1	$\label{lem:condition} \begin{tabular}{ll} Ultrastructure of the Jurassic serpulid tubes-phylogenetic and paleoecological implications \end{tabular}$
3	
4	Jakub Słowiński <sup>1</sup> , Olev Vinn <sup>2</sup> , Michał Zatoń <sup>1</sup>
5	
6	<sup>1</sup> University of Silesia in Katowice, Institute of Earth Sciences, Sosnowiec, Poland
7	<sup>2</sup> Department of Geology, University of Tartu, Tartu, Estonia
8	Corresponding author:
9	Jakub Słowiński <sup>1</sup>
10	Będzińska 60, Sosnowiec 41-205, Poland
11	E-mail address: jakub.slowinski@us.edu.pl
12	
13	Abstract
14	The ultrastructural diversity of the Middle and Late Jurassic serpulid tubes from t
15	Polish Basin has been investigated. The inspection of 12 taxa representing the two major

The ultrastructural diversity of the Middle and Late Jurassic serpulid tubes from the Polish Basin has been investigated. The inspection of 12 taxa representing the two major serpulid clades allowed for the identification of three ultrastructure kinds – irregularly oriented prismatic structure (IOP), spherulitic prismatic structure (SPHP), and simple prismatic structure (SP). Six of the studied species are single-layered and six species possess two distinct layers. Ultrastructural diversity corresponds to certain serpulid clades. The members of Filograninae have single-layered tube walls composed of possibly plesiomorphic, irregularly oriented prismatic structure (IOP). Two-layered tubes occur solely within the clade Serpulinae, where the external, denser layer is built of either the ordered spherulitic (SPHP) or simple prismatic microstructure (SP), and the internal layer is composed of irregularly oriented prismatic structure (IOP). Apart from phylogenetic signals provided by the tube ultrastructure, it can be used in analyzing paleoecological aspects of tube-dwelling polychaetes. Compared to the more primitive, irregularly oriented microstructures of Filograninae, the regularly oriented microstructures of Serpulinae need a higher level of biological control over biomineralization. The advent of the dense outer protective layer

(DOL) in serpulids, as well as the general increase in ultrastructure diversity, was likely a result of the evolutionary importance of the tubes for serpulids, enhanced also by the intensified predation during the Mesozoic Marine Revolution. Such diversity of the tube ultrastructural fabrics allowed for maximizing functionality by utilizing a variety of morphogenetic programs. The biomineralization system of serpulids remains more complex compared to other tube-dwelling polychaetes. Physiologically more expensive tube formation allows for mechanical strengthening of the tube by building robust, strongly ornamented tubes and firm attachment to the substrate. Contrary to sabellids, which perform a fugitive strategy, an increased tube durability allows serpulids a competitive advantage over other encrusters.

#### Introduction

Numerous polychaete families produce habitation tubes (Rouse and Pleijel, 2001). Tubicolous polychaetes may either agglutinate exogenous material, such as sand particles and shell fragments, using a proteinaceous cement to form a tube (e.g., Stewart et al., 2004; Zhao et al., 2005; Fournier, Etienne and Le Cam, 2010; Vinn and Luque, 2013) or produce secretions by themselves utilizing a variety of glands (e.g., Hausen, 2005; Tanur et al., 2010). Secreted tubes may be composed of organic substances, such as proteins and polysaccharides (e.g., Barnes, 1965; Chamoy et al., 2001; Nishi and Rouse, 2013), and mineral substances such as calcium carbonate (Weedon, 1994; Fischer, Pernet and Reitner, 2000; Vinn et al., 2008). Amongst tube-dwelling polychaetes producing hard, mineralized exoskeletons are the families Sabellidae, Cirratulidae, and Serpulidae. Whereas calcareous sabellids (Perkins, 1991; Vinn, ten Hove and Mutvei, 2008; Słowiński, Banasik and Vinn, 2023) and cirratulids (Reish, 1952; Fischer, Oliver and Reitner, 1989; Fischer, Pernet and Reitner, 2000; Taylor et al., 2010; Kočí et al., 2021, Guido et al. 2024) are restricted to one (Glomerula) and two genera (Dodecaceria, Diplochaetetes) in each family respectively, only serpulids dwell exclusively in tubes composed of calcium carbonate (Vinn et al., 2008).

Thus, having a hard, mineralized exoskeleton, serpulids are the most abundant polychaetes in the fossil record, appearing as far back as the Middle Permian (Sanfilippo et al., 2017; 2018) and Upper Permian (Ramsdale 2021). However, their hard parts are not strictly related to the organism's soft body and form only a habitation tube serving, among others, for protection against predators (e.g., Morton and Harper, 2009; Klompmaker, 2012). Although tubes' taxonomic usefulness is not without validity, as many genera possess easily recognizable tubes, which are diagnostic displaying different longitudinal and transverse elements (see e.g., Ippolitov et al., 2014), homeomorphy is still a common feature in serpulid

Okomentoval(a): [p1]: Guido, A., D'Amico, F., DeVries, T. J., Kočí, T., Collareta, A., Bosio, G., & Sanfilippo, R. (2024). Double-phased controlled and influenced biomineralization in marine invertebrates: The example of Miocene to recent reef-building polychaete cirratulids from southern Peru. Palaeogeography, Palaeoclimatology, Palaeoecology, 639, 112060.

Okomentoval(a): [p2]: Please to cite this work of our colleague Ramsdale, R. (2021). A possible serpulid tube worm of the genus Filograna from the upper Permian Cadeby Formation of South Yorkshire, UK. Proceedings of the Yorkshire Geological Society, 63(3), pygs2020-016.

tube morphology (e.g., Kupriyanova and Ippolitov, 2015) due to convergence and high ecophenotypic plasticity. This leads to certain discrepancies between paleontological and biological classification, the latter of which is based solely on the molecular data, soft parts, and their reciprocal relationships (Kupriyanova, Macdonald and Rouse, 2006; ten Hove and Kupriyanova, 2009). Moreover, it makes the linkage of contemporary species with their ancestors difficult due to the lack of fully reliable taxonomic tools.

Serpulid tubes can be composed of up to four layers; however, most of them are single-layered (Vinn et al., 2008). They exhibit different ultrastructural fabrics, depending on how calcium carbonate crystals are arranged and oriented (see Vinn et al., 2008). During the progressing expansion of the tube, an animal secretes consecutive growth lamellae, which correspond to a single growth episode. Such lamellae may be either straight, perpendicular to the direction of growth (e.g., Bałuk and Radwański, 1997) or much more frequently chevronshaped (see Weedon, 1994). Serpulid ultrastructures may provide phylogenetic signals in the case of Jurassic taxa but also serve as a record of the physiological changes of the worm during its entire lifespan, providing important ecological signatures that may be successfully implemented in paleoecological analyses. Moreover, serpulid tubes may exhibit additional skeletal structures such as tubulae, alveolar structures, or internal tube structures, helping in unraveling their true systematics (e.g., Thomas, 1940, Pillai, 1993; Pillai and ten Hove, 1994; Jäger, 2005).

Recent studies have proven a tube microstructure to be a useful tool in deciphering the true affinity of fossil serpulid (e.g., Bornhold and Milliman, 1973; Vinn, ten Hove and Mutvei, 2008; Vinn and Kupriyanova, 2011; Kupriyanova and Ippolitov, 2012; Vinn, 2013; Buckman, 2020; Buckman and Harries, 2020), and other calcareous tube-dwelling polychaete species (Vinn, ten Hove and Mutvei, 2008; Taylor et al., 2010; Słowiński, Banasik and Vinn, 2023). Anyhow, the majority of them were conducted based on single species, and/or single specimens.

In the present study, we performed a microstructural investigation of the Jurassic serpulid tubes based on the representatives of the two main serpulid clades according to the newest systematics (Kupriyanova, ten Hove and Rouse, 2023). We provide paleoecological implications of different microstructures and consequently biomineralization system and highlight some phylogenetic signals and evolutionary patterns of distinct microstructures.

