sub>-eq).

28 This study was performed in alfalfa fields of different ages (2, 3, 5 and 7 year old) during the growing  
29 season (from April to October) in a typical salinized meadow with temperate continental arid climate in  
30 the Northwest inland regions, China.

31 Despite its higher total GHG efflux (CO<sub>2</sub>-eq), the greater forage yield and water use efficiency with lower  
32 GEI<sub>hay</sub> and high CH<sub>4</sub> uptake in the 5-year alfalfa stand suggested an optimal alfalfa stand age of 5 years.

33 Results show that ST, SOC and RBM alone had positive effects (except RBM had no significant effect on  
34 CH<sub>4</sub> effluxes), but SWC and STN alone had negative effects on GHG fluxes. Furthermore, results  
35 demonstrate that in arid regions SWC superseded ST, SOC, STN and RBM as a key factor regulating  
36 GHG fluxes, and soil water stress may have led to a net uptake of CH<sub>4</sub> by soils and a reduction of N<sub>2</sub>O  
37 and CO<sub>2</sub> effluxes from alfalfa fields.

38 Our study has provided insights into the determination of alfalfa stand age and the understanding of  
39 mechanisms regulating GHG fluxes in alfalfa fields in the continental arid regions. This knowledge is  
40 essential to decide the alfalfa retention time by considering the hay yield, water use efficiency as well as  
41 GHG emission.

42 **Keywords** Arid climate, Stand age, Forage yield, Water use efficiency, Methane, Carbon dioxide, Nitrous  
43 oxide, Soil total nitrogen

44

## 45 INTRODUCTION

46 Alfalfa (*Medicago Sativa* L.) is the most widely grown perennial forage legume around the world (Yang  
47 et al., 2008). Since 2011, alfalfa planting has been gradually increased to meet the increasing demand for  
48 livestock production (Wang, Hansen & Xu, 2016). The increasing demand of alfalfa has great potential to

49 change the structure and function of the farming systems but has positive and negative influences. First,  
50 alfalfa with rhizobia, the root-dwelling symbiotic bacteria, can fix atmospheric  $N_2$  (Peterson & Russelle,  
51 1991) and thus reduce N fertilizer applications. Secondly, due to its strong ability to take up water (Wan  
52 et al., 2008), the alfalfa is mostly (> 75%) planted in the Northwest inland regions with an arid or semi-  
53 arid climate in or around the Loess Plateau of China (Hu & Cash, 2009; Wang, Hansen & Xu, 2016);  
54 however, alfalfa has a much higher water requirement than other crops which may deplete soil water and  
55 aggravate soil desiccation in long-term stands (Guan et al., 2013; Zhu et al., 2016). For example,  
56 McCallum, Peoples & Connor (2000) reported that in Australia soil profiles under alfalfa-based perennial  
57 fields remain consistently drier throughout the year compared with continuous annual cropping. After  
58 eight years of alfalfa cultivation on the north edge of Loess Plateau, China, four to five years are required  
59 to restore soil water condition to the initial level (Du, Wang & Long, 1999a), imposing a negative effect  
60 on the growth of subsequent crops. Moreover, the forage yield as well as the water use efficiency usually  
61 decrease after four to seven successive growing years (Du, Long & Wang, 1999b; Zhang et al., 2004;  
62 Cheng, Wan & Wang, 2005; Jia et al., 2009). Therefore, determination of the optimal cultivation ages of  
63 alfalfa is critical to avoid over-consumption of soil water by balancing the forage yield and water use  
64 efficiency.

65 Another concern with alfalfa is that the possible rhizobial denitrification may result in an increase of  
66 greenhouse gas (GHG, mainly the nitrous oxide  $N_2O$ , methane  $CH_4$ , and carbon dioxide  $CO_2$ ) emissions  
67 (O'Hara & Daniel, 1985). It is well known that  $N_2O$  effluxes are driven by nitrification (oxidation of  
68  $NH_4^+$  to  $NO_3^-$  via  $NO_2^-$ ) under aerobic conditions and denitrification (reduction of  $NO_3^-$  to  $N_2O$  and  $N_2$ )  
69 under anaerobic conditions (Ussiri & Lal, 2012; Oertel et al., 2016).  $CO_2$  release from soils is the  
70 subsequent results of soil respiration of both root and anaerobic and aerobic microbes (Oertel et al.,  
71 2016), where root respiration may contribute average up to about 50% of the total soil respiration  
72 depending on the season and vegetation type (Hanson et al., 2000).  $CH_4$  in soils is produced by  
73 methanogenesis under anaerobic conditions and is consumed by methanotrophic microorganisms that use  
74  $O_2$  and  $CH_4$  for their metabolism under aerobic conditions (Smith et al., 2003; Dutaur & Verchot, 2007;

75 Gao et al., 2014). Thus, whether soil is a net source or sink for CH<sub>4</sub> depends on the relative rates of  
76 methanogenic and methanotrophic activity (Tate, 2015; Tian et al., 2016). GHG fluxes are mediated by  
77 both biotic (e.g., microbial activity and root respiration) (Kitzler et al., 2006; Singh et al., 2010;  
78 Butterbach-Bahl et al., 2013) and abiotic (e.g., soil temperature, moisture, and soil carbon and nitrogen)  
79 factors (Kitzler et al., 2006; Singh et al., 2010; Butterbach-Bahl et al., 2013). Heretofore, how these  
80 factors influence GHG fluxes or which one is the most important factor influencing GHG fluxes in the  
81 arid continental regions is still not clear.

82 Previous studies on optimizing alfalfa stand age usually consider only two factors, the yield of alfalfa  
83 hay and water use efficiency, especially in the arid or semi-arid regions (e.g., Zhang et al., 2004; Cheng,  
84 Wan & Wang, 2005; Fan et al., 2016) but ignores greenhouse gas (GHG) effluxes. In contrast, other  
85 studies investigate GHG effluxes from alfalfa fields of different stand ages (e.g., Zhong, Nelson &  
86 Lemke, 2011; Uzoma et al., 2015; Burger et al., 2016) but do not measure alfalfa productivity;  
87 furthermore, those studies usually only consider N<sub>2</sub>O effluxes due to the N<sub>2</sub> fixation of alfalfa. So far only  
88 a few studies on GHG effluxes from alfalfa fields have included CH<sub>4</sub>, CO<sub>2</sub> and N<sub>2</sub>O in analyses (e.g.,  
89 Chaves et al., 2006; Ellert & Janzen, 2008).

90 In this study, we estimated the persistence of alfalfa in an arid continental region with respect to the  
91 tradeoffs between hay yield, water use efficiency, and GHG effluxes as affected by soil properties. The  
92 approach was to integrate the forage yield and water use efficiency with total GHG efflux (CO<sub>2</sub>-eq) and  
93 GHG efflux (CO<sub>2</sub>-eq) per unit hay yield. We then investigated the dynamics of soil properties (i.e., soil  
94 temperature, water content, organic carbon and total nitrogen), root biomass and CH<sub>4</sub>, CO<sub>2</sub> and N<sub>2</sub>O  
95 fluxes during the growing seasons, and finally analyzed the influence of soil properties and root biomass  
96 on GHG fluxes. Results from this study will improve our understanding in GHG effluxes in the arid areas  
97 and provide essential information to develop strategies for alfalfa field management.

