1 **A specimen of *Paralycoptera* Chang & Chou 1977 (Teleostei: Osteoglossoidei) from Hong**
2 **Kong (Guangdong, China) with a potential Late Jurassic age that extends the temporal**
3 **and geographical range of the genus**
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16 **Abstract**   We identify an osteoglossoid teleost fish - *Paralycoptera* - from Late Jurassic
17 volcaniclastic mudstones from the Lai Chi Chong Formation of Hong Kong, China. This
18 partially preserved postcranial skeleton represents the first Mesozoic fish from Hong Kong
19 and the most southerly *Paralycoptera* to date. A radiometric ~~date~~ age for the Lai Chi Chong
20 Formation of ~146 Ma implies a temporal range expansion for *Paralycoptera* of
21 approximately 40 million years back from the Early Cretaceous (~110 Ma). However, spores
22 found in the Formation suggest an Early Cretaceous age that is consistent with the existing
23 age assignment to *Paralycoptera*. We argue that the proposed temporal range extension is
24 genuine because it is based on recent ~~precise and accurate~~high precision radiometric age data,
25 but given the discrepancies with the biostratigraphic ages further investigation is needed to
26 confirm this. This study provides an important step towards revealing Hong Kong's Mesozoic
27 vertebrate fauna and understanding its relationship to well-studied mainland Chinese ones.
28
29 **Keywords**   Jurassic, fish, osteoglossomorph, *Paralycoptera*, Hong Kong, Lai Chi Chong
30 Formation, lacustrine, volcanism
31
32 **1   Introduction**
33
34 In the summer of 2013, a fish fossil - SHGM L275 - labelled as a plant was discovered in the
35 collections of the Stephen Hui Geological Museum (SHGM) at the University of Hong Kong
36 (HKU). The fossil (~2 cm long) is hosted within a small mudstone fragment (5 cm by 3 cm)
37 that was supposedly collected from the Lai Chi Chong Formation (荔枝莊組) of Lai Chi
38 Chong, ~~Sai Kung~~Tolo Channel, north-eastern New Territories, Hong Kong (**Fig. 1**). This

39 provenance information is based on the specimen label, which appears to be correct, given
40 that all fossils with the same catalogue number are lithologically similar and match the
41 locality's expected lithologies (see Section 4). It is not mentioned in the literature - probably
42 because of its incorrect specimen label - unlike a fossil fish specimen from associated rocks in
43 Shek Pik (石壁), Lantau Island, which has a passing mention in Lee *et al*. (1997) (**Fig. 1**). The
44 latter specimen is supposed to be in the SHGM collections, but as it could not be located, it is
45 assumed to have been lost. A fossil fish is also known from Lantau Peak (鳳凰山), Lantau
46 Island (C.M. Lee, personal communication, July 19th, 2014) (**Fig. 1**). The rocks from this site,
47 Shek Pik and the Lai Chi Chong Formation all belong to the Lantau Volcanic Group, so all
48 three fossils should have similar ages (Sewell *et al.* 2000) (**Fig. 1**). The fish fossil
49 assemblages of Lai Chi Chong and Lantau Island are therefore important to compare, but
50 unfortunately, the whereabouts of the Lantau Peak fossil is also unknown, so this comparison
51 is not currently possible.

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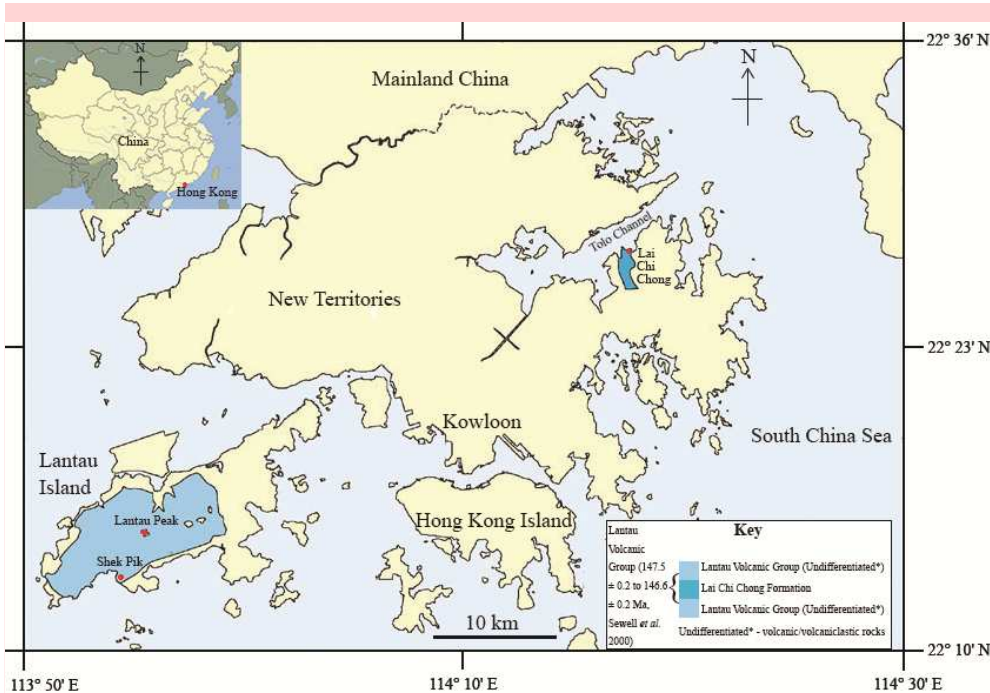
Figure 1: A map of Hong Kong showing the location of the Lai Chi Chong Formation in ~~Sai Kung~~Tolo Channel, as well as the broader Lantau Volcanic Group. The ~~approximate~~ locations of fossil fish discoveries in Hong Kong are also marked in red. Scale bar = 10 km (modified from Sewell *et al.* 2000).

58 The plant fossils discovered within the Lai Chi Chong Formation e.g. *Cyathidites*,
59 *Classopollis* and *Cicatricosisporites* suggest that it has an Early Cretaceous age (Lee *et al.*

Comment [s1]: The stratigraphy for Lantau Peak has been superseded by Campbell et al. (2007). Shek Pik and the Lai Chi Chong Fm belong to the Lantau Volcanic Group, but the upper part of Lantau Peak belongs to the Kau Sai Chau VG, (141 Ma) and the lower part to the Repulse Bay VG (143 Ma). See also Sewell et al. 2012, G-cubed for a composite section.

