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- 1 An Examination of the Impact of Olson's Extinction on Tetrapods from Texas
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Abstract

It has been suggested that the stransition between a pelycosaurian-grade synapsid dominated fauna of the Cisuralian (early Permian) and the therapsid dominated fauna of the Guadalupian (middle Permian) snegested that this jura of was accompanied by, and possibly driven by, a mass extinction dubbed Olson's Extinction. However, this interpretation of the record has recently been criticised 10 as being a result of inappropriate time-binning strategies: calculating species richness within 11 international stages or substages combines extinctions occurring throughout the late Kungurian 12 stage into a single event. To address this criticism, I examine the best record available for the 13 time of the extinction, the tetrapod-bearing formations of Texas, at a finer stratigraphic scale than 14 those previously employed. Species richness is calculated using four different time-binning 15 schemes: the traditional Land Vertebrate Faunachrons (LVFs); a re-definition of the LVFs using 16 constrained cluster analysis; individual formations treated as time bins; and a stochastic approach 17 assigning specimens to half-million-year bins. Diversity is calculated at the genus and species 18 level, both with and without subsampling, and extinction rates are also inferred. Under all time-19 binning schemes, both at the genus and species level, a substantial drop in diversity occurs during 20 the Redtankian LVF. Extinction rates are raised above background rates throughout this time, but 21 the biggest peak occurs in the Choza Formation (uppermost Redtankian), coinciding with the 22 disappearance from the fossil record of several of amphibian clades. This study, carried out at a 23 finer stratigraphic scale than previous examinations, indicates that Olson's Extinction is not an 24 artefact of the method used to bin data by time in previous analyses. 25

#### Introduction

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Key Words: Olson's Extinction; Tetrapods; Texas; Redtankian; Time Bins

28	A faunal turnover of tetrapods has long been recognised between the Cisuralian and
29	Guadalupian (early and middle Permian, respectively). The former is characterised by a diverse
30	array of amphibians, pelycosaurian-grade synapsids, particularly carnivorous sphenacodontids
31	and herbivorous edaphosaurids, and captorhinids, while the latter is dominated by therapsid
32	synapsids, with diversity of parareptiles increased and amphibian diversity substantially reduced
33	(Olson 1962, 1966; Kemp 2005; Sahney & Benton 2008; Ruta et al. 2011; Benton 2012; Benson
34	et al. 2013; Brocklehurst et al. 2013; 2017). These faunal changes were accompanied by
35	ecological shifts, including a transition towards more complex ecosystems with more trophic
36	levels (Olson 1966). However, the nature and progress of the transition is still strongly debated.
37	The possibility of a mass extinction accompanying this transition was first suggested by
38	Olson (1982), who noted a drop in the number of families across Cisuralian/Guadalupian
39	boundary. This drop was principally concentrated among amphibian families (amniote diversity
40	was shown to increase slightly). Sahney & Benton (2008) provided a more detailed examination
41	of diversity through the Permian, still at the family level but with temporal resolution at the stage
42	level. Decreases in both species richness, diversification rate and ecological diversity were
43	apparent through the Kungurian and Roadian (the last stage of the Cisuralian and the first of the
44	Guadalupian respectively). Sahney & Benton (2008) dubbed this mass extinction event "Olson's
45	extinction" and hypothesised that it might have been a causal factor in the faunal turnover
46	occurring at this time.
47	The hypothesis of Sahney & Benton has been criticised as being based on family-level
48	data that was not corrected for sampling heterogeneity (Benson & Upchurch 2013, Brocklehurst
49	et al. 2013). Nevertheless, subsequent studies both of tetrapods as a whole (Benton 2012, Benton
50	et al. 2013, Benson & Upchurch 2013, Brocklehurst et al. 2017) and subgroups within Tetrapoda
51	(Ruta & Benton 2008, Ruta et al. 2011, Brocklehurst et al. 2013, 2015), carried out at the species

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53	across the Kungurian/Roadian boundary.
54	Despite this, the theory of Olson's Extinction has been criticised in other ways. Benson &
55	Upchurch (2013) suggested that the mass extinction was an artefact of the geographically patchy
56	fossil record. The record from the Cisuralian is known almost entirely from palaeoequatorial
57	localities, particularly from North America and Europe, while that of the Guadalupian is
58	dominated by palaeotemperate localities from Russia and South Africa (Lucas 2004, Kemp
59	2006). Not only does this make it difficult to ascertain over what timescale the extinction took
60	place and to what extent the transition was a global event, but the apparent diversity drop might
61	simply represent a latitudinal diversity gradient (Benson & Upchurch 2013). In most modern
62	clades, diversity is higher in equatorial regions than temperate regions (Willig et al. 2003,
63	Hildebrand 2004), and so it was argued that the shift in sampling locality from more diverse to
64	less diverse latitudes might be the cause of the apparent decrease in species richness (Benson &
65	Upchurch 2013). Brocklehurst et al. (2017), however, argued against this point of view. It has
66	been noted that the latitudinal diversity gradient was not a constant feature through geological
67	time (Archibald et al. 2010, Rose et al. 2011, Yasuhara et al. 2012, Mannion et al. 2012, 2014),
68	and it was demonstrated that, in the few Permian time bins where tetrapod data was available
69	from both palaeoequatorial and palaeotemperate latitudes, the temperate latitudes exhibited
70	higher species richness after correcting for sampling (Brocklehurst et al. 2017).
71	Further criticism of Olson's extinction was put forward by Lucas (2017). Lucas argued
72	that the inference of a mass extinction across the Kungurian/Roadian at this time was an artefact
73	of two confounding factors. First, the majority of the studies cited used geological stage or
74	substages as their time bins, thus conflating the extinctions occurring throughout the Kungurian
75	into a single event. Second, they argued that incorrect ages were applied to numerous geological
76	formations, in particular the San Angelo and Chickasha formations of Texas and Oklahoma,