98

## 99 **METHODS**

**100 Study site and alfalfa field**

101 The study was carried out in Grassland Agricultural Trial Station of Lanzhou University (latitude 39°15'  
102 N, longitude 100°02' E), Gansu Province, China. The field ( $\approx$  280 ha) used in this study is 1,390 m above  
103 the sea level and classified as a typical salinized meadow with temperate continental arid climate in the  
104 Northwest inland regions (Zhu et al., 1997). The mean annual precipitation is about 123 mm with  $\geq$  65%  
105 occurring during the growing seasons from April to October (Nobuyuki et al., 2018). Irrigation is  
106 necessary and usually applied bimonthly during the growing seasons with a rate of 120 mm respectively  
107 in April, June or August in the study site. The annual mean air temperature is 7.6°C (from -28°C between  
108 December and February to 38°C between June and Mid-August). The soil pH value is about 8.0, and the  
109 soil at the study site is classified as Aquisalids according to USDA soil taxonomy (Zhu et al., 1997).

110 To optimize the alfalfa stand ages in relation to biomass, soil properties and GHG effluxes, we used  
111 a Randomized Complete Block design in 2014. A long-term established forage study with differing stand  
112 ages was used for this experiment, and data collected for this experiment took place over one year (Zhu et  
113 al., 1997; Nobuyuki et al., 2018). There were three blocks (about 2.3 ha for each block) and each block  
114 was evenly divided into four subblocks, four stand age treatments (i.e., 2, 3, 5 and 7 years old, sown in  
115 late August 2012, 2011, 2009 and 2007, respectively) were randomly assigned into each subblock. For  
116 each subblock, three sampling plots (30 m width and of 100 m length) were randomly set up for forage  
117 harvest, and soil and GHG sampling.

118

**119 Alfalfa biomass and soil property**

120 To determine alfalfa productivity of different stage ages, one quadrat (1 m  $\times$  1 m) in each sampling plot  
121 was randomly selected and the hay yield was measured by cutting above-ground biomass during early  
122 blooming periods (10 June, 20 July and 01 October). To measure the under-ground root biomass (RBM),  
123 another quadrat of the same size in each sampling plot was randomly selected, RBM was collected by  
124 digging 30 cm depth after gas collection (see next section for details). The harvested materials were oven-  
125 dried at 60°C for 48 hours, and then weighted.

126 To determine the soil characteristics in relation to field stage age, we also randomly selected two  
127 sampling sites in each sampling plot, and soil samples were collected at a 0-10 cm depth using the bucket  
128 auger (5 cm diameter) after gas collection (see next section for details). Soil samples were naturally dried  
129 then extracted by passing through a 0.25-mm sieve. Soil organic carbon (SOC) was measured by Chromic  
130 acid REDOX titration (Nelson & Sommers, 1996). Soil total nitrogen (STN) was determined following  
131 the methods of Bremner & Mulvaney (1982). Meanwhile, two cores (8.4 cm diameter × 6 cm length)  
132 were sampled by inserting soil profile of 0-10 cm depth in each quadrat and cores were dried at 105°C for  
133 48 hours. The soil water content (SWC) was then estimated as: (original wet weight - soil dry weight)/soil  
134 volume.

135

#### 136 **GHG efflux**

137 GHG effluxes from soils are more likely to occur in spring, summer and autumn than in winter (Liu,  
138 Wang & Xu, 2010), thus GHG samplings were only carried out during the growing seasons of April,  
139 June, July, August and October in 2014. Two sampling sites were randomly selected in each sampling  
140 plot on 13 April. Gases were sampled four times (i.e., 5:00, 10:00, 14:00 and 18:00) for three successively  
141 sunny days in each mid-month, after removing the above-ground plant and litter (Liu et al., 2017). The  
142 mean GHG fluxes during the three successive days were treated as the average daily fluxes for that  
143 month.

144 Gas was collected using a static opaque chamber (30 cm × 30 cm × 30 cm) (Liu et al., 2017). For  
145 each sampling event, four gas samples were taken within 30 minutes at a time interval of 10 minutes (i.e.,  
146 0, 10, 20 and 30 minutes). The chamber was also equipped with an electronic thermometer. The air  
147 temperature inside the chamber was recorded during gas sampling and applied to calculate gas flux (see  
148 below). Soil temperature (ST) was also measured by a mercury thermometer inserted 5 cm into the soil at  
149 the sampling site before and after gas sampling and the mean temperature of the two measurements was  
150 applied to detect its effect on GHG effluxes.

151 Gas concentration was measured within 24 hours, i.e., CH<sub>4</sub> and CO<sub>2</sub> were simultaneously analyzed

152 by a CH<sub>4</sub>/CO<sub>2</sub> Spektrum Analyser with syringe injection (Model No. 908-0011-0001, Los Gatos  
 153 Research, USA), and N<sub>2</sub>O was analyzed by a N<sub>2</sub>O Spektrum Analyser (Model No. 908-0015-0000, Los  
 154 Gatos Research, USA). According to Liu et al. (2017), the daily GHG fluxes were estimated as:  $\text{GHG}_{\text{daily}}$   
 155  $= (a \times \text{flux}_{7:00} + b \times \text{flux}_{12:00} + c \times \text{flux}_{16:00} + d \times \text{flux}_{18:00})$ , where  $a$ ,  $b$ ,  $c$  and  $d$  are the constant gas flux duration  
 156 (i.e.,  $a = 11$  h from 20:00 to 7:00,  $b = 5$  h from 7:00 to 12:00,  $c = 4$  h from 12:00 to 16:00, and  $d = 4$  h  
 157 from 16:00 and 20:00). The hourly GHG fluxes were thus estimated as:  $\text{GHG}_{\text{hourly}} = \text{GHG}_{\text{daily}}/e$ , where  $e =$   
 158 24 (number of hours per day); and the monthly GHG fluxes were then calculated as:  $\text{GHG}_{\text{monthly}} =$   
 159  $\text{GHG}_{\text{daily}} \times f$ , where  $f = 30$  or  $31$  (number of days per month between April and October 2014). The total  
 160 gas flux during growing seasons was the sum of monthly fluxes (from April to October). Gas fluxes in  
 161 May and September were not measured and thus gap-filled using linear interpolation of the arithmetical  
 162 means of gas fluxes for the two close months (Chen et al., 2013).

163 The flux of GHG describes the change of gas in unit time in the sampling box. Generally, a positive  
 164 value indicates gas effluxes, and a negative value suggests gas absorption. The specific formula is (Liu et  
 165 al., 2017):

$$166 \quad F = \rho \frac{V}{A} \cdot \frac{P}{P_0} \cdot \frac{T_0}{T} \cdot \frac{dCt}{dt}$$

167 where  $F$  is the gas flux (kg/m<sup>2</sup>/h),  $\rho$  is the gas density (kg/m<sup>3</sup>) under standard conditions ( $\rho_{\text{CO}_2} = 1.965$   
 168 kg/m<sup>3</sup>,  $\rho_{\text{CH}_4} = 0.715$  kg/m<sup>3</sup> and  $\rho_{\text{N}_2\text{O}} = 1.965$  kg/m<sup>3</sup> respectively for CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O),  $V$  is chamber  
 169 volume (m<sup>3</sup>),  $A$  is the base area of the chamber (m<sup>2</sup>),  $P$  is the atmospheric pressure (kPa) of the sampling  
 170 sites (approximately 85.48 kPa at 1,390 m above sea level),  $P_0$  is atmospheric pressure under standard  
 171 conditions (101.325 kPa),  $T_0$  is the temperature under standard conditions (273.15 K),  $T$  is the temperature  
 172 (K) inside the chamber, and  $dC/dt$  is the average rate of concentration change with time (ppm min<sup>-1</sup>).