**Comment [s2]:** Suggest that you show the location of Cheung Sheung.

Comment [s3]: Sai Kung is not labelled on the diagram, and it is more accurate to say Tolo Channel.

Comment [s4]: Apart for the fossil locality at Shek Pik, which is shown on the published geological map sheet 13, the locations of the other fish fossils are rather vague.

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1997). However, high precision U-Pb single crystal zircon dating of coarse crystal~~line~~ tuff from the upper Lai Chi Chong Formation suggests that the Formation is actually 146.6 ± 0.2 million years old, which corresponds to the Tithonian stage of the Late Jurassic (Campbell *et al.* 2007), some 40 million years earlier. The high sampling and analytical standards applied to obtaining the radiometric age for the ~~that were met in obtaining the absolute age at~~ Lai Chi Chong Formation (see Campbell *et al*. 2007 for details) suggests that its ~~numerical age~~ is unequivocal and that the plant fossil evidence deserves further detailed investigation.

This study focuses on the identification of SHGM L275 and understanding its ecology in the context of the palaeoenvironment of the Lai Chi Chong Formation, that has been inferred from its geology and plant fauna (Lee *et al.*1997) (**Fig. 2**).
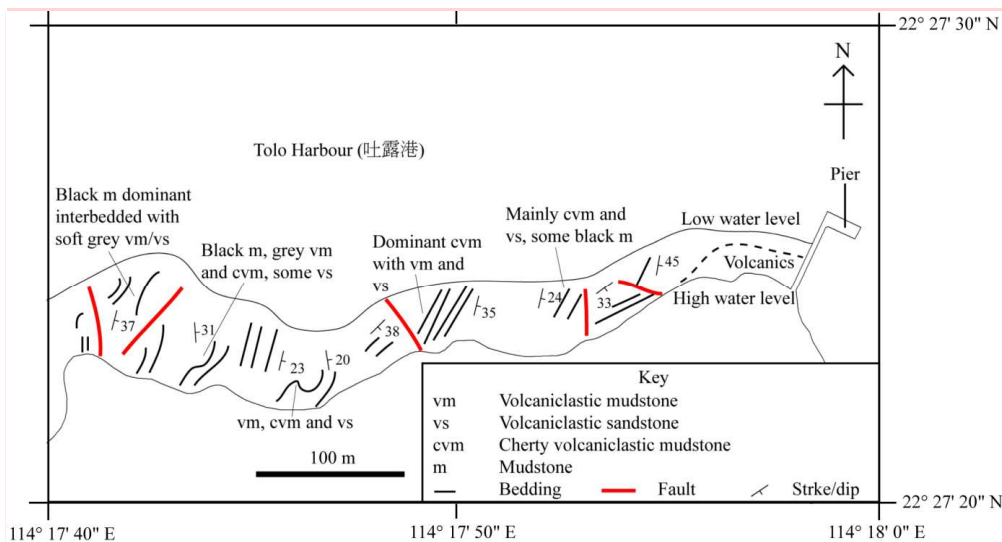
Figure 2: A simplified geological field sketch of the type locality of the Lai Chi Chong Formation at Lai Chi Chong, ~~Sai Kung~~Tolo Channel, NE New Territories, Hong Kong (simplified from Workman 1991).

2   Methods and Materials

The studied fossil specimen, SHGM L275, is a partially-preserved articulated bony fish skeleton that is missing its anterior portion (**Fig. 3**). The specimen shows the dorsal, anal and caudal fins and is preserved in a laminated, non-fissile, pale grey orange-spotted mudstone from the Lai Chi Chong Formation (**Fig. 3**). SHGM L275 is now deposited in the collections of the Stephen Hui Geological Museum (SHGM) at the University of Hong Kong. The specimen was prepared mechanically using a thin needle and was examined under a *Leica*

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S8APO stereomicroscope which has a magnification range of 10-80x. Photographs were taken of the specimen using a range of focal points with a Nikon D610 DSLR camera mounted to the stereomicroscope. The images were uploaded into the open-access computer software *CombineZP* (www.combinezp.software.informer.com/) to focus-stack them into fully-focused composite images. Based on a preliminary comparative study against Nelson (2006), SHGM L275 was diagnosed as an osteoglossomorph fish based on the possible presence of an epural and 15 principal branched caudal fin rays (Xu & Chang 2009). The specimen was then compared by standard methods with other Chinese Mesozoic osteoglossomorph fish from the collections of the Institute Vertebrate Paleontology and Paleoanthropology (IVPP; Beijing, China) and the Stephen Hui Geological Museum (SHGM (HKU), Hong Kong) (see **Supplemental Table S1** in the Supplementary Information). The specimen's features were then coded against character lists from osteoglossomorph-specific phylogenetic analyses (Shen 1996; Zhang 2006; Li & Wilson 1996; Wilson & Murray 2008; Xu & Chang 2009). The review of the osteoglossoid osteoglossomorph *Paralycoptera* by Xu & Chang (2009) was particularly important towards the referral of SHGM L275 to this genus because of its details on anatomical variation.
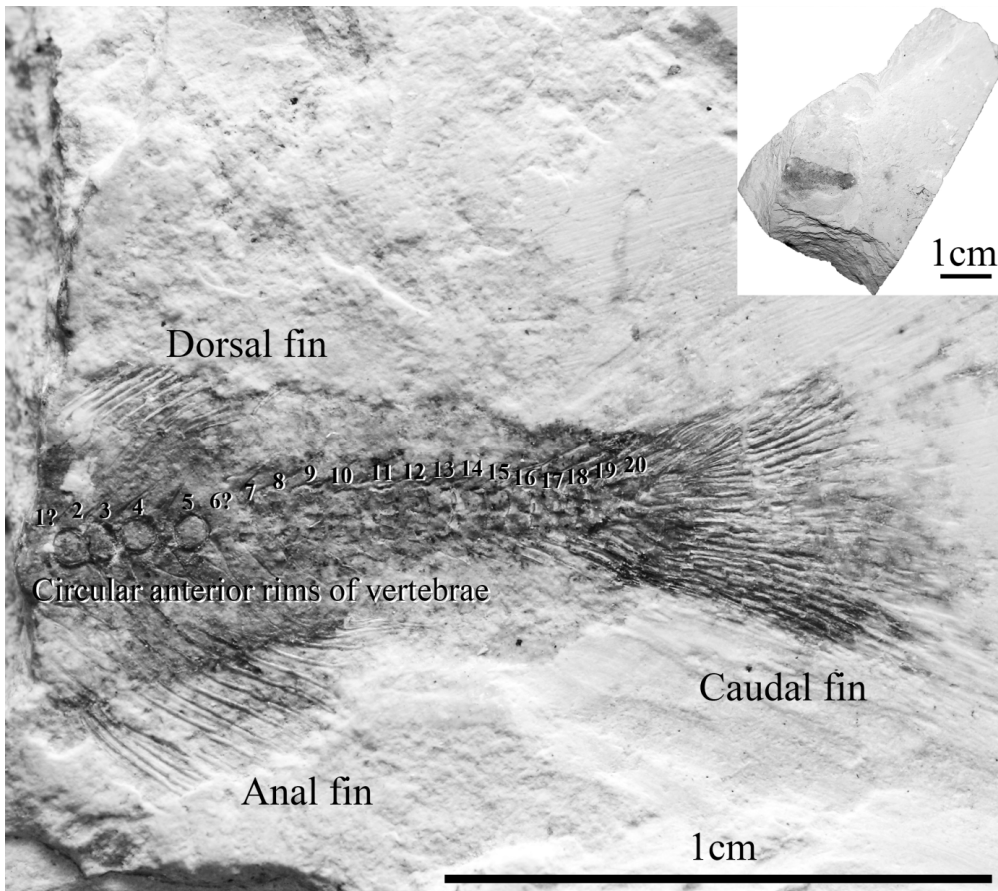
Figure 3: Magnified image (10.5x) of the specimen SHGM L275. The upper right image was taken before further preparation. The circular features in the anterior portion of the specimen appear to be the anterior rims of vertebrae. Identified vertebrae are numbered from 1 to 20, with 1 being an abdominal vertebra, and the rest (19) – caudal vertebrae. Scale bar = 1 cm.

