

New phiomorph rodents from the latest Eocene of Egypt, and the impact of Bayesian “clock”-based phylogenetic methods on estimates of basal hystricognath relationships and biochronology

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The Fayum Depression of Egypt has yielded fossils of hystricognathous rodents from multiple Eocene and Oligocene horizons that range in age from ~37 to ~30 Ma and document several phases in the early evolution of crown Hystricognathi and one of its major subclades, Phiomorpha. Here we describe two new genera and species of basal phiomorphs, *Birkamys korai* and *Mubhammys vadumensis*, based on rostra and maxillary and mandibular remains from the terminal Eocene (~34 Ma) Fayum Locality 41 (L-41). *Birkamys* is the smallest known Paleogene hystricognath, has very simple molars, and, like derived Oligocene-to-Recent phiomorphs (but unlike contemporaneous and older taxa) apparently retained $dP^4/4$ late into life, with no evidence for $P^4/4$ eruption or formation. *Mubhammys* is very similar in dental morphology to *Birkamys*, and also shows no evidence for $P^4/4$ formation or eruption, but is considerably larger. Phylogenetic analysis places *Birkamys* and *Mubhammys* either as sister taxa of extant *Thryonomys* to the exclusion of much younger relatives of that genus (using parsimony), or in much more basal positions within Hystricognathi, as sister taxa of Oligocene-to-Recent phiomorphs (using standard Bayesian and “tip-dating” approaches). Tip-dating recovers placements of *Birkamys* and *Mubhammys* that, when compared with parsimony-based trees, significantly reduce ghost lineages and are more consistent with the temporal appearance of phiomorph dental apomorphies in the African fossil record. We also employ tip-dating as a means for estimating the ages of early hystricognath-bearing localities, many of which are not well-constrained by geological, geochronological, or biostratigraphic evidence. By simultaneously taking into account phylogeny, evolutionary rates, and an underlying model of speciation and extinction, dating of tips with the fossilized birth-death prior allows paleontologists to move beyond vague and assumption-laden “stage of evolution” arguments in biochronology to provide relatively rigorous age assessments of poorly-constrained faunas. This approach should become increasingly robust as estimates are combined from multiple independent analyses of distantly related clades, is broadly applicable across the tree of life, and as such is deserving of paleontologists’ close

attention. Notably, in the example provided here, rodents from Libya and Namibia that are controversially considered to be of middle Eocene age are instead estimated to be of late Eocene and late Oligocene age, respectively. Finally, we reconstruct the evolution of first lower molar size among Paleogene African hystricognaths using a Bayesian approach; the results of this analysis reconstruct a rapid latest Eocene dwarfing event along the lineage leading to *Birkamys*.

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INTRODUCTION

The rodent clade Hystricognathi first appeared in the Eocene, and is now represented by three major groups with extant members — Hystricidae (Old World porcupines), Caviomorpha (New World hystricognaths), and Phiomorpha (African cane, dassie, and mole rats) (Singleton et al. 2006). The largest DNA datasets currently available place Hystricidae as the sister group of a Caviomorpha-Phiomorpha clade (e.g., Huchon et al., 2007; Meredith et al., 2011; Patterson & Upham, 2014). Despite their modern distribution, being restricted almost entirely to southern continents, phylogenetic evidence provided by later Paleogene Asian “baluchimyines” suggests that the stem lineage of Hystricognathi probably arose in Asia (Marivaux et al., 2004; Sallam et al., 2009), though no members of the group are definitively known from that continent before the Eocene-Oligocene boundary (~34 Ma: Marivaux et al., 2000; de Bruijn et al. 2003). In contrast, recent paleontological work in Tunisia and Peru has revealed that hystricognaths were present in Africa by ~39.5 Ma (Marivaux et al., 2014), and that caviomorphs were present in South America by ~41 Ma (Antoine et al., 2011)¹. The latter discovery is critically important for establishing that the common ancestor of Caviomorpha and Phiomorpha must be even older than 41 Ma, and that stem members of Phiomorpha were already diversifying at least four million years prior to the deposition of the earliest well-sampled hystricognathous rodent fauna from Africa, the ~37 Ma Birket Qarun Locality 2 in the Fayum Depression of Egypt (Sallam et al., 2009).