173 The total GHG efflux is estimated as the global warming potential (GWP) for a 100-year time  
 174 horizon, CO<sub>2</sub>-eq. One GWP of CH<sub>4</sub> accounts for 25CO<sub>2</sub>-eq and one GWP of N<sub>2</sub>O for 298 CO<sub>2</sub>-eq  
 175 (Intergovernmental Panel on Climate Change [IPCC], 2006). Water use efficiency (WUE) was calculated  
 176 according to Sun et al. (2018):  $\text{WUE} = \text{hay yield}/(\text{irrigation} + \text{precipitation} + \Delta\text{SWC}_{\text{October-April}})$ . The

177 precipitation and irrigation from April to October 2014 was 70 and 360 mm, respectively. GHG efflux  
178 intensity measuring the ratio of GHG effluxes per unit hay yield ( $GEI_{\text{hay}}$ ) was also estimated according to  
179 (Dyer et al., 2010):  $GEI_{\text{hay}}$  (kg CO<sub>2</sub>-eq/kg hay) = GHG efflux/hay yield.

180

## 181 **Statistical analysis**

182 All other statistical analyses were conducted using SAS 9.4 (SAS Institute Inc., Cary, NC, USA). Results  
183 of a Shapiro-Wilk test (UNIVARIATE Procedure) indicated that data collected from this study were  
184 normally distributed. The difference in hay yield and WUE, and total GHG efflux (CO<sub>2</sub>-eq) and  $GEI_{\text{hay}}$   
185 between different stand ages were analyzed using least significant difference test (LSD test, GLM  
186 Procedure). The correlations of soil properties (i.e., ST, SWC, SOC, STN) and RBM to GHG effluxes  
187 were determined (CORR Procedure). The variations of soil properties and RBM and hourly GHG fluxes  
188 in response to alfalfa stand age (y, year) and seasonal progress (m, month) were analyzed using a general  
189 linear model (GLM Procedure):  $\text{variation} = a + b \times m + c \times m^2 + d \times y + e \times y^2 + f \times m \times y$ , where  $a$  is intercept,  
190 and  $b$ ,  $c$ ,  $d$ ,  $e$  and  $f$  are estimated regression coefficients. The significant coefficients were only included  
191 in the final model. A stepwise multiple regression analysis was applied to determine the possible effects  
192 of soil properties and RBM on CH<sub>4</sub>, CO<sub>2</sub> and N<sub>2</sub>O fluxes (GLM Procedure) and the significant factors  
193 were only included in the final model. The proportional contributions of soil properties and RBM to CH<sub>4</sub>,  
194 CO<sub>2</sub> and N<sub>2</sub>O fluxes were then calculated as: the sum of squares for each test factor, divided by the total  
195 sum of squares then multiplied by the regression coefficient (i.e.,  $R^2$ ) of the model.

196

## 197 **RESULTS**

### 198 **Alfalfa biomass, GHG fluxes and soil property in relation to alfalfa stand age**

199 Both total annual hay yield and WUE significantly increased with the stand age from 2 to 5 years then  
200 significantly decreased after which time ( $LSD = 0.50$  and  $1.15$  respectively for hay yield and WUE,  $P <$   
201  $0.0001$ ) (Fig. 1). The first cutting respectively accounted for 56.1, 55.9, 55.2 and 59.5% of total annual

202 forage yield respectively from the 2-, 3-, 5- and 7-year-old fields, which was significantly greater than  
203 that of the second or third cutting ( $P < 0.05$ ).

204 The  $GEI_{\text{hay}}$  was significantly lower in 3- and 5-year-old fields than in 2- and 7-year-old ones ( $LSD =$   
205  $0.12$ ,  $P < 0.0001$ ) (Fig. 2A), and a significantly higher annual GHG efflux was detected in 5-year-old  
206 fields ( $LSD = 1.58$ ,  $P = 0.0030$ ) (Fig. 2B).

207 Alfalfa stand age had no significant effect on ST and SWC ( $F_{1,56} = 1.70$  and  $3.62$  respectively for ST  
208 and SWC,  $P > 0.05$ ) (Fig. 3A-B). However, ST significantly increased from mid-spring (April) until  
209 summer (July) ( $F_{1,57} = 322.89$ ,  $P < 0.0001$ ) and then significantly decreased after July ( $F_{1,57} = 358.83$ ,  $P <$   
210  $0.0001$ ) (Fig. 3A); while a reverse seasonal pattern was detected for SWC, i.e., it significantly decreased  
211 until July ( $F_{1,57} = 322.89$ ,  $P < 0.0001$ ) then significantly increased ( $F_{1,57} = 358.83$ ,  $P < 0.0001$ ) (Fig. 3B).

212 Both SOC and STN increased and peaked in 5-year-old fields ( $F_{1,55} = 5.28$  and  $13.76$  respectively for  
213 SOC and STN,  $P < 0.05$ ) then significantly decreased after which year ( $F_{1,55} = 5.75$  and  $14.08$  respectively  
214 for SOC and STN,  $P < 0.05$ ) (Fig. 3C-D). However, SOC significantly increased with seasonal progress  
215 and peaked in July ( $F_{1,55} = 13.09$ ,  $P = 0.0006$ ) after which month it significantly decreased ( $F_{1,55} = 13.61$ ,  
216  $P = 0.0005$ ) (Fig. 3C); but a reverse seasonal pattern was detected for STN, i.e., it significantly decreased  
217 from April to July ( $F_{1,55} = 7.78$ ,  $P = 0.0073$ ) then significantly increased ( $F_{1,55} = 9.35$ ,  $P = 0.0034$ ) (Fig.  
218 3D).

219 Both stand age and growing season initially promoted the RBM ( $F_{1,54} = 268.71$  and  $29.96$   
220 respectively for stand age and month,  $P < 0.0001$ ) (Fig. 3E). But the RBM started to decline after August  
221 ( $F_{1,54} = 269.42$ ,  $P < 0.0001$ ), and the decrease of RBM became slow ( $F_{1,54} = 24.00$ ,  $P < 0.0001$ ) due to  
222 significant positive interaction between stand age and seasonal progress ( $F_{1,54} = 23.79$ ,  $P < 0.0001$ ) (Fig.  
223 3E).