Material and its provenance

 **Okomentoval(a): [p3]:** Thomas, J. G. (1940). Pomatoceros, Sabella and Amphitrite. (*No Title*).

The examined fossils consist of 47 specimens of serpulids representing 12 taxa (table 1). The studied material has been selected from an ample collection (see Słowiński et al., 2020; 2022) concerning the well-preserved specimens, which allowed for a reliable ultrastructural investigation. All taphonomically altered tubes have been discarded from further studies. Almost all specimens encrust a variety of invertebrate fossils, oncoids, and hiatus concretions, and have been derived from various localities with Middle and Upper Jurassic (upper Bajocian – lower Kimmeridgian) deposits representing an array of paleoenvironments. The investigated material is stored at the Institute of Earth Sciences in Sosnowiec, abbreviated GIUS 8-3589 (Callovian of Zalas), GIUS 8-3730 (Gnaszyn Dolny), GIUS 8-3745 (Bolęcin), GIUS 8-3746 (Oxfordian of Zalas), GIUS 8-3747 (Małogoszcz), GIUS 8-3750 (Ogrodzieniec-Świertowiec), GIUS 8-3751 (Krzyworzeka and Żarki). In ascending stratigraphical order, the localities are as follows:

Ogrodzieniec-Świertowiec: serpulids from this outcrop heavily encrust oncoids of cyanobacterial genesis (Zatoń et al., 2012). The oncoids have been derived from condensed, sandy, and carbonate layers, being overlain by clay sediments. This and the subsequent three localities are situated within the area of the Polish Jura (Fig. 1). The deposits occur in the form of dark siliciclastic sediments belonging to the Częstochowa Ore-bearing Clay Formation (e.g., Majewski, 2000; Matyja and Wierzbowski, 2000; Zatoń et al., 2011; Leonowicz, 2015). The late Bajocian–early Bathonian age of the oncoid-bearing deposits was based on ammonites (Zatoń and Taylor, 2009a; Zatoń et al., 2012), dinoflagellate cysts, and foraminifers (Słowiński, 2019). Due to the origin of oncoids, the paleoenvironment is interpreted to be shallow marine, located within the photic zone, and slightly below a fairweather wave base (Zatoń et al., 2012).

Gnaszyn Dolny: serpulid tubes derived from this locality in majority encrust belemnite rostra and oyster shells. The fossils come from middle Bathonian mudstones originating in marine environment, below the storm wave base (Gedl and Kaim, 2012; Gedl et al., 2012).

Krzyworzeka: polychaetes from this site encrust carbonate hiatus concretions. The concretion-bearing sediments from this locality are confined to the upper Bathonian (Poulsen, 1998). The deposits from Krzyworzeka belong to the northern sedimentary region of the Polish Jura and originated in a calm outer shelf environment (Zatoń et al., 2011; Sadlok and Zatoń, 2020), located below the storm wave base (Różycki, 1953; Leonowicz, 2013).

Żarki: the deposits from which the serpulids encrusting hiatus concretions have been derived are confined to the northern sedimentary region of the Polish Jura, and are represented by the upper Bathonian mudstones (Zatoń and Taylor, 2009b). Sediments within this location

originated in a relatively calm but interrupted by episodic storm events (Leonowicz, 2015), marine paleoenvironment (Marynowski et al., 2007; Leonowicz, 2012; 2013).

Bolęcin: the outcrop is situated approximately 6 km east of Chrzanów. The serpulid tubes have been found within sandy limestones, presumably referring to the so-called "Balin Oolite." Serpulid tubes have been found fixed to a variety of fossils including bivalves, gastropods, and ammonites, as well as belemnite rostra. These condensed sediments (see Tarkowski et al., 1994; Mangold et al., 1996) deposited presumably in an open marine paleoenvironment and are late Bathonian-early Callovian in age (Mangold et al., 1996).

Zalas: the fossil material collected from this locality comes from two sites: 1) an active quarry with sandy limestones forming a hardground of middle-earliest late Callovian age (Giżejewska and Wieczorek, 1977; Dembicz and Praszkier, 2007); here, the serpulids encrust a wide variety of invertebrate fossils, notably large *Ctenostreon proboscideum* (J. Sowerby and G.B. Sowerby, 1820) bivalves. Hardground deposits originated in a deeper open marine setting corresponding to a shelf environment (Dembicz and Praszkier, 2007; Zatoń, Wilson and Zavar, 2011); 2) overlying, lower Oxfordian light limestones hosting sponge bioherms (Matyja, 2006; Matyszkiewicz, Kochman and Duś, 2012), to which serpulid tubes were attached.

Małogoszcz: serpulid fauna has been derived from an active quarry located in close proximity to the town of Małogoszcz. This site is located in the southwestern part of the Mesozoic margin of the Holy Cross Mountains. The tubes are attached to bivalve shells, particularly the oyster *Actinostreon gregareum* (J. Sowerby, 1815). Lower Kimmeridgian fossiliferous deposits are referred to as the Skorków Lumachelle (Kutek, 1994; Machalski, 1998; Zatoń and Machalski, 2013). These shell beds were deposited during episodic storms in a relatively shallow marine paleoenvironment (Machalski, 1998).

### Methods

Before the preparation of specimens for the SEM examination, they were coated with ammonium chloride and photographed using the Canon EOS camera. All previously selected samples were cut longitudinally, polished, and etched with a 5% solution of acetic acid for one minute prior to the SEM study. Part of these tube portions were oriented and mounted in epoxy resin before polishing. Subsequently, the serpulid tube microstructure investigations were performed on a SEM Zeiss EVO MA15 under a back-scattered electron detector (BSE) in a low vacuum regime (Department of Geology, University of Tartu, Estonia). The sections were also inspected under an environmental scanning electron microscope (ESEM) Philips

**Okomentoval(a):** [p4]: Please to specify how number etc....please to add inside the text

XL30, and a Quanta 250 scanning electron microscope using BSE imaging (both at the Institute of Earth Sciences in Sosnowiec). The beam voltage was operated at 20 kV.

#### Results

During the examination of 12 taxa within eight genera representing two major clades, three kinds of ultrastructure were distinguished: irregularly oriented prismatic structure (IOP), spherulitic prismatic structure (SPHP), and simple prismatic structure (SP) (Table 1).

## Filograninae (BI)

Members of the clade Filograninae, formerly referred to as BI, are represented in the herein material by two genera: *Metavermilia* and *Filogranula*, the latter one consisting of two species: *F. runcinata* (J. de C. Sowerby, 1829) and *F. spongiophila* (J. Słowiński, O. Vinn, M. Jäger et al., 2022).

The tube wall of *Metavermilia* cf. *striatissima* (F.T. Fürsich, T.J. Palmer, and K.L. Goodyear, 1994) is single-layered, and composed of an irregularly oriented prismatic structure (IOP) (Fig. 2). Minute (maximally a few µm), needle-like crystals are deployed more or less evenly within the entire tube wall. Their crystallization axes are longitudinal and lack a uniform orientation.

The tube of *Filogranula runcinata* is presumably single-layered and is composed of a spherulitic, regularly oriented prismatic microstructure (SPHP) (Fig. 3). It is formed by crystals of prismatic shape exhibiting a slightly spherulitic arrangement. Some internal parts of the tube are built of irregularly oriented, tiny, elongated crystals indicating an irregularly oriented prismatic structure (IOP). However, the boundary between the two putative layers is transitional and the latter microstructure occurs rather like inclusions in certain areas of the tube wall. Growth lines are apparent across almost the entire tube. Some minor, external parts of the section may appear like fine homogeneous granular microstructure; however, more likely it corresponds to the differences in preservation between the external and internal parts of the tube due to the diagenetic alteration of irregularly oriented prismatic microstructure.

The tube of *Filogranula spongiophila* consists of a single layer that is composed of a simple, regularly oriented prismatic structure (SP) (Fig. 4). Growth increments are visible across almost the whole tube (Fig. 4C-E).

Serpulinae, tribe Serpulini (AI)

This tribe is represented in the investigated material by one genus comprising two species: *Cementula spirolinites* (Münster in Goldfuss, 1831) and *Cementula radwanskae* (J. Słowiński, O. Vinn, M. Jäger et al., 2022).

The tube wall of *C. spirolinites* is single-layered and consists of a simple prismatic structure (SP) (Fig. 5). This microstructure is formed by parallel prismatic crystals oriented perpendicularly to each growth line showing incremental zonation. The crystals are arranged perpendicularly or obliquely to the tube wall.

The tube microstructure of the investigated specimens of *C. radwanskae* is altered in most places. However, some places exhibiting faint growth increments indicate possibly a simple (SP) or spherulitic prismatic structure (SPHP) constituting a single layer (Fig. 6). Certain areas with tiny, elongated irregularly oriented crystals presumably are an effect of a diagenetic distortion of the tube wall.