level and employing a variety of sampling correction methods, have identified diversity decreases

77	respectively. The ages of these formations have long been a point of contention. Early estimates
78	placed them in the latest Leonardian (late Kungurian) (Lucas & Heckert 2001, Lucas 2004), but
79	discovery of a specimen from Chickasha of the parareptile Macroleter, previously only known
80	from the Middle Permian of Russia, caused Reisz & Laurin (2001) to argue for an equivalency
81	between this formation and the Kazanian-aged (earliest Guadalupian) faunas of Russia. Lucas
82	(2002) rejected their arguments based on the ammonite fauna of the Blaine formation, a marine
83	formation immediately overlying the San Angelo, which he claimed supported a Leonardian age.
84	Rjesz & Laurin (2002) criticised the interpretation of Lucas, suggesting that a key taxon in the
85	arguments had a much longer range than suggested and highlighting previous studies of the
86	Blaine formation interpreting it as Guadalupian in age. Lozovsky (2003) also used ammonite
87	biostratigraphy to support a Roadian age for the Chickasha and San Angelo formations, and these
88	ages have been adopted in most subsequent studies (e.g. Sahney & Benton 2008, Benton 2012,
89	Brocklehurst et al. 2013, 2017). However, Lucas (2017) still supports a latest Kungurian age for
90	these two formations. He therefore suggested that an extinction across the Kungurian/Roadian
91	boundary cannot be assessed in a global framework, as there is no stratigraphic overlap between
92	the North American and Russian formations.
93	It is not the purpose of this paper to argue against these two criticisms of Lucas (2017).
94	Indeed, I am fully prepared to agree that time-binning strategies employing the geological stages
95	or substages, while often necessary for global analyses where the correlations between the
96	regional biostratigraphic schemes are inexact, have the potential to produce spurious results. Such
97	binning strategies produce time-averaged diversity estimates for a time bin that can differ from
98	the true standing diversity at any one time in the bin (Raup 1972, Lucas 1994, Foote 1994, Miller
99	& Foote 1996, Alroy 2010a, Gibert & Escarguel 2017). Instead it is my intention to approach the
100	question of Olson's extinction from a different angle, one that addresses the issues of binning
101	strategy while bypassing the disagreements surrounding the ages of the San Angelo and

Chickasha formations. In fact, the framework of this analysis is one suggested by Lucas himself (Lucas 2017): when the fossil record is geographically patchy with uncertain global correlations, it is better to study mass extinctions using the "best sections" method, focusing one or a few well sampled, stratigraphically dense fossiliferous sections to examine the progress of the extinction. While only providing a local perspective on the event under study, this method does allow more detailed analysis than is provided in global studies with coarse temporal resolution.

The "best section" of tetrapods in the Cisuralian is doubtless that of Texas, which represents a reasonably continuous sequence from the late Carboniferous until the end of the Cisuralian (Romer 1928, 1935, Hook 1989, Lucas 2006, 2017). A detailed examination of the Cisuralian tetrapod record from Texas, covering the stratigraphic sequence from the Pueblo Formation until the San Angelo Formation, allows much higher resolution than previous studies. Moreover, it renders the debate regarding the age of the San Angelo formation moot. The issue is no longer whether there is a Kungurian/Roadian boundary event, but instead whether an extinction event is identified between the Redtankian and Littlecrotonian land vertebrate faunachrons (biostratigraphic time bins based on the tetrapod fossil record, the former correlating in Texas with the Clear Fork froup, the latter with the San Angelo formation). The presence of an extinction event between these two faunachrons is assessed at both genus and species levels, with four different time-binning systems and results shown both with and without sampling correction.