224 The dynamics of  $CH_4$ ,  $CO_2$  and  $N_2O$  fluxes also largely depended on alfalfa stand age and season  
225 (Fig. 4).  $CH_4$  uptake was detected in the present study and it significantly increased when alfalfa aged up  
226 to 5 years old ( $F_{1,54} = 36.24$ ,  $P < 0.0001$ ) then significantly decreased ( $F_{1,54} = 42.06$ ,  $P < 0.0001$ ) (Fig.  
227 4A). While  $CH_4$  uptake significantly decreased from April to July ( $F_{1,54} = 149.67$ ,  $P < 0.0001$ ) and then

228 significantly increased after July ( $F_{1,54} = 149.05$ ,  $P < 0.0001$ ) (Fig. 4A). The CH<sub>4</sub> uptake was generally  
229 higher (-15.4 – -25.0 µg/m<sup>2</sup>/h) in 5-year-old fields for a given month.

230 The seasonal and annual dynamics of CO<sub>2</sub> and N<sub>2</sub>O effluxes were similar, i.e., the effluxes  
231 significantly increased when alfalfa aged to 5 years old ( $F_{1,55} = 15.62$  and  $15.35$  for CO<sub>2</sub> and N<sub>2</sub>O  
232 respectively,  $P < 0.001$ ) and the significantly decreased ( $F_{1,55} = 13.03$  and  $13.32$  for CO<sub>2</sub> and N<sub>2</sub>O  
233 respectively,  $P < 0.001$ ); similarly the effluxes significantly increased since April ( $F_{1,55} = 322.37$  and  
234  $195.10$  for CO<sub>2</sub> and N<sub>2</sub>O respectively,  $P < 0.001$ ) and then significantly decreased after July ( $F_{1,55} =$   
235  $363.12$  and  $200.82$  for CO<sub>2</sub> and N<sub>2</sub>O respectively,  $P < 0.001$ ) (Fig. 4B-C). The greatest effluxes of CO<sub>2</sub>  
236 (i.e., 551.1 mg/m<sup>2</sup>/h) and N<sub>2</sub>O (8.0 µg/m<sup>2</sup>/h) were estimated in 5-year-old fields in July.

237

### 238 **GHG efflux in relation to soil property and root biomass**

239 The CH<sub>4</sub> uptake significantly decreased with increasing ST and SOC but increased with increasing SWC  
240 and STN (Table 1). While CO<sub>2</sub> and N<sub>2</sub>O effluxes significantly increased with increasing ST and SOC but  
241 decreased with increasing SWC and STN (Table 1). RBM had no significant effect on CH<sub>4</sub>, whereas CO<sub>2</sub>  
242 and N<sub>2</sub>O effluxes significantly increased with the increase of RBM (Table 1).

243 When both soil property and RBM were considered, SWC was the only factor that significantly  
244 affected CH<sub>4</sub> fluxes (Table 2). While three factors (i.e., SWC, ST and SOC) significantly affected CO<sub>2</sub>  
245 effluxes, and four factors (i.e., SWC, ST, SOC and RMB) significantly affected N<sub>2</sub>O effluxes (Table 2).  
246 SWC accounted for ≥ 65% variation of CO<sub>2</sub> and N<sub>2</sub>O effluxes. ST explained about 15% variation of CO<sub>2</sub>  
247 effluxes, which was 4.3 times less than did SWC but 3.5 times more than did SOC. For N<sub>2</sub>O effluxes, ST  
248 only accounted for only < 5% variation, which was 7.3 and 2.1 times less than did SWC and SOC,  
249 respectively, and RBM accounted for only < 2% of variation (Table 2).

250

## 251 **DISCUSSION**

252 A number of empirical studies have determined the optimal stand age of alfalfa in the semi-arid Loess  
253 Plateau and Inner Mongolia regions, while different experimental designs, field management and

254 geographic locations could generate divergent conclusions. For example, when considering the forage  
255 yield only, the optimal stand age varies from 3 to 5 years depending on the annual precipitation (i.e., 300  
256 – 500 mm) (e.g., Du, Long & Wang, 1999b; Zhang et al., 2004; Cheng, Wan & Wang, 2005). Jia et al.  
257 (2009) suggested that the optimal stand age should be 7 years if considering hay yield only but could be  
258 up to 15 years when considering WUE alone. In the present study, we show that 5 years may be the  
259 optimal alfalfa stand, for two reasons. First, both the hay yield and WUE were significantly greater in 5-  
260 year-old alfalfa fields (Fig. 1). Second, although the significantly higher total annual GHG effluxes (Fig.  
261 2B; also see Fig. 4B-C), the total annual GHG efflux rate per unit hay yield (i.e.,  $GEI_{\text{hay}}$ ) was significantly  
262 lower (Fig. 2A) and the net  $CH_4$  uptake was higher in the 5-year-old alfalfa fields (Fig. 4A). Therefore,  
263 the net  $CH_4$  sink may have largely offset the alfalfa  $CO_2$  and  $N_2O$  effluxes in the arid continental regions  
264 (Fig. 4). To our knowledge, this is the first study providing evidence for the assessment of optimal crop  
265 stage age integrating the total annual GHG effluxes and  $GEI_{\text{hay}}$ .

266 Alfalfa stand age affecting soil GHG effluxes is mediated by changing soil properties. Many authors  
267 have demonstrated that successive cropping alfalfa will elevate soil nutrient due to sequestering carbon  
268 (C) and nitrogen (N) into soils and eliminating tillage (Halvorson Wienhold & Black, 2002; Liang et al.  
269 2003). Our results show that SOC and STN continuously increased when alfalfa aged up to 5 years old  
270 (Fig. 3C-D), which agreed with previous studies (Xu, 2014; Cao et al., 2012). Alfalfa RBM also had a  
271 similar seasonal and annual pattern as SOC (Fig. 3C and 3E). Because new root develops primarily in the  
272 spring and root biomass increases as more C is fixed by the greater leaf areas associated with plant  
273 regrowth, greater amounts of C are translocated to the root system (Lee & Jose, 2003; Jiang & Claude,  
274 2006). However, root development and C-fixation may decline after extended dry periods, which  
275 decreases SOC at the end of growing season (Fig. 3E). The lower STN content detected during the warm  
276 seasons may attribute to the higher uptake of soil inorganic N by the growing plants. The decreasing  
277 SOC, STN and RBM in the 7-year-old fields indicate a declining plant vitality (Zhang et al., 2004; Cheng,  
278 Wan & Wang, 2005; Jia et al., 2009; Xu, 2014) resulting in a lower forage yield (Fig. 1A).

279 Soil property change may influence the source and sink function of greenhouse gases (Oertel et al.,

280 2016). Amount the abiotic factors, soil temperature and moisture are the two major drivers regulating  
281 GHG effluxes mainly via soil respiration and microbial activity (Kitzler et al., 2006; Singh et al., 2010;  
282 Butterbach-Bahl et al., 2013). As reported in a study in alfalfa fields in the dry Loess Plateau in China  
283 (Xu, 2014), we found that increasing soil temperature promoted CO<sub>2</sub> and N<sub>2</sub>O effluxes and suppressed  
284 CH<sub>4</sub> uptake (Table 2; Figs. 3 and 4). These results agree with the general conclusions of previous studies  
285 (e.g., Kitzler et al., 2006; Singh et al., 2010; Butterbach-Bahl et al., 2013; Zhu et al., 2016). Therefore, it  
286 may be prevalent that CO<sub>2</sub> and N<sub>2</sub>O effluxes start to increase in spring and peak in summer (Fig. 4),  
287 because the soil warming promotes soil respiration rate via microbial activity (e.g., faster growth rate and  
288 substrate use rate) (Kitzler et al., 2006; Singh et al., 2010; Xu, 2014).

