

Triassic pentadactyl tracks from the Los Menucos Group (Río Negro province, Patagonia Argentina): possible constraints on the autopodial posture of Gondwanan trackmakers

Paolo Citton ^{Corresp., 1,2}, Ignacio Díaz-Martínez ^{1,2}, Silvina de Valais ^{1,2}, Carlos Cónsole-Gonella ^{1,3}

¹ Consejo Nacional de Investigaciones Científicas y Técnicas (CONICET), Buenos Aires, Argentina

² Instituto de Investigación en Paleobiología y Geología (IIPG), Universidad Nacional de Río Negro, General Roca, Argentina

³ Instituto Superior de Correlación Geológica (INSUGEO), Universidad Nacional de Tucumán, Tucumán, Argentina

Corresponding Author: Paolo Citton
Email address: pcitton@unrn.edu.ar

The Los Menucos locality in Patagonia, Argentina, bears a well-known ichnofauna mostly documented by small therapsid footprints. Within this ichnofauna, large pentadactyl footprints are also represented but to date were relatively underinvestigated. These footprints are here analyzed and discussed based on palaeobiological indications (i.e. trackmaker identification). High resolution digital photogrammetry method was performed to achieve a more objective representation of footprint three-dimensional morphologies. The footprints under study are compared with *Pentasauropus* from the Upper Triassic lower Elliot Formation (Stormberg Group) of the Karoo Basin (Lesotho, southern Africa). Some track features suggest a therapsid-grade synapsid as the potential trackmaker, to be sought among anomodont dicynodonts (probably Kannemeyeriiformes). While the interpretation of limb posture in the producer of *Pentasauropus* tracks from the Los Menucos locality agrees with those described from the dicynodont body fossil record, the autopodial posture does not completely. The relative distance between the impression of the digital (ungual) bases and the distal edge of the pad trace characterizing the studied tracks likely indicates a subunguligrade foot posture in static stance, but plantiportal during the dynamics of locomotion. The reconstructed posture might have implied an arched configuration of the articulated metapodials and at least of the proximal phalanges, as well as little movement capabilities of the metapodials. Usually, a subunguligrade-plantiportal autopod has been described for huge animals to obtain an efficient management of body weight. Nevertheless, this kind of autopod is described here for large but not gigantic animals, as the putative trackmakers of *Pentasauropus* were. This attribution implies that such an autopodial structure was promoted independently from the body size in the putative trackmakers. By an evolutionary point of view, subunguligrade-plantiportal autopods not necessarily must be related with by an increase in body size, but

rather the increase in body size requires a subunguligrade or unguligrade, plantiportal foot. Chronostratigraphically, *Pentasauropus* was reported from Upper Triassic deposits of South Africa and United States, and from late Middle Triassic and Upper Triassic deposits of Argentina. A Late Triassic age is here proposed for the *Pentasauropus*-bearing levels of the Los Menucos Group.

Triassic pentadactyl tracks from the Los Menucos Group (Río Negro province, Patagonia, Argentina): possible constraints on the autopodial posture of Gondwanan trackmakers

Paolo Citton^{1,2}, Ignacio Díaz-Martínez^{1,2}, Silvina de Valais^{1,2}, Carlos Cónsole-Gonella^{1,3}

¹CONICET - Consejo Nacional de Investigaciones Científicas y Técnicas

²Instituto de Investigación en Paleobiología y Geología (IIPG), General Roca, Río Negro province, Argentina

³Instituto Superior de Correlación Geológica (INSUGEO), Miguel Lillo 205, Tucumán, Argentina

Corresponding Author:

Paolo Citton

Av. Roca 1242, General Roca, Río Negro province, 8332, Argentina

Email address: pcitton@unrn.edu.ar

Abstract

The Los Menucos locality in Patagonia, Argentina, bears a well-known ichnofauna mostly documented by small therapsid footprints. Within this ichnofauna, large pentadactyl footprints are also represented but to date were relatively underinvestigated. These footprints are here analyzed and discussed based on palaeobiological indications (i.e. trackmaker identification). High resolution digital photogrammetry method was performed to achieve a more objective representation of footprint three-dimensional morphologies. The footprints under study are compared with *Pentasauropus* from the Upper Triassic lower Elliot Formation (Stormberg Group) of the Karoo Basin (Lesotho, southern Africa). Some track features suggest a therapsid-grade synapsid as the potential trackmaker, to be sought among anomodont dicynodonts (probably Kannemeyeriiformes). While the interpretation of limb posture in the producer of *Pentasauropus* tracks from the Los Menucos locality agrees with those described from the dicynodont body fossil record, the autopodial posture does not completely. The relative distance between the impression of the digital (ungual) bases and the distal edge of the pad trace characterizing the studied tracks likely indicates a subunguligrade foot posture in static stance, but plantiportal during the dynamics of locomotion. The reconstructed posture might have implied an arched configuration of the articulated metapodials and at least of the proximal phalanges, as well as little movement capabilities of the metapodials. Usually, a subunguligrade-plantiportal autopod has been described for huge animals to obtain an efficient management of body weight. Nevertheless, this kind of autopod is described here for large but not gigantic animals, as the putative trackmakers of *Pentasauropus* were. This attribution implies that such an autopodial structure was promoted independently from the body size in the putative trackmakers.

By an evolutionary point of view, subunguligrade-plantiportal autopods not necessarily must be related with ~~by~~ an increase in body size, but rather the increase in body size requires a subunguligrade or unguligrade, plantiportal foot. Chronostratigraphically, *Pentasauropus* was reported from Upper Triassic deposits of South Africa and United States, and from late Middle Triassic and Upper Triassic deposits of Argentina. A Late Triassic age is here proposed for the *Pentasauropus*-bearing levels of the Los Menucos Group.

INTRODUCTION

Tetrapod tracks are valuable fossils informing about anatomy (e.g. Carpenter, 1992), functional adaptations (e.g. Baird, 1980), motion (e.g. Avanzini, Piñuela & García-Ramos, 2011) and ethology (e.g. Lockley et al., 2016) of extinct animals, greatly expanding the potential of information that is often precluded from the body-fossil record. The detailed analysis of tetrapod footprints is therefore significant for integrating and revising data derived from the tetrapod body-fossil record.

The scientific study of tetrapod footprints in Argentina is relatively recent compared to that of Europe (Duncan, 1831; Kaup, 1835a, b) and North America (Hitchcock, 1836), dating back to the first half of the twentieth century (von Huene, 1931). One of the most important contribution to tetrapod ichnology in Argentina is that of Casamiquela (1964), who devoted himself to the study of Triassic and Jurassic tetrapod tracks from Patagonia. Later, other contributions focused on important Triassic ichnofaunas from other regions of Argentina have been published (Romer, 1966, Bonaparte, 1966; Leonardi, 1994; Melchor & de Valais, 2006). Among the Triassic vertebrate ichnofaunas, the Los Menucos ~~one~~, which is dominated by small therapsid footprints, was ~~largely~~ studied (Casamiquela, 1964, 1975, 1987; Leonardi & de Oliveira, 1990; Leonardi, 1994; Domnanovich & Marsicano, 2006; Melchor & de Valais, 2006; de Valais, 2008; Domnanovich et al., 2008; Díaz-Martínez & de Valais, 2014). The bulk of this ichnofauna was originally attributed to different ichnotaxa by Casamiquela (1964, 1975), but after the revision made by Melchor & de Valais (2006), most of the ichnogenera erected by

Casamiquela have been considered as synonyms of *Dicynodontipus*. Moreover, an indetermined chirotheroid track (de Valais, 2008), a single track referred to as *Rhynchosauroides*, and large pentadactyl footprints mentioned as *Pentasauropus* sp. (Domnanovich et al., 2008) have been reported from the Los Menucos area. From the same locality, several slabs with pentadactyl tracks comparable to those described by Domnanovich et al. (2008) were collected many years ago but remained unpublished until now.

An ichnological analysis based on this material is here proposed and discussed in terms of the palaeobiology, identity and autopodial anatomy of the trackmaker. Besides, a brief discussion of the chronostratigraphy of this record is provided.

INSTITUTIONAL ABBREVIATIONS

LES - Laboratoire de Paléontologie, Institut de Sciences de l'Evolution of the University of Montpellier II collection, Montpellier, France; MPCA - Museo Provincial Carlos Ameghino, Cipolletti, Río Negro province, Argentina; MMLM - Museo Municipal de Los Menucos, Los Menucos, Río Negro province, Argentina; MMLM (ex MRPV) - Museo Provincial María Inés Kopp, Valcheta, Río Negro province, Argentina.

MATERIAL AND METHODS

The present study is based on the direct examination of track-bearing slabs MPCA 27029-1 with three pes-manus couples, two of which incomplete (concave epireliefs), MPCA 27029-2 with a single left pes-manus couple (convex hyporeliefs), MPCA 27029-3 with two pes-manus couples and an incomplete pes (convex hyporeliefs), MPCA 27029-4 with a single pes-manus couple (convex hyporeliefs), MPCA 27029-5 with a single track (convex hyporeliefs), MPCA 27029-9

with five pes-manus couples, three of which incomplete (convex hyporeliefs), MPCA 27029-16 with three pes-manus couples, one of which incomplete, and five incomplete tracks (convex hyporeliefs), MPCA 27029-21 with two pes-manus couples and four tracks (convex hyporeliefs), MPCA 27029-33 with two pes-manus couples (convex hyporeliefs), MMLM 1 with two pes-manus couples (convex hyporeliefs), MMLM 2 with two incomplete pes-manus couples (convex hyporeliefs), and MMLM 075-1 (ex MRPV 1987P.V.06 *in* Domnanovich et al., 2008, hereafter MMLM 075-1) with two incomplete pes-manus couples (concave epireliefs). Except for the specimen MMLM 075-1, the material under study was to date unpublished. A few other slabs, both with and without label, are stored at the MPCA but were not considered in this study due to poor preservation of the tracks. In total, about 60 footprints were analyzed. For each slab, tracks were numbered using Arabic numerals and, when referring to single tracks in the text, they are indicated as /number following the slab label (e.g. MPCA 27029-1/4 where MPCA 27029-1 and number 4 indicate slab and single track, respectively). The studied material mainly consists of isolated sets or incomplete trackways.

The provenance of the track-bearing slabs can be traced back to the Felipe Curuil ex quarry, Yancaqueo farm, east of the town of Los Menucos (Domnanovich et al., 2008), but the exact stratigraphic repositioning of the material is currently prevented and inherent data are lacking in the literature.

In order to characterize the microfacies of the trampled layers, two thin sections were obtained from the slab MPCA 27029-19 (its footprints are poorly preserved and not included in this study), both parallel and perpendicular to the trampled surface. For the description of the thin sections, Mackenzie, Donaldson & Guilford (1982), and Scasso & Limarino (1997) were taken as a reference. Thin sections are presently stored at the MPCA and labelled as MPCA

27029/19.1 (parallel to the trampled surface) and MPCA 27029/19.2 (perpendicular to the trampled surface).

Measurements related to trackmaker body dimension were obtained from slabs MPCA 27029-1, MPCA 27029-9, MPCA 27029-16 and MPCA 27029-21. From single tracks, which are mainly represented by digit traces, measurements of footprint width were taken. Also, track features and differential depth of impressions in some cases allowed to recognize the footprint identity, the side of the trackway when incompletely preserved, or tracks belonging to different trackways (e.g. MMLM 075-1), and element orientations. Track measurements were performed according to guidelines introduced by Leonardi (1987).

Track outlines were represented through interpretive drawings. High resolution digital photogrammetry was undertaken to achieve a more objective representation of track three-dimensional morphology, according to a recently described standard protocol for ichnological studies (Falkingham et al., 2018). To model the studied specimens, the software package Agisoft PhotoScan Pro (Educational License), which enables creating 3D textured meshes by means of semi-automatic processing of images (Mallison & Wings, 2014), was used.

The images selected for the photogrammetric process were acquired using a Nikon Coolpix P520 camera with 4.3-7.6 focal length, resolution 4896x3672 and pixel size ranging from 1.25x1.25 μm and 1.27x1.27 μm . Main processing parameters are reported in Table 1. In order to correctly scale the calculated model, a metric reference marker was applied on the surface. Three-dimensional models were converted to colour topographic profiles using the software Paraview (version 5.4.1).

GEOLOGICAL SETTING

Continental deposits of Triassic age in Argentina accumulated in different basins in western and northwestern regions (Mendoza, San Juan, San Luis and La Rioja provinces) as well as in Patagonia (northern sector of the Santa Cruz province and Río Negro provinces). These elongated, narrow rift basins with prevalent NW-SE and NNW-SSE trends were developed during Permian and Triassic periods and **witness** the breakup of the western margin of south-west Gondwana (Kokogian et al., 1999; Franzese & Spalletti, 2001; Barredo et al., 2012).

The Triassic tetrapod track record of southern South America is exclusive to three basins, namely the Ischigualasto-Villa Unión Basin (San Juan and La Rioja provinces), the Cuyo Basin (Mendoza and San Juan provinces) and the Los Menucos basin (Río Negro province) (e.g. Melchor, Genise & Poiré, 2001; Melchor & de Valais, 2006; de Valais, 2008 and references therein). According to Spalletti (1999), in the northern basins (i.e. Ischigualasto-Villa Unión and Cuyo) the sedimentation encompasses the Lower to Upper Triassic, while in the Los Menucos **Basin the sedimentation took place in the Late Triassic.**

After the first works of Stipanovic (1967), Stipanovic et al. (1968) and Stipanovic & Methol (1972, 1980), the Los Menucos Group (also as ‘Complejo Los Menucos’ - Los Menucos Complex *sensu* Cucchi, Busteros & Lema, 2001) was established by Labudía & Bjerg (2001) to indicate **dacitic to rhyolitic ignimbrites, mesosilicic lavas** and subordinate Triassic sedimentary rocks exposed around Los Menucos town, in the north-western sector of the North Patagonian Massif (Río Negro province, Argentina; Fig. 1A).

Within the Los Menucos Group, two lithostratigraphic units were defined, namely the Vera Formation at the base and the Sierra Colorada Formation on top (Labudía & Bjerg, 2001, 2005, and references therein; Fig. 1B). **The Vera Formation** is mainly composed of volcanic and continental deposits laid down inside small basins bordered by regional and local faults with

attitudes NE-SW, E-W and NW-SE (Labudía & Bjerg, 2001, 2005). The Vera Formation is mainly represented by brownish to yellowish conglomerates, white to greenish sandstones and reddish brown to red pelites, with which volcanic ashes, tuffs and tuffites, dacitic pyroclastic flow products and volcanic breccias are intercalated (Labudía & Bjerg, 2001, 2005). Sedimentation took place mainly in alluvial plain, floodplain, ephemeral river and small lacustrine palaeoenvironments (Labudía & Bjerg, 2005), under seasonal climate condition with alternating periods of dry and wet conditions (Gallego, 2010). Sedimentary and volcanoclastic levels within the Vera Formation are characterized by a very rich palaeoflora, the so-called “*Dicroidium* type flora” (Stipanovic, 1967; Stipanovic & Methol, 1972; Artabe, 1985a, b; Labudía et al., 1995; Labudía & Bjerg, 2001, 2005) and by an abundant tetrapod ichnofauna, preserved on sandstones with poorly sorted grains and with a variable content of tuffaceous breccias (Melchor & de Valais, 2006). Finds of skeletal fauna are scarce and are so only represented by remains of an amiiiform fish (Bogan, Taverne & Agnolin, 2013).

The Sierra Colorada Formation is essentially made of ignimbritic volcanic rocks (Labudía & Bjerg, 2001, 2005), dated at 222 ± 2 Ma with the Rb/Sr isochron method (Norian, Late Triassic; Rapela et al., 1996) and at 206.9 ± 1.2 Ma with the Ar/Ar method (Rhaetian, Late Triassic; Lema et al., 2008). Unfortunately, these datations do not radiometrically constrain the Vera Formation, for which a Late Triassic age was historically proposed on the basis of the “*Dicroidium* flora” and the tetrapod ichnofauna.

More recent results indicated an age of 257 ± 2 Ma (Wuchiapingian, Late Permian) for a rhyolitic ignimbrite, 252 ± 2 Ma (Changhsingian, Late Permian) for an andesite, and 248 ± 2 Ma (Olenekian, Early Triassic) for a dacitic ignimbrite (Luppo et al., 2017) of the Los Menucos Group. These new data predate the main volcanic activity to an about a 10 Ma period between

the Late Permian and the Early Triassic, ~~making the Los Menuecos Group~~ coeval with the La Esperanza Plutono-Volcanic Complex (Gonzalez et al., 2017; Luppo et al., 2017).

Sedimentological observations

Track-bearing slabs consist of yellowish to greenish, medium to mainly coarse grained and poorly sorted volcanoclastic sandstone lacking of sedimentary structures in hand samples, neither on the surface or cross-section.

The observed texture ranges from inequigranular/equigranular (Figs. 2A, 2B) to predominantly equigranular (Fig. 2C). Phenocrysts, mainly subhedral and anhedral, range in dimension from 0,5 mm to 1,5 mm and show in one case incipient orientation. Phenocrysts are represented mainly by plagioclase, quartz, alkaline feldspar, biotite, amphibole (hornblende), orthopyroxene (enstatite) and calcite floating in a mafic, glassy matrix.