#### Materials and Methods

121 Data

Data on the number of specimens of tetrapod species in each time bin was assembled from a variety of sources. The primary literature and the paleobiology database, downloaded from the fossilworks website (http://fossilworks.org) on October 2017, were the principal sources, but were supplemented by observation of specimens in museum collections and also by

data sent from some museums (Museum of Comparative Zoology, Harvard; Field Museum of Natural History, Chicago; American Museum of Natural History, New York; Yale Peabody Museum, New Haven; University of California Museum of Palaeontology, Berkeley; Sam Noble Oklahoma Museum of Natural History, Norman). The data was examined at both species and genus level. While it has often been the preference to examine data at the species-level (Sepkoski [1984] argued that as the species are the real "units" of evolution, it is at that level that evolution should be studied), Lucas (2017) suggested that the genus is preferable for early Permian tetrapods to avoid the influence of large numbers of singletons (single-specimen taxa), which under poor sampling produce a great deal of "noise" in the evolutionary signal (Alroy 1998, Foote 2000). The final datasets are provided in Data S1 and S2 

#### Time bins

Four methods were used to define time bins, each successively dividing the early Permian into smaller portions of time. The first set of bins used are the land vertebrate faunachrons (LVFs): the biostratigraphic bins based on the first and last appearances of key tetrapod genera (Lucas 1998). As these are biostratigraphic bins, their boundaries should correspond to major periods of turnover among tetrapods, and so the diversity estimates within each faunachron should provide a better approximation of the standing diversity at any point in time than using the international stages.

The second binning scheme used represents a redefinition of the land vertebrate faunochrons using a clustering approach. CONISS is a constrained clustering analysis, which groups stratigraphic sections into hierarchical clusters based on the taxonomic distances between, while maintaining the order of the stratigraphic sequence (Grimm 1987). The taxonomic distances between the formations were calculated using Alroy (2015a)'s modification of the Forbes metric, applying the RAC correction suggested by Brocklehurst et al. (2018) to account

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150	for differences in the evenness of the relative abundance distributions, which under incomplete
151	sampling can bias the distances observed. The CONISS analysis was carried out in R version
152	3.3.2 (R core team 2016), using functions from the package rioja (Juggins 2009). The boundaries
153	of the original LVFs were then shifted to ensure that formations which were clustered together
154	were grouped in the same bin.
155	The third binning scheme simply treats each formation as a time bin. The lithostratigraphy
156	was devised by Plummer & Moore (1921) and dated based on the marine strata which intercalate
157	with the terrestrial strata. This provides a finer resolution than the land vertebrate faunachrons (11
158	bins rather than 5).
159	The fourth and final binning scheme uses a stochastic approach. The ages of the top and
160	bottom of each formation were used as maximum and minimum bounds on the ages of each
161	specimen known from within that formation. The period of time under study was split into half-
162	million-year time bins, and each specimen was assigned at random to one of the bins between its
163	maximum and minimum age brackets. 100 such datasets were generated, and the analyses of
164	diversity and extinction rate were applied to all 100. Such stochastic methods have been shown to
165	provide more accurate estimates of standing diversity than binning approaches, even when the
166	origination and extinction are biased towards coinciding with the boundaries of bins (Gibert &
167	Escarguel 2017)
168	For all four binning schemes, the absolute ages were derived from Lucas (2017). Thus,
169	the Littlecrotonian LVF and the San Angelo formation are deemed to be latest Kungurian rather
170	than Roadian. For most of the binning schemes, this does not make a difference as mentioned in
171	the introduction) when analysing only the Texas "best section", than the question of whether an
172	extinction event is identified between the Redtankian and Littlecrotonian is more relevant than

the precise timing of the boundary. Where the absolute ages do make a difference is in calculating

extinction rates using the stochastic binning scheme. By compressing the Redtankian and

Littlecrotonian into a smaller period of time, the density of the specimens sampled is increased.

This will lower extinction rates estimated under the gap-fillers method (see below): counts of
two-timers will increase and counts of part-timers and gap-fillers will decrease (Alroy 2014). The
use of a Kungurian age for the Littlecrotonian is therefore more conservative, biasing against the
inference of a mass extinction.

### Diversity and Rate Estimates

For each time bin in each binning scheme, diversity (species richness) estimates were calculated using two methods. The first is a taxic diversity estimate, a simple count of the number of species observed in each time bin without sampling correction. The second employs shareholder quorum subsampling (SQS; Alroy 2010), which standardises the coverage (the proportion of the rank abundance distribution sampled) in each time bin. Coverage is measured using Good's U (the proportion of singletons relative to the total sample size). Diversity was estimates at four levels of coverage: 0.6-0.9 at intervals of 0.1 (a quorum of 0.6 allowed diversity to be calculated in all time bins in all binning schemes). SQS diversity estimates were calculated in R using version 3.3 of the function available on the website of John Alroy (http://bio.mq.edu.au/~jalroy/SQS.html).