289 Unlike temperature, moisture influences GHG fluxes via changing GHG diffusion rate and oxygen  
290 availability or regulating microbial communities because they require water for physiological activities  
291 (Singh et al., 2010). However, each soil type may have a specific soil moisture that optimizes GHG fluxes  
292 (Schindlbacher et al., 2004). When moisture exceeds the optimum level, gas transport is restricted  
293 (Schaufler et al. 2010; Kim et al. 2012), leading to anaerobic conditions; whereas suboptimal moisture  
294 levels will limit GHG fluxes due to water stress of soil microbes (Schindlbacher et al., 2004; Kitzler et al.,  
295 2006). By following this line, it may be predicted that at the arid conditions such as that of our experiment  
296 where SWC is below the optimum level, increasing SWC will promote CH<sub>4</sub> uptake and CO<sub>2</sub> and N<sub>2</sub>O  
297 emission as it elevates diffusivity of oxygen (O<sub>2</sub>) in soils which is essential for soil respiration and  
298 bacterial nitrification and methanotrophy under aerobic conditions. Our results partially support the  
299 assumptions. Increased SWC induced higher CH<sub>4</sub> uptake (Figs. 3B and 4A; Tables 1 and 2), agreeing  
300 with Dutaur and Verchot (2007) that methanotrophy is a dominant process in upland dry soils and there is  
301 thus a net uptake of CH<sub>4</sub> by soils. However, our results show that CO<sub>2</sub> and N<sub>2</sub>O emission decreased with  
302 increasing of SWC (Figs. 3B and 4B; Tables 1 and 2). Therefore, it is supposed that different GHGs have  
303 various thresholds of SWC invoking gas emission in arid regions, which are warranted for future studies.

304 Beside the main abiotic drivers of soil moisture and temperature, agricultural GHG fluxes are  
305 directly mediated by biotic factors including root respiration and microbial activity (Kitzler et al., 2006;

306 Singh et al., 2010; Butterbach-Bahl et al., 2013) of which are regulated by root biomass and soil nutrient  
307 such as carbon and nitrogen (e.g., Schindlbacher et al., 2004; Wang, Peng & Fang, 2010; Oertel et al.,  
308 2016). Indeed, some researchers have reported the positive correlations between soil respiration, RBM  
309 and SOC (Lee & Jose, 2003; Jiang & Claude, 2006) and between N<sub>2</sub>O efflux and SOC (Xu, 2014). Our  
310 results indicate that increasing SOC and RBM significantly elevated CO<sub>2</sub> and N<sub>2</sub>O effluxes (Table 1).  
311 According to Xu (2014), the greater GHG effluxes in 5-year-old fields (Figs. 2B, 3B, and 3C) may  
312 partially attribute to the higher microbial abundance and activity and root respiration owing to the higher  
313 soil nutrient and RBM (Fig. 3).

314 When compared the annual dynamics of STN (Fig. 3D) with that of N<sub>2</sub>O effluxes (Fig. 4C), we may  
315 assume that increasing STN (organic and inorganic N) in soils may enhance N<sub>2</sub>O effluxes via the  
316 biological processes of nitrification or denitrification (Xu, 2014; MacDonald, Farrell & Baldock, 2016;  
317 Oertel et al., 2016). However, our findings do not support the above notion, rather STN had a  
318 significantly negative effect on N<sub>2</sub>O effluxes (Table 1). In agricultural systems, plants only take up  
319 inorganic N (i.e., NO<sub>3</sub><sup>-</sup>-N and NH<sub>4</sub><sup>+</sup>-N) (Schmidt, Nasholm & Rentsch, 2014), but may use organic N  
320 through the processes of mineralization (bacteria digest organic material and release NH<sub>4</sub><sup>+</sup>-N) and  
321 nitrification (bacteria convert NH<sub>4</sub><sup>+</sup>-N to NO<sub>3</sub><sup>-</sup>-N) (Schmidt, Nasholm & Rentsch, 2014; Fernandez &  
322 Kaiser, 2018). The causes of negative effect of STN on N<sub>2</sub>O effluxes may be that the increasing STN  
323 (Fig. 3D) promotes the uptake of NH<sub>4</sub><sup>+</sup>-N and NO<sub>3</sub><sup>-</sup>-N by plants for growth (Ghimire, Norton & Pendall,  
324 2013), which reduces NH<sub>4</sub><sup>+</sup>-N available for microbial nitrification and results in less N<sub>2</sub>O effluxes.  
325 Considering the climate conditions in this study, the net N<sub>2</sub>O effluxes in the alfalfa fields may attribute to  
326 the processes of nitrification under aerobic conditions.

327 Although ST, SWC, SOC, STN and RBM alone had significant positive or negative effect on GHG  
328 effluxes (except RBM had no significant effect on CH<sub>4</sub> effluxes, Table 1), STN had little impact on GHG  
329 effluxes, and increasing RBM could significantly elevate N<sub>2</sub>O efflux, but its impact was very small, i.e.  
330 explained < 2% variation (Table 2). Agreeing with that of Schindlbacher et al. (2004) and Oertel et al.  
331 (2016), the flux rates of N<sub>2</sub>O and CO<sub>2</sub> largely depend on ST, SWC and SOC (Table 2). Generally, as

332 discussed above an increase of soil temperature will lead to greater effluxes and soil respiration rates as a  
333 positive feedback response of increased microbial metabolism (Kitzler et al., 2006; Singh et al., 2010;  
334 Butterbach-Bahl et al., 2013; Oertel et al., 2016). However, Fowler et al. (2009) stated that the positive  
335 temperature impact could be limited by soil water stress, as water is needed as a transport medium for  
336 nutrients required by microbes. In the current study, N<sub>2</sub>O and CO<sub>2</sub> effluxes were more sensitive to SWC  
337 than to ST and SOC, and CH<sub>4</sub> efflux responded only to SWC (Table 2).

338

## 339 **CONCLUSIONS**

340 Based on the forage yield, WUE, GHG efflux (CO<sub>2</sub>-eq) and GEI<sub>hay</sub>, we found that in the arid inland regions  
341 the optimal alfalfa stand age is 5 years. This knowledge is helpful in decision of alfalfa retention time based  
342 on the maximum benefit by considering the hay yield, water use efficiency, GHG emission as well as the  
343 cost of field establishment of alfalfa. Our results also indicate that in the arid regions with higher soil water  
344 stress, SWC overrides ST, SOC and RBM as a key factor regulating GHG fluxes and increasing SWC leads  
345 to an increase of net uptake of CH<sub>4</sub> by soils and a reduction of N<sub>2</sub>O and CO<sub>2</sub> effluxes from the alfalfa fields.  
346 Irrigation is required for alfalfa growing in the arid regions, while it also has significant impacts on GHG  
347 emission (Dutaur & Verchot, 2007; Schaufler et al., 2010; Ussiri & Lal, 2013; Burger et al., 2016).  
348 Therefore, future researches on the dynamics of GHG fluxes affected by irrigation (i.e., frequency, timing,  
349 and amount of irrigation water used) are warranted to develop strategies for GHG mitigation, increasing  
350 alfalfa forage yield and prolonging alfalfa persistence in the continental arid regions.