#### Serpulinae, tribe Ficopomatini (AII)

196

197 198

199

200

201

202

203204

205

206207

208

209

210

211212

213

214

215216

217

218219

220

221

222

223

224225

226

227

228229

The tribe Ficopomatini, formerly referred to as AII, includes here 5 genera with 7 species. Propomatoceros lumbricalis (E.F. Schlotheim, 1820) possesses a tube wall consisting of two or three layers (Fig. 7). The majority of the examined specimens have twolayered tubes. The external part is formed by a spherulitic prismatic structure (SPHP), which consists of parallel crystals of a slightly spherulitic arrangement. The inner part has an irregularly oriented prismatic microstructure (IOP) comprising elongated, but relatively short, tiny crystals with inconsistent orientation axes. In some areas of the tubes' wall, the IOP structure is diagenetically altered superficially resembling a fine homogeneous granular structure. The putative three-layered tube wall (Fig. 7D) is built of the external and internal (uncertain) parts composed of a spherulitic prismatic structure (SPHP). The middle part shows a fine homogeneous granular microstructure, which may be either a primary microstructure or similarly to other specimens, an obliterated IOP structure. In contrast to two-layered P. lumbricalis tubes, no areas exhibiting irregularly oriented prismatic structures have been found. As such, it is not evident, whether these microstructures are primary or have undergone diagenetic alteration. The external parts exhibit incremental zonation with prominent chevronshaped growth lines. The crystals are continuously and regularly positioned through successive growth increments.

In the herein material *Nogrobs* is represented by three species: *N*. aff. *quadrilatera* (A. Goldfuss, 1831), *N*. aff. *tetragona* (J. de C. Sowerby, 1829), and *N*? aff. *tricristata* (A. Goldfuss, 1831), which, however, share all the major microstructural characters. The tube walls of all the species have two layers, which are separated from each other with a sharp

boundary. The external layer is composed of a simple prismatic structure (SP), whereas the internal layer is composed of very thin, short, needle-like crystals, which make up an irregularly oriented prismatic structure (IOP) (Fig. 8).

Mucroserpula sp. (H. Regenhardt, 1961) possesses a two-layered tube wall (Fig. 9). The internal layer is composed of an irregularly oriented prismatic structure (IOP) consisting of bunches of densely packed, short but elongated crystals lacking uniform orientation axes. The external layer is composed of a regular spherulitic prismatic structure (SPHP), which constitutes the major part of the tube wall. Within this layer, crystals are oriented perpendicularly with respect to each incremental zone and have a somewhat prismatic arrangement. The chevron-shaped growth lines are very well-visible alongside almost the whole tube length. The border between the two layers is somewhat transitional.

The tube wall of *Placostegus planorbiformis* (Münster in Goldfuss, 1831) is single-layered, and composed of a simple prismatic structure (SP) (Fig. 10). All crystals are more or less parallel within each incremental zone. Within the outermost part of the tube, the crystals are oriented perpendicularly to the tube wall.

The specimens of undetermined serpulid, Serpulidae sp. possess a two-layered tube wall (Fig. 11). The external part features a regularly oriented simple prismatic structure (SP). The internal, thinner part, separated from the external with a sharp boundary is formed by an irregularly oriented prismatic structure (IOP) composed of densely packed, minute, elongated crystals.

#### Discussion

## Tube ultrastructure evolution and its phylogenetic constraints

Three distinct kinds of ultrastructure have been identified within 12 taxa corresponding to the two of three main clades of serpulids – Filograninae and Serpulinae (Kupriyanova, ten Hove and Rouse, 2023). These microstructures comprise irregularly oriented prismatic structure (IOP), spherulitic prismatic structure (SPHP), and simple prismatic structure (SP) (see Table 1), which are among the most prevalent microstructure kinds in fossil serpulids (Vinn et al., 2008; Vinn, 2020). Six of these taxa have single-layered tubes and six are two-layered, of which one taxon perhaps may possess either two or three ultrastructural layers (see discussion). The majority of both Cenozoic and contemporary serpulids are single-layered, only about one-third of serpulid species have at least two or up to four distinct ultrastructural layers (Vinn et al., 2008). During the Jurassic, the percentage of multi-layered serpulid species was lower, constituting approximately 25% (Vinn and Furrer,

2008). It may have resulted from the evolutionary pattern where more complex tube walls with at least two layers were more common from the beginning of the Cenozoic onward – the growing complexity of microstructures combined with relatively fast biomineralization likely enhanced the strengthening abilities of the tubes. It can also be explained by the vagaries of the fossil record resulting in incomplete preservation of the outermost layers, which might be similar to the case of calcareous sabellids (Vinn, ten Hove and Mutvei, 2008; Słowiński, Banasik and Vinn, 2023).

 Sanfilippo (1998) presented a concept to utilize the ultrastructural diversity of serpulids to assess their generic determination. Further studies revealed that microstructure characters may be species-dependent (Vinn, 2007; Vinn et al., 2008; Kupriyanova and Ippolitov, 2015), which limits the application of tube microstructures in deciphering the generic affiliation of serpulids. Regardless, ultrastructural fabrics of tubes may still be used to approach the relationships between distinct kinds of structures throughout the evolution of the main serpulid clades (e.g., Vinn et al., 2008; Vinn and Kupriyanova, 2011; Vinn, 2013; Ippolitov and Rzhavsky, 2014; 2015a; b).

The most recent phylogenetic analyses indicate that serpulids are split into three major clades – Filograninae, Spirorbinae, and Serpulinae, the last of which is further subdivided into two tribes: Serpulini and Ficopomatini (Kupriyanova, ten Hove and Rouse, 2023). The family Serpulidae until recently was maintained to comprise two major clades referred to as B and A, separated into BI, BII, AI, and AII (Kupriyanova, Macdonald and Rouse, 2006, Kupriyanova et al., 2009; Kupriyanova and Nishi, 2010). Formerly recognized clades BI and BII generally may be related to Filograninae and Spirorbinae, whereas AI and AII, both settled within Serpulinae, correspond to the tribes Serpulini and Ficopomatini, respectively. Apart from Spirorbinae (former BII), members of all clades are present in the material studied.

Clade Filograninae is possibly rooted even in the Permian with primitive tubes such as *Filograna* (Sanfilippo et al., 2017; Ramsdale, 2021), and abundantly represented in the fossil record during the Mesozoic by many strongly ornamented tubes with several keels, e.g., *Vermiliopsis*, *Metavermilia*, and *Filogranula*, the latter two of which are present in the material investigated. *Metavemilia* cf. *striatissima* is single-layered and possesses irregularly oriented prismatic microstructure (IOP), which has been also found in the recent *Metavermilia multicristata* (Vinn et al., 2008). It is characteristic of this clade and is also the most common microstructure in the recent serpulids encompassing ca 60% of species (Vinn, 2007). IOP microstructure most often builds the wall of single-layered tubes, or alternatively the inner part of multi-layered tubes (see Vinn et al., 2008).

The representatives of Filograninae are generally characterized in the fossil record by more or less uniform fabrics of tube microstructures forming irregularly oriented structures, notably irregularly oriented prismatic structure (IOP). Secretion of such microstructures seems to be governed by the lower biological control of an animal, compared to more advanced microstructures. The first confirmed appearance of such ultrastructure dates back to the Early Jurassic, possibly being present already in the Triassic (Vinn, Jäger and Kirsimäe, 2008). Such a simpler way of tube formation may indicate its plesiomorphic character and earlier origin.

298 299

300

301

302 303

304

305 306

307

308 309

310

311312

313 314

315

316

317 318

319

320 321

322

323

324

325

326327

328

329

330

331

Another representative of this clade in the investigated material is the genus Filogranula comprising the recently discovered new species F. spongiophila (Słowiński, Vinn, Jäger et al., 2022), which has been studied with respect to ultrastructure for the first time, and by F. runcinata. The tube wall of F. spongiophila is single-layered and has a simple prismatic structure (SP), whereas F. runcinata possesses one well-preserved layer composed of spherulitic prismatic structure (SPHP) and a dubious IOP layer consisting of somewhat distorted, occasionally occurring areas of tiny, irregularly oriented crystals. As a result, the tube wall is interpreted as single-layered, composed of SPHP structure. Such microstructure occurs commonly in the clade Serpulinae comprising both formerly established "A" clades. The systematic position of the fossil Filogranula seems to be complicated, as the Recent Filogranula is considered to be settled within Filograninae, relatively closely related to Vermiliopsis and Metavermilia (Kupriyanova, ten Hove and Rouse, 2023). While the Cretaceous Filogranula cincta (A. Goldfuss, 1831) is classified as belonging to Filograninae too, the Jurassic Filogranula runcinata was considered to be placed in Serpulinae, within the tribe Ficopomatini (see Ippolitov et al., 2014; Kočí and Jäger, 2015), due to its prismatic microstructure, characteristic of the serpulid tubes appearing transparent in the Recent species, such as e.g., Placostegus. On the other hand, prismatic structures are not a fully satisfactory character in differentiating between former A and B clades. Ippolitov and Rzhavsky (2014; 2015a; b) found out that some spirorbins may also have their tubes composed of regularly oriented prismatic microstructures, in spite of the fact, that Spirorbinae is phylogenetically closely related to Filograninae, which commonly form irregularly oriented structures.