¹ We exclude from our discussion two molars from the Silica North and Silica South localities in the Sperrgebiet area of Namibia that were assigned by Pickford *et al.* (2008) to the otherwise late Eocene genus *Protophiomys*. Pickford *et al.* (2014) argue that Silica North and Silica South are Bartonian (~38-41.3 Ma; Gradstein et al., 2012) in age, but the biochronological evidence that we present later in this contribution suggests a much younger age (late Oligocene). In the absence of much more complete material, we must view the great age proposed for the Sperrgebiet *Protophiomys* specimens with skepticism.

A recent study of molecular divergence estimates that took into account much of this new fossil evidence (Patterson & Upham, 2014) placed the caviomorph-phiomorph split at ~42 Ma, the divergence of that clade from Hystricidae at ~44.9 Ma, and the origin of crown Phiomorpha at ~36.3 Ma. These estimates would suggest that the origin and initial diversification of crown Hystricognathi is not yet documented in the fossil record of any landmass, but that the origin of crown Phiomorpha should have occurred very close in time to the deposition of Locality BQ-2. Despite this, species that are known from ~39 to ~37 Ma African sites [Djebel el Kébar in Tunisia (Marivaux et al., 2014), Bir el Ater in Algeria (Jaeger et al., 1985), and BQ-2 in Egypt (Sallam et al., 2009)] — i.e., from a time period that would, given Patterson & Upham's divergence estimates, postdate the caviomorph-phiomorph split by 3-5 Ma, and the origin of crown Hystricognathi by 6-8 Ma — have not been placed as stem phiomorphs in previous phylogenetic analyses (Sallam et al., 2009, 2011, 2012; Coster et al., 2010; Antoine et al., 2014), but instead are consistently placed as stem hystricognaths, or as stem members of the Caviomorpha-Phiomorpha clade. If these results are correct, then stem phiomorphs simply have not yet been sampled in the middle Eocene and early late Eocene sites of northern Africa.

One possible explanation for this incongruency is that early phiomorphs have not yet been sampled due to a geographic bias, because all of the key sites documenting early hystricognath evolution in Africa are from the northernmost part of the continent. A reasonable alternative hypothesis, given the surprising discovery of ~41 Ma caviomorphs and the poor early African record of this group, is that phylogenetic signal has been obscured by homoplasy between basal caviomorphs and more advanced stem phiomorphs, and some or all of the earliest African hystricognaths are actually basal stem phiomorphs that retain primitive morphology similar to that of the caviomorph-phiomorph ancestor. The possibility of early homoplasy between

caviomorphs, phiomorphs, and the Asian “baluchimyine” radiation must be seriously entertained, because at present phylogenetic analyses of basal hystricognaths depend almost entirely on dental characters, many of which are known to have undergone remarkably rapid evolution in some early hystricognath lineages (notably Gaudeamuridae; Sallam et al., 2011). Compounding this problem is the fact that any phylogenetic arrangement of basal hystricognaths implies middle Eocene colonizations of large and rodent-less (in the case of South America) or hystricognath-less (in the case of Afro-Arabia) continents, both of which might have spurred rapid early diversification (and potentially rapid morphological change) associated with filling of open niche space.

Here we describe two new phiomorph genera and species from the latest Eocene Quarry L-41, in the Fayum area of northern Egypt (Fig. 1), that are the oldest to show suppression and non-formation of $P^4/4$, one of the key dental synapomorphies of crown Phiomorpha. We include these and other basal African, Asian, and South American hystricognaths in a series of parsimony and Bayesian phylogenetic analyses, including Bayesian “tip-dating” analyses (Beck and Lee, 2014; Close et al., 2015; Dembo et al., 2015; Ronquist et al., 2012) that are able to take into account information about the ages of fossil taxa, rates of morphological evolution, and models of speciation and extinction, and as such are potentially ideally suited to test relationships given the challenging circumstances presented by basal hystricognaths.