351

## 352 **ACKNOWLEDGEMENTS**

353 Financial support for this research was provided by Strategic Priority Research Program of Chinese  
354 Academy of Science (XDA20100102), National Key Basic Research Program of China  
355 (2014CB138706), Changjiang Scholars and Innovative Research Team in University (IRT17R50), and  
356 The 111 project (B12002). We thank two anonymous reviewers and the handling editor for their

357 constructive comments on earlier versions, which have significantly improved the paper. We also are  
358 grateful to Dr Charles West (Texas Tech University) for his valuable comments and time spent editing the  
359 English of a previous version of the paper.

360

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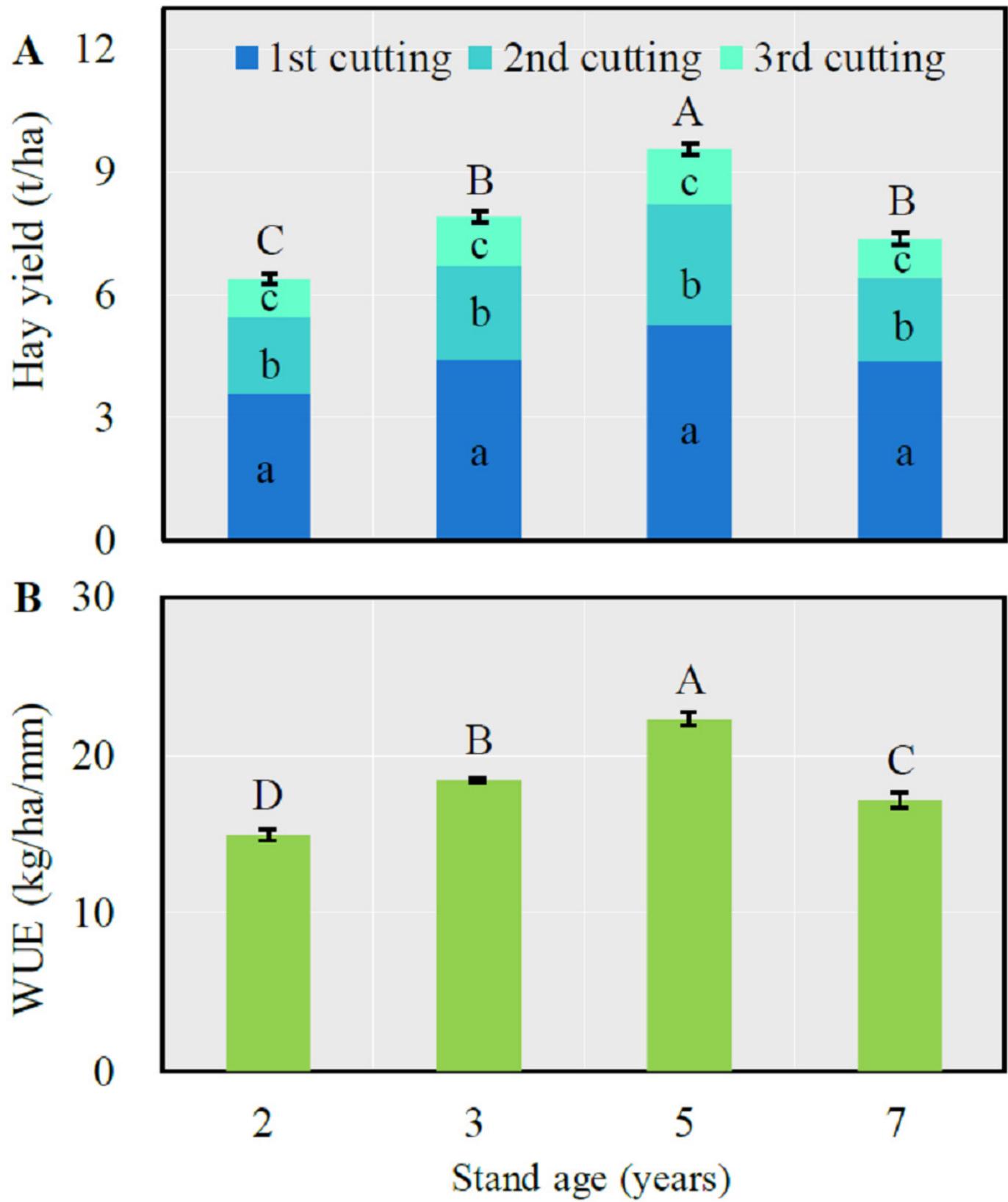
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# Figure 1

Mean ( $\pm$  SE) annual hay yield (A) and water use efficiency (WUE) (B) in alfalfa fields of different stand ages.

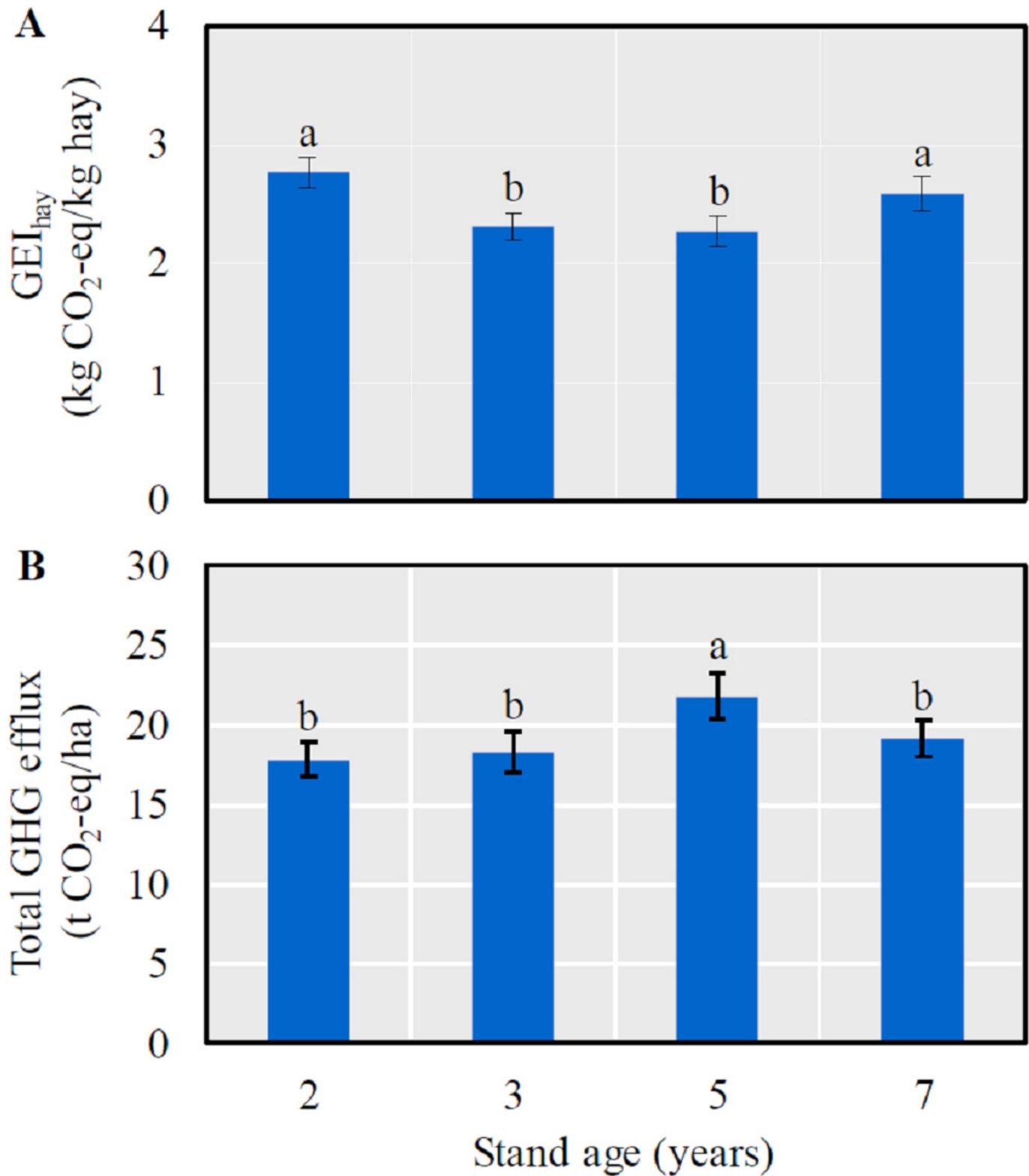
For the total hay yield (A) or WUE (B), columns with the same uppercase letters are not significantly different ( $P > 0.05$ ). For the hay yield of each cutting, columns with the same lowercase letters are not significantly different ( $P > 0.05$ ).