3   Results

The specimen SHGM L275 is a rather small fish. The preserved part corresponds to the caudal portion of the fish, with the head and abdominal portion missing. The total length of the preserved part, including the caudal fin, is approximately 18 mm. Twenty vertebrae are identified in this portion (**Fig. 3**) between the anterior part of the dorsal and anal fins and the caudal fin, with the anteriormost preserved one being an abdominal vertebra, and the rest (19) – caudal vertebrae. The number of caudal vertebrae is comparable to many stem osteoglossomorphs, like *Huashia gracilis and Jinanichthys longicephalus* (Wilson & Murray 2008). Most of the vertebral centra are dorsoventrally deeper than anteroposteriorly long,

which may allow easier lateral movements during propulsion, as in most fish. In the anterior part of the specimen, there are four circular features directly on the vertebral column (**Fig. 3**) - these are the anterior rims of the vertebrae. This feature is also identified in the osteoglossoid osteoglossomorph *Paralycoptera wui*, IVPP V2989.100 (**Fig. 4**), and in other studied osteoglossomorph specimens like *Yanbiania wangqingica*, IVPP V6767-1, and *Tongxinichthys microdus*, IVPP 2332.1 (Wilson & Murray 2008).
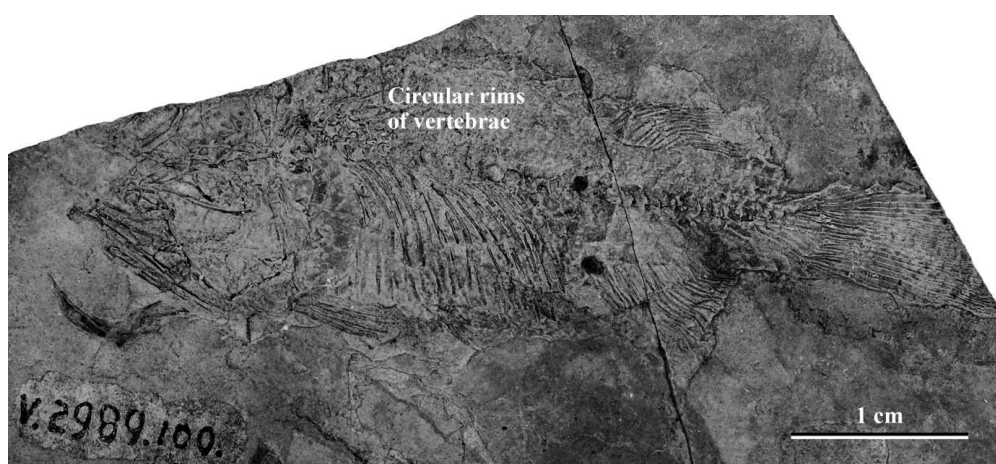

Figure 4: *Paralycoptera*, IVPP V2989.100, has a partially disarticulated vertebral column that reveals numerous circular vertebral rims (most of them are impressions), as in SHGM L275. Scale bar = 1 cm.