Serpulinae constitutes two well-supported tribes – Serpulini and Ficopomatini (Kupriyanova, ten Hove and Rouse, 2023). The first unequivocal members of Serpulinae very likely existed since the Triassic (Assmann, 1937) if not even since the Permian (Sanfilippo et al., 2017; 2018). In contrast to Filograninae, which have mostly non-oriented structures,

serpulids within this lineage have a more advanced biomineralization system and are capable of forming complex microstructures, such as e.g., oriented, lamello-fibrillar, or oriented-fibrillar (see Vinn et al., 2008; Vinn and Kupriyanova, 2011). During the Jurassic, serpulid microstructure diversity was relatively modest (ten kinds; see Schlögl et al. 2018 Vinn, 2020) compared to an array of contemporary and Cenozoic ones (Vinn, 2007, Kočí et al 2023). The most sophisticated microstructures in serpulids evolved in the Eocene (e.g., Vinn, 2008; Buckman, 2020, Kočí 2022). Although such complex microstructures as lamello-fibrillar or regularly ridged prismatic appeared quite late in the evolution of serpulids, the oriented prismatic structures evolved at the latest during the Middle Jurassic. Simple and spherulitic prismatic microstructures are very common and characteristic of the Jurassic Serpulinae and are considered to be apomorphic. Oriented prismatic structures are unknown from any genus of Filograninae; however, they were found to persist in the Cretaceous clade Spirorbinae (Ippolitov and Rzhavsky, 2015a; b), which may mean, that either it is plesiomorphic, or the prismatic structures have evolved at least twice.

 The tribe Serpulini is represented in the Jurassic by characteristically coiled *Spiraserpula* and related *Cementula*, which is present in the herein material. *C. spirolinites* and *C. radwanskae* are both single-layered and possess their tube walls built of simple prismatic microstructure (SP), though in the latter it is somewhat obliterated. Simple and spherulitic prismatic microstructures exhibit the predominantly uniform orientation of calcium carbonate crystals which determines the optical transparency of the tubes (Zibrowius and ten Hove, 1987), supported also by the dense arrangement of crystals and their large size (Vinn and Kupriyanova, 2011). In contrast, the recent tubes with IOP microstructure most often are optically opaque (Vinn et al., 2008).

All the main members of the Jurassic representatives of Ficopomatini are present in the investigated material. They contain genera with various morphotypes comprising a robust, single-keeled *Propomatoceros*, a three-keeled *Mucroserpula*, quadrangular tubes attributed most often to *Nogrobs*, and a well-defined, planispirally coiled *Placostegus*. Serpulids in this tribe commonly possess two layers of tube walls having the outer dense layer built of ordered prismatic structures (simple or spherulitic), and the internal layer composed of irregularly oriented prismatic structure. Two-layered tubes appeared in the Jurassic (Vinn and Furrer, 2008) and have their external layer denser than the internal, which is composed of thinner mineral microstructure. Dense outer protective layers (DOL) have been found in tube walls of serpulids inhabiting diverse environments (see Vinn and Kupriyanova, 2011). Consequently, rather than being dependent on the environment, the advent of DOLs during the Jurassic

Okomentoval(a): [p5]: Schlögl, J., Kočí, T., Jäger, M., Segit, T., Sklenář, J., Sadki, D., ... & Tomašových, A. (2018). Tempestitic shell beds formed by a new serpulid polychaete from the Bajocian (Middle Jurassic) of the Central High Atlas (Morocco). *PalZ*, 92, 219-240.

**Okomentoval(a):** [p6]: This work is For Paleocene microstructures

Okomentoval(a): [p7]: Please cite this work KOČÍ, T., MILÀN, J., & JÄGER, M. (2023). Neovermilia gundstrupensis sp. nov.(Polychaeta, Serpulidae) from the Selandian (middle Paleocene) of Fyn, Denmark. Bulletin of the Geological Society of Denmark, 72.

Okomentoval(a): [p8]: Please to cite these microstructures here: Kočí, T., Goedert, J. L., & Buckeridge, J S. (2022). Eocene tube-dwelling annelids (Polychaeta: Sedentaria) from the Black Hills, western Washington State: the first record of Neodexiospira from North America. PalZ, 96(4), 631-653.

appears to be a significant evolutionary adaptation of serpulids. Except for one species in the clade Filograninae, dense, outer layers exist exclusively in the clade Serpulinae (Vinn and Kupriyanova, 2011).

Propomatoceros lumbricalis and Mucroserpula possess here two well-distinguishable layers: internal IOP and external SPHP. A single Propomatoceros specimen may have three layers, which, however, is ambiguous, as its internal layer may be diagenetically obliterated, and therefore all the specimens are considered two-layered. As discussed previously (Ippolitov et al., 2014; Słowiński et al., 2022), it cannot be ruled out, that at least in some populations, Mucroserpula and Propomatoceros may represent the same genus. The diagnostic characters of both genera are highly transitional, subjective, and dependent to a large extent on different variables, such as ontogenetic stage or diagenesis (see Słowiński et al., 2022). Nonetheless, the microstructural fabrics of the two genera in our investigation are essentially the same and typical of Ficopomatini.

Nogrobs consists here of three species: N. aff. quadrilatera, N? aff. tricristata, and N. aff. tetragona. All of them are two-layered composed of internal IOP and external SP microstructure. Apart from Nogrobs, Jurassic quadrangular fossils of serpulids that share morphological similarities with this genus were attributed by various authors to different, partially uncertain genera - Tetraserpula, Tetraditrupa, Glandifera, Tubulostium, Tectorotularia, and Ditrupula. Kupriyanova and Ippolitov (2015) examined and reviewed a number of extant and fossil taxa having tusk-shaped, tetragonal in cross-section tubes, concluding, that these morphologically similar forms belong to several different genera, most likely being an effect of convergence. Not surprisingly, these recent taxa cannot be synonymized with the fossil Nogrobs. Nevertheless, the authors claim that at least the majority of fossil Nogrobs species may be members of a single clade, as opposed to morphologically-related recent species. According to its external tube layer composed of simple prismatic microstructure (responding to a transparent tube), the three investigated species fit well Ficopomatini, confirming previous ultrastructural studies (Kupriyanova and Ippolitov, 2012). It has to be noted, however, that the recent Nogrobs grimaldii (P. Fauvel, 1909) has been found to possess an opaque tube (Kupriyanova and Nishi, 2011).

The only exception not having a two-layered tube wall within this tribe is *Placostegus*, represented here by *P. planorbiformis*, which has a single layer made entirely of a simple prismatic structure (SP). Similarly, contemporary *P. tridentatus* (J.C. Fabricius, 1779) possesses a tube wall composed of a simple prismatic microstructure resulting in a completely transparent tube (ten Hove and Kupriyanova, 2009: 8, fig. 1F; Vinn and Kupriyanova, 2011).

The undetermined serpulid taxon Serpulidae sp., described recently by Słowiński et al. (2022) has been studied here with respect to its microstructure. Although the investigated specimens externally resemble a few genera, such as *Propomatoceros*, *Placostegus*, and *Metavermilia*, they do not exactly fit any of those taxa (see Słowiński et al., 2022). The specimens possess a tube wall composed of two layers – external simple prismatic (SP), and internal irregularly oriented prismatic microstructure (IOP), which is of a very close resemblance to other members of Ficopomatini, notably *Nogrobs*.

#### Ecological implications and comparisons with sabellids

Serpulids perform biologically controlled biomineralization where their cellular activity regulates the nucleation and extracellular growth of the calcium carbonate crystals by the ion uptake from the surrounding water (see Neff, 1971a; b) using a secretory epithelium, mediated and controlled by the organic matrix (Vinn, Kirsimäe and ten Hove, 2009; Vinn, 2021a). As a result, serpulids accomplish a specified crystal orientation within the tube wall, which is reflected by their variety of ultrastructural fabrics - the growth direction of crystals may be anisotropic, semi-oriented, or oriented (Weedon, 1994; Vinn et al., 2008). Apart from a matrix-mediated crystallization (Vinn, 2021a), an alternative explanation has been proposed recently to define the crystal orientation mechanism, which is based on a variable application of the serpulid's collar rotational force between the formation of ordered and unordered microstructures (Buckman and Harries, 2020). Considering the fact that the same growth increments may occur across zones with diverse ultrastructures makes (Vinn, 2021a; b) this model strongly disputable.