## Figure 2

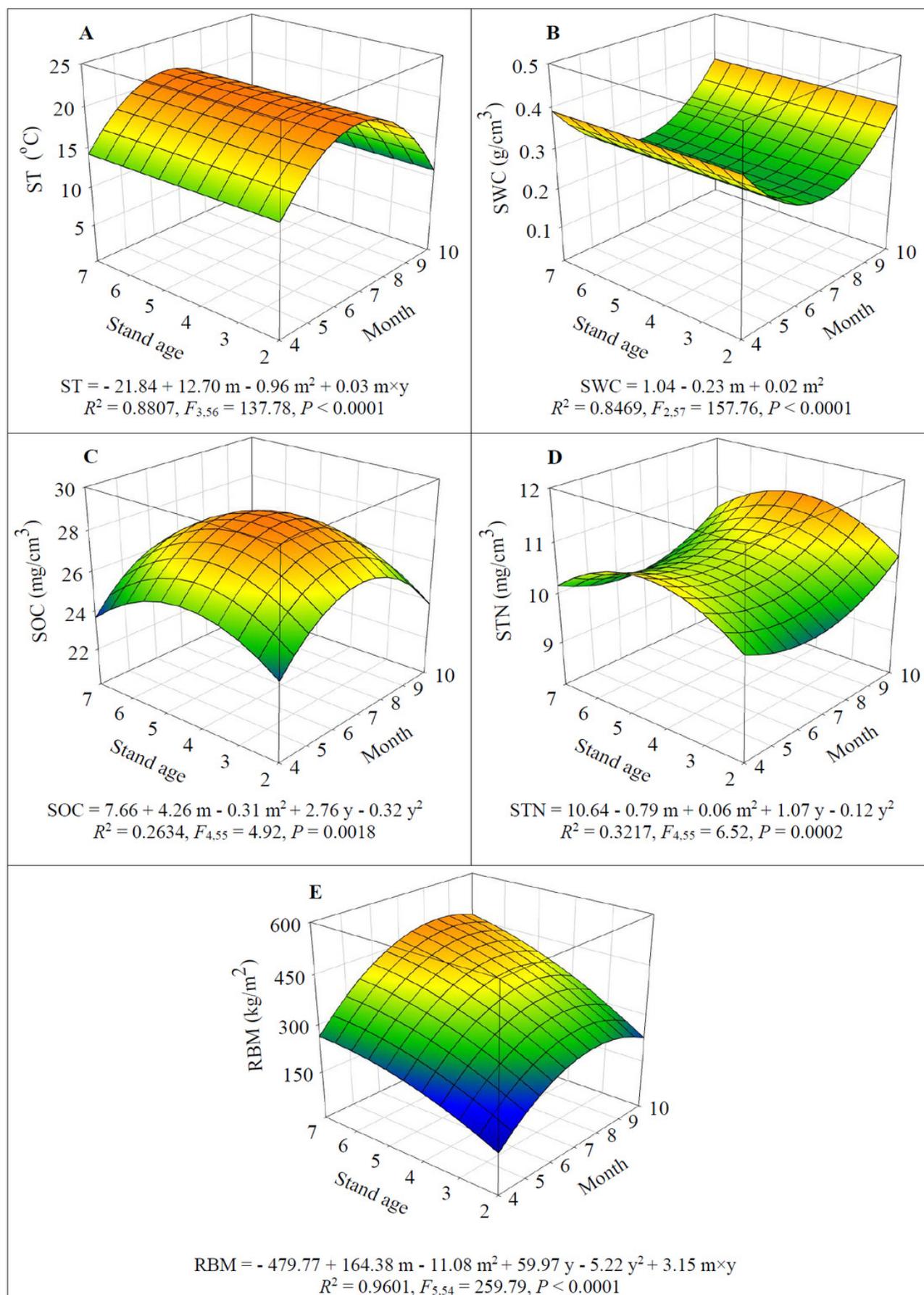
The mean ( $\pm$  SE) GHG efflux ( $\text{CO}_2$ -eq) per unit hay yield ( $\text{GEI}_{\text{hay}}$ ) (A) and total GHG efflux ( $\text{CO}_2$ -eq) (B) in alfalfa fields of different stand ages.

Columns with the same lowercase letters are not significantly different ( $P > 0.05$ ).



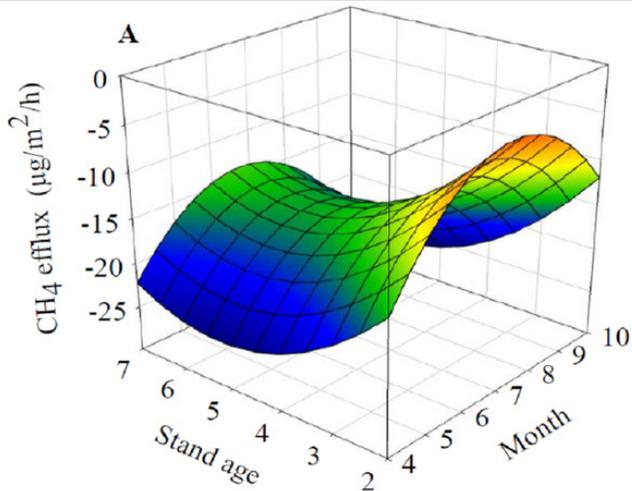
## Figure 3

Dynamics of soil temperature (ST) (A), soil water content (SWC) (B), soil organic carbon (SOC) (C), soil total nitrogen (STN) (D), and root biomass (RBM) (E) depending on alfalfa stand age (years,  $y$ ) and growing season (month,  $m$ ).