The dominant epiclastic texture observed at the base of the trampled surface (thin section MPCA 27029/19.2), mainly represented by fragments of quartz and some lithics displaying attrition and rounded to sub-angular shape, suggesting sedimentary reworking of an original tuff of probable dacitic composition. The texture observed in the thin section MPCA 27029/19.1 instead indicates ~~a scarce~~ sedimentary reworking (Fig. 2D). In section, a faint normal gradation can be observed most likely indicating short sedimentation events; on the whole, the track-bearing slabs can be related to a proximal fluvial environment

TRACK RECORD

Track preservation

Specimen MMLM 075-1 is composed of four slabs, two as casts -with negative **epichnial** tracks, labeled as MMLM 075-1/1a, /2 and /3a- and two as their moulds -with positive hypichnial

234 tracks, labeled as MMLM 075-1/1b and 3/b. There are no evidences of any layer between the
 235 casts and the moulds and the shape of both concave epireliefs and convex hyporeliefs are exactly
 236 complementary (Fig. 3). Therefore, and taking into account that the tracks preserved similarly
 237 (i.e. sub-circular/sub-ovoidal to pointed digit impressions; roughly sub-circular to elliptical pad
 238 tracks; very thin displacement rims in the pad and well-marked in the digit impressions; Figs. 4-
 239 9), ~~in our opinion~~ the concave epireliefs are true tracks (*sensu* Marty, Falkingham & Richter,
 240 2016) and the convex hyporeliefs are their natural casts (*sensu* Marty, Falkingham & Richter,
 241 2016).

242 In general, the tracks studied here are moderately well preserved (grade 1 *sensu*
 243 Belvedere & Farlow, 2016), and the true tracks are not elite tracks (*sensu* Lockley, 1991). In
 244 addition, they are not modified true tracks (*sensu* Marty, Falkingham & Richter, 2016) because
 245 they lack evidence of physiochemical (e.g. weathering) and/or biological influences after they
 246 were made. Thereby, the shape of these tracks is mainly conditioned by the substrate consistency
 247 (grain size and water content). Recently, Falk et al. (2017) performed neoichnological
 248 experiments that compared the shape of tracks impressed in three different sediments (fine,
 249 medium and coarse sand) with different moisture contents (wet, moist and dry). They concluded
 250 that wet and dry coarse sediments preserve tracks without fine details, but moisture coarse
 251 sediment might preserve the overall track shape and details as claw impressions. As has been
 252 previously commented, the tracking surface is a medium to coarse sandstone, and tracks have
 253 depth digit impressions with extruded rims.

254 Therefore, and according to ~~Falk et al. (2017) experiments~~, the trackmakers most likely
 255 walked on humid, not waterlogged nor dry, coarse sediments with a moderately plastic
 256 behaviour, able to record the main anatomical features of the autopods.

Track description

The material are manus and pes tracks with very low dimensional **heteropody**, mainly preserved as tetradactyl impressions, although pentadactyl tracks are also present (MPCA 27029-1/4/6, MPCA 27029-2/2, MPCA 27029-4/2, MPCA 27029-16/10, MPCA 27029-33/2, MMLM 075-1) (Figs. 4, 5E-5H, 7A-7D, 8), as well as tridactyl ones displaying only the central digits (MPCA 27029-9/2, MPCA 27029-16/8, MPCA 27029-21/3, MMLM 1/1, MMLM 2/3) (Figs. 6G-6H, 7, 9). Morphologically, manus and pes tracks are strongly symmetrical. Digit traces are commonly arranged to shape an arcuate pattern that is convex anteriorly, according to which the digit III trace (the central one) or digit III and IV traces are the most projecting. Variability affecting the number of digits can occur on the same slab (e.g. MPCA 27029-21, MMLM 2; Figs. 7E-7H, 9E-9H). In the material under study the degree of curvature of the arcuate pattern is variable and appears more pronounced in some smaller tracks (e.g. MPCA 27029-16/7/9/10; Figs. 7A-7D) than in larger ones (e.g. MPCA 27029-1/4, MMLM 075-1, MMLM 2/2; Figs. 4, 8E-8H, 9E-9H). In the smaller tracks (e.g. MPCA27029-16), the morphology of digit traces, their relative spacing and orientation, as well as the position of pes and manus impression is comparable with that of the larger tracks. When present, also the sole pad trace resembles that observed in the footprints of larger dimension. Thus, apart from the degree of curvature, the general morphology remains consistent despite dimensional differences (see Figs. 4A-4D and 7A-7D).

Digit traces can be characterized by a sub-circular/sub-ovoidal morphology (e.g. MPCA 27029-9, MPCA 27029-16, MPCA 27029-21; Figs. 6E-6H, 7), while in other cases they are markedly pointed (e.g. MPCA 27029-2, MPCA 27029-4, MPCA 27029-5, MPCA 27029-33; Figs. 4E-4H, 5E-5H, 6A-6D, 8A-8D). These two morphologies can co-exist on the same slab and within the same set or trackway, thus pertaining to the spectrum of internal variability of the

material under study. When pointed, the most medial digit traces (i.e. digit I or II imprints and, to a lesser extent, digit III and IV imprints), both of manus and pes tracks, can be affected by drag marks. These extramorphological features (see Peabody, 1948) qualitatively range from weakly hinted and short (e.g. MPCA 27029-1/6, MPCA 27029-2/1, MPCA 27029-3/1/2, MPCA 27029-4/2, MPCA 27029-5, MPCA 27029-9/4/6; Figs. 4, 5, 6) to highly sharp and long (e.g. MMLM 075-1, MMLM 1 and MMLM 2; Figs. 8E-8H, 9A-9D).

Central digits are commonly the most deeply and uniformly impressed, both in manus and pes tracks (e.g. MPCA 27029-4, MPCA 27029-5, MMLM 1; Figs. 5F-5G, 6B-6C, 9B-9C). When a certain degree of variability is observed, digit III and IV imprints are the most deeply imprinted (e.g. MPCA 27029-1, MPCA 27029-3, MPCA 27029-16; Figs. 4B-4C, 5B-5C, 7B-7C), followed by digit II and I imprints. The digit V trace, when preserved, is shorter and closer to the pad trace than the other digit traces and is only faintly imprinted (e.g. MMLM 075-1, but see **MPCA 27029-16/9/10 for a different configuration of depth of impression**; Figs. 7B-7C, 8F-8G).

Behind the digit traces, a roughly sub-circular to elliptical sole pad trace can be preserved (e.g. MPCA 27029-1/2/4/5/6, MPCA 27029-5, MPCA 27029-16/10, MMLM 075-1, MMLM 1/2/4, MMLM 2/1; Figs. 4A-4D, 6A-6D, 7A-7D, 8E-8H, 9). The sole pad trace lies at a short distance from the base of the central digit traces and commonly approximates the most medial and lateral digit imprints (e.g. MPCA 27029-1; Figs. 4A-4D). Commonly, the sole pad trace is separated from central digit traces ahead by a non impressed area, which appears as a groove or as a ridge depending on the mode of preservation, tapering towards the most medial and lateral digit imprints. This should not be confused with displacement areas of similar morphology, which are instead related to digit traces (i.e. thrust of digit pushing the sediment backwardly; Fig.

10), where this area is not impressed (e.g. MPCA 27029-2, MPCA 27029-3/1/4, MPCA 27029-4, MPCA 27029-9/3/5, MPCA 27029-21/4, MPCA 27029-33/1/2/3, MMLM 2/3; Figs. 4E-4H, 5, 6E-6H, 7E-7H, 8A-8D, 9E-9H).

The sole pad trace is more deeply impressed in its central portion; depth of impression slightly decreases toward the lateral and distal portion (i.e. close to the ~~non-impressed~~ area behind digit traces, MPCA 27029-1, MMLM 075-1, MMLM 1; Figs. 4B-4C, 8F-8G, 9B-9C).

When possible, we tried to define the orientation of the footprint axis with respect to the trackway midline. The axis of pes tracks is in some cases rotated inwardly with respect to the trackway midline but it can also be parallel to the trackway midline (e.g. MPCA 27029-9, 27029-16), while manus tracks show a wider range of variability, being both inwardly and outwardly rotated with respect to the hypothetical trackway midline (e.g. MPCA 27029-1 and MMLM2, respectively). When possible, measurements and ratios were taken; measurements were

performed taking into account digit III as the homologous point, both for manus and pes tracks.

Results are reported in Table 2 and Table 3.

Remarks. The footprints from the Los Menucos ichnosite are characterized by having the following features: homopodic manus and pes tracks with low dimensional heteropody, up to five digit imprints aligned, forming an anteriorly convex arch, a sole pad trace more impressed centrally or centro-laterally. On the basis of these general features, the specimens from Los Menucos are tentatively referred to as *Pentasauropus*.

The ichnogenus *Pentasauropus* Ellenberger, 1970 was established on the basis of material collected and described some years before (Ellenberger, 1955) from the Upper Triassic lower Elliot Formation (Stormberg Group) of the Karoo Basin of Lesotho (Southern Africa). Five ichnospecies were originally included in the ichnogenus, namely *Pentasauropus erectus*,

Pentasauropus incredibilis, *Pentasauropus maphutsengi*, *Pentasauropus morobongensis* and *Pentasauropus motlejoi*, which remained unchanged in the subsequent formal listing (Ellenberger, 1970, 1972). Material from the Ellenberger collection referred to this ichnogenus, is housed at the University of Montpellier (France) and represented by six casts originally mentioned as *Pentasauropus incredibilis* (LES 054 1-3, LES 054 4), *Pentasauropus morobongensis* (LES 005) *Tetrasauropus gigas* (LES 038), plus some missing specimens (see D’Orazi Porchetti & Nicosia, 2007, and reference therein for a complete assessment of inventory numbers).

After the original and subsequent publications of Ellenberger (1955, 1970, 1972), the ichnogenus was considered as valid by Olsen & Galton (1984), Lockley & Meyer (2000), D’Orazi Porchetti & Nicosia (2007), Bordy, Abrahams & Sciscio (2017), and Hunt, Lucas & Klein (2018). D’Orazi Porchetti & Nicosia (2007) emended the diagnosis of the ichnogenus to appoint the type ichnospecies and considered the five ichnospecies as synonyms of *Pentasauropus incredibilis*. Differences in track pattern were considered as originated by dimensional constraints and/or behavioural factors, and the main footprint characters (e.g. number and arrangement of digits, heteropody) do not justify an ichnospecies separation (D’Orazi Porchetti & Nicosia, 2007). Moreover, agreeing with Lockley & Meyer (2000), the same authors assigned tracks originally referred to as *Tetrasauropus gigas* to *Pentasauropus*.

In agreement with the emended ichnogeneric diagnosis by D’Orazi Porchetti & Nicosia (2007), the arcuate pattern of manus and pes tracks derived from the five equally spaced claw or ungual traces (those of imprints of digit II, III and IV are the largest). In other cases a roughly rounded sole pad is observed behind claw or ungual traces (LES 053 A, B, C in Ellenberger, 1972, pl. IV and V and LES 038). According to D’Orazi Porchetti & Nicosia (2007), the axis of

the pes impressions is always inwardly rotated, while that of the manus impression can range from slightly inwardly rotated (LES 052 B and LES 053 A) to slightly outwardly rotated (LES 038 and LES 052 A). Although long and complete trackways are not represented, this character seems to characterize also the studied material based on the reconstruction of an hypothetical midline (e.g. Fig. 4A-D, footprints 4 and 6; Fig. 5A-D, footprint 2). In some cases, short stride length in relation to overall footprint dimension could indicate a primary overstepping. However, in our opinion this trackway character cannot be ensured on the basis of the material under study and most probably, taking into account the complexity of the dynamic of locomotion, cannot be inferred only from the stride lengths and footprint dimensions. Thus, for the time being, we prefer not to stress the interpretation about the overstepping.

Outside of South Africa, tracks tentatively referred to the ichnogenus were reported from Upper Triassic Chinle Group of Utah (Lockley & Hunt, 1995; Hunt-Foster et al., 2016) and Colorado (Gaston et al., 2003, fig. 12B), both USA. Moreover, tracks possibly referable to *Pentasauropus* were found in the Gettysburg Shale of the Gettysburg Basin of the Newark Supergroup (Baird pers. comm in Olsen & Galton, 1984). In Argentina, apart from the report from the Triassic Vera Formation (Río Negro province, Domnanovich et al., 2008), tracks referred to as *Pentasauropus* were described from the Carnian Portezuelo Formation ('Type Q2' sensu Marsicano & Barredo, 2004). In addition, tracks with similar morphology to *Pentasauropus* were also reported from the Middle Triassic Cerro de Las Cabras Formation (Mendoza province, as cf. *Pentasauropus* in de Valais, Melchor & Bellosi, 2006) and from the Portezuelo Formation (San Juan province) as 'huellas cuadrúpedas tipo C' (de Valais, 2008) but, for the time being, this material remains in open nomenclature.

Zoological attribution

Several attempts to identify the trackmaker of *Pentasauropus* have been made. The ichnogenus was originally attributed to amphibians, basal melanorosaurid, ornithischian, anapsid and basal sauropod (Ellenberger & Ellenberger, 1958: p. 67; Ellenberger, 1970, 1972). Moreover, Haubold (1974, 1984) referred *Pentasauropus* to a sauropod or therapsid trackmaker. A dicynodont was also proposed as producer by Olsen & Galton (1984), Anderson, Anderson & Cruickshank (1998) and Lockley & Meyer (2000). Galton & Heerden (1998) attributed *Pentasauropus* to large anomodont dicynodonts. D'Orazi Porchetti & Nicosia (2007) accepted the attribution to a dicynodont, observing a good match between the skeletal autopodia of Triassic dicynodonts and the structure of digital impressions of the manus and pes, apart from the strong homopody and the limb posture (see also Walter, 1986). Recently, the kannemeyeriiform dicynodont *Pentasaurus goggai* from the lower Elliot Formation of South Africa has been referred as probable trackmaker of *Pentasauropus* tracks from the same lithostratigraphic unit (Kammerer, 2018).

The studied material from the Los Menucos locality presents some features that allow corroboration of the therapsid interpretations about trackmaker identity. At the same time, attempts to identify the putative trackmaker opens the way for new inferences about the posture of the autopodia.

Limb posture kept by *Pentasauropus* trackmakers during the step cycle can be tentatively inferred from the trackway pattern (Peabody, 1948, 1959; Kubo & Benton, 2009; Kubo & Ozaki, 2009), even if this interpretation is often far from being simple and linear (Crompton & Jenkins, 1973). For example, it must be noted that non therapsid-synapsids with sprawling posture can left trackways in which the left and right tracks lie near to the axial midline (i.e. narrow internal trackway width mirroring a semi-erect to erect posture of the trackmaker) by adopting side to

side flexion of the trunk (Hopson, 2015). The same was described for a therapsid trackway by Smith (1993). Some degree of lateral undulation of the vertebral column causing swing of the hips has been also described on the basis of skeletal remains (Fröbisch, 2006: p. 1305). However, in sprawling trackmakers adopting trunk flexion and producing narrow trackways, tracks are mainly inwardly oriented with respect to the trackway midline (see Hopson, 2015, fig. 8.1). In the studied material, the orientation of footprint axis (passing through digit III) is parallel to the travel direction. This feature, combined with the extremely narrow internal trackway width measured from pes tracks (see Table 2 and Figs. 4A-4D, 5A-5D, 6E-6H, 7A-7D), allow to exclude a sprawling posture and most likely indicate a semi-erect posture for the trackmaker hind limbs. A lateral trunk undulation during the step cycle could have been also adopted by the *Pentasauropus* trackmaker and it would account for the variable pace angulation measured in manus tracks, possibly coupled with low trackmaker velocity.

A more upright posture with respect to that of non-therapsid synapsids is also indirectly sustained by the symmetry of manus and pes tracks. This character mirrors a symmetry of the trackmaker's autopods, a character combined with the acquisition of upright posture and limbs parallel to the sagittal plane of the trackmaker during locomotion (Romer, 1956; Hopson, 1995).

Number of digit imprints, symmetry of manus and pes track, and the morphology of unguals enable us to corroborate previous interpretations and suggest a dicynodont as the most probable trackmaker of *Pentasauropus*. Also the limb posture as supposed from tracks sufficiently matches those discussed for Triassic dicynodonts by Fröbisch (2006).

Pentasauropus producers, based on the type specimens, were characterized by a very low dimensional heteropody and by morphological homopody. Thus, mainly based on these characters, a quite confident match can be found, among dicynodonts, with Kannemeyeriiformes

[see, for example, the descriptions and reconstructions of the autopods of *Dinodontosaurus* by Morato (2006:fig. 30), *Tetragonias njalilus* by Cruickshank (1967:fig. 17) and by Fröbisch (2006:fig. 9)]. Kannemeyeriiform manual and pedal skeletal elements are in fact generally homomorphic and conservative across the clade, as recently stated by Kammerer (2018). Variation in manus morphology among kannemeyeriiforms is only limited to minor differences in ungual shape (Lucas, 2002). Digit traces are here considered roughly compatible with broad ungual phalanges characterized by rounded tips but the exact shape cannot be determined. The morphological variability of ungual traces most likely ~~depends from~~ substrate conditions at the time of impression and, as discussed below, from the dynamics of the locomotion of the producers.