The serpulid biomineralization system differs from that of other tube-dwelling polychaetes. By the formation of the cylindrical and parable layers (see Jäger, 1983) serpulids are capable of forming multi-layered tubes arranged in distinct microstructures (see Vinn et al., 2008). Importantly, the parable layer is formed by adding secretory increments to the rim of the worm's aperture allowing it to actively modify the external morphology according to the tube's sculpture, but also depending on the temporary ecological requirements. Such a solution results in a wide array of serpulid tube characters including attachment structures and base widenings of the tube, which combined with the ability to form several layers greatly improves the durability of the tube. Fossil cirratulids (Vinn, 2009; Taylor et al., 2010) and sabellids (Vinn, ten Hove and Mutvei, 2008; Słowiński, Banasik and Vinn, 2023) possess single-layered tube wall composed of a spherulitic prismatic structure. The presence of only cylindrical layer in sabellids strongly impairs their biomineralization abilities. Unlike

serpulids having mostly chevron-shaped growth lines (see Weedon, 1994), they form their tubes by secreting calcareous material along distinct growth lamellae oriented parallel to the tube wall (Vinn, ten Hove and Mutvei, 2008; Słowiński et al., 2022). Subsequent increments are added to the internal surface of the tube, and therefore *Glomerula* is unable to modify its simple tube architecture. A much longer secretion zone in sabellids compared to serpulids enables them for a fast calcification incurring comparably lower physiological costs, which allows them to considerably prolong their tubes. On the other hand, such fast dispersal may be required by the diminishing inside of the tube, insufficient for the growing worm.

Interestingly, cirratulids perform a double-phased, combined controlled, and influenced biomineralization. The products of an influenced biomineralization such as agglutinated xenolithic granules within a calcareous matrix, as well as the anisotropic orientation of different-sized crystals, indicate, that cirratulids govern even a weaker than sabellids control over biomineralization (Guido et al., 2024).

434 435

436

437

438 439

440

441 442

443

444445

446

447

448

449

450

451

452

453 454

455

456 457

458

459

460

461

462

463 464

465

466 467

A much more advanced biomineralization system of serpulids and the resulting variety of microstructures are an important evolutionary adaptation for this taxon, which is an obligatory tube-dweller. The emergence of multi-layered tubes with dense, outer protective layers during the Jurassic could have been triggered by the intensified predation during the Mesozoic Marine Revolution (Vermeij, 1977). With the skeleton development, serpulids could perform competitive strategies against other organisms instead of avoiding competition or temporary paleoenvironmental vagaries. Gradual increase in abundance of serpulids during the Mesozoic and Cenozoic (see Ippolitov et al., 2014 for a review), and their advantage over other organisms allowed for higher plasticity and divergence of different morphotypes (Vinn et al., 2023 (in press)). Different morphogenetic programs in turn enhanced their functional utilization not only by mechanical strengthening of the tubes but also by optimizing their space by planispiral coiling, providing a higher feeding tier, avoiding being overgrown or coated by sediment thanks to upward growth. Free-living serpulids with tusk-shaped tubes were presumably adapted to live in or on the surface of soft sediment during episodes of increased sedimentation rates (see Vinn et al., 2023 for a review (in press)). Sabellid biomineralization system allowed them to quickly elongate their tubes but these polychaetes did not exhibit such diverse morphotypes. Instead, they could grow irregularly away from the initial point of encrustation and perform an opportunistic, fugitive strategy (see Taylor, 2016; Słowiński, Banasik and Vinn, 2023). Additionally, it could be the outcome of biomineralization's lack of significance in this group as calcareous sabellids were restricted to a single genus in the family.

# Conclusions

The first thorough assessment of the ultrastructural diversity of Middle and Late Jurassic serpulid tubes has been conducted. The obtained data reveal a characteristic of Jurassic serpulids, a relatively low diversity of ultrastructural fabrics, which generally correspond to certain clades. Amongst 12 taxa representing two (Filograninae and Serpulinae) of the three main serpulid clades, six of them possess tube walls composed of a single layer, and six are two-layered. There are certain evolutionary trends in tube ultrastructures. The representatives of the clade Filograninae are single-layered and have their tube walls built of a primitive, irregularly oriented prismatic microstructure (IOP). The majority of members of possibly apomorphic clade Serpulinae possess two-layered tube walls, where the denser, external layer is composed of oriented, prismatic microstructures (either spherulitic (SPHP) or simple(SP)), and the internal is irregularly oriented prismatic (IOP). The exceptions are *Placostegus planorbiformis* and the genus *Cementula*, which are single-layered, and built of simple prismatic structure (SP).

Serpulid tube ultrastructures reflect their biomineralization abilities providing important paleoecological signatures. Formation of the regularly oriented microstructures of Serpulinae requires a higher biological control over biomineralization compared to the more primitive, anisotropic microstructures of Filograninae. The development of serpulid ultrastructure diversity was likely triggered by the evolutionary importance of the tubes for this group. The formation of dense outer protective layers in serpulids was presumably induced by increased predation rates during the Mesozoic Marine Revolution. A variety of microstructure kinds and the ability to form multi-layered tubes allowed serpulids to employ different morphogenetic programs, which had an impact on their functional utilization.

The differences in the biomineralization system between serpulids and other tube-dwelling polychaetes resulted from the importance of tubes for the former taxon. The complex biomineralization system of serpulids resulted in a multiplicity of forms and an ability to form robust, strongly ornamented tubes, which mechanically strengthened their durability. Although higher energy expenditure of skeletal secretion decreased the rate of tube formation, the solid attachment and skeleton robustness allowed for a competitive advantage over other encrusters. Conversely, sabellid primitive biomineralization abilities presumably were elicited by the unimportance of skeleton for this taxon resulting in a simplicity of forms.

502 On the other hand, it enabled a fast spreading over substrate shunning competition and 503 disadvantageous conditions by utilizing an opportunistic strategy of quick tube elongation. 504 Acknowledgments 505 We would like to warmly thank Marian Külaviir (University of Tartu) and Arkadiusz 506 Krzątała (University of Silesia) for their assistance with SEM. 507 508 509 References Assmann P. 1937. Revision der Fauna der Wirbellosen der oberschlesischen Trias. 510 Abhandlungen der Preußischen Geologischen Landesanstalt 170: 1–134. 511 Bałuk W, Radwański A. 1997. The micropolychaete Josephella commensalis sp.n. 512 513 commensal of the scleractinian coral Tarbellastraea reussiana (Milne-Edwards & Haime, 1850) from the Korytnica Clays (Middle Miocene; Holy Cross Mountains, Central 514 515 Poland). Acta Geologica Polonica 47: 211-224. Barnes RD. 1965. Tube-building and feeding in chaetopterid polychaetes. Biological 516 517 Bulletin 129: 217-233. DOI: 10.2307/1539840. Bornhold BD, Milliman JD. 1973. Generic and environmental control of carbonate 518 519 mineralogy in serpulid (polychaete) tubes. Journal of Geology 81: 363–373. DOI: 520 10.1086/627876. Buckman JO. 2020. The tube of Ditrupa bartonensis (Annelida, Serpulidae), from the 521 Eocene of southern England: Observations on microstructure and its significance. 522 Palaeontologia Electronica 23: a37. DOI: 10.26879/1098. 523 524 Buckman JO, Harries DB. 2020. Reef forming Serpula vermicularis from Scotland 525 and Ireland: Tube structure, composition and implications. Zoologischer Anzeiger 288: 53-65. 526 DOI: 10.1016/j.jcz.2020.07.003. 527 Chamoy L, Nicola M, Ravaux J, Quennedey B, Gaill F, Delachambre J. 2001. A novel chitin-binding protein from vestimentiferan Riftia pachyptila interacts specifically with β-528 chitin. Journal of Biological Chemistry 276: 8051-8058. DOI: 10.1074/jbc.M009244200. 529 Dembicz K, Praszkier T. 2007. Kelowej południowo-wschodniej części Jury 530 Krakowsko-Częstochowskiej. Tomy Jurajskie 4: 71-76. 531 Fabricius O. 1779. Reise nach Norwegen mit Bemerkungen aus der Naturhistorie und 532 533 Oekonomie. Carl Ernst Bohn: Hamburg. 388 pp.