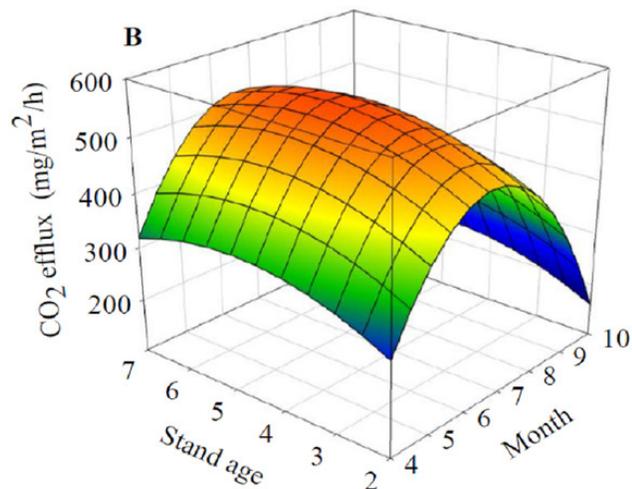
## Figure 4

Dynamics of GHG fluxes depending on alfalfa stand age (years,  $y$ ) and growing season (month,  $m$ ): CH<sub>4</sub> (A), CO<sub>2</sub> (B) and N<sub>2</sub>O (C).



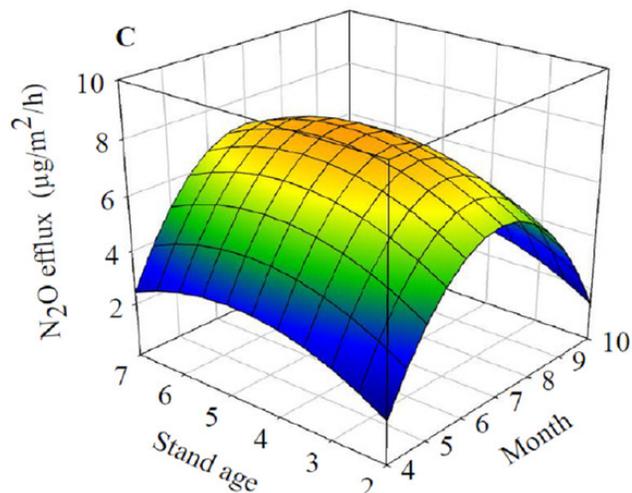
$$\text{CH}_4 = -47.73 + 16.47 m - 1.1063 m^2 - 8.86 y + 0.93 y^2 - 0.22 m \times y$$

$$R^2 = 0.8491, F_{5,54} = 60.77, P < 0.0001$$



$$\text{CO}_2 = -925.13 + 379.97 m - 28.50 m^2 + 85.21 y - 0.85 y^2$$

$$R^2 = 0.8846, F_{4,55} = 105.35, P < 0.0001$$



$$\text{N}_2\text{O} = -22.48 + 7.28 m - 0.52 m^2 + 2.08 y - 0.21 y^2$$

$$R^2 = 0.7990, F_{4,55} = 54.60, P < 0.0001$$

**Table 1** (on next page)

Soil CH<sub>4</sub> efflux (μg/m<sup>2</sup>/h), CO<sub>2</sub> efflux (mg/m<sup>2</sup>/h) and N<sub>2</sub>O efflux (μg/m<sup>2</sup>/h) correlated to soil temperature (ST, °C), soil water content (SWC, g/cm<sup>3</sup>), soil organic carbon (SOC, mg

<sup>ns</sup> non-significant correlation, \* < 0.05, \*\* < 0.01, \*\*\* < 0.001.

- 1 **Table 1** Soil CH<sub>4</sub> efflux (μg/m<sup>2</sup>/h), CO<sub>2</sub> efflux (mg/m<sup>2</sup>/h) and N<sub>2</sub>O efflux (μg/m<sup>2</sup>/h) correlated to soil  
 2 temperature (ST, °C), soil water content (SWC, g/cm<sup>3</sup>), soil organic carbon (SOC, mg/cm<sup>3</sup>), soil total nitrogen  
 3 (STN, mg/cm<sup>3</sup>) and root biomass (RBM, g/m<sup>2</sup>).

GHG	ST	SWC	SOC	STN	RBM
CH <sub>4</sub>	0.5027***	-0.6106***	0.3152*	-0.3882**	-0.1533 <sup>ns</sup>
CO <sub>2</sub>	0.8756***	-0.8054***	0.5958***	-0.3333**	0.3852**
N <sub>2</sub> O	0.8186***	-0.8200***	0.6210***	-0.2573*	0.5077***

- 4 <sup>ns</sup> non-significant correlation, \* < 0.05, \*\* < 0.01, \*\*\* < 0.001.

**Table 2** (on next page)

The final optimal general linear models (GLMs) and the contribution of factors .

Factors include soil temperature (ST, °C), soil water content (SWC, g/cm<sup>3</sup>), soil organic carbon (SOC, mg/cm<sup>3</sup>), soil total nitrogen (STN, mg/cm<sup>3</sup>) and root biomass (RBM, g/m<sup>2</sup>) to CH<sub>4</sub> flux (µg/m<sup>2</sup>/h), CO<sub>2</sub> flux (mg/m<sup>2</sup>/h) and N<sub>2</sub>O flux (µg/m<sup>2</sup>/h).

1 **Table 2** The final optimal general linear models (GLMs) and the contribution of factors.

GHG	Factor	df	Type I SS	Contribution (%)	<i>F</i>	<i>P</i>
CH <sub>4</sub>	SWC	1	1104.51	37.28	34.47	< 0.0001
	Error	58	1858.44	62.72		
CO <sub>2</sub>	SWC	1	696778.94	64.86	224.75	< 0.0001
	ST	1	158944.53	14.80	51.27	< 0.0001
	SOC	1	44915.24	4.18	14.49	0.0004
	Error	56	173616.20	16.16		
N <sub>2</sub> O	SWC	1	251.37	67.23	212.86	< 0.0001
	ST	1	16.38	4.36	13.87	0.0005
	SOC	1	34.43	9.21	29.15	< 0.0001
	RBM	1	6.75	1.81	5.72	0.0202
	Error	55	64.95	17.37		

Final models:

$$\text{CH}_4 \text{ flux} = 3.31 - 60.11 \text{ SWC} \quad (R^2 = 0.3728)$$

$$\text{CO}_2 \text{ flux} = - 29.50 - 473.10 \text{ SWC} + 17.45 \text{ ST} + 10.48 \text{ SOC} \quad (R^2 = 0.8384)$$

$$\text{N}_2\text{O flux} = - 2.27 - 11.98 \text{ SWC} + 0.24 \text{ SOC} + 0.21 \text{ ST} + 0.37 \times 10^{-2} \text{ RBM} \quad (R^2 = 0.8263)$$

- 2 Factors include soil temperature (ST, °C), soil water content (SWC, g/cm<sup>3</sup>), soil organic carbon (SOC, mg/cm<sup>3</sup>), soil  
3 total nitrogen (STN, mg/cm<sup>3</sup>) and root biomass (RBM, g/m<sup>2</sup>) to CH<sub>4</sub> flux (μg/m<sup>2</sup>/h), CO<sub>2</sub> flux (mg/m<sup>2</sup>/h) and N<sub>2</sub>O  
4 flux (μg/m<sup>2</sup>/h).  
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