Among the other possible producers already mentioned in the literature, are excluded: i) an amphibian trackmaker, for the pentadactyl manus and general morphology; ii) an anapsid trackmaker, for the trackway configuration and footprint axis orientation; iii) a sauropodomorph and sauropod trackmaker, for the trackway configuration, footprint axis orientation and morphological homopody.

The posture of the autopodia of the *Pentasauropus* trackmaker that can be inferred from the ichnological material differs from that inferable from the description by Cruickshank (1967) description. In those *Pentasauropus* tracks that show an impression of the sole or palm, a negligible distance between the distal margin of the sole/palm pad trace and the proximal margin of the central digit traces, further reducing towards digits I and V, was observed (Table 3). This feature most likely indicates that not all of the foot bones contacted the ground during locomotion, at the same time constraining the orientation of metapodial and basipodial elements, and most likely also that of the more proximal phalanges, in the articulated autopod. Thus, the

reconstruction proposed here contemplates an inclined position of pedal and manual elements in the autopods of *Pentasauropus* producers. The sub-circular to elliptical pad trace behind the digit traces is consequently considered compatible with an extended fleshy pad below the basipodials and likely metapodials of the autopods of the producer.

DISCUSSION

The autopod posture of *Pentasauropus* trackmaker

Compared to the ichnogenus *Pentasauropus*, tracks from Los Menucos Group have allowed us to verify previous ichnological interpretations based on the reference material from Lesotho and have enabled us to corroborate the identification of a putative trackmaker and its limb posture. Moreover, the studied material sheds light on the trackmaker autopod posture. The smaller tracks (e.g. MPCA27029-16) are here interpreted to have been left by a juvenile trackmaker and allowed to appreciate that track morphology and structure are uniform in different ontogenetic stages of the same type of producer.

Number of digit imprints, symmetry of manus and pes track, morphology of ungual traces, limb posture and morphologically homopodic manus and pes tracks indicate the producer of *Pentasauropus* to be sought among dicynodonts of the clade Kannemeyeriiformes (Fröbisch, 2009). Within this panorama and accepting the proposed palaeozoological attribution, *Pentasauropus* tracks represent a valuable datum, further enriching the dicynodont record of the Triassic of Argentina (Cox, 1962; Bonaparte, 1969, 1971, 1981; Lucas, 1998, 2010, 2018; Rogers et al., 2001; Zavattieri & Arcucci, 2007; Fröbisch, 2009; Domnanovich & Marsicano, 2012; Abdala et al., 2013; Mancuso et al., 2014).

The studied tracks enabled us to improve the knowledge of the therapsid faunas from south-western Gondwana, especially about their locomotion and functionality of fore and hind autopods. The inferred limb posture of the *Pentasauropus* trackmaker finds a match with the osteological data provided by the therapsid record (e.g. King, 1981a; Fröbisch, 2006) and allows to corroborate interpretations derived from body-fossils. Meristic and qualitative track characters and trackway parameters, if jointly considered, suggest that the *Pentasauropus* trackmaker had a semi-erect to erect posture, especially the hind limbs (Fig. 11A-B).

Contrarily to what was stated ~~in the past about therapsid posture~~ (Charig, 1980; Bonaparte, 1982), therapsid-grade limb osteology was characterized by several important modifications, which indicate a more parasagittal stance of the limbs (Romer, 1956; Boonstra, 1967; Jenkins, 1971), especially if compared with the prevalent sprawling posture of non-therapsid synapsids (Romer, 1956; Hopson, 2015). Modifications of the scapula and the glenoid have allowed the elbow to rotate inwardly, bringing the humerus closer to the sagittal plane (Walter, 1986). The iliac blade was expanded anteriorly and allowed the insertion of a larger **iliofemoralis muscle**, enabling femoral retraction (Romer, 1956; Walter, 1986). Moreover, the femoral head folded medially and enabled a more parasagittal position of the propodial (Romer, 1922; Walter, 1986).

Concerning the dynamics of locomotion in non-mammalian therapsids, Kemp (1978) proposed a **dual-gait condition**, intermediate between the plesiomorphic gait of amniotes (Sumida & Modesto, 2001) and the mammalian erect gait, based on the therocephalian *Regisaurus jacobi*. This condition has proved to be not possible for derived dicynodonts, such as *Kingoria nowacki* (King, 1985) and for kannemeyeriiformes dicynodonts, all characterized by an ankle joint inhibiting extensive rotational movements needed for dual-gait locomotion (Fröbisch,

2006). In dicynodonts, the forelimb step cycle was performed in an abducted (i.e. sprawling) posture, whereas the hind limb step cycle passed from a primitive abducted posture in earlier members, such as *Robertia broomiana* (see King, 1981b) to an adducted (i.e. erect) posture in more derived taxa (Walter, 1986), such as *Dicynodon trigonocephalus* and *Tetragonias njalilus* (e.g. King, 1981a; Fröbisch, 2006).

The autopod posture proposed for the studied tracks quite differ from the information and reconstructions derived from the body-fossil record. As stated before, the alleged autopodial structure inferred from *Pentasauropus* tracks is dictated by the relative distance between the base of digital (ungual) traces and the distal edge of the sub-circular pad trace, which has been referred to a fleshy pad behind the basipodial. The observed track morphology seems to imply that, except for acropodial and the fleshy pad, no other bony elements of the producer's autopods were imprinted on the substrate, consequently indicating that they were likely raised in position.

Such a configuration is considered valid for the foot bones in a static-state and would fall at least within a subunguligrade posture, implying that the phalanges were the only bony pedal elements contacting the ground in a static stance. However, if the three-dimensional footprint morphology is considered (i.e. ungual traces and pad trace behind them) concurrently with spatial data regarding pad trace/digit trace distance (Table 2), it is evident that the unguals were not the only pedal elements performing the cycle of locomotion. Thus, the foot cannot be regarded as subunguligrade from a dynamic point of view. During locomotion, the body weight of *Pentasauropus* producers was not carried only by phalanges but, most likely, the entire foot supported the load (Figs. 11, 12). The fleshy pad behind the basipodials actively contacted the ground most likely during the touch-down and weight-bearing phase, as was already inferred from footprint depth of impression in other producers (e.g. Romano, Citton & Nicosia, 2016;

Citton et al., 2017). Thus, from a functional standpoint, the autopod posture of the *Pentasauropus* trackmaker can be regarded as plantiportal (*sensu* Michilsens et al., 2009). Such a posture could have been accompanied by an arched configuration of the articulated metapodials and at least of the proximal phalanges (Kümmell & Frey, 2012) (Fig. 11C-D). Metacarpals forming an arched configuration when articulated were described in a specimen of *Tetragonias njalilus* (Cruickshank, 1967), and this kind of configuration could have been accompanied by little movement capabilities (Rubidge & Hopson, 1996) of the metapodials and could have dictated the observed relative position of the ungual traces. A manual/pedal structure like the one here hypothesized could have maintained a large surface in contact with the ground by means of cartilaginous elements and fleshy cushions on which the basipodials rested, ensuring a supportive role of the whole autopods during the cycle of locomotion and particularly during the maximum load. Digit traces were formed by **acropodials** deeply penetrating into the substrate during the final weight-bearing phase, kick-off and thrust. This could explain the different depth of the impression that is observed in completely preserved tracks. Among digits the series II-IV, and with a lesser extent digit I, played a major role in performing the end of the cycle of locomotion. Drag traces affecting the most medial digits could be formed during the recovery of the autopod at the end of the step.

A functionally plantiportal posture has been described in several mammals regardless of body-weight (e.g. South-American coati, armadillo, coypu, among others; see Michilsens et al., 2009) but also can represent a functional strategy, co-~~occurrent~~ with a graviportal structure of the limbs. A subunguligrade-plantiportal foot implies a complex set of associated characters in the autopodial anatomy of the *Pentasauropus* producer. Taking into account the large but not gigantic dimensions of the putative trackmakers (e.g. Morato, 2006,

estimates a body mass between 23- 32 kg for juvenile of *Dinodontosaurus* and exceeding 300 kg for adults, while Mancuso et al., 2014, indicate a body mass of 362 kg for *Dinodontosaurus platyceps* and 170 kg for *Dinodontosaurus brevirostris*), most likely the subunguligrade-plantiportal autopod posture was promoted in these dicynodonts regardless of the body-dimension, not necessarily implying an increase in body size but being a pre-requisite for lineages experiencing such an evolutionary path.

The Vera Formation and the track record: chronostratigraphical observations

As before stated, *Pentasauropus* or *Pentasauropus*-like footprints were reported to date mainly from Upper Triassic units. In Argentina, *Pentasauropus* tracks were reported both from Upper Triassic unit (e.g. Portezuelo Formation) and from late Middle Triassic unit (Cerro de Las Cabras Formation). In Lesotho (Southern Africa) *Pentasauropus* was reported from the lower Elliot Formation (Stormberg Group), which lies above the Carnian Molteno Formation. The lower Elliot Formation was considered Upper Triassic by Ellenberger (1970), Norian-Rhaetian by Olsen & Galton (1984) and Norian by Knoll (2004), Lucas & Hancox (2001) and Lucas (2018), based on fossil remains, both bones and traces. Recently, the Elliot Formation (lower and upper) was discussed by means of magnetostratigraphy, and fixed as Upper Triassic - Lower Jurassic by Sciscio et al. (2017). The same authors confirmed a Norian-Rhaetian age for the lower Elliot Formation and correlated the unit with the Los Colorados Formation in the Ischigualasto-Villa Union Basin of Argentina (Sciscio et al., 2017). A Late Triassic age of *Pentasauropus*-bearing levels of the Vera Formation is here proposed.

At the same time, the recent datations provided by Luppó et al. (2017) contrast with the Late Triassic age historically proposed for the whole Vera Formation and in particular for the deposits bearing the ‘*Dicroidium*’ flora and *Dicynodontipus*. On the basis of the new isotopic

ages, Luppó et al. (2017) concluded that the levels bearing the ‘*Dicroidium*’-type flora (Artabe, 1985a, b) are intercalated between deposits dated 252 ± 2 Ma (Changhsingian, Late Permian) and 248 ± 2 Ma (Olenekian, Early Triassic). These authors also suggested that the stratigraphic position of the deposits exposed in the Tchering quarry, west of Los Menucos town, where *Dicynodontipus* (*sensu* Melchor & de Valais, 2006) come from, is not yet completely clear. Nevertheless, this quarry is spatially close to the outcrops where geochronological data were provided by Luppó et al. (2017). On the other hand, the Yancaqueo farm from which the *Pentasauropus* footprints come, is located east of Los Menucos town and lacks detailed geochronological and geological studies.

Thus, taking into account these new data and the chronostratigraphical distribution of *Dicynodontipus* (e.g. Haubold, 1983; Ceoloni et al., 1988; Retallack, 1996; de Klerk, 2002; Marsicano et al., 2004; Hunt & Lucas, 2007; Klein & Lucas, 2010; Costa da Silva, Sedor & Sequeira Fernandes, 2012; Fichter & Kunz, 2013; Díaz-Martínez et al., 2015), the previously proposed Late Triassic age for the whole Vera Formation (i.e. strata bearing *Dicynodontipus* and *Pentasauropus*, respectively) is here questioned.

CONCLUSIONS

Large pentadactyl tracks from the Upper Triassic Vera Formation of the Los Menucos Group (Río Negro province, North Patagonia, Argentina) were studied and discussed in terms of palaeobiological attribution.

The tracks are currently referred to as *Pentasauropus* Ellenberger 1970, an ichnotaxon established from the Upper Triassic lower Elliot Formation (Stormberg Group) of Karoo Basin (Lesotho, Southern Africa).

Material under study allowed to more effectively appreciate ichnotaxon variability and proved to be significant for a better definition of the locomotor dynamics of the producer and particularly of its foot anatomy.

Track and trackway parameters indicate a dicynodont as the most probable producer, and a relationship with the South-American members of the clade Kannemeyeriiformes is proposed.

An affinity between the Gondwanan therapsid ichnofauna and that from South Africa is evident, as well as functional features of the autopods of the producer are considered significantly similar and may be related to the same autopodial anatomy shared by the clade.

The autopod posture for the *Pentasauropus* trackmaker has been interpreted as subunguligrade in static posture and plantiportal during locomotion. A large pad of connective tissue behind the basipodials and partially metapodials can be proposed for the heavy-footed producers of *Pentasauropus*. The cushion allowed to decrease the stress transferred to the bones and spread it on a larger area during the touch-down and weight-bearing phase of the locomotion cycle.

Finally, a Late Triassic age for the *Pentasauropus*-bearing levels of Vera Formation is confirmed, based on the age of other lithostratigraphic units bearing *Pentasauropus* in South Africa and United States. At the same time, a detailed stratigraphic study of the lower, bearing-*Dicynodontipus*, strata of the Vera Formation is needed to corroborate palaeontological and geochronological data and to account the validity of the Vera Formation as lithostratigraphic unit.

ACKNOWLEDGEMENTS

I. Cerda and C. Muñóz of the Museo Provincial Carlos Ameghino (Cipolletti, Río Negro province, Argentina), D. Ramos and S. Mercado of the Museo Municipal de Los Menucos (Los Menucos, Río Negro province, Argentina) and R. Rial of the Museo Provincial María Inés Kopp (Valcheta, Río Negro province, Argentina) are kindly acknowledged for having made access to Los Menucos material possible and for their assistance during museum operations. Umberto Nicosia is warmly thanked for having reviewed an advanced draft of the manuscript. Christian F. Kammerer is sincerely thanked for having discussed some aspects of the autopod osteology in kannemeyeriiformes. Spencer Lucas, Hendrik Klein, Paula Dentzien-Dias and the Academic Editor Kenneth De Baets are acknowledged for having provided revisions that significantly improved the manuscript.

REFERENCES

- Abdala F, Marsicano CA, Smith RMH, Swart R. 2013. Strengthening Western Gondwanan correlations: A Brazilian Dicynodont (Synapsida, Anomodontia) in the Middle Triassic of Namibia. *Gondwana Research* 23 (3): 1151–1162. DOI: 10.1016/j.gr.2012.07.011
- Anderson JM, Anderson HM, Cruickshank ARI. 1998. Late Triassic ecosystems of the Molteno/Lower Elliot biome of Southern Africa. *Palaeontology* 41: 387–421.
- Artabe AE. 1985a. Estudio sistemático de la taoflora triásica de Los Menucos, provincia de Río Negro, Argentina. Parte I. *Ameghiniana* 22(1-2): 3–22.
- Artabe AE. 1985b. Estudio sistemático de la taoflora triásica de Los Menucos, provincia de Río Negro, Argentina. Parte II. *Ameghiniana* 22(3-4): 197–212.
- Avanzini M, Piñuela L, García-Ramos JC. 2011. Late Jurassic footprints reveal walking kinematics of theropod dinosaurs. *Lethaia* 45(2): 238–252. DOI: 10.1111/j.1502-3931.2011.00276.x

- Baird D. 1980. A prosauropod dinosaur trackway from the Navajo Sandstone (Lower Jurassic) of Arizona. In: Jacobs LL, ed. *Aspect of Vertebrate History: essay in honor of Harris Colbert*. Flagstaff Museum of Northern Arizona Press, 219–230.
- Barredo S, Chemale F, Ávila JN, Marsicano C, Ottone EG, Ramos VA, 2012. Tectono–sequence stratigraphy and U–Pb zircon ages of the Rincón Blanco depocenter, northern Cuyo rift, Argentina. *Gondwana Research* 21: 624–636. DOI: 10.1016/j.gr.2011.05.016
- Belvedere M, Farlow JO. 2016. A Numerical Scale for Quantifying the Quality of Preservation of Vertebrate Tracks. In: Falkingham PL, Marty D, Richter A, eds. *Dinosaur tracks: the next steps*. Bloomington: Indiana University, Press, 93–98.
- Bogan S, Taverne L, Agnolin F. 2013. First Triassic and oldest record of a South American amiiiform fish: *Caturus* sp. from the Los Menucos Group (lower Upper Triassic), Río Negro province, Argentina. *Geologica Belgica* 16(3): 191–195.
- Bonaparte JF. 1966. Cronología de algunas formaciones triásicas argentinas. *Revista de la Asociación Geológica Argentina* 21: 20–38.
- Bonaparte JF. 1969. Dos nuevas “faunas” de reptiles triásicos de Argentina. In: *First Gondwana Symposium*, Mar del Plata 1967: Abstract book, 283–306.
- Bonaparte JF. 1971. Annotated list of the South American Triassic tetrapods. *Symposium on Gondwana Stratigraphy. Proceedings and Papers, Pretoria* 2: 665–682.
- Bonaparte JF. 1981. Nota sobre una nueva fauna del Triasico Inferior del Sur de Mendoza, Argentina, correspondiente a la zona de *Lystrosaurus* (Dicynodontia-Proterosuchia). In: *Congreso Latino-Americano de Paleontologia*, Abstract book, 277–288.
- Bonaparte JF. 1982. Faunal replacement in the Triassic of South America. *Journal of Vertebrate Paleontology* 2: 362–371. DOI: 10.1080/02724634.1982.10011938