535 campagnes de l'Hirondelle et de la Princesse-Alice, ou déposéesdans la Musée 536 Océanographique de Monaco. Bulletin de l'Institute Oceéanographique 142: 1–76. Fischer R, Oliver CG, Reitner J. 1989. Skeletal structure, growth, and paleoecology of 537 the patch reef-building polychaete worm Diplochaetetes mexicanus Wilson, 1986 from the 538 Oligocene of Baja California (Mexico). Geobios 22: 761-775. DOI: 10.1016/S0016-539 6995(89)80071-3. 540 Fischer R, Pernet B, Reitner J. 2000. Organomineralization of Cirratulid tubes – Fossil 541 and Recent examples. Facies 42: 35-50. DOI: 10.1007/BF02562565. 542 543 Fournier J, Etienne S, Le Cam JB. 2010. Inter- and intraspecific variability in the chemical composition of the mineral phase of cements from several tube-building 544 polychaetes. Geobios 43: 191-200. DOI: 10.1016/j.geobios.2009.10.004. 545 Fürsich FT, Palmer TJ, Goodyear KL. 1994. Growth and disintegration of bivalve-546 547 dominated patch-reefs in the Portlandian (Upper Jurassic) of southern England. Palaeontology 37: 131-171. 548 Gedl P, Kaim A. 2012. An introduction to the palaeoenvironmental reconstruction of 549 the Bathonian (Middle Jurassic) ore-bearing clays at Gnaszyn, Kraków-Silesia Homocline, 550 Poland. Acta Geologica Polonica 62: 267–280. DOI: 10.2478/v10263-012-0014-y. 551 Gedl P, Kaim A, Leonowicz P, Boczarowski A, Dudek T, Kędzierski M, Rees J, 552 Smoleń J, Szczepanik P, Sztajner P, Witkowska M, Ziaja J. 2012. Palaeoenvironmental 553 554 reconstruction of Bathonian (Middle Jurassic) ore-bearing clays at Gnaszyn, Kraków-Silesia 555 Homocline, Poland. Acta Geologica Polonica 62: 463-484. DOI: 10.2478/v10263-012-0026-556 557 Giżejewska M, Wieczorek J. 1977. Remarks on the Callovian and lower Oxfordian of the Zalas Area (Cracow Upland, Southern Poland). Bulletin de l'Académie Polonaise des 558 Sciences, Série des Sciences de la Terre 24: 167–175. 559 560 Goldfuss A. 1831. Petrefacta Germaniae. I. Divisio secunda: Radiariorum Reliquiae -Strahlenthiere der Vorwelt. Arnz and Co.: Düsseldorf. 165–240. 561 Guido A, D'Amico F, DeVries TJ, Kočí T, Collareta A, Bosio G, Sanfilippo R. 2024. 562 Double-phased controlled and influenced biomineralization in marine invertebrates: The 563 example of Miocene to recent reef-building polychaete cirratulids from southern Peru. 564 Palaeogeography, Palaeoclimatology, Palaeoecology 639: 112060. DOI: 565

10.1016/j.palaeo.2024.112060.

566

Fauvel P. 1909. Deuxiéme note préliminaire sur les Polychétes provenantdes

Hydrobiologia 535: 25-35. DOI: 10.1007/s10750-004-4442-x. 568 569 Hove HA ten, Kupriyanova EK. 2009. Taxonomy of Serpulidae (Annelida, Polychaeta): the state of affairs. Zootaxa 2036: 1-126. 570 Ippolitov AP, Rzhavsky AV. 2014. Tube morphology, ultrastructures and mineralogy 571 in recent Spirorbinae (Annelida: Polychaeta: Serpulidae). I. General introduction. Tribe 572 Paralaeospirini. Invertebrate Zoology 11: 293–314. DOI: 10.15298/invertzool.11.2.01. 573 Ippolitov AP, Rzhavsky AV. 2015a. Tube morphology, ultrastructures and mineralogy 574 575 in recent Spirorbinae (Annelida: Polychaeta: Serpulidae). II. Tribe Spirorbini. Invertebrate 576 Zoology 12: 61–92. DOI: 10.15298/invertzool.12.1.03. Ippolitov AP, Rzhavsky AV. 2015b. Tube morphology, ultrastructures and mineralogy 577 in recent Spirorbinae (Annelida: Polychaeta: Serpulidae). III. Tribe Circeini. Invertebrate 578 Zoology 12: 151–173. DOI: 10.15298/invertzool.12.2.03. 579 580 Ippolitov AP, Vinn O, Kupriyanova EK, Jäger M. 2014. Written in stone: history of 581 serpulid polychaetes through time. Memoirs of Museum Victoria 71: 123–159. DOI: 10.24199/j.mmv.2014.71.12. 582 Jäger, M. 1983. Serpulidae (Polychaeta sedentaria) aus der norddeutschen höheren 583 Oberkreide - Systematik, Stratigraphie, Ökologie. Geologisches Jahrbuch, Reihe A 68: 3-584 585 Jäger M. 2005. Serpulidae und Spirorbidae (Polychaeta sedentaria) aus Campan und 586 587 Maastricht von Norddeutschland, den Niederlanden, Belgien und angrenzenden Gebieten. 588 Geologisches Jahrbuch A 157: 121-249. Klompmaker AA. 2012. Drill hole predation on fossil serpulid polychaetes, with new 589 590 data from the Pliocene of the Netherlands. Palaeogeography, Palaeoclimatology, Palaeoecology 321–322: 113–120. DOI: 10.1016/j.palaeo.2012.01.024. 591 Kočí T, Bosio G, Collareta A, Sanfilippo R, Ekrt B, Urbina M, Malinverno E. 2021. 592 First report on the cirratulid (Annelida, Polychaeta) reefs from the Miocene Chilcatay and 593 Pisco Formations (East Pisco Basin, Peru). Journal of South American Earth Sciences 107: 594 595 103042. DOI: 10.1016/j.jsames.2020.103042. 596 Kočí, T., Goedert, J. L., & Buckeridge, J. S. 2022. Eocene tube-dwelling annelids 597 (Polychaeta: Sedentaria) from the Black Hills, western Washington State: the first record of Neodexiospira from North America. PalZ, 96(4), 631-653. 598 599 Kočí T, Jäger M. 2015. Filogranula cincta (Goldfuss, 1831), a serpulid worm

(Polychaeta, Sedentaria, Serpulidae) from the Bohemian Cretaceous Basin. Sborník

Hausen H. 2005. Comparative structure of the epidermis in polychaetes (Annelida).

567

Historia Naturalis 71: 293–300. DOI: 10.14446/AMNP.2015.293. 602 Kočí, T., Milàn, J., & Jäger, M. 2023. Neovermilia gundstrupensis sp. 603 604 nov.(Polychaeta, Serpulidae) from the Selandian (middle Paleocene) of Fyn, 605 Denmark. Bulletin of the Geological Society of Denmark, 72. Kupriyanova EK, ten Hove HA, Rouse GW. 2023. Phylogeny of Serpulidae 606 (Annelida, Polychaeta) Inferred from Morphology and DNA Sequences, with a New 607 Classification. Diversity 15: 398. DOI: 10.3390/d15030398. 608 609 Kupriyanova EK, ten Hove HA, Sket B, Trontelj P, Zakšek V, Rouse GW. 2009. 610 Evolution of a unique freshwater cave-dwelling serpulid polychaete Marifugia cavatica Absolon and Hrabě, 1930. Systematics and Biodiversity 7: 389–401. 611 Kupriyanova EK, Ippolitov AP. 2012. Are Mesozoic shallow-water serpulids 612 (Annelida, Polychaeta) ancestors of the Recent deep-sea fauna? Abstracts of the 13th Deep-613 Sea Biology Symposium, Wellington, New Zealand. 614 615 Kupriyanova EK, Ippolitov AP. 2015. Deep-sea serpulids (Annelida: Polychaeta) in tetragonal tubes: on a tube convergence path from the Mesozoic to Recent. Zootaxa 4044: 616 151-200. DOI: 10.11646/zootaxa.4044.2.1. 617 Kupriyanova EK, Macdonald TA. Rouse GW. 2006. Phylogenetic relationships within 618 619 Serpulidae (Sabellida, Annelida) inferred from molecular and morphological data. Zoologica Scripta 35: 421-439. DOI: 10.1111/j.1463-6409.2006.00244.x. 620 621 Kupriyanova EK, Nishi E. 2010. Serpulidae (Annelida, Polychaeta) from Patton-Murray Seamounts, Gulf of Alaska, North Pacific Ocean. Zootaxa 2665: 51-68. DOI: 622 623 10.11646/zootaxa.2665.1.3. Kupriyanova EK, Nishi E. 2011. New records of the deep-sea Nogrobs grimaldii 624 (Serpulidae: Annelida). Marine Biodiversity Records 4: 1-4. DOI: 625 10.1017/S1755267211000674. 626 Kutek J. 1994. Jurassic tectonic events in south-eastern cratonic Poland. Acta 627 Geologica Polonica 44: 167-222. 628 Leonowicz P. 2012. Sedimentology and ichnology of Bathonian (Middle Jurassic) ore-629

bearing clays at Gnaszyn, Kraków-Silesia Homocline, Poland. Acta Geologica Polonica 62:

Národního Muzea v Praze, řada B, Přírodní védy. Acta Musei Nationalis Pragae, Series B,

601

630

631

281-296.