The stochastic binning method allows the implementation of the more precise and accurate methods of calculating extinction rates using the gap fillers method (Alroy 2014). Since this method is based on estimating sampling from the patterns of occurrences in a moving "window" covering four time bins, it is impractical to apply it to the short time series produced by the three other binning strategies. The gap-fillers method was implemented, applying the "two for one" correction (Alroy 2015b) to increase precision, using custom functions written in R. As suggested by Alroy (2014), sampling heterogeneity was accounted for by classical rarefaction

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198	(standardising the sample size by number of occurrences) rather than by standardising coverage.
199	10000 subsampling iterations were carried out, each drawing five occurrences per time bin.
200	Results
201	Redefined Land Vertebrate Faunachrons
202	When clustering the formations using CONISS, a number of changes are made to the
203	boundaries of the LVFs (Fig. 1). The Littlecrotonian and Redtankian remain as they were defined
204	previously. The lower boundary of the Mitchellcreekian is shifted downwards to include the Belle
205	Plains formation, found to cluster more closely with the Clyde than the Admiral formation. The
206	Admiral formation itself clusters with the Putnam formation, and so the Seymourian LVF is
207	redefined to include these two. Thus, the Coyotean LVF contains only the Pueblo and Moran
208	formations.
209	Diversity estimates
210	Raw, uncorrected species and genus-level diversity estimates indicate a substantial fall in
211	diversity between the Redtankian and Littlecrotonian, based on all four time binning schemes
212	(Fig. 2). The finer resolution time bins (formation-level and half-million-year time bins) indicate
213	that the Arroyo formation represents the peak richness, and number of genera and species
214	declined throughout the Redtankian.
215	When the data are binned by the Land Vertebrate Faunachrons (whether original or
216	redefined), subsampling by SQS supports the Littlecrotonian as the time of lowest diversity (Figs.
717	3.4) The status of the Redtankian as a diversity peak is less clear: when the original LVFs are

used, the Mitchellcreekian is found to contain a similar richness to the Redtankian (Fig. 3).

However, the redefined LVFs indicate a substantial increase between these two bins (Fig. 4).

The higher-resolution-binning schemes both indicate the drop in subsampled diversity occurs throughout the Redtankain (Figs. 5-6). The Arroyo formation produced the highest species and genus richness of this faunachron, and the diversity deceases in the Vale formation and reaches a trough in the Choza formation. When subsampling is applied, species and genus richness is found to increase slightly between the Choza and San Angelo formations.

#### Extinction rates

Three peaks in extinction rate are identified at the both at the genus and species level: at the top of the Belle Plains, Arroyo and Choza formations (the latter being the largest) (Fig. 7).

During the time covered by the Vale formation, extinction rates fall, but remain above background levels. The principal difference between the species and genus curves is the relative height of the Belle Plains extinction peak; at the species level it is higher than the Arroyo peak.

### Discussion

Having argued that an extinction of tetrapods across the Kungurian/Roadian boundary (due to the inappropriate time-binning strategies used in other diversity studies and the disagreement over the age of the San Angelo and Chickasha formations), Lucas (2017) briefly examined the possibility of a mass extinction between the Redtankian and Littlecrotonian LVFs in the "best section" of Texas. Although he noted a peak in extinction rates during the Redtankian and a decrease in genus richness during the Littlecrotonian, he was dubious over the reality of a mass extinction. First, he suggested that families previously suggested to be major components of the extinction, Edaphosauridae and Ophiacodontidae (Brocklehurst et al. 2013), had already disappeared prior to the end of the Redtankian. Lucas also examined diversity changes through the Redtankian using the specimen lists compiled by Olson (1958, 1989) for the Arroyo, Vale and



242 Choza formations, demonstrating that diversity was decreasing throughout the Redtankian, rather 243 than there being a single decline at the end of the LVF.

All diversity estimates presented here support a decrease in species and genus richness between the Redtankian and Littlecrotonian. The diversity estimates at finer stratigraphic scales support the observations of Lucas (2017): the decline occurs throughout the Redtankian from a peak in the Arroyo formation to a trough in the Choza formation, followed by a slight, but not substantial, recovery in the San Angelo formation. The same inferences may be made from extinction rates. The rates are noticeably higher in the Arroyo formation than the background rates experienced for most of the early Permian. Only once prior to this are extinction rates reliably inferred to reach similar levels: at the end of the Belle Plains formation. The extinction rates experienced in the Choza Formation are considerably higher than any other time in the early Permian.

Does this period of elevated extinction rates and declining diversity constitute a mass extinction? Lucas (2017) argued not, since it was a prolonged decline throughout the Redtankian LVF. Unfortunately, there is no set definition of a "mass extinction", and while the general consensus does seem to be elevated extinction over a short period of time, there is no indication of how short a time that should be. Discussion of mass extinctions in the scientific literature have included events where extinction rates were substantially higher than background rates over periods of millions of years. For example, discussion of the late Devonian mass extinction (one of the "big five" mass extinctions) has in the past suggested a duration of up to three million years (Racki 2005); the end Triassic extinction (another of the big five) is thought to represent periods of elevated extinction rate bracketing the entire Rhaetian stage (Ward et al. 2001, 2004), a duration of almost seven million years based on the most recent timescale of the International Commission on Stratigraphy. Moreover, if one is to follow the stratigraphic ages espoused by Lucas (1998, 2002, 2004, 2006, 2017), the Redtankian would be compressed into a period

covering less than four million years. During these four million years, extinction rates remain consistently higher than background levels. The Arroyo formation records a substantial increase in extinction, and the Choza formation records extinction rates that have more-than doubled those of the Arroyo, higher than in any other formation. The number of tetrapod species observed in the Choza formation is less than a quarter of those observed in the Arroyo formation, and subsampling does not diminish the extent of the diversity loss.