- 646 Boonstra LD. 1967. An early stage in the evolution of the mammalian quadrupedal walking gait.
647 *Annals of the South African Museum* 50: 27–42.
- 648 Bordy EM, Abrahams M, Sciscio L. 2017. The Subeng vertebrate tracks: stratigraphy,
649 sedimentology and a digital archive of a historic Upper Triassic palaeosurface (lower
650 Elliot Formation), Leribe, Lesotho (southern Africa). *Bollettino della Società*
651 *Paleontologica Italiana* 56(2): 181–198. DOI: 10.4435/BSPI.2017.12
- 652 Carpenter K. 1992. Behavior of hadrosaurs as interpreted from footprints in the “Mesaverde”
653 Group (Campanian) of Colorado, Utah, and Wyoming. *Contributions to Geology,*
654 *University of Wyoming* 29(2): 81–96.
- 655 Casamiquela RM. 1964. Estudios icnológicos. Problemas y métodos de la icnología con
656 aplicación al estudio de pisadas mesozoicas (Reptilia, Mammalia) de la Patagonia.
657 Talleres Gráficos Colegio Industrial Pío IX, Buenos Aires.
- 658 Casamiquela RM. 1975. Nuevo material y reinterpretación de las icnitas mesozoicas
659 (Neotriásicas) de Los Menucos, Provincia de Rio Negro (Patagonia). In: *1° Congreso*
660 *Argentino de Paleontología y Biostratigrafía*, Tucumán, Argentina 1974: Abstract book,
661 555–580.
- 662 Casamiquela RM. 1987. Novedades en icnología de vertebrados en la Argentina. In: *10°*
663 *Congreso Brasileiro de Paleontologia*, Rio de Janeiro, Brazil: Abstract book 1, 445–456.
- 664 Ceoloni P, Conti MA, Mariotti N, Nicosia U. 1988. New Late Permian tetrapod footprints from
665 Southern Alps. *Memorie della Società Geologica Italiana* 34: 45–65.
- 666 Charig AJ. 1980. Differentiation of lineages among Mesozoic tetrapods. *Mémoires de la Société*
667 *Géologique de France N.S.* 59: 207–210.
- 668 Citton P, Carluccio R, Nicolosi I, Nicosia U. 2017. Re-evaluation of *Chelichnus tazelwürmi*, a

- 669 non mammalian therapsid-grade track from the Upper Permian Arenaria di Val Gardena.
- 670 *Historical Biology*, DOI: 10.1080/08912963.2017.1370586
- 671 Costa da Silva R, Sedor FA, Sequeira Fernandes AC. 2012. Fossil footprints from the Late
- 672 Permian of Brazil: an example of hidden biodiversity. *Journal of South American Earth*
- 673 *Sciences* 38: 31–43. DOI: 10.1016/j.jsames.2012.05.001
- 674 Cox CB. 1962. Preliminary diagnosis of *Ischigualastia*, a new genus of dicynodont from
- 675 Argentina. *Breviora* 156: 8–9.
- 676 Crompton AW, Jenkins Jr. FA. 1973. Mammals from Reptiles: a review of mammalian origins.
- 677 *Annual Review of Earth Planetary Science*, 1: 131–155. DOI:
- 678 10.1146/annurev.ea.01.050173.001023
- 679 Cruickshank ARI. 1967. A new dicynodont genus from the Manda Formation of Tanzania
- 680 (Tanganyika). *Journal of Zoology* 153: 163–208. DOI: 10.1111/j.1469-
- 681 7998.1967.tb04059.x
- 682 Cucchi R, Busteros A, Lema H. 2001. Hoja Geológica 4169 - II, Los Menucos, Provincia de Río
- 683 Negro. IGRM-SEGE-MAR Boletín 265: 1–105. Buenos Aires.
- 684 D’Orazi Porchetti S, Nicosia U. 2007. Re-examination of some large Early Mesozoic tetrapod
- 685 footprints from the African collection of Paul Ellenberger. *Ichnos* 14(3-4): 219–245.
- 686 DOI: 10.1080/10420940601049990
- 687 de Klerk WJ. 2002. A dicynodont trackway from the *Cistecephalus* Assemblage Zone in the
- 688 Karoo, east of Graaff-Reinet, South Africa. *Palaeontologia Africana* 38: 73–91.
- 689 <http://hdl.handle.net/10539/16345>
- 690 de Valais S. 2008. Icnología de tetrápodos triásicos y jurásicos de Argentina: aportes al origen de
- 691 las aves y los mamíferos. D. Phil Thesis, Universidad de Buenos Aires, Facultad de
- 692 Ciencias Exactas y Naturales.

- de Valais S, Melchor RN, Bellosi E. 2006. New large vertebrate footprints from the Cerro de Las Cabras Formation (Middle Triassic), Mendoza province. *Ameghiniana* 43 (suppl.): 34R.
- Díaz-Martínez I, Castanera D, Gasca JM, Canudo JL. 2015. A reappraisal of the Middle Triassic chirotheriid *Chirotherium ibericus* Navás, 1906 (Iberian Range NE Spain), with comments on the Triassic tetrapod track biochronology of the Iberian Peninsula. *PeerJ* 3:e1044. DOI 10.7717/peerj.1044
- Díaz-Martínez I, de Valais S. 2014. Estudio de la variabilidad en la conservación de huellas de tetrápodos del Triásico Superior de Los Menucos, Río Negro, Argentina. *Ameghiniana* 52 (suppl. 1): 8.
- Domnanovich NS, Marsicano C. 2006. Tetrapod footprints from the Triassic of Patagonia: reappraisal of the evidence. *Ameghiniana* 43(1): 55–70.
- Domnanovich NS, Marsicano C. 2012. The Triassic dicynodont *Vinceria* (Therapsida, Anomodontia) from Argentina and a discussion on basal Kannemeyeriiformes. *Geobios* 45: 173–186. DOI: 10.1016/j.geobios.2011.03.003
- Domnanovich NS, Tomassini R, Manera de Bianco T, Dalponte M. 2008. Nuevos aportes al conocimiento de la icnofauna de tetrápodos del Triásico Superior de Los Menucos (Complejo Los Menucos), provincia de Río Negro, Argentina. *Ameghiniana* 45(1): 221–224.
- Duncan H. 1831. An account of the tracks and footmarks of animals found impressed on sandstone in the quarry of Corncockle Muir, in Dumfriesshire. *Transactions of the Royal Society of Edinburgh* 11: 194–209. DOI: 10.1017/S0080456800021906

- 714 Ellenberger P. 1955. Note préliminaire sur les pistes et les restes osseux de Vertébrés du
715 Basutoland (Afrique du Sud). *Comptes Rendus Hebdomadaires des Séances de*
716 *l'Académie des Sciences* 240: 889–891.
- 717 Ellenberger P. 1970. Les niveaux paléontologiques de première apparition des mammifères
718 primordiaux en Afrique du Sud et leur ichnologie. Etablissement de zones
719 stratigraphique détaillées dans le Stormberg du Lesotho (Afrique du Sud) (Trias superior
720 a Jurassique). *Proceedings and Papers II Gondwana Symposium* 1970: 343–370.
- 721 Ellenberger P. 1972. Contribution à la classification des Pistes de Vertébrés du Trias: Les types
722 du Stormberg d'Afrique du Sud (I). *Paleovertebrata, Memoire Extraordinaire*,
723 Montpellier: 152 p.
- 724 Ellenberger F, Ellenberger P. 1958. Principaux types de pistes de vertébrés dans les couches du
725 Stormberg au Basutoland (Afrique du Sud) (Note préliminaire). *Comptes-Rendus*
726 *sommaire des Séances de la Société géologique de France*: 65–67.
- 727 Falk AR, Hasiotis ST, Gong E, Lim J-D, Brewer ED. 2017. A new experimental setup for
728 studying avian neoichnology and the effects of grain size and moisture content on tracks:
729 trials using the domestic chicken (*Gallus gallus*). *Palaaios* 32 (11): 689–707.
- 730 Falkingham PL, Bates KT, Avanzini M, Bennett M, Bordy E, Breithaupt BH, Castanera D,
731 Citton P, Díaz-Martínez I, Farlow JO, Fiorillo AR, Gatesy SM, Getty P, Hatala KG, Hornung
732 JJ, Hyatt JA, Klein H, Lallensack JN, Martin AJ, Marty D, Matthews NA, Meyer ChA, Milàn
733 J, Minter NJ, Razzolini NL, Romilio A, Salisbury SW, Sciscio L, Tanaka I, Wiseman ALA,
734 Xing LD, Belvedere M. 2018. A standard protocol for documenting modern and fossil
735 ichnological data. *Palaeontology* (2018): 1–12. DOI: 10.1111/pala.12373

- Fichter J, Kunz R. 2013. “Dinosauiromorph” tracks from the Middle Buntsandstein (Early Triassic: Olenekian) of Wolfhagen, northern Hesse, Germany. *Comunicações Geológicas* 100 (1): 81–88. <http://www.lneg.pt/iedt/unidades/16/paginas/26/30/141>
- Franzese JR, Spalletti LA. 2001. Late Triassic-Early Jurassic continental extension in southwestern Gondwana: tectonic segmentation and pre-break-up rifting. *Journal of South American Earth Sciences* 14: 257–270. DOI: 10.1016/S0895-9811(01)00029-3
- Fröbisch J. 2006. Locomotion in derived dicynodonts (Synapsida, Anomodontia): a functional analysis of the pelvic girdle and hind limb of *Tetragonias njalilus*. *Canadian Journal of Earth Sciences* 43: 1297–1308. DOI: 10.1139/e06-031
- Fröbisch J. 2009. Composition and similarity of global anomodont-bearing tetrapod faunas. *Earth-Science Reviews* 95: 119–157. DOI:10.1016/j.earscirev.2009.04.001
- Gallego OF. 2010. A new crustacean clam shrimp (Spinicaudata: Eosestheriidae) from the Upper Triassic of Argentina and its importance for ‘conchostracan’ taxonomy. *Alcheringa: An Australasian Journal of Palaeontology* 34: 179–195. DOI: 10.1080/03115510903546152
- Galton PM, Heerden van J. 1998. Anatomy of the prosauropod dinosaur *Blikanasaurus cromptoni* (Upper Triassic, South Africa), with notes on the other tetrapods from the Lower Elliot Formation. *Paläontologische Zeitschrift* 72: 163–177. DOI: 10.1007/BF02987824
- Gaston R, Lockley MG, Lucas SG, Hunt AP. 2003. *Grallator*- dominated fossil footprints assemblages and associated enigmatic footprints from the Chinle Group (Upper Triassic), Gateway Area, Colorado. *Ichnos* 10: 153–163. DOI: 10.1080/10420940390256258
- González SN, Greco GA, Sato AM, Llambías EJ, Basei MAS, Gonzalez PD, Díaz PE. 2017. Middle Triassic trachytic lava flows associated with coeval dyke swarm in the North

Patagonian Massif: a postorogenic magmatism related to extensional collapse of the
Gondwanide orogen. *Journal of South American Earth Sciences* 75: 134–143. DOI:
10.1016/j.jsames.2017.02.007

Haubold H. 1974. *Die fossilien Saurierfährten*. A. Ziemsen, Wittenberg.

Haubold H. 1983. Archosaur evidence in the Buntsandstein (Lower Triassic). *Acta
Palaeontologica Polonica* 28 (1-2): 123–132.

Haubold H. 1984. *Saurierfährten*. Die Neue Brehm-Bucherei, A. Ziemsen Verlag, Wittenberg
Lutherstadt.

Hitchcock E. 1836. Ornithichnology - description of the foot marks of birds (Ornithichnites) on
New Red Sandstone in Massachusetts. *American Journal of Science* 29: 307–340.

Hopson JA. 1995. Patterns of evolution in the manus and pes of nonmammalian therapsids.
Journal of Vertebrate Paleontology 15: 615–639. DOI:
10.1080/02724634.1995.10011252

Hopson JA. 2015. Fossils, trackways, and transitions in locomotion: a case study of *Dimetrodon*.
In: Dial KP, Shubin N, Brainerd EL, eds. *Great transformations in vertebrate evolution*.
Chicago: The University of Chicago Press, 125–141. DOI:
10.7208/chicago/9780226268392.001.0001

Huene F. von. 1931. Die fossilen Fährten im Rhät von Ischigualasto in Nordwest-Argentinien.
Palaeobiologica 4: 99–112.

Hunt AP, Lucas SG. 2007. The Triassic tetrapod track record: ichnofaunas, ichnofacies and
biochronology. *New Mexico Museum of Natural History and Science Bulletin* 41:78–87.

Hunt AP, Lucas SG, Klein H. 2018. Late Triassic nonmarine vertebrate and invertebrate trace
fossils and the pattern of the Phanerozoic record of vertebrate trace fossils. In: Tanner

- LH, ed. *The Late Triassic World*. Cham: Springer International Publishing, 447–544.
DOI: 10.1007/978-3-319-68009-5_12
- Hunt-Foster RK, Lockley MG, Milner ARC, Foster JR, Matthews NA, Breithaupt BH, Smith JA.
2016. Tracking dinosaurs in BLM canyon country, Utah. *Geology of the Intermountain
West* 3: 67–100.
- Jenkins FA, Jr. 1971. The postcranial skeleton of African cynodonts. *Peabody Museum of
Natural History, Yale University, Bulletin* 36: 1–216.
- Kammerer CF. 2018. The first skeletal evidence of a dicynodont from the lower Elliot Formation
of South Africa. *Palaeontologia africana* 52: 102–128.
<http://wiredspace.wits.ac.za/handle/10539/24148>
- Kaup JJ. 1835a. Über Thierfährten bei Hildburghausen. *Neues Jahrbuch für Mineralogie,
Geologie und Paläontologie* 1835: 327–328.
- Kaup JJ. 1835b. Fährten von Beuteltieren. In *Das Tierreich*: 246–248.
- Kemp TS. 1978. Stance and gait in the hindlimb of a therocephalian mammal-like reptile.
Journal of Zoology 186: 143–161. DOI: 10.1111/j.1469-7998.1978.tb03362.x
- King GM. 1981a. The functional anatomy of a Permian dicynodont. *Philosophical Transaction
of the Royal Society of London B* 291: 243–322. DOI: 10.1098/rstb.1981.0001
- King GM. 1981b. The post cranial skeleton of *Robertia broomiana*, an early dicynodont
(Reptilia, Therapsida) from the South African Karoo. *Annals of South African Museum*
84: 203–231.
- King GM. 1985. The postcranial skeleton of *Kingoria nowacki* (von Huene) (Therapsida:
Dicynodontia). *Zoological Journal of the Linnean Society* 84: 263–298. DOI:
10.1111/j.1096-3642.1985.tb01801.x

- 805 Klein H, Lucas SG. 2010. Tetrapod footprints – their use in biostratigraphy and biochronology of
806 the Triassic. In: Lucas SG, ed. *The Triassic Timescale. Geological Society, London,*
807 *Special Publications* 334: 419–446. DOI: 10.1144/SP334.14
- 808 Knoll F. 2004. Review of the tetrapod fauna of the “Lower Stormberg Group” of the main Karoo
809 Basin (southern Africa): implication for the age of the Lower Elliot Formation. *Bulletin*
810 *de la Societe géologique de France* 175(1): 73–83. DOI: 10.2113/175.1.73
- 811 Kokogian DA, Spalletti L, Morel E, Artabe A, Martínez RN, Alcober OA, Milana JP, Zavattieri
812 AM, Papù OH. 1999. Los depósitos continentales triásicos. In: Caminos R, Panza J, eds.
813 *Geología Argentina, Instituto de Geología y Recursos Minerales.* Buenos Aires, Anales,
814 29(15), 377–398.
- 815 Kubo T, Benton MJ. 2009. Tetrapod postural shift estimated from Permian and Triassic
816 trackways. *Palaeontology* 52: 1029–1037. DOI: 10.1111/j.1475-4983.2009.00897.x
- 817 Kubo T, Ozaki M. 2009. Does pace angulation correlate with limb posture? *Palaeogeography,*
818 *Palaeoclimatology, Paleoecology* 275: 54–58. DOI: 10.1016/j.palaeo.2009.02.001
- 819 Kümmell S, Frey E. 2012. Digital arcade in the autopodia of Synapsida: standard position of the
820 digits and dorsoventral excursion angle of digital joints in the rays II–IV.
821 *Palaeobiodiversity and Palaeoenvironments* 92: 171–196. DOI: 10.1007/s12549-012-
822 0076-6
- 823 Labudía CH, Bjerg EA. 2001. El Grupo Los Menucos: redefinición estratigráfica del Triásico
824 Superior del Macizo Nordpatagonico. *Revista de la Asociación Geologica Argentina* 56:
825 404–407.