Aside from the tip-dating method’s obvious utility for phylogenetic reconstruction, we note that for species whose temporal ranges are poorly constrained by geological data, tip-dating takes into account both phylogenetic position and rate of morphological evolution to provide age estimates for those species, and this information provides a relatively rigorous testable hypothesis for the ordering of hystricognath-bearing faunas of Eocene and Oligocene age in

Africa and Asia (i.e., DT1 and DT2 (Dur at-Talah, Libya), Lokone in Kenya, Silica North in Namibia, Paali Nala C2 and contemporaneous sites in the lower part of the Chitarwata Formation, Pakistan) that was not previously possible.

MATERIAL AND METHODS

Taxonomy

The electronic version of this article in Portable Document Format (PDF) will represent a published work according to the International Commission on Zoological Nomenclature (ICZN), and hence the new names contained in the electronic version are effectively published under that Code from the electronic edition alone. This published work and the nomenclatural acts it contains have been registered in ZooBank, the online registration system for the ICZN. The ZooBank LSIDs (Life Science Identifiers) can be resolved and the associated information viewed through any standard web browser by appending the LSID to the prefix <http://zoobank.org/>. The LSID for this publication is: urn:lsid:zoobank.org:pub:9DB0476B-E752-4EA1-8745-8C92E429C65B. The online version of this work is archived and available from the following digital repositories: PeerJ, PubMed Central and CLOCKSS.

Terminology, measurements, and CT-scanning

Dental terminology follows Marivaux et al. (2004) (Fig. 2). Teeth are referred to as I, P, and M (for incisors, premolars, and molars, respectively), with upper and lower teeth designated by superscript and subscript numbers (respectively) for locus (e.g., the second lower molar is referred to as M₂). Dental measurements were taken with a micrometer mounted in the lens of a

Meiji binocular microscope. Specimens were scanned using a Nikon XT H 225 ST micro-CT scanner housed at at Duke University's Shared Materials Instrumentation Facility and three-dimensional reconstructions were rendered in Avizo v. 8. Digital surface models (in Stanford "ply" format) of all specimens described here are available for viewing and direct download at www.morphosource.org.

Some of the Fayum rodent species described by Wood (1968) have been revised by Holroyd (1994) as a part of her Ph.D. dissertation; taxonomic names that she considered to be invalid or incorrect are placed in quotation marks pending formal revision. Fossils are housed at the Egyptian Geological Museum (CGM) and Duke Lemur Center Division of Fossil Primates (DPC); a collection of casts is also housed in the Mansoura University Vertebrate Paleontology Center (MUVF) cast collection.

Phylogenetic Analysis

Matrix. The matrix employed here is that of Sallam et al. (2012), which was built first on the original matrix of Marivaux et al. (2004), and was then modified by Sallam et al. (2009, 2011, 2012). The matrix contains 118 characters, mostly from the dentition, of which 77 were treated as ordered in all analyses; 97 of the characters are parsimony informative. Three additional early African species were added: "*Protophiomys*" *tunisiensis* from the late middle Eocene (Bartonian) of Jebel el Kébar, Tunisia (Marivaux et al., 2014); *Turkanamys hexalophus*, from the Oligocene Lokone Hill sites in the Turkana Basin, northern Kenya (Marivaux et al., 2012); and *Prepomonomys bogenfelsi*, from the Silica North site in the Sperrgebiet area of Namibia (Pickford et al., 2008), which is of contentious age, either Bartonian (Pickford et al., 2014) or

significantly younger (Coster et al., 2012a; Marivaux et al., 2014). In all analyses, the early middle Eocene “chapattimysid” *Birbalomys* was designated as the outgroup.

Parsimony analyses (see Dataset S1) were run in PAUP 4.0b10 (Swofford, 1998) using the heuristic search algorithm, random addition sequence, and tree bisection-and-reconnection branch swapping across 10,000 replicates. Characters were not weighted. Bootstrap support was also calculated in PAUP, based on 1,000 pseudoreplicates.