It is worth noting at this point that mass extinctions appearing in the fossil record as prolonged declines is an issue that has a long history of discussion in the published literature, going back to the work of Signor & Lipps (1982). The fact that the last appearance of a taxon in the fossil record is not its last true appearance, combined with differential preservation probabilities of different taxa, causes a set of species, which in reality died out nearly simultaneously, to appear to have died out over a longer period of time (Butterfield 1995), a phenomenon dubbed the Signor-Lipps effect. Lucas (2017) acknowledged the Signor-Lipps effect in his introduction but did not mention it in his discussion of specific extinction events. He also employed no sampling correction when examining diversity and extinction rate, instead arguing that "whatever biases exist may be roughly equivalent in the Permian tetrapod record across times and localities" (Lucas 2017, p. 35). Unfortunately, this is simply not true: there is a wealth of literature detailing analyses of the quality of the fossil record of Paleozoic tetrapods, all suggesting the opposite and emphasising the need for sampling correction (Benson & Upchurch 2013; Brocklehurst et al. 2013, 2014, 2017; Verriere et al. 2016).

Another argument put forward by Lucas (2017) to show that Olson's Extinction does not qualify as a genuine mass extinction is that many of the clades previously deemed to have died out at this time actually disappeared before the end of the Redtankian, and the number of actual casualties of the event, at the family level, was very restricted. Brocklehurst et al. (2013) previously noted Edaphosauridae and Ophiacodontidae as "casualties", but Lucas (2017)

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countered that the former's last appearance is from the Arroyo formation rather than the end of the Redtankian, and that the latter is not known from beyond the Mitchellcreekian LVF. In the case of the Ophiacodontidae, this is actually not the case, and the family survived into the Redtankian. Lucas (2017) based his assertion on the last record of Ophiacodon, and the abundant record of the Varanosaurus, represented in the Arroyo Formation by the species V. acutirostris (Broili 1904, Case 1907, 1910, Romer & Price 1940) and V. witchitaensis (NB Pers. Obs), was discounted as representing a taxon of uncertain assignment. However, almost three decades of study, both anatomical and cladistic, support the ophiacodontid affinity of Varanosaurus (Sumida 1989, Berman et al 1995, Benson 2012, Brocklehurst et al. 2016), and I see no reason not to count it as the youngest record of Ophiacodontidae. On the subject of Edaphosauridae, only one species of Edaphosaurus is known from the Arroyo formation (E. pogonias), but it still represents one of the most abundant herbivores in this fauna (Data S1). Neural spine material of Edaphosaurus is also known from the Hennessey Formation (Daly 1973), a Redtankian aged formation in Oklahoma. It is clear, therefore, that both Ophiacodontidae and Edaphosauridae survived into the Redtankian. While they may not have survived beyond the lowest of the Redtankian formations, this does not remove them from the Olson's Extinction casualty list. As already discussed, extinction rates were raised considerably above background levels throughout the Redtankian, and extinctions of the taxa of the Arroyo formation should be included in event. Even if we are to limit our discussion to clades which went extinct at the end of the Choza

Even if we are to limit our discussion to clades which went extinct at the end of the Choza Formation, there are still multiple clades above the genus level which may be included in the list of casualties of Olson's extinction, mostly amphibians. Probably the most prominent are the Eryopidae, since they represent one of the few cases where we have data on their disappearance from both palaeoequatorial (USA) and palaeotemperate localities (Brocklehurst et al. 2017). Eryopids represent among the most abundant of the large amphibians throughout the Cisuralian, and *Eryops* itself survives until the Choza Formation (Data S1). Crucially, two eryopid species

are known from the latest Kungurian of Russia: Clamorosaurus borealis and C. nocturnus from the Inta formation (Gubin 1983). Eryopids are not known beyond the Kungurian in either the 318 palaeoequatorial or palaeotemperate latitiudes beyond this time (Brocklehurst et al. 2017). The 319 Trimerorhachidae and Lysorophia are two more clades highly abundant throughout the 320 Cisuralian, but which do not survive beyond the Choza formation (Data S1). Both are also known 321 from the Redtankian aged Hennessey formation in Oklahoma, but not from the Littlecrotonion 322 Chickasha formation (Brocklehurst et al. 2017). The Choza Formation represents the greatest 323 peak in extinction rate in the entire Cisuralian in this particular section, both at the genus and 324 species level, with extinction rates more than double the next highest peak. Therefore, even if one 325 discounts the losses occurring earlier in the Redtankian, it is difficult to deny the presence of a 326 severe extinction event across the Redtankian/Littlecrotonian boundary. 327

#### Conclusions

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No matter what time-binning scheme is employed, no matter whether the data is examined at the species or genus level, and no matter whether the data is corrected for sampling or not, a substantial mass extinction event is observed in tetrapods during the Redtankian Land

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Vertebrate Fanuachron.