- 826 Labudía CH, Bjerg EA. 2005. Geología del Grupo Los Menucos, Comarca Nordpatagónica,
827 Argentina. In: *16° Congreso Geológico Argentino*, La Plata, Argentina: Abstract book,
828 233–238.
- 829 Labudía CH, Llambías EJ, Rapela CW, Artabe AE. 1995. El Triásico de Los Menucos: procesos
830 volcánicos y sedimentarios. In: *2° Reunión del Triásico de Cono Sur*, Bahía Blanca,
831 Argentina: Abstract book, 17–21.
- 832 Lema H, Busteros A, Giacosa RE, Cucchi R. 2008. Geología del complejo volcánico Los
833 Menucos en el área tipo, Río Negro. *Revista de la Asociación Geológica Argentina* 63:
834 3–13.
- 835 Leonardi G. 1987. *Glossary and manual of tetrapod footprint palaeoichnology*. Brasilia:
836 Departamento Nacional da Produção Mineral.
- 837 Leonardi G. 1994. *Annotated atlas of South America tetrapod footprints (Devonian-Holocene)*.
838 Ministerio de Minas y Energia, Companhia de Pesquisa de Recursos Minerais, Brazil.
- 839 Leonardi G, de Oliveira FH. 1990. A revision of the Triassic and Jurassic tetrapod footprints of
840 Argentina and a new approach on the age and meaning of the Botucatu Formation
841 footprints (Brazil). *Revista Brasileira de Geociencias* 20: 216–229.
- 842 Lockley MG. 1991. *Tracking Dinosaurs. A new look at an ancient world*. Cambridge University
843 Press, Cambridge.
- 844 Lockley MG, Hunt AP. 1995. *Dinosaur tracks and other fossil footprint of the Western United*
845 *States*. Columbia University Press, New York.
- 846 Lockley MG, McCrea RT, Buckley LG, Lim JD, Matthews NA, Breithaupt BH, Houck KJ,
847 Gierlinski GD, Surmik D, Kim KS, Xing L, Kong DY, Cart K, Martin J, Hadden G.
848 2016. Theropod courtship: large scale physical evidence of display arenas and avian-like

scrape ceremony behaviour by Cretaceous dinosaurs. *Scientific Reports* 6:18952. DOI: 10.1038/srep18952

Lockley MG, Meyer CA. 2000. *Dinosaur Tracks and other fossil footprints of Europe*. Columbia University Press, New York.

Lucas SG. 1998. Global Triassic tetrapod biostratigraphy and biochronology. *Palaeogeography, Palaeoclimatology, Palaeoecology* 143: 347–384. DOI: 10.1016/S0031-0182(98)00117-5

Lucas SG. 2002. A new dicynodont from the Triassic of Brazil and the tetrapod biochronology of the Brazilian Triassic. *New Mexico Museum of Natural History and Science Bulletin* 21: 131–141.

Lucas SG. 2010. The Triassic timescale based on nonmarine tetrapod biostratigraphy and biochronology. In: Lucas SG, ed. *The Triassic timescale*. Geological Society, London, Special Publications 334: 447–500. DOI: 10.1144/SP334.15

Lucas SG. 2018. Late Triassic Terrestrial Tetrapods: biostratigraphy, biochronology and biotic events In: Tanner LH, ed. *The Late Triassic World*. Cham: Springer International Publishing, 351–406. DOI: 10.1007/978-3-319-68009-5_12

Lucas SG, Hancox J. 2001. Tetrapod-based correlation of the nonmarine Upper Triassic of Southern Africa. *Albertiana* 25: 5–9.

Luppo T, López De Luchi MG, Rapalini AE, Martínez Dopico CI, Fanning CM. 2017. Geochronologic evidence of a large magmatic province in northern Patagonia encompassing the Permian-Triassic boundary. *Journal of South American Earth Sciences* (2017): 1–10. DOI: 10.1016/j.jsames.2018.01.003

Mackenzie WS, Donaldson CH, Guilford C. 1982. *Atlas of Igneous Rocks and Their Textures*. John Wiley, New York.

872 Mallison H, Wings O. 2014. Photogrammetry in Paleontology – a practical guide. *Journal of*
873 *Paleontological Techniques* 12: 1–31.

874 Mancuso AC, Gaetano LC, Leardi JM, Abdala F, Arcucci AB. 2014. The Chañares Formation: a
875 window to a Middle Triassic tetrapod community. *Lethaia* 47: 244–265. DOI:
876 10.1111/let.12055

877 Marsicano CA, Arcucci AB, Mancuso A, Caselli AT 2004. Middle Triassic tetrapod footprints of
878 southern South America. *Ameghiniana* 41 (2): 171–184.

879 Marsicano CA, Barredo SP. 2004. A Triassic tetrapod footprint assemblage from southern South
880 America: palaeobiogeographical and evolutionary implications. *Palaeogeography,*
881 *Palaeoclimatology, Palaeocology* 203: 313–335. DOI: 10.1016/S0031-0182(03)00689-8

882 Marty D, Falkingham PL, Richter A. 2016. Dinosaur Track Terminology: A Glossary of Terms.
883 In: Falkingham PL, Marty D, Richter A, eds. *Dinosaur tracks: the next steps*.
884 Bloomington: Indiana University, Press, 399–402.

885 Melchor RN, de Valais S. 2006. A review of Triassic tetrapod track assemblages from Argentina.
886 *Palaeontology* 49(2): 355–379. DOI: 10.1111/j.1475-4983.2006.00538.x

887 Melchor RN, Genise JF, Poiré DGG. 2001. Icnología de los depósitos continentales triásicos. In:
888 Artabe AE, Morel EM, Zamuner AB, eds. *El Sistema Triásico en Argentina*. Fundación
889 Museo de La Plata ‘Francisco Pascasio Moreno’, La Plata, 101–135.

890 Michilsens F, Aerts P, Van Damme R, D’Aout K. 2009. Scaling of plantar pressures in
891 mammals. *Journal of Zoology* 279: 236–242. DOI: 10.1111/j.1469-7998.2009.00611.x

892 Morato L. 2006. *Dinodontosaurus* (Synapsida, Dicynodontia): reconstituições morfológicas e
893 aspectos biomecânicos. Master Thesis, Universidade Federal do Rio Grande do Sul do
894 Porto Alegre, Instituto de Geociências, Área de Concentração Paleontologia.

895 Olsen PE, Galton PM. 1984. A review of the reptile and amphibian assemblages from the
896 Stormberg of southern Africa, with special emphasis on the footprints and the age of the
897 Stormberg. *Paleontologica Africana* 25: 87–110. <http://hdl.handle.net/10539/16131>

898 Peabody FE. 1948. Reptile and Amphibian trackways from the Lower Triassic Moenkopi
899 Formation of Arizona and Utah. *Bulletin of the Department of Geological Sciences,*
900 *Berkeley and Los Angeles* 27: 295–468.

901 Peabody FE. 1959. Trackways of living and fossil salamanders. *University of California*
902 *Publications in Zoology* 63: 1–72.

903 Rapela CW, Pankhurst RJ, Llambías EJ, Labudia CH, Artabe AE. 1996. Gondwana magmatism
904 of Patagonia: inner cordilleran calc-alkaline batholiths and bimodal volcanic provinces.
905 In: 3° *International Symposium on Andean Geodynamics, Saint Malo*, Extended Abstract
906 book, 791–794.

907 Retallack GJ. 1996. Early Triassic therapsid footprints from the Sidney Basin, Australia.
908 *Alcheringa* 20: 301–314. DOI: 10.1080/03115519608619473

909 Romano M, Citton P, Nicosia U. 2016. Corroborating trackmaker identification through footprint
910 functional analysis: the case study of *Ichniotherium* and *Dimetropus*. *Lethaia* 49: 102–
911 116. DOI: 10.1111/let.12136

912 Rogers RR, Arcucci A, Abdala F, Sereno PC, Forster CA, May CL. 2001. Paleoenvironment and
913 taphonomy of the Chañares Formation tetrapod assemblage (Middle Triassic),
914 northwestern Argentina: spectacular preservation in volcanogenic concretions. *Palaios* 16
915 (5): 461–481. DOI: 10.1669/0883-1351(2001)016<0461:PATOTC>2.0.CO;2

916 Romer AS. 1922. The locomotor apparatus of certain primitive and mammal-like reptiles.
 917 *Bulletin of the American Museum of Natural History* 46: 517–606.
 918 <http://hdl.handle.net/2246/929>

919 Romer AS. 1956. *Osteology of the Reptiles*. Malabar (FL): Krieger Publishing Company.

920 Romer AS. 1966. The Chañares (Argentina) Triassic reptile fauna; I. Introduction. *Breviora* 247:
 921 1–14.

922 Rubidge BS, Hopson JA. 1996. A primitive anomodont therapsid from the base of the Beaufort
 923 Group (Upper Permian) of South Africa. *Zoological Journal of the Linnean Society* 117:
 924 115–139. DOI: 10.1111/j.1096-3642.1996.tb02152.x

925 Scasso RA, Limarino CO. 1997. Petrología y diagénesis de rocas clásticas. *Asociación Argentina*
 926 *de Sedimentología, Publicación Especial* 1, 259 pp., Buenos Aires.

927 Sciscio L, de Kock M, Bordy E, Knoll F. 2017. Magnetostratigraphy over the Triassic-Jurassic
 928 boundary in the main Karoo Basin. *Gondwana Research* 51: 177–192. DOI:
 929 10.1016/j.gr.2017.07.009

930 Smith RMH. 1993. Sedimentology and ichnology of floodplain paleosurfaces in the Beaufort
 931 Group (Late Permian), Karoo Sequence, South Africa. *Palaios* 8:339–357. DOI:
 932 10.2307/3515265

933 Spalletti LA. 1999. Cuencas triásicas del Oeste argentino: origen y evolución. *Acta Geológica*
 934 *Hispanica* 32: 29–50.

935 Stipanovic PN. 1967. Consideraciones sobre las edades de algunas fases magmáticas del
 936 Neopaleozoico y Mesozoico. *Revista de la Asociación Geológica Argentina* 22: 101–133.

937 Stipanovic PN, Methol EJ. 1972. Macizo de Somún Curá. In: Leanza AF, ed. *Geología Regional*
 938 *Argentina*. Academia Nacional de Ciencias de Córdoba: 581–600. Córdoba.

- Stipanovic PN, Methol EJ. 1980. Comarca Norpatagónica. In: *Geología Regional Argentina*. Academia Nacional de Ciencias de Córdoba: 1071–1097. Córdoba.
- Stipanovic PN, Rodrigo F, Baulies OL, Martínez CG. 1968. Las formaciones pre-senonianas en el denominado Macizo Nordpatagónico y regiones adyacentes. *Revista de la Asociación Geológica Argentina* 23: 67–98.
- Sumida SS, Modesto SP. 2001. A phylogenetic perspective on locomotory strategies in early amniotes. *American Zoologist* 41: 586–597. DOI: 10.1093/icb/41.3.586
- Walter LR. 1986. The limb posture of kannemeyeriid dicynodonts: functional and ecological consideration. In: Padian K, ed. *The Beginning of the Age of Dinosaurs*. Cambridge University Press, Cambridge: 89–97.
- Zavattieri AM, Arcucci AB. 2007. Edad y posición estratigráfica de los tetrapodos del cerro Bayo de Potrerillos (Triásico), Mendoza, Argentina. *Ameghiniana* 44(1): 133–142. http://www.scielo.org.ar/scielo.php?script=sci_arttext&pid=S0002-70142007000100009&lng=es&nrm=iso

Figure captions

Figure 1 The Los menucos area. (A) Location map and geological sketch of Los Menucos area (from Labudía & Bjerg, 2005, redrawn and slightly modified). White star indicates Estancia Yancaqueo, from which the *Pentasauropus* footprints come. (B) Simplified stratigraphic section of the Los Menucos Group (from Labudía & Bjerg, 2005, redrawn and slightly modified).

Figure 2 Thin sections (MPCA 27029/19.1 and MPCA 27029/19.2) of track-bearing slab MPCA 27029-19. Inequigranular, epiclastic texture with anhedral and subhedral phenocrysts at the

base (A) and middle portion (B) of the track-bearing slab MPCA 27029-19. (C, D) Equigranular less epiclastic texture indicating a minor sedimentary reworking of the trampled surface.

Figure 3 Tracks mode of preservation. Convex hyporeliefs (A, C) fitting with concave epireliefs (B, D) preserved on slab MMLM 075-1 (true tracks and natural casts, respectively).

Figure 4 Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MPCA 27029-1; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs MPCA 27029-2; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E). In (A)-(D), footprint 2 and 5, note the non-impressed area between the sole pad trace and the base of digit traces. In (E)-(H) note the displacement areas behind digit traces, interpreted as the result of the thrust of digit pushing the sediment backwardly.

Figure 5 Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MPCA 27029-3; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs MPCA 27029-4; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E). Note the digit trailing marks slightly affecting the digit traces of footprint 2 in (A)-(D), which are absent in footprints showed in (E)-(H) where digit traces are roughly sub circular in morphology.

Figure 6 Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MPCA 27029-5; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs

MPCA 27029-9; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E).

Figure 7 Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MPCA 27029-16, produced by a juvenile trackmaker; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs MPCA 27029-21; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E). The general morphology and structure of footprints 6-10 in (A)-(D), left by a juvenile trackmaker, is identical to that characterizing larger footprints, even when preserved only as digit traces.

Figure 8 Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MPCA 27029-33; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs MMLM 075-1; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E). In (E)-(H) note the long and sharp digit trailing marks.

Figure 9 Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MMLM 1; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs MMLM 2; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E). In (A)-(D) note the long and sharp digit trailing marks affecting footprints 2 and 4, resembling those of Figure 8E-H.

Figure 10 Morphological and extramorphological features identified on the studied material. (A)

Manus track MPCA 27029/2 and (B) interpretative drawing. (C) Pes track MPCA 27029-1/5

and (D) interpretative drawing. Extramorphological features are in blue and grey, morphological features are in black.

Figure 11 Limb and autopod posture in the *Pentasauropus* trackmaker. Simplified reconstruction of limb posture in back (A) and lateral (B) views. Simplified reconstruction of zeugopodials and hind autopod in lateral (C) and bottom (D) views. In colour the possible extension of the fleshy cushion on which the basipodials rested, ensuring support during locomotion. See the supplementary video to get a more complete view of the reconstruction. Artwork by Fabio Manucci.

Figure 12 Speculative *in vivo* reconstruction (based on *Dinodontosaurus*) of a kannemeyeriiformes dicynodont, a most probable producer of *Pentasauropus* tracks. Reconstruction in back (A) and lateral (B) view of the trackmaker walking in amble gait. See the supplementary video to get a more complete view of the reconstruction. Artwork by Fabio Manucci.

Table captions

Table 1 Photogrammetric report. Main processing parameters of the photogrammetric models (from Agisoft Photoscan Professional reports).

Table 2 Mean measurements (in cm) of track and trackway parameters. ETW, external trackway width; Fl, footprint length; Fw, footprint width; GAD, gleno-acetabular distance: (a), ‘primitive’ alternate pace (the trunk length of the producer is underestimated); (b), alternate pace; (c), amble (a, b, c, considering primary overlap *sensu* Leonardi, 1987); ITW, internal trackway width; Mpa, manus pace angulation; Mpl, manus pace length; Msl, manus stride length; Ppa, pes pace angulation; Ppl, pes pace length; Psl, pes stride length; Psl/GAD, pes

1029 stride length/gleno-acetabular distance ratio; ETW/SL, external trackway width/stride length
1030 ratio.

1031 Table 3 Sole pad-ungual trace distance. Distance (in cm) between the distal margin of the sole
1032 pad trace and the proximal margin of the digit traces in complete *Pentasauropus* footprints.
1033 The measurements most likely indicate a raised and inclined position of the metapodial
1034 elements of fore and hind foot in the *Pentasauropus* trackmaker.

Figure 1

The Los menucos area.

(A) Location map and geological sketch of Los Menucos area (from Labudía & Bjerg, 2005, redrawn and slightly modified). White star indicates Estancia Yancaqueo, from which the *Pentasauropus* footprints come. (B) Simplified stratigraphic section of the Los Menucos Group (from Labudía & Bjerg, 2005, redrawn and slightly modified).