632 Leonowicz P. 2013. The significance of mudstone fabric combined with 633 palaeoecological evidence in determining sedimentary processes—an example from Middle Jurassic of southern Poland. Geological Quarterly 57: 243–260. DOI: 10.7306/gq.1092. 634 Leonowicz P. 2015. Ichnofabrics of shallow-marine mudstone, the result of changing 635 environmental conditions: an example from the Middle Jurassic ore-bearing clay from 636 southern Poland. Facies 61: 11. DOI: 10.1007/s10347-015-0438-4. 637 Machalski M. 1998. Oyster life positions and shell beds from the Upper Jurassic of 638 Poland. Acta Palaeontologica Polonica 43: 609-634. 639 Majewski W. 2000. Middle Jurassic concretions from Częstochowa (Poland) as 640 641 indicators of sedimentation rates. Acta Geologica Polonica 50: 431-439. Mangold C, Marchand D, Thierry J, Tarkowski R. 1996. Les Ammonites de l'Oolite 642 de Balin (Pologne); nouvelles données et réinterprétation stratigraphique. Reuve de 643 644 Palèobiologie 15: 55–77. Marynowski L, Zatoń M, Simoneit BRT, Otto A, Jędrysek MO, Grelowski C, 645 646 Kurkiewicz S. 2007. Compositions, sources and depositional environments of organic matter from the Middle Jurassic clays of Poland. Applied Geochemistry 22: 2456–2485. DOI: 647 648 10.1016/j.apgeochem.2007.06.015. Matyja BA. 2006. Stop A17—Zalas Quarry. Callovian transgressive to condensed 649 pelagic deposits, Lower to lowermost Middle Oxfordian deposits of sponge megafacies. In: 650 Wierzbowski A, Aubrecht R, Golonka J, Gutowski J, Krobicki M, Matyja BA, Pieńkowski G, 651 652 Uchman A, eds. Jurassic of Poland and Adjacent Slovakian Carpathians. Field Trip 653 Guidebook of 7th International Congress on the Jurassic System Poland, Kraków, September 6-18, 2006, 70-72. Kraków, Poland. 654 655 Matyja BA, Wierzbowski A. 2000. Ammonites and stratigraphy of the uppermost Bajocian and Lower Bathonian between Częstochowa and Wieluń, Central Poland. Acta 656 Geologica Polonica 50: 191-209. 657 Matyszkiewicz J, Kochman A, Duś A. 2012. Influence of local sedimentary conditions 658 on development of microbialites in the Oxfordian carbonate buildups from the southern part 659 of the Kraków-Częstochowa Upland (South Poland). Sedimentary Geology 263-264: 109-660 132. 661 Morton B, Harper EM. 2009. Drilling predation upon Ditrupa arietina (Polychaeta: 662 Serpulidae) from the Mid-Atlantic Açores, Portugal. Açoreana, Suplemento, Açores 6: 157-663 664 165.

666	serpulid polychaete worm, Pomatoceros caeruleus. Zeitschrift für Zellforschung und
667	Mikroskopische Anatomie 120: 160–186. DOI: 10.1007/BF00335534.
668	Neff JM. 1971b. Ultrastructure of calcium phosphate-containing cells in the serpulid
669	polychaete worm Pomatoceros caeruleus. Calcified Tissue Research 7: 191–200. DOI:
570	10.1007/BF02062606.
571	Nishi E, Rouse GW. 2013. First whale fall chaetopterid; a gigantic new species of
572	Phyllochaetopterus (Chaetopteridae: Annelida) from the deep sea off California. Proceedings
573	of the Biological Society of Washington 126: 287–298. DOI: 10.2988/0006-324X-126.4.287.
574	Perkins, T.H. 1991: Calcisabella piloseta, a new genus and species of Sabellinae
575	(Polychaeta: Sabellidae). Bulletin of Marine Science 48: 261–267.
576	Pillai TG. 1993. A review of some Cretaceous and Tertiary serpulid polychaetes of the
577	genera Cementula and Spiraserpula Regenhardt 1961, Laqueoserpula Lommerzheim 1979
578	and Protectoconorca Jäger 1983. Paläontologische Zeitschrift 67: 69-88. DOI:
579	10.1007/BF02985871.
680	Pillai TG, ten Hove HA. 1994. On recent species of Spiraserpula Regenhardt, 1961, a
581	serpulid polychaete genus hitherto known only from Cretaceous and Tertiary fossils. $Bulletin$
582	of the Natural History Museum, London (Zoology) 60: 39–104.
583	Poulsen N. 1998. Upper Bajocian to Callovian (Jurassic) dinoflagellate cysts from
584	central Poland. Acta Geologica Polonica 48: 237–245.
585	Ramsdale R. 2021. A possible serpulid tube worm of the genus Filograna from the
86	upper Permian Cadeby Formation of South Yorkshire, UK. Proceedings of the Yorkshire
87	Geological Society 63: pygs2020-016. DOI: 10.1144/pygs2020-016.
888	Regenhardt H. 1961. Serpulidae (Polychaeta sedentaria) aus der Kreide Mitteleuropas,
589	ihre ökologische, taxionomische und stratigraphische Bewertung. Mitteilungen aus dem
590	Geologischen Staatsinstitut in Hamburg 30: 5–115.
591	Reish DJ. 1952. Discussion of the Colonial Tube-Building Polychaetous Annelid
592	Dodecaceria fistulicola Ehlers. Bulletin of the Southern California Academy of Sciences 51:
593	103–107.
594	Rouse GW, Pleijel F. 2001. <i>Polychaetes</i> . Oxford: Oxford University Press. 354 pp.
595	Różycki SZ. 1953. Górny dogger i dolny malm Jury Krakowsko-Częstochowskiej.
96	Przegląd Instytutu Geologicznego 17: 1–420.
97	Sadlok G, Zatoń M. 2020. Ichnology of the Middle Jurassic hiatus concretions from
392	Poland: implications for their formation, exhumation, and palaegenvironment

Neff JM. 1971a. Ultrastructural studies of the secretion of calcium carbonate by the

700 6. 701 Sanfilippo R. 1998. Tube morphology and structure of the bathyal Maditerranean serpulid Hyalopomatus variorugosus Ben-Eliahu et Fiege, 1996 (Annelida, Polychaeta). 702 Rivista Italiana di Paleontologia e Stratigrafia 104: 131–138. DOI: 10.13130/2039-703 4942/6124. 704 Sanfilippo R, Rosso A, Reitano A, Insacco G. 2017. First record of sabellid and 705 serpulid polychaetes from the Permian of Sicily. Acta Palaeontologica Polonica 62: 25-38. 706 707 DOI: 10.4202/app.00288.2016. 708 Sanfilippo R, Rosso A, Reitano A, Viola A, Insacco G. 2018. New serpulid polychaetes from the Permian of western Sicily. Acta Palaeontologica Polonica 63: 579-584. 709 DOI: 10.4202/app.00448.2017. 710 Schlotheim EF von 1820. Die Petrefactenkunde auf ihrem jetzigen Standpunkte. 711 712 Gotha: Becker. 437 pp. 713 Schlögl, J., Kočí, T., Jäger, M., Segit, T., Sklenář, J., Sadki, D., ... & Tomašových, A. 714 2018. Tempestitic shell beds formed by a new serpulid polychaete from the Bajocian (Middle 715 Jurassic) of the Central High Atlas (Morocco). PalZ, 92, 219–240. 716 Słowiński J. 2019. Środkowojurajskie onkoidy z okolic Ogrodzieńca. Unpublished 717 M.Sc. Thesis, University of Silesia, Sosnowiec. 57 pp. 718 719 Słowiński J, Banasik K, Vinn O. 2023. Insights into mineral composition and 720 ultrastructure of Jurassic sabellid tubes (Annelida, Polychaeta): the evolution of sabellid calcification and its palaeoecological implications. Lethaia 56: 1-12. DOI: 721 722 10.18261/let.56.3.8. Słowiński J, Surmik D, Duda P, Zatoń M. 2020. Assessment of serpulid-hydroid 723 association through the Jurassic: A case study from the Polish Basin. PLOS ONE 15: 724 e0242924. DOI: 10.1371/journal.pone.0242924. 725 Słowiński J, Vinn O, Jäger M, Zatoń M. 2022: Middle and Late Jurassic tube-dwelling 726 polychaetes from the Polish Basin: diversity, palaeoecology and comparisons with other 727 assemblages. Acta Palaeontologica Polonica 67: 827-864. DOI: 10.4202/app.01006.2022. 728 Sowerby J. 1815. The Mineral Conchology of Great Britain. London: The Author. 234 729 730 pp. Sowerby J, Sowerby GB. 1820. The Genera of Recent and Fossil Shells, For the Use 731

of Students in Conchology and Geology. London: G.B. Sowerby. 275 pp.