Bayesian phylogenetic analyses (see Dataset S2) were run in MrBayes 3.2.5 (Ronquist et al., 2012). The M_k model for morphological data (Lewis, 2001) was used, coding was set to “variable”, and gamma-distributed rate variation across characters was assumed. Markov Chain Monte Carlo (MCMC) chains were run for 25 million generations, with two independent runs, each with one cold chain and three heated chains (temp=0.02), sampling every 1000 generations. The first 25% of the resulting 25,000 samples were discarded as the “burn-in” period, and the remaining trees were summarized using an “allcompat” (majority-rule plus compatible groups) consensus tree. Convergence was assessed by checking both effective sample sizes and the average standard deviation of split frequencies in the final generation.

Bayesian “tip-dating” analyses (see Dataset S3) were also run in MrBayes 3.2.5 (Ronquist et al., 2012). We employed the IGR (independent gamma rates) relaxed clock model and the fossilized birth-death prior on branch lengths, with “samplestrat” set to “fossiltip” (indicating that tips left no descendants). We ran several analyses with various perturbations of the priors “speciationpr” (the prior on the net speciation rate), “igrvarpr” (the prior on the variance of the gamma distribution from which branch lengths are drawn), and “clockratepr” (the prior on the base substitution rate, measured in number of changes per site per Ma) (Huelsenbeck et al., 2015), all

of which yielded remarkably similar “allcompat” topologies, divergence dates, tip dates, and support values — however many of these analyses did not show adequate evidence for convergence across all parameters, as judged by effective sample sizes and potential scale reduction factors. Ultimately the combination of priors that yielded the strongest evidence for convergence across all parameters was $\text{speciationpr}=\text{exp}(50)$, $\text{clockratepr}=\text{normal}(0.25,0.05)$, and $\text{igrvarpr}=\text{exp}(3)$, with flat beta priors on fossilizationpr and extinctionpr , and we present the results from that analysis. “Sampleprob” (the percentage of extant species sampled in the analysis) was set to 0.005, as only extant *Thryonomys* (African cane rat) was sampled from the entire sample of extant hystricognaths. The root node was constrained to fall within a uniform prior on node age from 47.8 Ma (the oldest possible age of the species in the matrix) to 55 Ma (beyond which no ctenohystricans, or even demonstrable crown rodents, have been found in the fossil record; e.g., Marivaux et al., 2004). Two analyses were run: in the first analysis (referred to as TD1), each tip was calibrated with a uniform prior on age, employing minimum and maximum estimates based on the currently accepted upper and lower bounds of magnetostratigraphic or geological stages or ages to which fossils have been assigned (i.e., in MrBayes, $\text{calibrate taxon}=\text{uniform}(\text{minimum age, maximum age})$; see Appendix S1, which provides justification for the minimum and maximum ages for each taxon). Fourteen of the species in the analysis are from the Fayum succession, and we followed the magnetostratigraphic correlation of the Fayum beds to the Geomagnetic Polarity Timescale that was preferred by Seiffert (2006) and Seiffert et al. (2008). In addition to topology and attendant support and parameter estimates, this first analysis importantly also output point age estimates for each tip species from within its uniform prior, taking into account the the base clockrate and the amount of change expected along the terminal branch leading to the tip. However as would be expected given such parameters, the

point age estimates for species from a single locality were not all the same, as most are assumed to be given that they are from the same stratum or tightly constrained interval (also assuming that time-averaging in an accumulation is negligible). Because tips from the same locality should ideally “line up” so that branch lengths are not artificially long or short (thereby implying artificially slow or fast rates of evolution), a second analysis (TD2) was run with the point age estimates for species from each locality (i.e., the estimates calculated in TD1) averaged and used as fixed dates [i.e., in MrBayes, calibrate taxon=fixed(mean age for locality based on TD1 estimates)]. For both analyses, the MCMC chains were run for 50 million generations, with two independent runs, each with one cold chain and three heated chains (temp=0.01), sampling every 1000 generations. The first 25% of the resulting 50,000 samples were discarded as the “burn-in” period, and the remaining trees were summarized using an “allcompat” (majority-rule plus compatible groups) consensus tree.