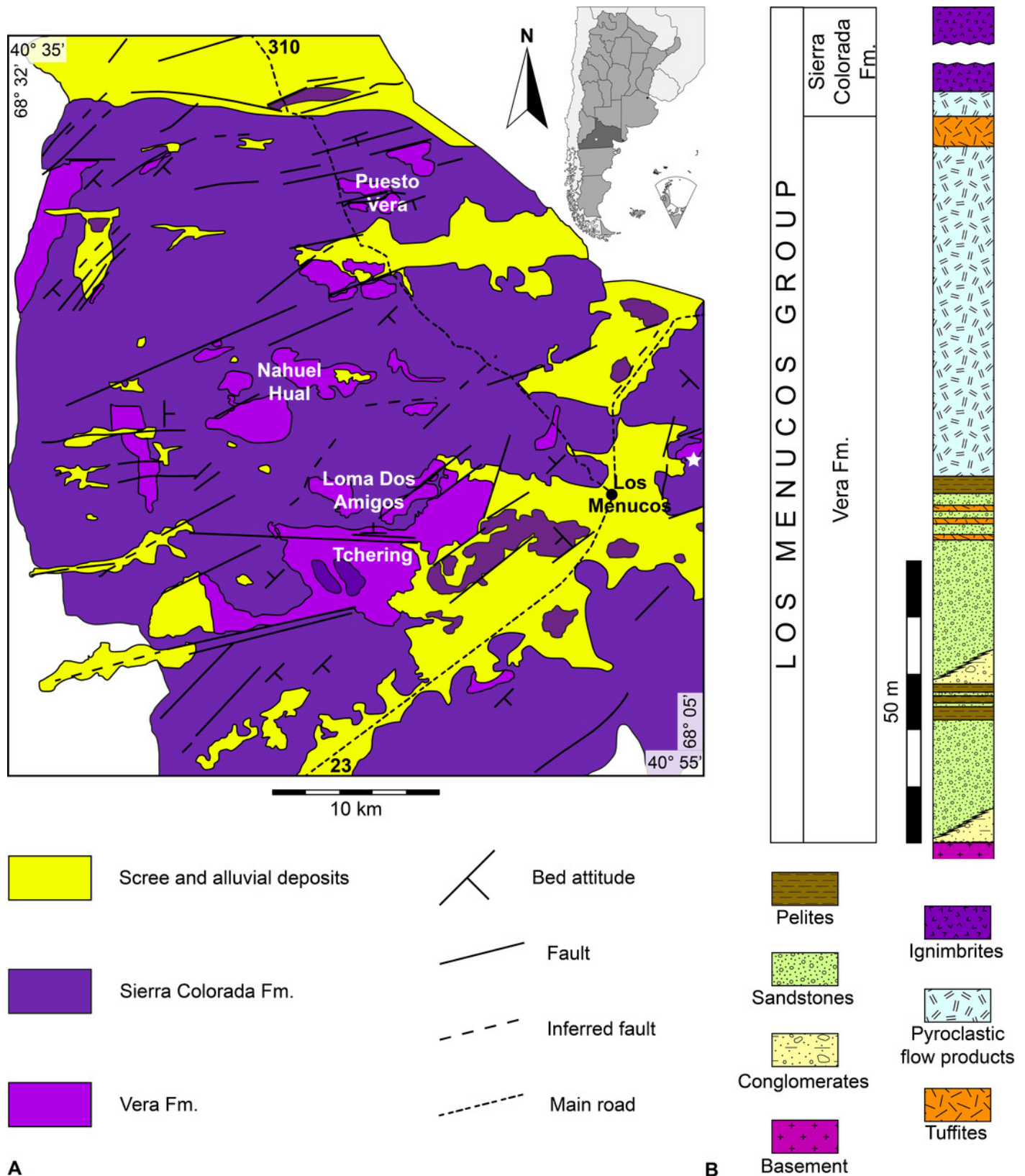


Figure 2

Thin sections (MPCA 27029/19.1 and MPCA 27029/19.2) of track-bearing slab MPCA 27029-19.

Inequigranular, epiclastic texture with anhedral and subhedral phenocrysts at the base (A) and middle portion (B) of the track-bearing slab MPCA 27029-19. (C, D) Equigranular less epiclastic texture indicating a minor sedimentary reworking of the trampled surface.

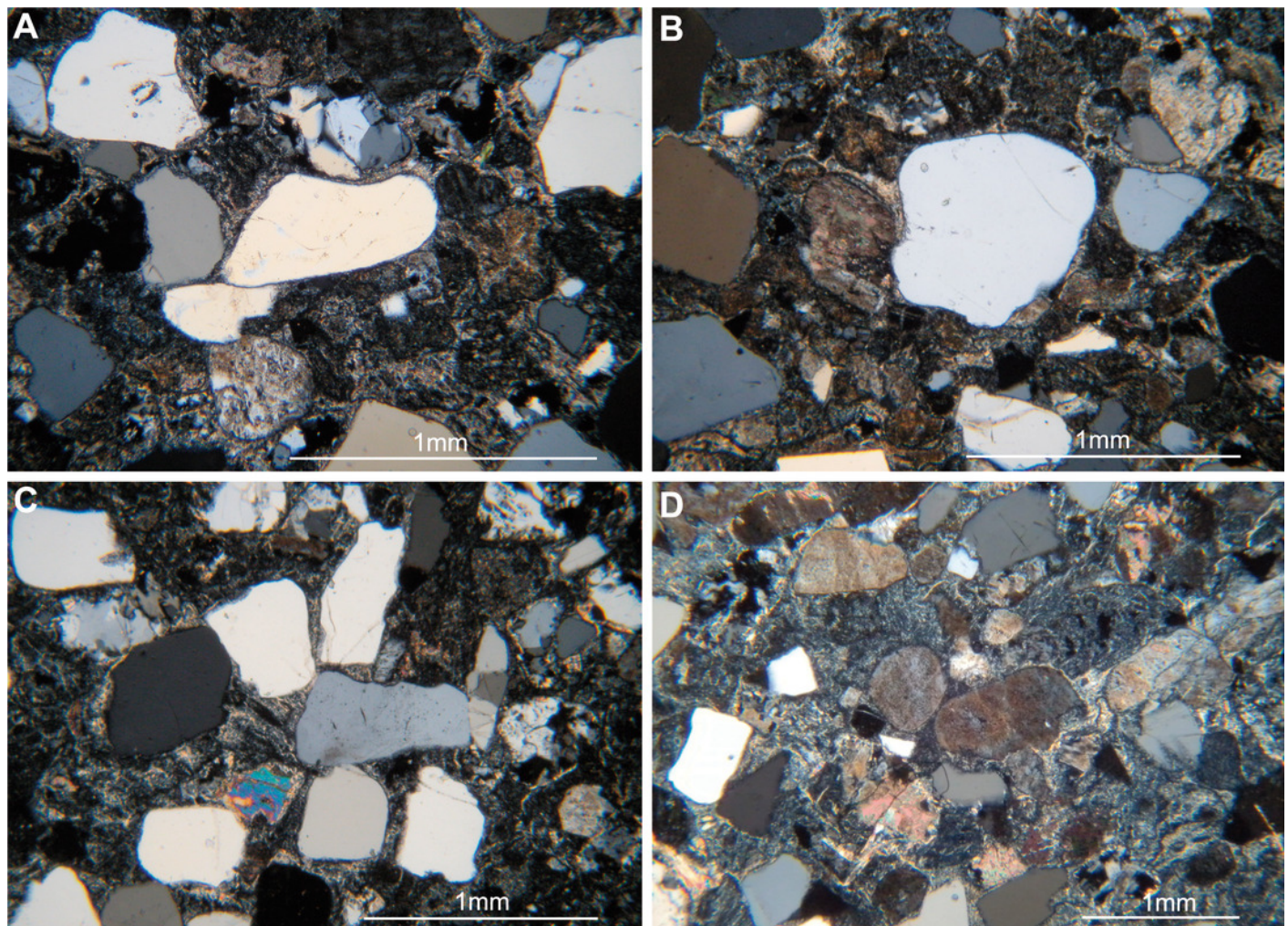


Figure 3

Tracks mode of preservation.

Convex hyporeliefs (A, C) fitting with concave epireliefs (B, D) preserved on slab MMLM 075-1 (true tracks and natural casts, respectively).

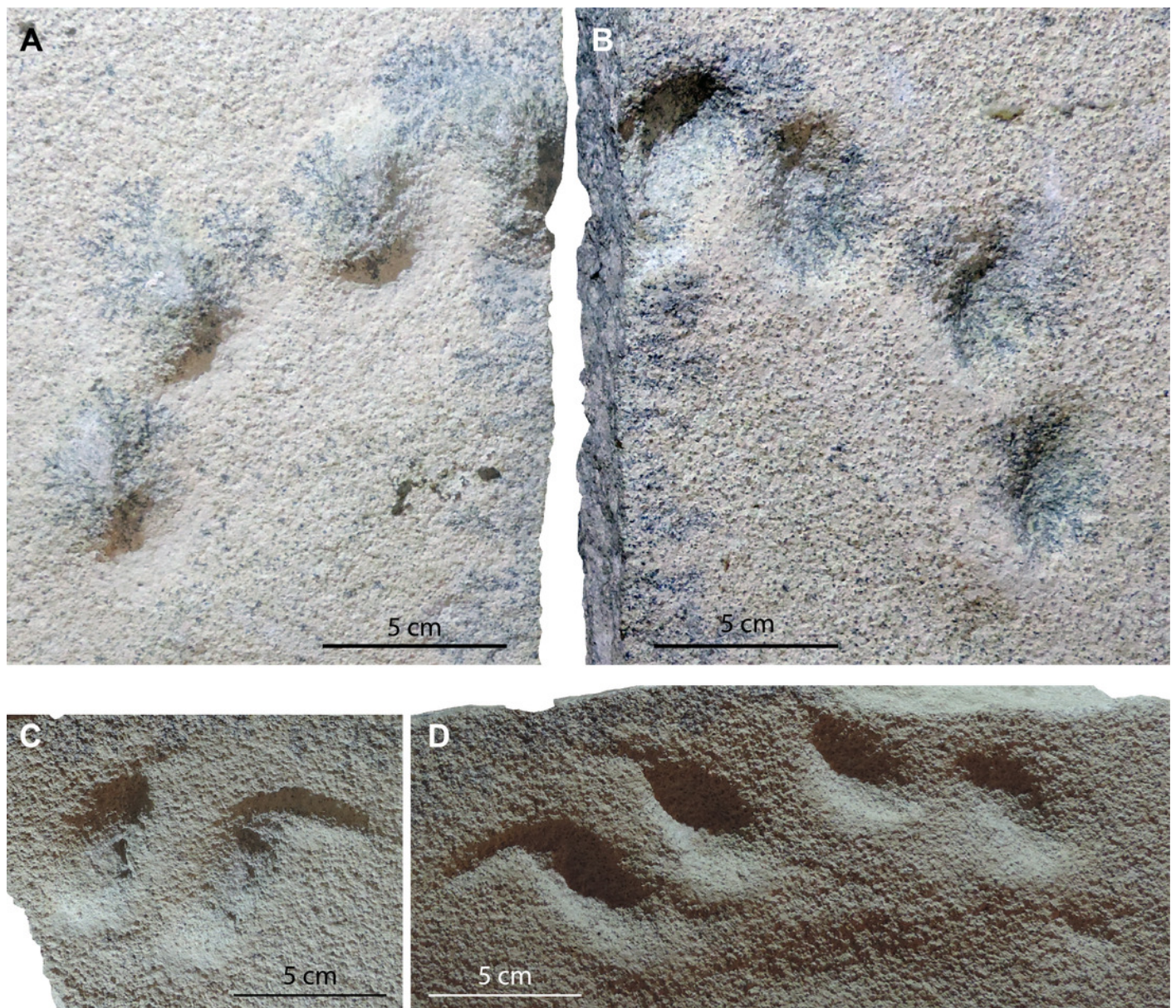


Figure 4

Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MPCA 27029-1; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs MPCA 27029-2; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E). In (A)-(D), footprint 2 and 5, note the non-impressed area between the sole pad trace and the base of digit traces. In (E)-(H) note the displacement areas behind digit traces, interpreted as the result of the thrust of digit pushing the sediment backwardly.

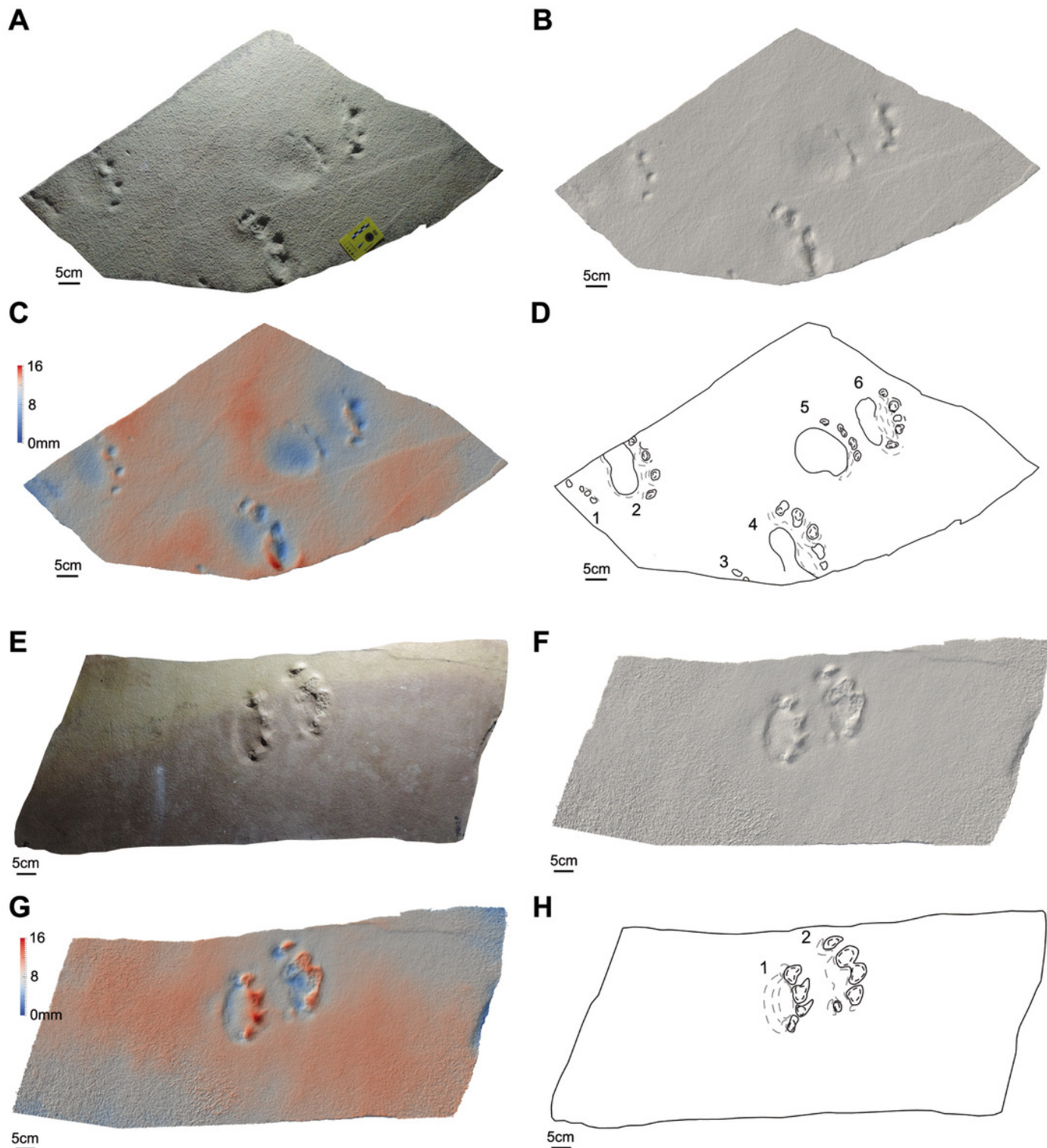


Figure 5

Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MPCA 27029-3; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs MPCA 27029-4; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E). Note the digit trailing marks slightly affecting the digit traces of footprint 2 in (A)-(D), which are absent in footprints showed in (E)-(H) where digit traces are roughly sub circular in morphology.

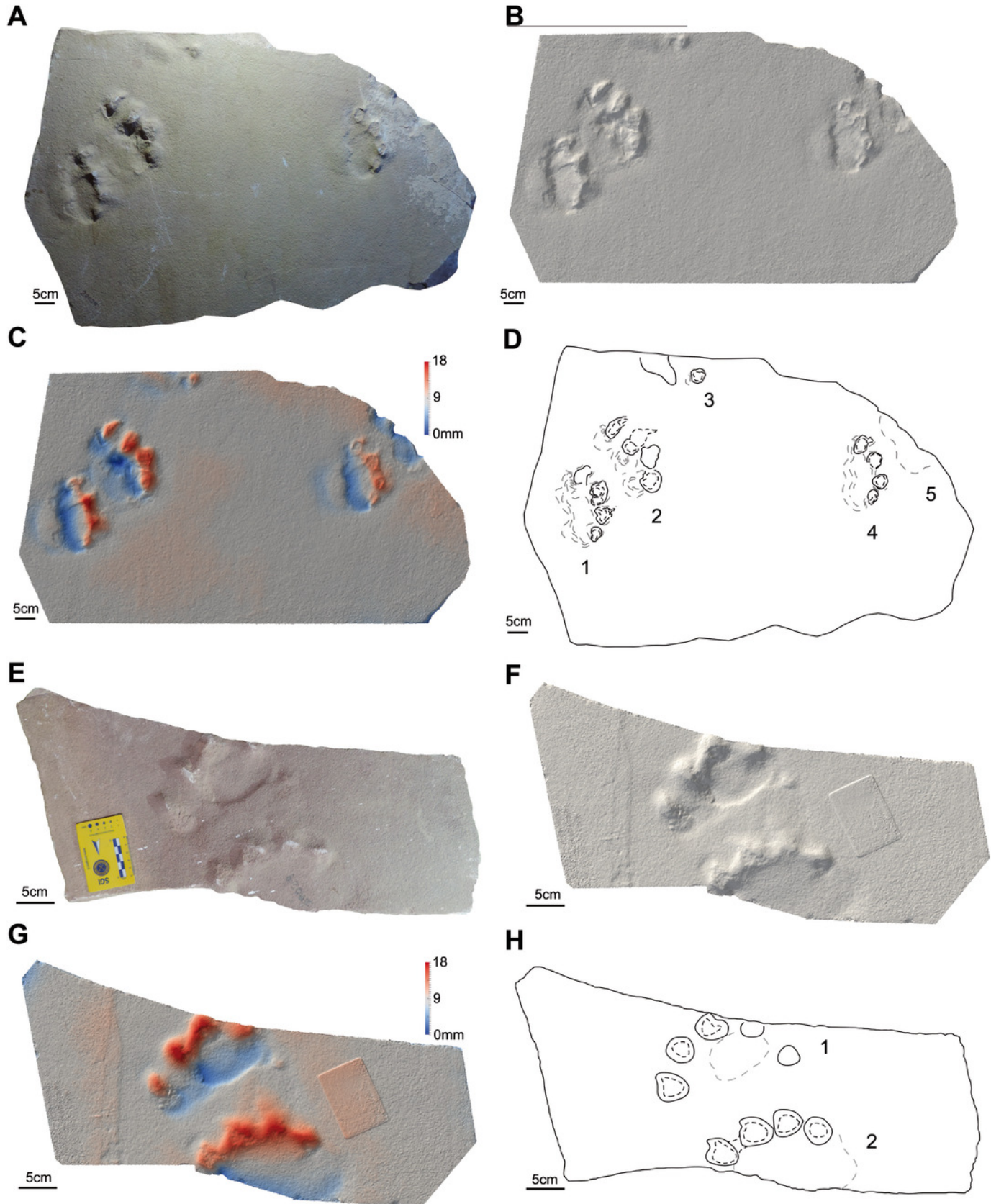


Figure 6

Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MPCA 27029-5; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs MPCA 27029-9; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E).