In SHGM L275, the anal fin is larger than the dorsal fin like those in *Paralycoptera wui* (Chang & Chou 1977; Xu & Chang 2009). Seventeen fin rays were observed in the anal fin whilst 10 were observed in the dorsal fin, although the actual number of fin rays may be higher because the anterior ends of both fins are incomplete (**Fig. 5**). However, the fin ray counts - as they are - are the same as those of *Paralycoptera wui* IVPP V2989.100 and .105, although the fins of the latter specimen are also incomplete, as in SHGM L275. For both the anal and dorsal fins, the lengths of the fin rays are longer in the anterior portion of the fin than in the posterior portion giving them a sub-triangular shape. The preserved anterior margins of the anal and dorsal fins are opposite to each other and are rather close to the caudal fin suggesting that the dorsal fin is posteriorly situated along the fish. Such features, together with the shape of the fins, are seemingly similar to the posterior portion of *Lycoptera*, but in the latter taxon the size difference between the anal and dorsal fin is not significant compared to SHGM L275. Between the fins and the vertebrae, pterygiophores supporting the fin rays are observed (**Fig. 5**). The number of pterygiophores is more or less the same as the number of fin rays because the ends of each pterygiophore preserved leads to the base of a fin ray.
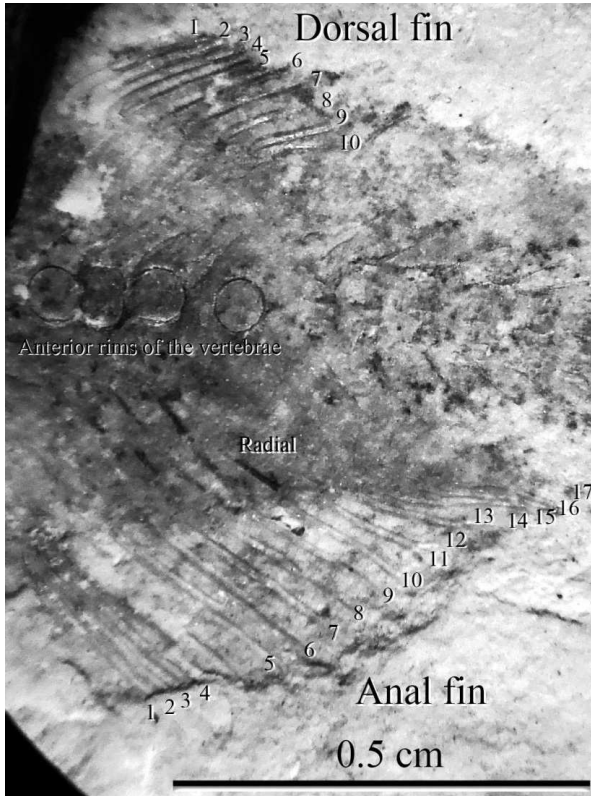
Figure 5: Magnified image (10.5x) of the anterior portion of SHGM L275 showing the position of the fin rays in the anal and dorsal fins. Scale bar = 0.5 cm.

In the caudal skeleton of SHGM L275 (**Fig. 6**), six hypurals were identified. The first one is posteriorly broader, giving a fan-like shape, whereas the second is comparatively narrow. These hypurals articulate with the first ural centrum, and support the rays of the lower lobe of the caudal fin. Under the first hypural, the parhypural, articulating with the preural centrum 1, also has a somewhat fan-shaped broader posterior portion. The second ural centrum is triangular in shape and is slightly upturned towards the upper lobe of the fin. The third to sixth hypurals are rectangular rod-shaped, articulating with the second ural centrum, and supporting the rays of the upper lobe of the caudal fin. Comparing ural centrum 2 with ural centrum 1 and neighbouring vertebral centra, ural centrum 2 is anteroposteriorly longer than dorsoventrally deep whilst the others are dorsoventrally deeper than anteroposteriorly long. In the area above the ural centra 1 and 2, traces of uroneurals can be seen, though it is difficult to estimate their number (possibly two or three). The anterior tip(s) of the uroneurals extend to the posterodorsal end of the preural centrum 1. An epural is probably present above the

uroneurals. No urodermals were found.

Even though the caudal fin rays are poorly-preserved, the caudal fin appears to be symmetrical because the vertebral column only bends slightly towards the upper lobe. Thus, the specimen is likely to possess a homocercal tail, which is a trait of all teleostean fish (Nelson 2006). We were able to find out the approximate counts of the caudal fin rays: 17 principal fin rays are recognized, seven branched rays with one unbranched ray at the upper margin in the upper lobe, and eight branched rays with one unbranched ray at the lower margin in the lower lobe. Besides, about 5-6 and 3-4 short, procurrent rays are observed in front of the upper and lower lobe respectively. Five neural spines on the 1st-5th preural centra and four haemal spines under 2nd-5th preural centra are prolonged, the posterior ones of which are in support of the procurrent rays. The ural centrum 1, perhaps, also carries a short neural spine (**Fig. 6**).

A

B
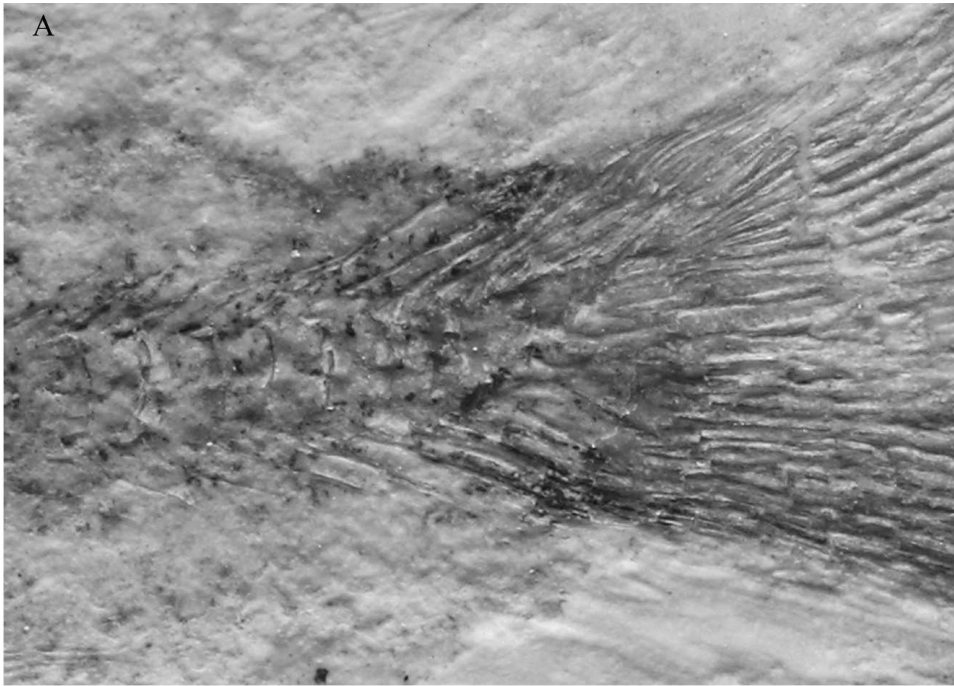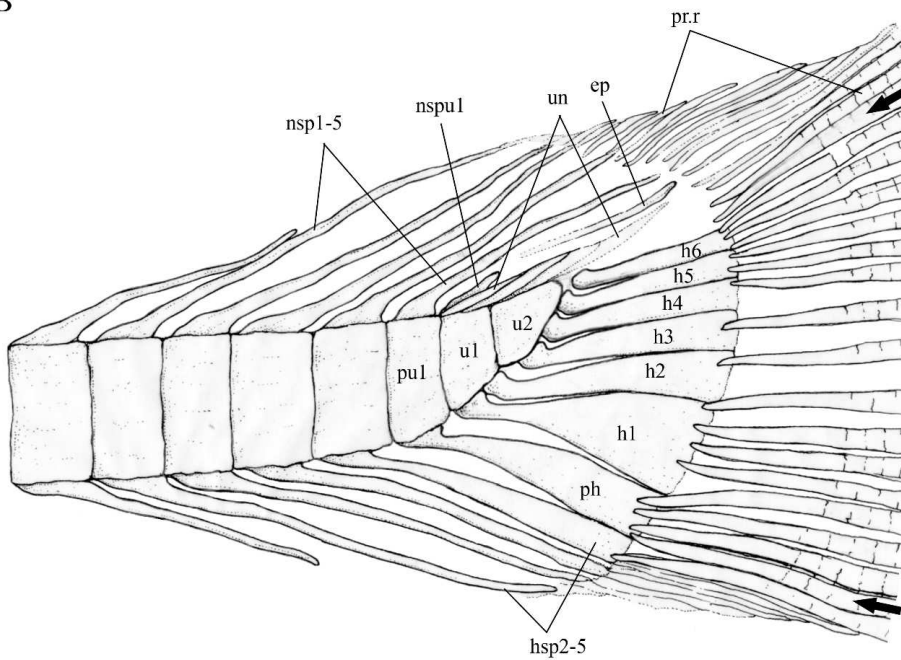
Figure 6: Magnified image of the caudal skeleton and bases of caudal fin rays in SHGM