Bayesian ancestral reconstruction of first lower molar size in early hystricognaths

We collected length and width measurements on the first lower molars of early hystricognaths in our character-taxon matrix, either directly (in the case of Fayum species) or from published data (in the case of species for which we only had casts), with the goal of reconstructing the evolution of first lower molar area (natural log; see Dataset S4) within a Bayesian context using the *Continuous* module in BayesTraits v. 2 (Pagel, 2002; Pagel & Meade, 2013). Using the “allcompat” consensus derived from the tip-dating analysis (TD2) described above as the input tree, we first ran maximum likelihood analyses of the data set under random walk and directional models, with and without the phylogenetic scaling parameters delta, kappa, and lambda, to determine which model/parameter combination had the highest log likelihood. The random walk

model with the lambda scaling parameter was effectively indistinguishable from the directional model with the lambda scaling parameter, so we present results from both of these analyses. Model files were first created by running MCMC chains for 10,050,000 generations, with the first 50,000 discarded as burn-in. These model files were then employed in longer (20,050,000 generations, first 50,000 discarded as burn-in) MCMC chains for ancestral state reconstructions, in which ancestral values were estimated for all internal nodes in the tree.

RESULTS

Systematic Paleontology

Class MAMMALIA Linnaeus, 1758

Order RODENTIA Bowdich, 1821

Infraorder HYSTRICOGNATHI Tullberg, 1899

Parvorder PHIOMORPHA Lavocat, 1962

Family *Incertae sedis*

***Birkamys*, new genus** (Figs. 3-5, Table 1) urn:lsid:zoobank.org:act:[in process]

Type and only known species

Birkamys korai, new species urn:lsid:zoobank.org:act:[in process]

Etymology

Combination of *birka*, Arabic word for lake or swamp, in reference to the L-41 deposits and *mys*, Greek for mouse.

214 *Diagnosis*

215 As for the type and only known species.

216 ***Birkamys korai*, new species** urn:lsid:zoobank.org:act:[in process]

217 (Figs. 3-5, Table 1)

218 *Etymology*

219 In honor of Professor Mahmoud Kora of Mansoura University, for his important contributions to
220 the study of stratigraphy and paleontology in Egypt.

221 *Holotype*

222 CGM 66000, rostrum with right and left upper incisors and dP³-M³ (Fig. 3; measurements in
223 Table 1).

224 *Referred specimens*

225 DPC 9276, left maxilla with M² and M³ (Fig. 4E-H); DPC 15625, left maxilla with M¹ (Fig. 4J-
226 L); DPC 17457, right maxilla with dP³-M³; DPC 22737, right mandible with dP₄-M₃ (Fig. 5).

227 *Type locality*

228 Locality 41 (L-41), 46 meters above the base of the lower sequence of the Jebel Qatrani
229 Formation. The fine green claystone at L-41 contains 12% postdepositional sodium chloride and
230 is unique among Fayum fossil quarries, most of which occur in sands and gravels. The sediments
231 at L-41 might have been laid down in a freshwater lake that was periodically flooded, resulting
232 in large accumulations of vertebrate carcasses (Simons et al. 1998). Over the last three decades,
233 work at L-41 has produced a wide variety of fish, amphibian, reptile, bird and mammal taxa.
234 There is no clear sorting of fossil mammals on the basis of size, and the locality contains not

only large-bodied hyracoids (Rasmussen and Gutiérrez 2010) and anthracotheriid artiodactyls, but also very small primates (Simons 1997; Simons et al. 2001; Seiffert et al. 2005), macroscelideans (Simons et al. 1991), tenrecoids (Seiffert and Simons 2000; Seiffert et al. 2007), bats (Gunnell et al. 2008), and rodents (Holroyd, 1994; Lewis and Simons, 2006; Sallam et al. 2011, 2012). Hundreds of rodent specimens are known from L-41, but the only clade represented is Hystricognathi, whereas both Hystricognathi and Anomaluroidea occur at the ~37 Ma Locality BQ-2 (Sallam et al., 2009, 2010a, b).