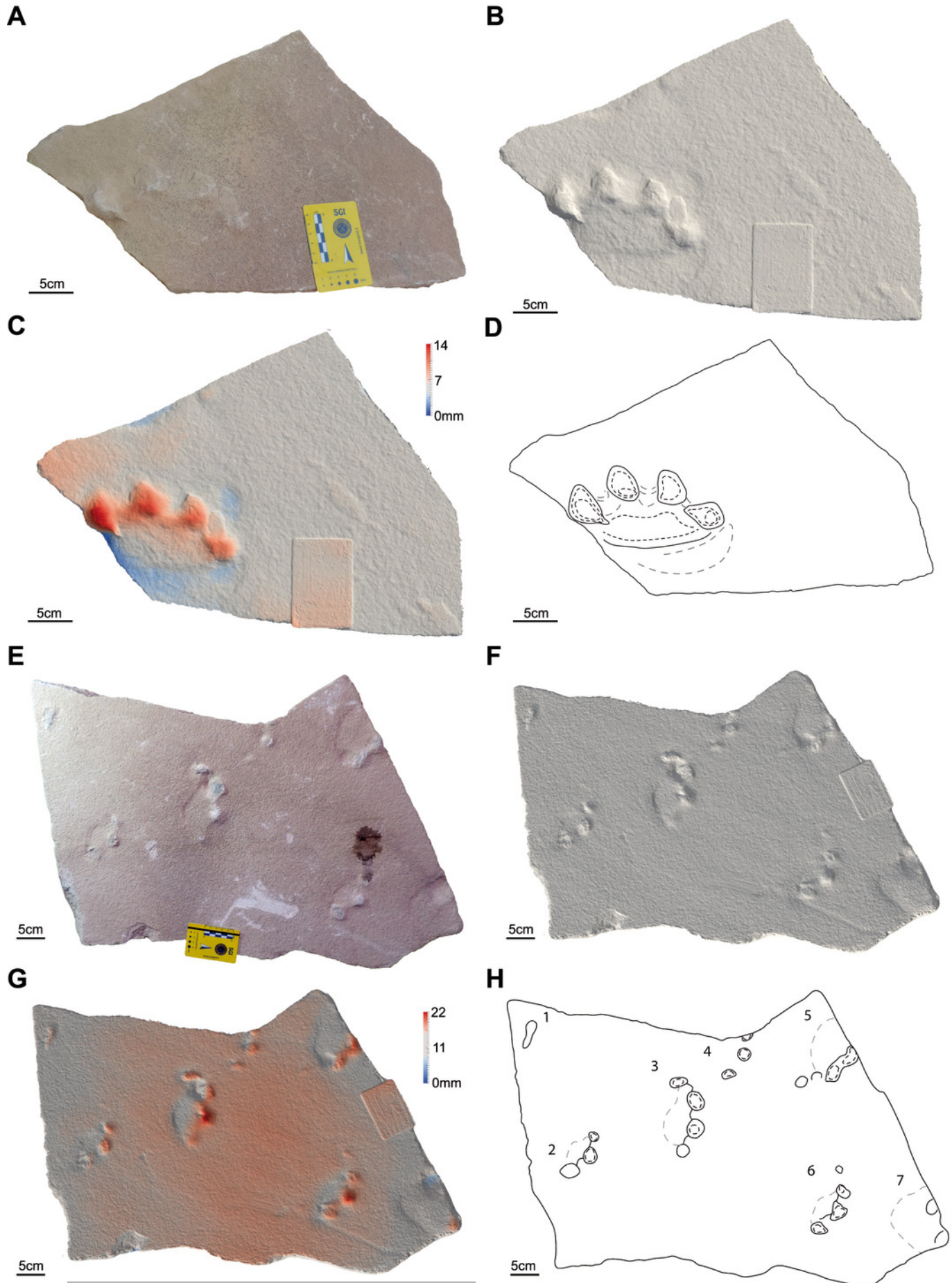


Figure 7

Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MPCA 27029-16, produced by a juvenile trackmaker; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs MPCA 27029-21; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E). The general morphology and structure of footprints 6-10 in (A)-(D), left by a juvenile trackmaker, is identical to that characterizing larger footprints, even when preserved only as digit traces.

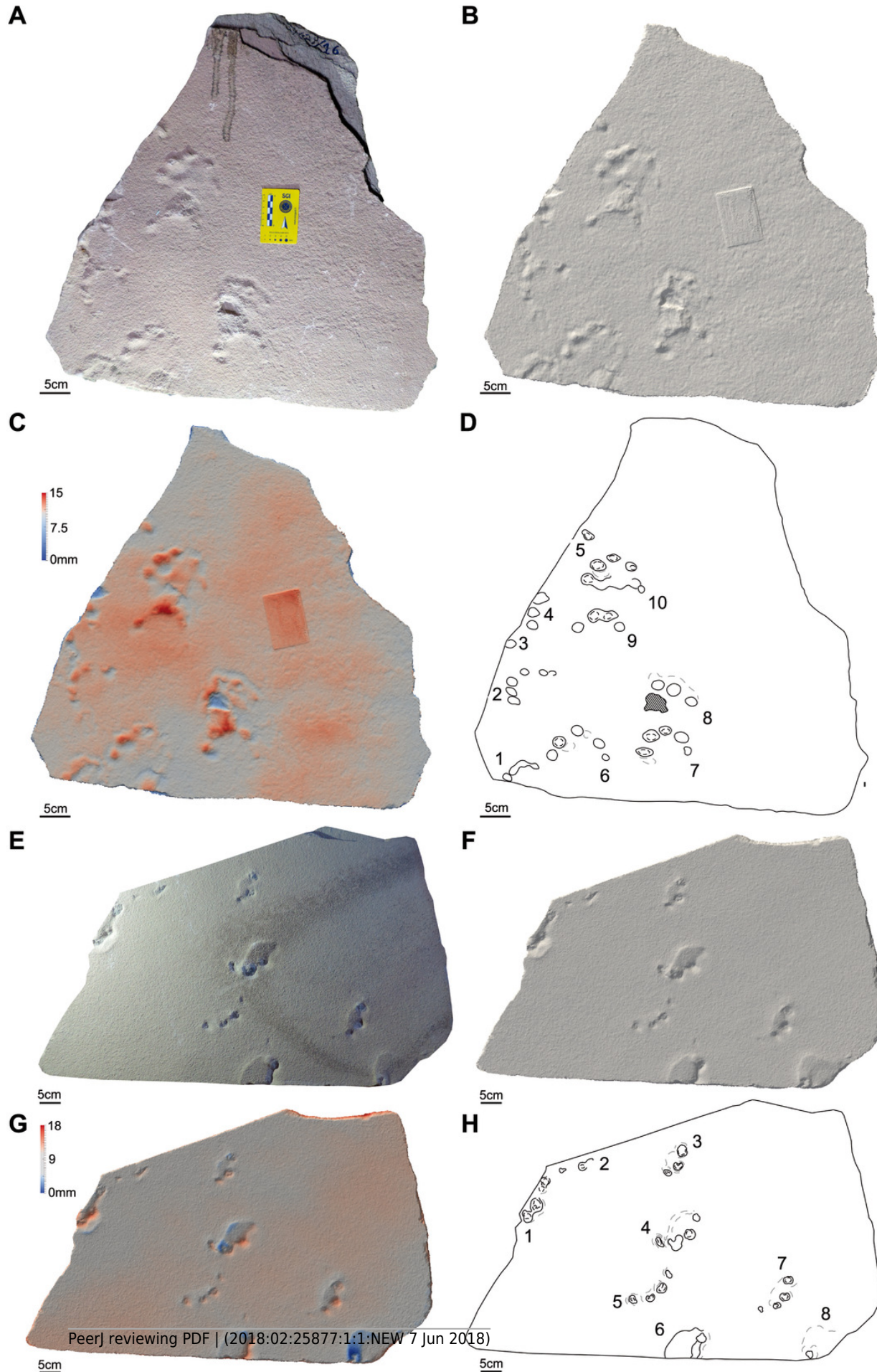


Figure 8

Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MPCA 27029-33; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs MMLM 075-1; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E). In (E)-(H) note the long and sharp digit trailing marks.

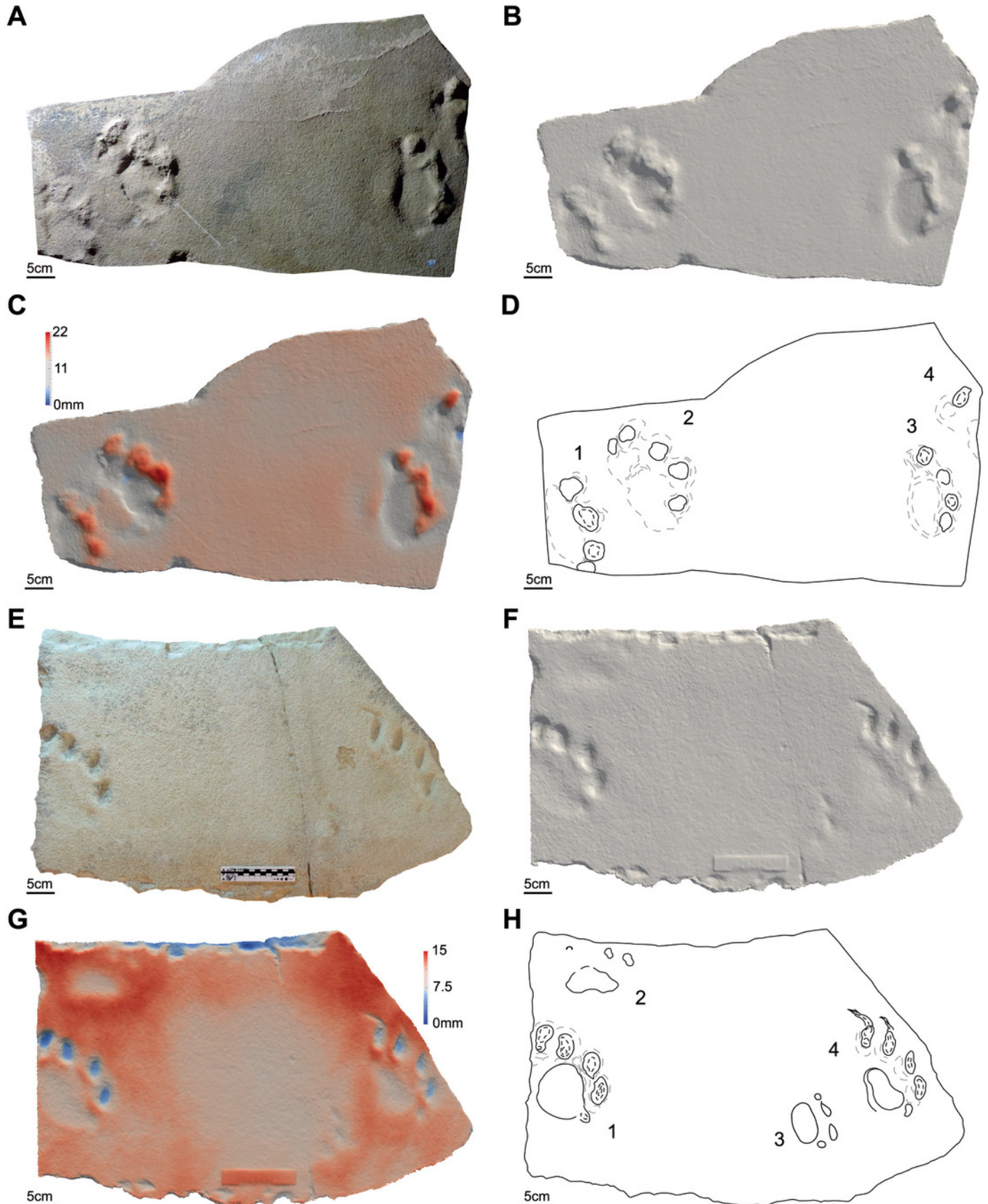


Figure 9

Photos, three-dimensional models and interpretative drawings of the studied material.

(A) Track-bearing slabs MMLM 1; (B) Solid three-dimensional model of (A); (C) Colour topographic profile and (D) interpretative drawing of (A). (E) Track-bearing slabs MMLM 2; (F) Solid three-dimensional model of (E); (G) Colour topographic profile and (H) interpretative drawing of (E). In (A)-(D) note the long and sharp digit trailing marks affecting footprints 2 and 4, resembling those of Figure 8E-H.

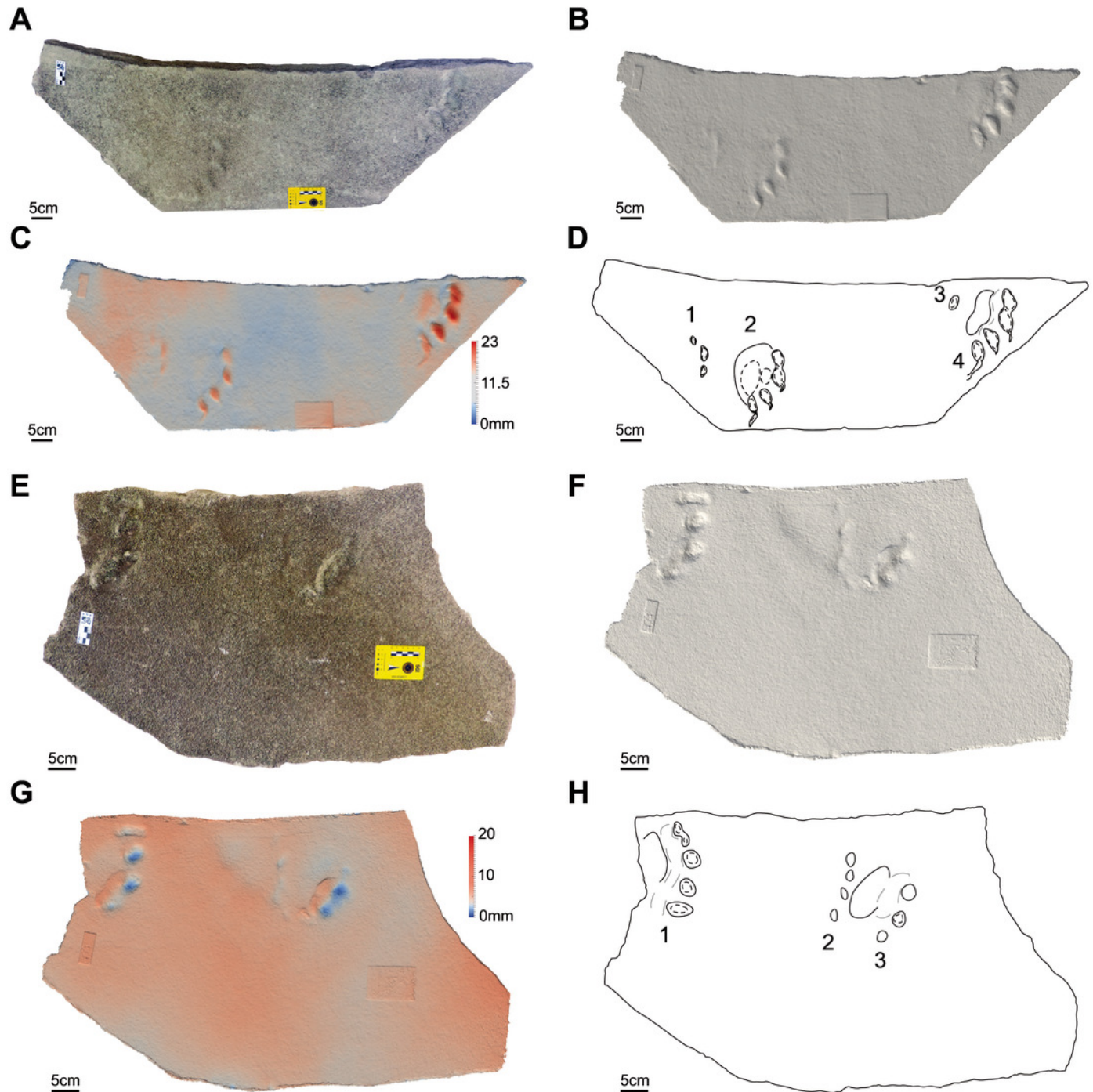


Figure 10

Morphological and extramorphological features identified on the studied material.