L275, the arrows point to the outermost (unbranched) principal caudal fin rays. Abbreviations: ep, epural; h1-6, hypurals 1-6; hsp2-5, haemal spines on preural centrum 2-5; nsp1-5, neural spines on preural centrum 1-5; nspu1, neural spine on u1; ph, parhypural; pr.r, procurrent rays; pu1, preural 1; u1, u2, ural centra 1 and 2; un, uroneurals. Scale = 1 mm.

Based on the features described above, especially that a possible epural is present, the number of branched caudal fin rays is 15 and the dorsal fin is posteriorly situated, SHGM L275 most likely belongs to the order Osteoglossiformes (Shen 1997, Xu & Chang 2009), under the superorder Osteoglossomorpha (Greenwood *et al*. 1966).

Comment [s6]: Not in references

SHGM L275 was added to the osteoglossomorph phylogenetic data matrices of Shen (1996), Zhang (2006), Wilson & Murray (2008), and Xu & Chang (2009) (**Table 1**) and in all four analyses the taxon that has the most similar codings was *Paralycoptera*. However, for the Zhang (2006) matrix, SHGM L275 has more closely matched codings to *Singida* than to *Paralycoptera*. The Eocene temporal range of *Singida* (Xu & Chang 2009) is at odds with the Late Jurassic age of SHGM L275, but it might be possible that the new specimen supports an extremely large range extension. However, SHGM L275 is referable to *Paralycoptera* based on additional details of the caudal skeleton: the two hypurals in the lower lobe of *Paralycoptera* are separated and unfused (Shen 1996) like in SHGM L275, whilst those in *Singida* are partially fused (Murray & Wilson 2005). In addition, *Singida* has a falcate anal fin instead of the triangular one in *Paralycoptera* (Murray & Wilson 2005) and SHGM L275.

Table 1: The applicable characters from Shen (1996), Zhang (2006), Wilson & Murray (2008) and Xu & Chang (2009) to SHGM L275, coding similarities with the most closely-matched genus – *Paralycoptera* (○ = matched; x = not matched). For the codings of the individual studies please see **Supplemental Tables S2-S5** in the Supplemental Information).

Osteoglossomorph study	Equivalent character numbers						
Shen (1996)	28	29	33	34	35	36	
Zhang (2006)	47	48	49	53	54	61	60
Wilson & Murray (2008)	69	67	68	71		65	64
Xu & Chang (2009)	54	55	56			62	
SHGM L275 compared to *Paralycoptera*	○	x	x	○	○	○	○

In comparing SHGM L275 and *Paralycoptera* based on the above analyses (**Tables 1, S2-S5**), there were a few character state discrepancies. These unmatched characters include: (1) the condition of the neural spine on ural centrum 1 – whether the neural spine is complete or rudimentary, and (2) the number of epurals. According to Wilson & Murray (2008), the neural

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spine on the first ural centrum of *Paralycoptera* should be absent or rudimentary, whereas Shen (1996) and Xu & Chang (2009) observe a completely developed neural spine. Zhang (2006) is uncertain about the relative development of this spine, but in SHGM L275 a rudimentary neural spine is present. The number of epurals present in *Paralycoptera* remains controversial. Shen (1996) identified a single epural in *Paralycoptera* whereas Xu & Chang (2009) noted its absence. In specimens IVPP V2989.65, .100 and .105 of *Paralycoptera*, we also identified no epurals, like Xu & Chang (2009). An 'x' was been marked in **Table 1** for this character, even though the character state used by Xu & Chang (2009) - 'one or absent' - should justify the use of a '○' mark instead. We therefore advocate the separation of this state in future work in accordance with Greenwood (1970) and the epural characters of Shen (1996), Zhang (2006) and Wilson & Murray (2008). There is a possible epural in SHGM L275. Zhang (2006) and Wilson & Murray (2008) both record uncertainty in the number of epurals in *Paralycoptera*. The first preural centrum of SHGM L275 has a complete neural spine, as identified in *Paralycoptera* by all four aforementioned analyses, but Xu & Chang (2009) mistakenly recorded a 'rudimentary or absent' neural spine in their data matrix. Excluding the aforementioned discrepancies, the four studies otherwise converge on SHGM L275 being a specimen of *Paralycoptera*. However, Xu & Chang's (2009) observations of individual anatomical variation within *Paralycoptera* actually explain the differences in the caudal skeleton observed by Shen (1996), Zheng (2006) and Wilson & Murray (2008). This therefore confirms that SHGM L275 is a specimen of *Paralycoptera* (**Fig. 7**), which in our opinion negates the need for a numerical phylogenetic analysis. Xu & Chang (2009) synonymised the genus into one species *P. wui* whose features in SHGM L275 are:

> (1) a completely developed neural spine on the first preural centrum;
> (2) two or three uroneurals;
> (3) four upper hypurals and two lower hypurals, and
> (4) all hypurals are independent.
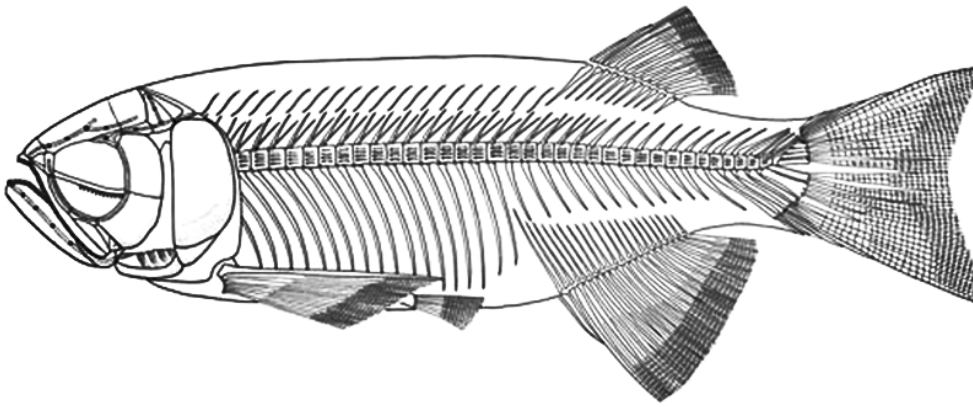
Comment [s7]: Not in references

Figure 7: Reconstruction of *Paralycoptera* (Xu & Chang, 2009; used with the permission of the authors).

Systematic Palaeontology

SUBDIVISION TELEOSTEI MÜLLER, 1846
SUPERORDER OSTEOGLOSSOMORPHA GREENWOOD *ET AL.*, 1966
ORDER OSTEOGLOSSIFORMES REGAN, 1909
SUBORDER OSTEOGLOSSOIDEI REGAN, 1909
GENUS †*PARALYCOPTERA* CHANG & CHOU, 1977
†*PARALYCOPTERA* sp. CHANG & CHOU, 1977

4 Discussion

4.1 Ecology of Paralycoptera

Paralycoptera is a member of both northern China's *Lycoptera* Fauna and south-eastern China's *Mesoclupea* Fauna (Chang & Jin 1996). It has been discovered in Jilin, Liaoning, Shandong, Zhejiang and Fujian provinces (Xu & Chang 2009) and now in Guangdong Province too (this study) (**Fig. 8**). This geographical range is impressive given that the northern part of China has been separated from the south by the Qinling-Dabie Shan orogenic belt since the Late Triassic (Hacker *et al.* 2004), and the 20° difference in latitude between the southernmost and northernmost localities - Hong Kong, Guangdong Province and Tonghua, Jilin Province respectively - a distance of over 2000 km. This geographic distribution may imply that *Paralycoptera* was adaptable to a wide range of environments compared to other members of the two faunas. However, climate variability over this geographical area was not very significant in the Late Mesozoic - climate change towards more temperate and humid

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conditions is reflected by geochemical weathering indices (Ohta *et al.* 2014) with
temperatures between 5 °C and 25 °C reconstructed from oxygen isotope data from
sedimentary rocks in north-eastern China (Wang *et al.* 2013). However, occasional semi-arid
periods are indicated by the appearance of the arid plants *Ephedripites* and *Classopollis* in
Hong Kong (Lee *et al.*1997), as well as oxygen isotope data from sedimentary rocks in
north-eastern China, including from Jilin and Liaoning provinces (Wang *et al.* 2013).
Therefore, *Paralycoptera* most likely lived in areas with a tropical-subtropical climate similar
to many modern osteoglossoids, such as *Scleropages formosus* (Kottelat 2011).

Figure 8: The locations of *Paralycoptera* discoveries within China (Locations from Chang &
Miao, 2004).

Paralycoptera localities were all continental basins (**Fig. 8**) where fluvial or lacustrine
deposits dominated (Chang & Jin 1996) and these have similar lithologies (see **Table 2**).
Vigorous tectonic activity and episodes of volcanism were common in these localities during
the late Mesozoic (Chang & Jin 1996; Chang & Chou 1977; Li & Li 2007). The Lai Chi

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Chong Formation of Hong Kong consists of mainly tuff and tuffaceous sedimentary rocks (Lai *et al.* 1996). A shallow freshwater lake environment subject to the influence of volcanic activity is indicated by fluvial-lacustrine and volcaniclastic sedimentary facies (Strange Shaw & Addison,1990; Workman 1991; Sewell *et al.* 2000), predominantly turbidites (Lai *et al.* 1996), and the discovery of terrestrial freshwater plant fossils including *Equisetites*, *Cladophlebis exiliformis*, *Gleichenites gladiatus* and *Carpolithus* (Lee *et al.*1997). According to Lin & Lee (2012), the 'parallel laminated fine sandstone and mudstone' facies is the most likely origin of SHGM L275 as the only light-coloured mudstone unit is confined to this facies (grey volcaniclastic mudstone from the western portion of the Lai Chi Chong locality; **Fig**. **2**). This facies contains fine-grained, cross-laminated, white and grey coloured mudstone representing a depositional environment below the wave base, where suspension currents might affect deposition (Lin & Lee 2012). The similarities in the palaeoenvironment between Lai Chi Chong and existing *Paralycoptera* localities (Workman 1991; Chang *et al.* 2008; Hu *et al.* 2012; P.J. Chen 1983) provides additional support for the inference that *Paralycoptera* from Lai Chi Chong lived in shallow freshwater lakes near areas of active volcanism. One potential hypothesis to explain the association of *Paralycoptera* discovery sites with volcanism is that *Paralycoptera* may have thrived on the higher nutrient levels in the lake caused by the influx of volcaniclastic material, and/or the warmer water temperatures provided by thermo-tectonic activities. The sedimentary rocks preserved at Lai Chi Chong frequently show syn-sedimentary structures including microfaults, slumps, convolute bedding, load and flame structures, suggesting the occurrence of mass flows that might have been triggered by episodic volcanic and seismic activity directly related to the local subduction tectonic setting (Sewell *et al.* 2000). This implies that the habitat of *Paralycoptera* was subjected to episodic catastrophic events and was not a prolonged quiet, tranquil water body. This habitat is possibly similar to the turbid and swift-water habitat of Hiodontiformes - a closely related group to Osteoglossiformes (Gray 1988). These episodic conditions could indicate that *Paralycoptera* had a high tolerance to environmental stress (highly variable sediment and nutrient input and possible changes in water temperature). However, the association of the fish with volcanism may more simply reflect the higher fossil preservation potential by volcaniclastic sediments, especially give that only one specimen is known among the strata so far. Crucially, the laminated mudstone that SHGM L275 is preserved in represents a relatively stable rather than unstable depositional setting. This also fits the living environments of most modern osteoglossoid fish which tend to prefer still water bodies e.g. *Pantodon buchholzi* and *Scleropages formosus* (Moelants 2010; Kottelat 2011). It therefore seems more plausible that *Paralycoptera* lived in relatively stable water body like their modern counterparts and probably migrated in times of environmental stress (no evidence of mass fish mortality in the rocks showing synsedimentary structures).