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340	References
341	Alroy J. 1998. Equilibrial diversity dynamics in North American mammals. Pp. 233-287 in M. L
342	McKinney and J. A. Drake, eds. Biodiversity dynamics: turnover of populations, taxa and
343	communities. Columbia University Press, New York.
344	Alroy J. 2010a Geographical, environmental and intrinsic biotic controls on Phanerozoic marine
345	diversification. Palaeontology 53:1211-1235.
346	Alroy J. 2010b. Fair sampling of taxonomic richness and unbiased estimation of origination and
347	extinction rates. The Paleontological Society Papers 16: 55-80.
348	Alroy J. 2014. Accurate and precise estimates of origination and extinction rates. Paleobiology
349	40:374–397.
350	Alroy J. 2015a A new twist on a very old binary similarity coefficient. Ecology 96:575-586
351	Alroy J. 2015b A more precise speciation and extinction rate estimator. Paleobiology 41: 633-639
352	Archibald SB, Bossert WH, Greenwood DR, Farrell BD. 2010. Seasonality, the latitudinal
353	gradient of diversity, and Eocene insects. Paleobiology 36:374-398.
354	Benson RB. 2012. Interrelationships of basal synapsids: cranial and postcranial morphological
355	partitions suggest different topologies. Journal of Systematic Palaeontology 10: 601-624.
356	Benson RBJ, Upchurch P. 2013 Diversity trends in the establishment of terrestrial vertebrate
357	ecosystems: interactions between spatial and temporal sampling biases. Geology 41:43-
358	46.
359	Benton MJ. 2012. No gap in the Middle Permian record of terrestrial vertebrates. Geology
360	40:339-342.

361	Berman DS, Reisz RR, Bolt JR, Scott D. 1995. The cranial anatomy and relationships of the
362	synapsid Varanosaurus (Eupelycoauria: Ophiacodontidae) from the early Permian of
363	Texas and Oklahoma. Annals of the Carnegie Museum 64:99-133
364	Brocklehurst N, Fröbisch J. 2014. Current and historical perspectives on the completeness of the
365	fossil record of pelycosaurian-grade synapsids. Paleogeography, Paleoclimatology,
366	Paleoecology 399:114-126
367	Brocklehurst N, Kammerer CF, Fröbisch J. 2013 The early evolution of synapsids and the
368	influence of sampling on their fossil record. Paleobiology 39:470-490.
369	Brocklehurst N, Ruta M, Müller J, Fröbisch J. 2015 Elevated extinction rates as a trigger for
370	diversification rate shifts: early amniotes as a case study. Scientific Reports 5:17104.
371	Brocklehurst N, Reisz RR, Fernandez V, Fröbisch J. 2016. A re-description of 'Mycterosaurus'
372	smithae, an Early Permian eothyridid, and its impact on the phylogeny of pelycosaurian-
373	grade synapsids. PloS one, 11: e0156810.
374	Brocklehurst N, Day MO, Rubidge BS, Fröbisch J. 2017 Olson's extinction and the latitudinal
375	biodiversity gradient of tetrapods in the Permian. Proceedings of the Royal Society B 284
376	20170231
377	Brocklehurst N, Day MO, Fröbisch J. 2018. Accounting for differences in species frequency
378	distributions when calculating beta diversity in the fossil record. Methods in Ecology and
379	Evolution: In Press
380	Broili F. 1904. Permische Stegocephalen und Reptilien aus Texas. Palaeontographica 51:1-120
381	Butterfield NJ. 1995. Secular distribution of Burgess-Shale-type preservation. Lethaia 28: 1-13.

382	Case EC. 1907. Revision of the Pelycosauria of North America. Carnegie Institution of
383	Washington 55: 3-176
384	Case EC. 1910. New or Little Known Reptiles and Amphibians from the Permian of Texas. Order
385	of the Trustees, American Museum of Natural History.
386	Daly E. 1973. A Lower Permian vertebrate fauna from southern Oklahoma. Journal of
387	Paleontology 47: 562-589
388	Foote M. 1994. Temporal variation in extinction risk and temporal scaling of extinction metrics.
389	Paleobiology 20:424–444.
390	Foote 2000. Origination and extinction components of taxonomic diversity: general problems.
391	Paleobiology 26:74–102.
392	Gibert C, Escarguel G. 2017. Evaluating the accuracy of biodiversity changes through geologic
393	times: from simulation to solution. Paleobiology 43:667-692
394	Grimm EC. 1987. CONISS: a FORTRAN 77 program for stratigraphically constrained cluster
395	analysis by the method of incremental sum of squares. Computers & geosciences 13:13-
396	35.
397	Gubin YM. 1983. The first eryopids from the Permian of the Eastern European platform.
398	Paleontologicheskii Zhurnal 4:110-115
399	Hillebrand H. 2004 On the generality of the latitudinal diversiy gradient. American Naturalist
100	163:192-211
101	Hook RW. 1989. Stratigraphic distribution of tetrapods in the Bowie and Wichita groups, Permo-
102	Carboniferous of north-central Texas. In: Hook, R.W. (Ed.), PermoCarboniferous