(A) Manus track MPCA 27029/2 and (B) interpretative drawing. (C) Pes track MPCA 27029-1/5 and (D) interpretative drawing. Extramorphological features are in blue and grey, morphological features are in black.

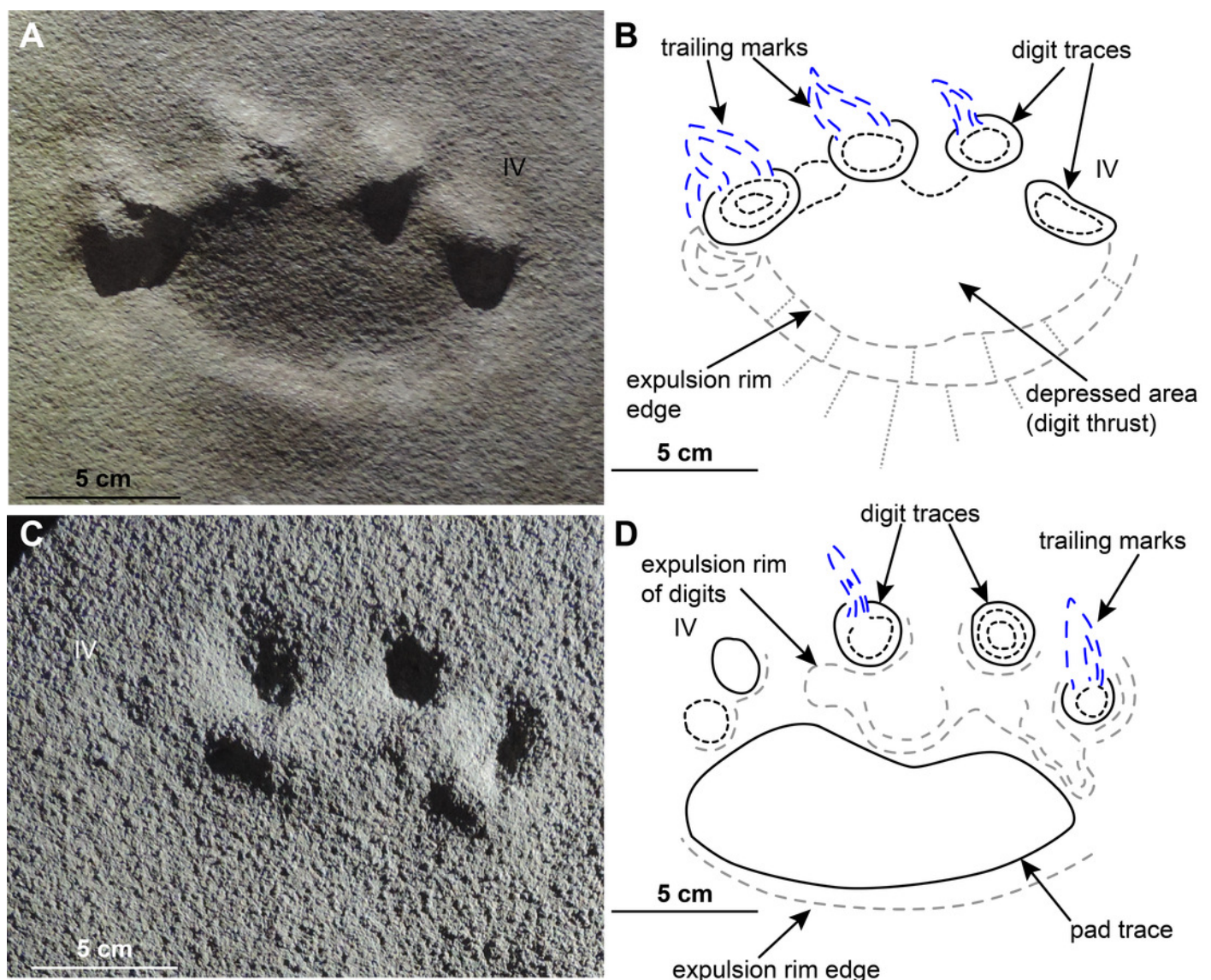


Figure 11

Limb and autopod posture in the *Pentasauropus* trackmaker.

Simplified reconstruction of limb posture in back (A) and lateral (B) views. Simplified reconstruction of zeugopodials and hind autopod in lateral (C) and bottom (D) views. In colour the possible extension of the fleshy cushion on which the basipodials rested, ensuring support during locomotion. See the supplementary video to get a more complete view of the reconstruction. Artwork by Fabio Manucci.

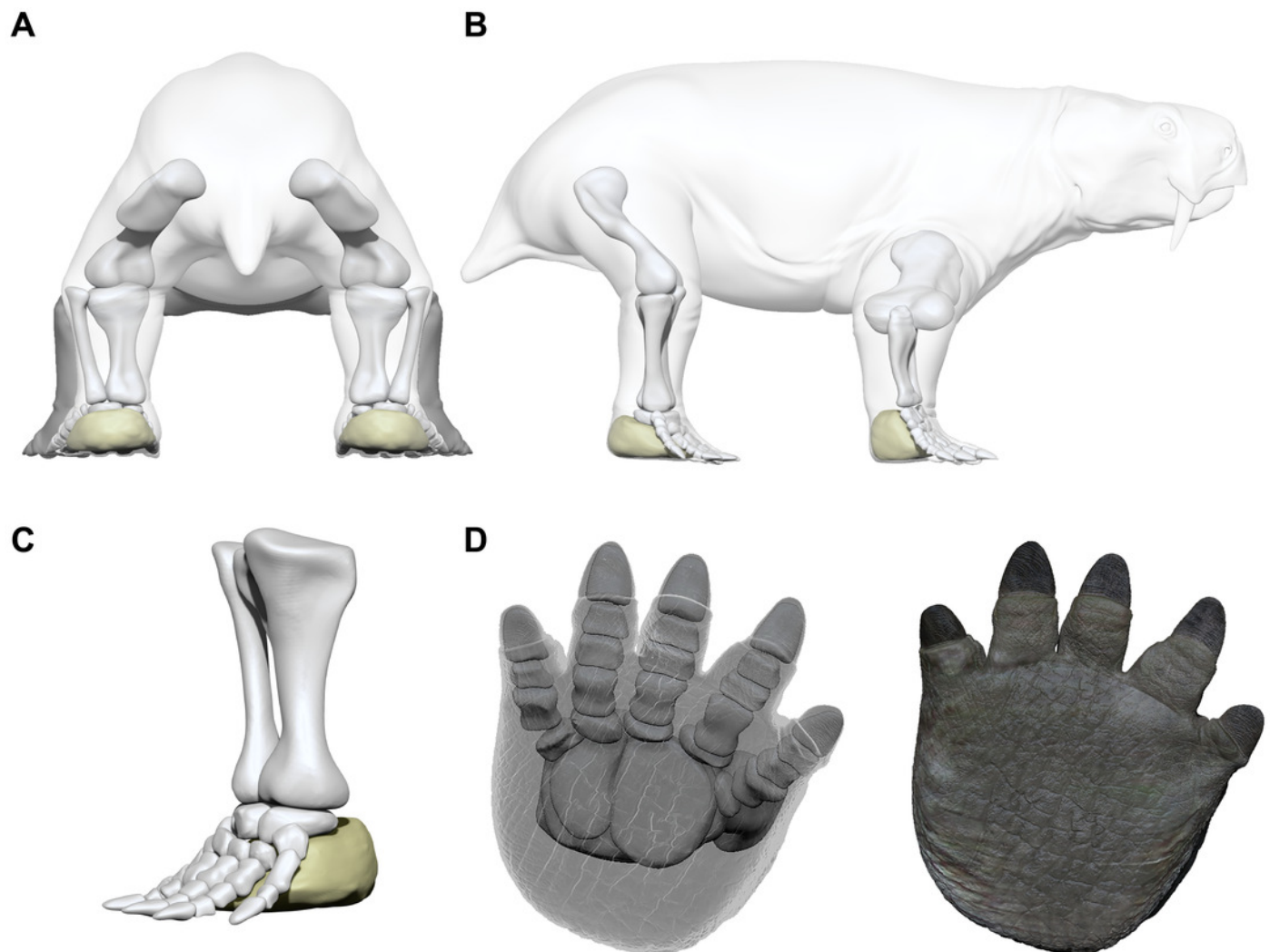


Figure 12

Speculative *in vivo* reconstruction (based on *Dinodontosaurus*) of a kannemeyeriiformes dicynodont, a most probable producer of *Pentasauropus* tracks.

Reconstruction in back (A) and lateral (B) view of the trackmaker walking in amble gait. See the supplementary video to get a more complete view of the reconstruction. Artwork by Fabio Manucci.

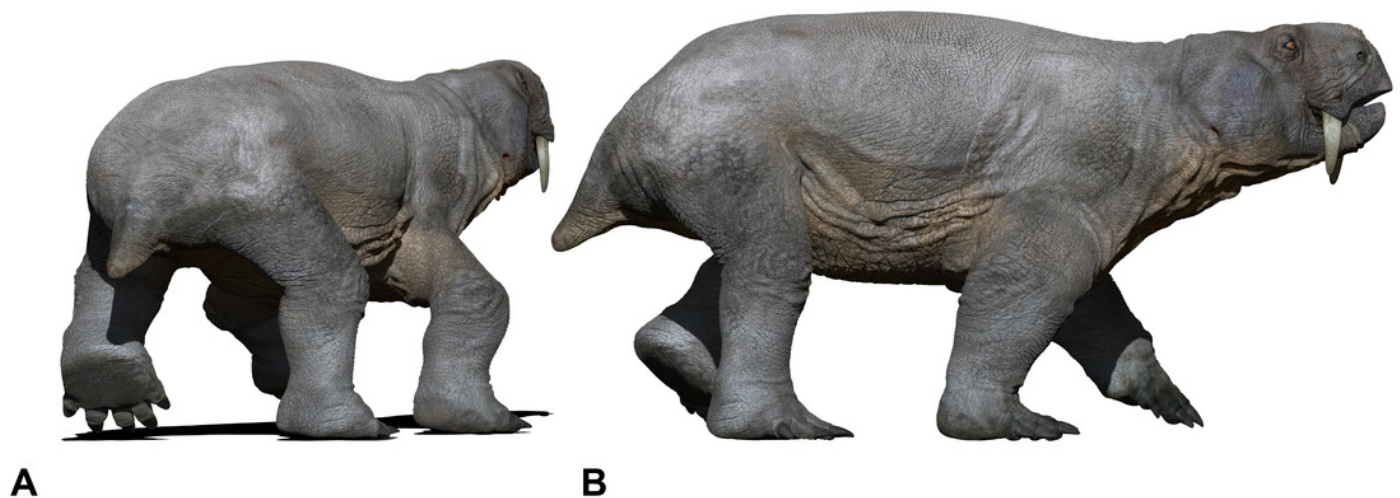


Table 1(on next page)

Photogrammetric report.

Main processing parameters of the photogrammetric models (from Agisoft Photoscan Professional reports).

3D Model	Number of images	Camera altitude (cm)	Ground resolution (mm/pix)	RMS reprojection error	Mean Key point size (pix)	Scale bars total error (m)
MPCA 27029-1	61	55	0.108	0.145637 (0.595447 pix)	3.99373	0.000211567
MPCA 27029-2	38	49.6	0.142	0.211741 (0.749639 pix)	3.83623	0.000120665
MPCA 27029-3	36	63.8	0.183	0.221627 (0.663238 pix)	3.1918	0.000101555
MPCA 27029-4	36	26.2	0.0753	0.225287 (0.695116 pix)	3.21872	0.000145928
MPCA 27029-5	25	30.2	0.0867	0.186287 (0.548475 pix)	3.41294	0.000125956
MPCA 27029-9	36	31.9	0.0916	0.206052 (0.638116 pix)	3.27202	0.000154204
MPCA 27029-16	30	52.3	0.13	0.254353 (0.874324 pix)	3.64063	0.000179922
MPCA 27029-21	59	47.3	0.127	0.273984 (0.529254 pix)	2.28824	0.000184177
MPCA 27029-33	74	50.2	0.144	0.25616 (0.993387 pix)	3.90857	0.000121254
MMLM 075-1	54	37.8	0.0898	0.196238 (0.739726 pix)	3.75732	0.000118491
MMLM 1	77	42.5	0.101	0.222075 (0.673395 pix)	3.09592	0.00332392
MMLM 2	52	33	0.0949	0.234591 (0.852769 pix)	3.78559	5.99994e-05

Table 2 (on next page)

Mean measurements (in cm) of track and trackway parameters.

ETW, external trackway width; Fl, footprint length; Fw, footprint width; GAD, gleno-acetabular distance: (a), 'primitive' alternate pace (the trunk length of the producer is underestimated); (b), alternate pace; (c), amble (a, b, c, considering primary overlap *sensu* Leonardi, 1987); ITW, internal trackway width; Mpa, manus pace angulation; Mpl, manus pace length; Msl, manus stride length; Ppa, pes pace angulation; Ppl, pes pace length; Psl, pes stride length; Psl/GAD, pes stride length/gleno-acetabular distance ratio; ETW/SL, external trackway width/stride length ratio.

Specimen	Fl	Fw	Mpl	Ppl	Mpa	Ppa	Msl	PsI	ET W	ITW	ETW/SL	GAD	PsI/GA D
MPCA 27029-1	10.6 12.8 8.9	10.3 12.2	42.2 33.3	41.7 40.5	101°	99°	59.5	62.4	50.0	3.65	0.82	46.2 (a) 47.8 (b) 91.6 (c)	1.01
MPCA 27029-2	/	15.7 13.5	/	/	/	/	/	/	/	/	/	/	/
MPCA 27029-3	/	12.7 14.7	/	/	/	/	/	63.5	/	/	/	/	/
MPCA 27029-4	/	12.6 13.3	/	/	/	/	/	/	/	/	/	/	/
MPCA 27029-5	6.6	11.6	/	/	/	/	/	/	/	/	/	/	/
MPCA 27029-9	/	7.4 10.9	28.5	37.0 32.5	/	81°	/	45.5	36	7.36	0.79	40.0 (a) 46.9 (b) 63.1 (c)	0.91
MPCA 27029-16	5.6	8.1 7.1	21.0 15.0	22.5	100°	/	28.5	/	23.0	4.40	0.81	26.0 (a) 33.1 (b) 37.4 (c)	/
MPCA 27029-21	/	10.7	/	37.0 34.0 34.5	/	70°	/	37.5 41	38.5	5.85	1.03	/	/
MPCA 27029-33	/	15.4 12.8	/	/	/	/	/	63.5	/	/	/	/	/
MMLM 075-1	10.2 11.8	13.6 9.8	/	/	/	/	/	57.3	/	/	/	/	/
MMLM 1	10.2 10.9	13.5 14	/	/	/	/	52.0	/	/	/	/	/	/
MMLM 2	/	12.2 11.4	/	/	/	/	40.5	/	/	/	/	/	/

Table 3(on next page)

Sole pad-ungual trace distance.

Distance (in cm) between the distal margin of the sole pad trace and the proximal margin of the digit traces in complete *Pentasauropus* footprints. The measurements most likely indicate a raised and inclined position of the metapodial elements of fore and hind foot in the *Pentasauropus* trackmaker.

MPCA 27029-1					
	Digit I	Digit II	Digit III	Digit IV	Digit V
Footprint 2	1.35	1.41	1.84	1.22	/
Footprint 4	1.68	1.66	1.84	1.91	/
Footprint 5	/	1.26	1.45	1.52	1.20
Footprint 6	1.2	1.25	2.18	2.24	1.33
MPCA 27029-5					
	Digit I	Digit II	Digit III	Digit IV	Digit V
Footprint 1	/	0	0.96	1.13	0.4
MPCA 27029-16					
	Digit I	Digit II	Digit III	Digit IV	Digit V
Footprint 10	0.92	0.9	1.1	0.94	0.79
MMLM 075-1					
	Digit I	Digit II	Digit III	Digit IV	Digit V
Footprint 1	1.65	2.09	2.94	2.56	1.45
Footprint 3	/	1.41	1.68	2.19	1.63
Footprint 4	1.95	2.25	2.45	2.25	1.58
MMLM 1					
	Digit I	Digit II	Digit III	Digit IV	Digit V
Footprint 2	1.88	2.53	2.73	2.25	/
Footprint 4	1.59	1.86	2.25	2.12	/
MMLM 2					
	Digit I	Digit II	Digit III	Digit IV	Digit V
Footprint 1	1.66	2.37	2.02	2.27	/