Table 2: Lithological characteristics of the formations preserving *Paralycoptera*.

Formation	Province	Age	Major lithology (those that yield *Paralycoptera* are in bold font)
Lai Chi Chong	Guangdong	~146 Ma, Tithonian, Late Jurassic (Campbell *et al.* 2007)	Light grey thickly laminated **tuffaceous mudstone**, massive black cherty mudstone, alternatively light and dark thickly laminated and cross-bedded coarse sandstone, conglomerate; greenish grey fine ash crystal tuff and rhyolite (Lai *et al.* 1996; Lin & Lee 2012)
Fenshuiling	Shandong	Late Jurassic to Early Cretaceous (Li 1998)	Mudstone, shale, siltstone, sandstone, conglomerate and tuff (Wang 1985)
Guantou	Zhejiang	~110 Ma, Early Cretaceous (Xu & Chang 2009)	Purplish grey, greyish green and greyish yellow **tuffaceous siltstone**, dark grey **mudstone**, purple **sandstone**; andesite and tuff breccia (Q.S. Chen 1983; Hu *et al.* 2012)
Hengtongshan	Jilin	Early Cretaceous (Han *et al.* 2013)	Black mudstone, oil shale and tuffite (Han *et al.* 2013)
Baiyashan	Fujian	Early Cretaceous (Zhang 2009)	Purplish red conglomerate, siltstone and sandstone (Zhang 2009)

4.2 Geographical distribution of Paralycoptera and the biogeography of the Mesoclupea Fauna

The discovery of *Paralycoptera* in Hong Kong extends the geographical range (**Fig. 8**) of the genus ~700 km further south of the previously most southerly locality in the Baiyashan Formation of Fujian Province (Xu & Chang 2009). This implies that *Paralycoptera* was much more widespread than previously thought and suggests that the genus may also be present in other similarly-aged lacustrine deposits in southeastern China (**Fig. 8**). *Paralycoptera* is a typical member of the *Mesoclupea* Fauna (Chang & Jin 1996) so it is possible that the other members of this fauna such as *Mesoclupea*, *Sinamia* and *Paraclupea* could be found in Hong Kong in the future.

4.3 Age of the Lai Chi Chong Formation and osteoglossomorph evolution and biogeography

Another implication of SHGM L275 arises from the age of the Lai Chi Chong Formation. A Jurassic age was originally proposed by Workman (1991) based on the identification of the fossil plants *Cladophlebis* and *Equisetites*. However, subsequent studies of spore fossils (including *Cicatricosisporites*, *Klukisporites*, *Cyaathidites*, *Classopollis* and *Pinuspollenite*) from exposures of the Formation at Cheung Sheung (嶂上), ~~Sai Kung~~ - ~2.5 km south of Lai Chi Chong - suggest that the Formation was deposited between the Valanginian to Barremian stages of the Early Cretaceous (Lee *et al.* 1997) (**Table S6**). This age determination is closer to the Aptian age of other *Paralycoptera* specimens found elsewhere in China, based on absolute dating of volcanic units (Xu & Chang 2009) (**Table 2**). However, as mentioned in the introduction (see section 1), an Early Cretaceous age is not corroborated by the Late Jurassic radiometric age of the Formation (Campbell *et al.* 2007). The small degree of uncertainty in the radiometric data (see Campbell *et al.* 2007) suggests that the Formation - and so SHGM L275 - date to the Tithonian stage of the Late Jurassic (**Fig. 9**), but its discrepancy with the biostratigraphic ages warrants further investigation. Thus, a Late Jurassic age is cautiously assigned to SHGM L275 pending the discovery of an *in situ* specimen - the formation and locality information of SHGM L275 are based on its specimen label only, as it was not possible to verify this with the fossil's discoverer because they are not known. This conclusion has a profound impact on the origins of osteoglossomorphs as it shows that *Paralycoptera* and *Lycoptera* were contemperanoeous. This new evidence therefore rejects Chang & Chou's (1977) hypothesis that *Lycoptera* gave rise to *Paralycoptera*. According to Chang & Chou (1977), *Paralycoptera* is more derived than *Lycoptera*, despite their many similarities, because of features including a dorsoventrally deeper body and a larger anal fin compared to dorsal fin. However, as found in this study, the age of *Paralycoptera* is comparable to the age of *Lycoptera*, so the many morphological similarities between *Paralycoptera* and *Lycoptera* described in Chang & Chou (1977) were likely results of convergence, given their relatively disparate phylogenetic relationships (Xu & Chang 2009).

Given the freshwater habitats of osteoglossomorphs, migration across an oceanic barrier was unlikely, so these fish should have a Pangean origin (Xu & Chang 2009). However, the location of their origins, whether in Africa or Asia, has been debated (Wilson & Murray 2008; Xu & Chang 2009). The Late Jurassic occurrence of *Paralycoptera* in Hong Kong provides additional evidence to support the hypothesis (Xu & Chang 2009) that osteoglossomorphs originated from eastern Asia, as the oldest represenatives of this clade are all known from the Late Jurassic of China, e.g. *Lycoptera* and *Tongxinichthys* (Chang & Jin 1996) instead of Africa, which instead has members with more derived anatomical traits (Xu & Chang 2009).