## PaerJ

403	Vertebrate Paleontology, Lithostratigraphy and Depositional Environments of North-
404	central Texas. Society of Vertebrate Paleontology, Austin, 47-53.
405	Juggins, S., 2009. Rioja: an R package for the analysis of quaternary science data, Version 0.5-3.
406	Kemp TS. 2006 The origin and early radiation of the therapsid mammal-like reptiles: a
407	palaeobiological hypothesis. Journal of Evolutionary Biolology 19:1231-1247.
408	Lozovsky, V.R. 2003. Correlation of the continental Permian of norther Pangea: a review.
409	Bolletina della Societa Italiano Volume Especiale 2:239-244.
410	Lucas, S.G., 1994. Triassic tetrapod extinctions and the compiled correlation effect. Canadian
411	Society of Petroleum Geology Memoir 17, 869–875.
412	Lucas, S.G., 1998b. Toward a tetrapod biochronology of the Permian. New Mexico Museum of
413	Natural History Science Bulletin 12, 71-91.
414	Lucas, S.G. 2002. The reptile Macroleter: First vertebrate evidence for correlation of Upper
415	Permian continental strata of North America and Russia: Discussion. Geological Society
416	of America Bulletin 114:1174-1175
417	Lucas SG. 2004. A global hiatus in the Middle Permian tetrapod fossil record. Stratigraphy 1:47-
418	64.
419	Lucas SG. 2006. Global Permian tetrapod biostratigraphy and biochronology. In: Lucas, S.G.,
420	Cassinis, G., Schneider, J.W. (Eds.), Non-marine Permian Biostratigraphy and
421	Biochronology. Geological Society London Special Publications Vol. 265. 65-93
422	Lucas SG. 2017. Permian tetrapod extinction events. Earth Science Reviews 170:31-60
423	Lucas SG, Heckert SB. 2001 A global hiatus in the record of Middle Permian tetrapods. Journal
124	of Vertebrate Paleontology 21:75.

425	Mannion PD, Benson RBJ, Upchurch P, Butler RJ, Carrano WT, Barrett PM. 2012 A temperate
426	palaeodiversity peak in Mesozoic dinosaurs and evidence for Late Cretaceous
427	geographical partitioning. Global Ecolology and Biogeography Letters 21:898-908.
428	Mannion PD, Upchurch P, Benson RBJ, Goswami A. 2014 The latitudinal biodiversity gradient
429	through deep time. Trends in Ecology and Evolution 29:42-50.
430	Miller AI, and Foote M. 1996. Calibrating the Ordovician radiation of marine life: implications
431	for Phanerozoic diversity trends. Paleobiology, 22:304-309.
432	Olson, E.C., 1958. Fauna of the Vale and Choza: 14 summary, review, and integration of geology
433	and the faunas. Fieldiana. 10, 397-448.
434	Olson EC. 1962 Late Permian terrestrial vertebrates, USA and USSR. Transactions of the
435	American Philosophical Society 52:1-224.
436	Olson EC. 1966 Community evolution and the origin of mammals. Ecology 47:291-302.
437	Olson EC. 1982 Extinctions of Permian and Triassic nonmarine vertebrates. GSA Special Papers
438	190:501-512.
439	Olson EC. 1989. The Arroyo Formation (Leonardian: lower Permian) and its vertebrate fossils.
440	Texas Memorial Museum Bulletin 2: 1-25
<b>4</b> 41	Plummer FB, Moore RC. 1921. Stratigraphy of the Pennsylvanian formations of northcentral
142	Texas. University of Texas Bulletin 2132: 1-237
143	Racki G. 2005. Toward understanding Late Devonian global events: few answers, many
144	questions. Developments in Palaeontology and Stratigraphy 20: 5-36.
145	Raup DM. 1972. Taxonomic diversity during the Phanerozoic. Science 177:1065–1071.