Palaeobiodiversity and Palaeoenvironments 100: 757-771. DOI: 10.1007/s12549-019-00410-

699

733	Sowerby J de C. 1829. The Mineral Conchology of Great Britain. London: The
734	Author. 230 pp.
735	Stewart RJ, Weaver JC, Morse DE, Waite JH. 2004. The tube cement
736	of Phragmatopoma californica: a solid foam. Journal of Experimental Biology 207: 4727-
737	4734. DOI: 10.1242/jeb.01330.
738	Tanur AE, Gunari N, Sullan RMA, Kavanagh CJ, Walker GC. 2010. Insights into the
739	composition, morphology, and formation of the calcareous shell of the serpulid Hydroides
740	dianthus. Journal of Structural Biology 169: 145–160. DOI: 10.1016/j.jsb.2009.09.008.
741	Tarkowski R, Thierry J, Marchand D, Mangold C, Delance JH, Garcia JP, Laurin B.
742	1994. L'"oolithe de Balin" (Pologne méridionale). Nouvelles observations et interprétations
743	stratigraphiques. The "Balin Oolithe" (Southern Poland). New observations and
744	stratigraphical interpretations. Geobios 27: 461-467.
745	Taylor PD. 2016. Competition between encrusters on marine hard substrates and its
746	fossil record. <i>Palaeontology</i> 59: 481–497. DOI: 10.1111/pala.12239.
747	Taylor PD, Vinn O, Kudryavtsev A, Schopf JW. 2010. Raman spectroscopic study of
748	the mineral composition of cirratulid tubes (Annelida, Polychaeta). Journal of Structural
749	Biology 171: 402–405. DOI: 10.1016/j.jsb.2010.05.010.
750	Thomas, J. G. 1940. Pomatoceros, Sabella and Amphitrite. (No Title).
751	Vermeij GJ. 1977. The Mesozoic marine revolution: evidence from snails, predators
752	and grazers. Paleobiology 3: 245–258.
753	Vinn O. 2007. Taxonomic implications and fossilization of tube ultrastructure of some
754	Cenozoic serpulids (Annelida, Polychaeta) from Europe. Neues Jahrbuch für Geologie und
755	Paläontologie Abhandlungen 244: 115–128. DOI: 10.1127/0077-7749/2007/0244-0115.
756	Vinn O. 2008. Tube ultrastructure of the fossil genus Rotularia Defrance, 1827
757	(Polychaeta, Serpulidae). Journal of Paleontology 82: 206–212. DOI: 10.1666/06-125.1.
758	Vinn O. 2009. The ultrastructure of calcareous cirratulid (Polychaeta, Annelida) tubes.
759	Estonian Journal of Earth Sciences 58: 153-156. DOI: 10.3176/earth.2009.2.06.
760	Vinn O. 2013. SEM Study of semi-oriented tube microstructures of Serpulidae
761	(Polychaeta, Annelida): implications for the evolution of complex oriented microstructures.
762	Microscopy Research and Technique 76: 453-456. DOI: 10.1002/jemt.22186.
763	Vinn O. 2020. Biomineralization of Polychaete Annelids in the Fossil Record.
764	Minerals 10: 858. DOI: 10.3390/min 10100858.
765	Vinn O. 2021a. Biomineralization in Polychaete Annelids: A Review. <i>Minerals</i> 11:

766 1151. DOI: 10.3390/min11101151.

769 Vinn O, Furrer H. 2008. Tube structure and ultrastructure of serpulids from the Jurassic of France and Switzerland, its evolutionary implications. Neues Jahrbuch für 770 Geologie und Paläontologie, Abhandlungen 250: 129-135. DOI: 10.1127/0077-771 7749/2008/0250-0129. 772 Vinn O, ten Hove HA, Mutvei H. 2008. On the tube ultrastructure and origin of 773 calcification in sabellids (Annelida, Polychaeta). Palaeontology 51: 295–301. DOI: 774 10.1111/j.1475-4983.2008.00763.x. 775 776 Vinn O, ten Hove HA, Mutvei H, Kirsimäe K. 2008. Ultrastructure and mineral composition of serpulid tubes (Polychaeta, Annelida). Zoological Journal of the Linnean 777 Society 154: 633-650. DOI: 10.1111/j.1096-3642.2008.00421.x. 778 Vinn O, Jäger M, Kirsimäe K. 2008. Microscopic evidence of serpulid affinities of the 779 780 problematic fossil tube "Serpula" etalensis from the Lower Jurassic of Germany. Lethaia 41: 417-421. DOI: 10.1111/j.1502-3931.2008.00093.x. 781 Vinn O, Jäger M, Słowiński J, Zatoń M. 2023. Convergent evolution of encrusting 782 calcareous tubeworms. *Palaeoworld* (in press). DOI: 10.1016/j.palwor.2023.04.001. 783 Vinn O, Kirsimäe K, ten Hove HA. 2009. Tube ultrastructure of *Pomatoceros* 784 americanus (Polychaeta, Serpulidae): Implications for the tube formation of serpulids. 785 Estonian Journal of Earth Sciences 58: 148-152. DOI: 10.3176/earth.2009.2.05. 786 787 Vinn O, Kupriyanova EK. 2011. Evolution of a dense outer protective tube layer in 788 serpulids (Polychaeta, Annelida). Carnets de Géologie-Notebooks on Geology Letter 789 CG2011/05 (CG2011\_ L05) 137-147. DOI: 10.4267/2042/43896. 790 Vinn O, Luque J. 2013. First record of a pectinariid-like (Polychaeta, Annelida) agglutinated worm tube from the Late Cretaceous of Colombia. Cretaceous Research 41: 791 107-110. DOI: 10.1016/j.cretres.2012.11.004. 792 Weedon MJ. 1994. Tube microstructure of Recent and Jurassic serpulid polychaetes 793 and the question of the Palaeozoic 'spirorbids'. Acta Palaeontologica Polonica 39: 1-15. 794 Zatoń M, Kremer B, Marynowski L, Wilson MA, Krawczyński W. 2012. Middle 795 796 Jurassic (Bathonian) encrusted oncoids from the Polish Jura, southern Poland. Facies 58: 57-77. DOI: 10.1007/s10347-011-0273-1. 797

Zatoń M, Machalski M. 2013. Oyster-microbial rolling stones from the Upper Jurassic

(Kimmeridgian) of Poland. Palaios 28: 839-850. DOI: 10.2110/palo.2013.025.

Vinn O. 2021b. The Role of Aragonite in Producing the Microstructural Diversity of

Serpulid Skeletons. *Minerals* 11: 1435. DOI: 10.3390/min11121435.

767

768

798

800 Zatoń M, Machocka S, Wilson MA, Marynowski L, Taylor PD. 2011. Origin and paleoecology of Middle Jurassic hiatus concretions from Poland. Facies 57: 275–300. DOI: 801 10.1007/s10347-010-0244-y. 802 Zatoń M, Marynowski L, Bzowska G. 2006. Konkrecje hiatusowe z iłów rudonośnych 803 804 Wyżyny Krakowsko-Częstochowskiej. Przegląd Geologiczny 54: 131–138. Zatoń M, Taylor PD. 2009a. Microconchids (Tentaculita) from the Middle Jurassic of 805 Poland. Bulletin of Geosciences 84: 653-660. DOI: 10.3140/bull.geosci.1167. 806 807 Zatoń M, Taylor PD. 2009b. Middle Jurassic cyclostome bryozoans from the Polish Jura. Acta Palaeontologica Polonica 54: 267–288. DOI: 10.4202/app.2008.0088. 808 809 Zatoń M, Wilson MA, Zavar E. 2011. Diverse sclerozoan assemblages encrusting large bivalve shells from the Callovian (Middle Jurassic) of southern Poland. 810 Palaeogeography, Palaeoclimatology, Palaeoecology 307: 232–244. DOI: 811 10.1016/j.palaeo.2011.05.022. 812 813 Zhao H, Sun C, Stewart RJ, Waite JH. 2005. Cement proteins of the tube-building polychaete Phragmatopoma californica. Journal of Biological Chemistry 280: 42938-42944. 814 DOI: 10.1074/jbc. 815 816 Zibrowius H, ten Hove HA. 1987. Neovermilia falcigera (Roule, 1898) a deep- and cold-water serpulid polychaete common in the Mediterranean Plio-Pleistocene. Bulletin of 817

Biological Society of Washington 7: 259–271.