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Comment [s10]: Suggest that this location is shown on Fig. 1.

Comment [s11]: It might be helpful to place this discussion in its stratigraphic context. Lai Chi Chong Fm is stratigraphically overlain by the Long Harbour Formation (see Sewell, Tang & Campbell 2012, Geochemistry, Geophysics, Geosystems) which has also been precisely dated at 142.8±0.2 Ma (Davis, Sewell & Campbell 1997). This gives additional support to the accuracy and reliability of the radiometric ages.

Figure 9: Jurassic fish localities in SE Asia and the localities of *Paralycoptera* (Modified from Chang & Miao, 2004).

4.4 New phylogenetic characters

In our study a numerical phylogenetic analysis was not performed because existing data made

it possible to unequivocally assign SHGM L275 to *Paralycoptera*. However, in the course of this study it was noted that *Singida* was not easily distinguishable from *Paralycoptera* on the basis of existing characters relating to the posterior skeleton. Anatomical characteristics such as the degree of fusion in the hypurals and the shape of the anal fin that were not included in exisiting phylogenetic character lists would therefore be useful to include in future phylogenies:

- Hypurals in the lower lobe: [0] = independent; [1] = partially fused; [2] = fully fused.
- Anal fin shape: [0] = triangular; [1] = falcate.

4.5 Limitations and future work

The taxonomic identification of SHGM L275 was difficult because the fossil is incomplete, and is the only specimen of its kind from Hong Kong. Thus, further discoveries of *Paralycoptera* in the city (in Lai Chi Chong and on Lantau Island) would help to facilitate further anatomical comparisons with mainland Chinese specimens providing additional insights into anatomical variation in this taxon. To resolve the current discrepancies between the biostratigraphic and radiometric ages of the Lai Chi Chong Formation, and confirm the proposed temporal range extension for *Paralycoptera*, a reappraisal of current biostratigraphic evidence is required. Radiometric dating of fossil-bearing strata within the Formation will be particularly valuable, if suitable rocks can be identified in the future. However, both of these aspects are beyond the scope of this paper to address further. More detailed petrological analysis of the matrix of SHGM L275 using scanning electron microscopy would be valuable for corroborating its facies assignment (the 'parallel laminated fine sandstone and mudstone' facies with the scheme of Lin & Lee (2012) and the volcaniclastic sedimentary facies of Workman (1991)) and facilitating comparisons with the sedimentary facies of other *Paralycoptera* localities in mainland China, such as in Liaoning province (P.J. Chen 1983). These facies investigations, in addition to comparisons between the floras at these different localities, will be important towards elucidating the palaeoenvironment of *Paralycoptera* (and its co-inhabitants) in greater detail, particularly in relation to neighbouring volcanic activity. Future fossil collection and petrological analysis of non-Lai Chi Chong Formation Lantau Volcanic Group sediments, such as those on Lantau Island, will improve our understanding of local variations in the palaeoenvironment of *Paralycoptera*, and will potentially provide evidence of how this taxon (and its co-inhabitants) responded to the well-documented episodes of Middle Jurassic to Early Cretaceous volcanism in Hong Kong (Sewell *et al.* 2000). The latter narrative therefore makes Hong Kong an ideal place to understand the biotic response of Mesozoic fossils to significant environmental stress, so it is hoped that this will lead to further development of palaeontological studies in Hong Kong.

Comment [s12]: A detrital zircon age on the matrix plus one on the 'parallel laminated fine sandstone and mudstone' would also lend supporting evidence to the age assignment.

Comment [s13]: Fossils from upper Lantau Peak are likely to belong to either the Repulse Bay VG (143Ma) or the Kau Sai Chau VG (141Ma). Those on Sunset Peak are likely to be Repulse Bay VG (see Campbell et al. 2007).

5   Conclusions

A fossil fish, SHGM L275, from Lai Chi Chong, Hong Kong was rediscovered in the fossil collections of Stephen Hui Geological Museum at the University of Hong Kong. This specimen is identified as *Paralycoptera* based on the following four anatomical characteristics:

(1) a completely developed neural spine on the first preural centrum;
(2) two or three uroneurals;
(3) four upper hypurals and two lower hypurals, and
(4) all hypurals are independent.

The discovery of *Paralycoptera* in Late Jurassic-aged strata in Hong Kong - the city's only Mesozoic vertebrate – appears to extend the temporal range of the genus back by ~40 million years. However, discrepancies between the biostratigraphic and radiometric ages of the strata, which belongs to the Lai Chi Chong Formation, warrants a cautious treatment of the proposed temporal range extension, pending further geochronological investigation. However, our discovery unequivocally extends *Paralycoptera*'s geographical range approximately 700 km southwards, potentially affecting the *Mesoclupea* Fish Fauna. In the context of the geological literature on the Lai Chi Chong Formation and our knowledge of the fossil's matrix, it is suspected that *Paralycoptera* lived in freshwater lakes in close proximity to volcanic environments that experienced episodic earthquakes and volcanic eruptions that greatly affected the lake's regime. This palaeoenvironment appears to match those of other *Paralycoptera* localities in mainland China inspiring the conclusion that this taxon was potentially tolerant of high environmental stresses and may even have thrived on higher nutrient levels and changeable water temperatures in the lake, during times of volcanic activity.

This study makes an important contribution to our understanding of Hong Kong's fossil heritage, given that the city has a relatively poor fossil record and limited sedimentary rock exposures (Lee *et al.* 1997). This study is the first on Hong Kong fossils in over 15 years (Lee *et al.* 1997) so it is hoped that it can help to promote further interest in Hong Kong's palaeontology, particularly given the rare opportunity to study the biotic response to long-lived and accurately-dated Mesozoic volcanic events.

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Author contributions:

TKT and MP collected and analysed the data and also wrote the paper, MMC had some input in the description of the specimen. MP designed the project.

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Comment [s14]: Could not locate reference in the text or supplementary information.

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