446	Reisz RR, Laurin M. 2001. The reptile Macroleter: First vertebrate evidence for correlation of
447	Upper Permian continental strata of North America and Russia. Geological Society of
448	America Bulletin, 113:1229-1233.
449	Reisz RR, Laurin M. 2002. The reptile Macroleter: First vertebrate evidence for correlation of
450	Upper Permian continental strata of North America and Russia: Discussion and reply.
451	Geological Society of America Bulletin, 114:1174-1175.
452	Romer AS. 1928. Vertebrate faunal horizons in the Texas Permo-Carboniferous red beds.
453	University of Texas Bulletin 2801, 67–108.
454	Romer AS. 1935. Early history of Texas redbeds vertebrates. Geological Society of America
455	Bulletin 46, 1597–1658.
456	Romer AS. Price LI. 1940. Review of the Pelycosauria. Geological Society of America Special
457	Papers Number 28
458	Rose PJ, Fox DL, Marcot J, Badgley C. 2011 Flat latitudinal gradient in Paleocene mammal
459	richness suggests decoupling of climate and biodiversity. Geology 39:163-166.
460	Ruta M, Benton MJ. 2008 Calibrated diversity, tree topology and the mother of mass extinctions
461	the lesson of temnospondyls. Palaeontology 51:1261-1288.
462	Ruta M, Cisneros JC, Liebrecht T, Tsuji LA, Müller J. 2011 Amniotes through major biological
463	crises: faunal turnover among parareptiles and the end-Permian mass extinction.
464	Palaeontology 54:1117-1137.
465	Sahney S, Benton MJ. 2008 Recovery from the most profound mass extinction of all time.
466	Proceedings of the Royal Society B 275:759-765.

467	Sepkoski Jr. JJ. 1984. A kinetic model of Phanerozoic taxonomic diversity. III. Post Paleozoic
468	families and mass extinctions. Paleobiology 10, 246-267.
469	Signor PW., Lipps JH. 1982. Sampling bias, gradual extinction patterns, and catastrophes in the
470	fossil record. In L. T. Silver and P. H. Schultz (ed.) Geological implications of impacts of
471	large asteroids and comets on the Earth. Geological Society of America Special
472	Publication, vol. 190: 291-296.
473	Sumida SS. 1989. Reinterpretation of vertebral structure in the Early Permian pelycosaur
474	Varanosaurus acutirostris (Amniota, Synapsida). Journal of Vertebrate Paleontology 9:
475	451-458
476	Verrière A, Brocklehurst N, Fröbisch J. 2017. Assessing the completeness of the fossil record:
477	comparison of different methods applied to parareptilian tetrapods (Vertebrata:
478	Sauropsida). Paleobiology 42:680-695
479	Ward PD, Haggart JW, Carter ES, Wilbur D, Tipper HW, Evans T. 2001. Sudden productivity
480	collapse associated with the Triassic-Jurassic boundary mass extinction. Science
481	292:.1148-1151.
482	Ward PD, Garrison GH, Haggart JW, Kring DA, Beattie MJ. 2004. Isotopic evidence bearing on
483	Late Triassic extinction events, Queen Charlotte Islands, British Columbia, and
484	implications for the duration and cause of the Triassic/Jurassic mass extinction. Earth and
485	Planetary Science Letters 224: 589-600.
486	Willig MR, Kaufman DM, Stevens RD. 2003 Latitudinal gradients of biodiversity: pattern,
487	process, scale, and synthesis. Annual Review of Ecology, Evolution and Systematics
488	34:273-309.

489	Yasuhara M, Hunt G, Dowsett HJ, Robinson MM, Stoll DK. 2012 Latitudinal species diversity
490	gradient of marine zooplankton for the last three million years. Ecology Letters 15:1174-
491	1179.
492	Figure Captions
493	Fig. 1: The time bins used in the diversity analysis
494	Legend: A) The cluster dendrogram indicating the grouping of the formations by CONISS; B)
495	The tetrapod bearing formations in texas; C) the Land Vertebrate Faunachrons (LVFs)
496	redefined by CONISS; D) the original LVFs
497	Fig. 2: Taxic Diversity Estimates
498	Legend: Diversity estimates without correcting for sampling, using four different methods of
499	time-binning the data. A) Species level diversity estimate; B) Genus level diversity
500	estimate.
501	Fig 3: Subsampled diversity estimates (Original Land Vertebrate Faunachrons)
502	Legend: Numbers of species (A) and genera (B) in each land vertebrate faunachron (original
503	definitions), corrected for sampling heterogeneity using shareholder quorum subsampling.
504	Legend indicates quorum level.
505	Fig 4: Subsampled diversity estimates (redefined Land Vertebrate Faunachrons)

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506	Legend: Numbers of species (A) and genera (B) in each land vertebrate faunachron (definitions
507	based on CONISS), corrected for sampling heterogeneity using shareholder quorum
508	subsampling. Legend indicates quorum level.
509	Fig 5: Subsampled diversity estimates (Formations)
510	Legend: Numbers of species (A) and genera (B) in each formation, corrected for sampling
511	heterogeneity using shareholder quorum subsampling. Legend indicates quorum level.
512	Fig 6: Subsampled diversity estimates (half-million-year time bins)
513	Legend: Means of the numbers of species (A) and genera (B) found in each half-million-year
514	time bin in each of the 100 stochastic distributions of specimens, corrected for sampling
515	heterogeneity using shareholder quorum subsampling. Legend indicates quorum level.
516	Fig 7: Extinction Rates
517	Legend: Median (thick black lines) of the extinction rates calculated for each half-million-year
518	time bin in each of the 100 stochastic distributions of specimens at the genus (A) and
519	species (B) levels. Dashed lines indicate standard error around the median.