Age and Formation

Latest part of late Eocene (latest Priabonian, ~34 Ma), lower sequence of Jebel Qatrani Formation, northern Egypt.

Diagnosis

Birkamys korai differs from early Oligocene *Phiomys andrewsi* in having smaller molars; in retaining deciduous premolars; in having relatively short metalophulid IIs on dP₄-M₃; in having a relatively small M₃ when compared to M₂; and in lacking a mesostyle and mesolophule on the upper molars. Differs from contemporaneous and sympatric *Acritophiomys bowni* in having smaller teeth; in showing no evidence for replacement of deciduous premolars; in having a relatively weak anterior cingulid, no metalophulid I or II, and no mesolophid or mesostylid on dP₄; in having no hint of an M₁₋₂ anteroconid; in consistently lacking M₁₋₂ mesostylids and mesolophids, and having relatively short metalophulid IIs; in having relatively large M₁₋₂ protoconids; in having a relatively small M₃ when compared to M₂; in lacking a dP₄ mesolophule; in consistently lacking M¹⁻² mesostyles, mesolophules, and pericingula; in having relatively large M¹⁻² metaconules; and in lacking enamel wrinkling and crenulation. Differs from

257 Oligocene *Turkanamys hexalophus* from Kenya in having smaller molars; in showing no
 258 evidence for replacement of deciduous premolars; in having relatively well-developed anterior
 259 cingulids on M_{1-2} ; in having a small, rather than large, metastylid on M_{1-2} ; in lacking mesostylids
 260 and well-developed metalophulid IIs on M_{1-2} ; in lacking a connection of the entoconid and
 261 hypoconid on the M_{1-2} ; in having a relatively small M_3 when compared to M_2 ; in lacking M^{1-2}
 262 mesostyles and mesolophules; and in lacking enamel wrinkling and crenulation. Differs from
 263 "*Phiomys*" *hammadai* from the late Eocene of Libya in having smaller molars; in showing no
 264 evidence for replacement of deciduous premolars; in lacking a dP_4 mesolophid, mesostylid,
 265 metalophulid I, and metalophulid II; in having a more distinct anterocingulid but lacking a
 266 mesostylid, mesolophid, and a well-developed metalophulid II on M_{1-2} ; in having a relatively
 267 small M_3 when compared to M_2 ; in lacking M^{1-2} mesostyles and mesolophules, and having
 268 relatively large metaconules; and in lacking enamel wrinkling and crenulation. Differs from early
 269 Oligocene *Neophiomys paraphiomyoides* from Egypt and Libya in lacking a distinct
 270 metalophulid II and having a complete ectolophid on dP_4 ; in having a relatively small M_3 when
 271 compared to M_2 ; in lacking M^{1-2} mesostyles, mesolophules, and posterior arms of paracones; and
 272 in having relatively large M^{1-2} metaconules. Differs from early Oligocene "*Phiomys*" *lavocati*
 273 from younger quarries in the Fayum succession in showing no evidence for replacement of
 274 deciduous premolars; in having a dP_4 protoconid that is more mesially placed relative to the
 275 metaconid; in having M_{1-2} protoconids that are relatively large when compared with metaconids;
 276 in having a dP^4 metaloph that is connected to the metaconule, rather than distally oriented; in
 277 having a dP^4 mure; in having M^{1-2} mures and metaconules that are submerged into the mures; in
 278 having anterior arms of the M^{1-2} hypocones that are relatively well-developed; in having an M^1
 279 metaloph that is connected to both the metaconule and posteroloph; and in lacking posterior arms

of M^{1-2} paracones. Differs from late Eocene *Talahphiomys lavocati* from Libya in having a dP_4 protoconid that is more mesially placed relative to the metaconid; in lacking a dP_4 mesostylid; in having a more distinct M_{1-2} anterocingulid; in having a dP^4 metaloph that is connected to the metaconule, rather than distally oriented; in having a dP^4 mure; in lacking an M^{1-2} mesostyle; in having M^{1-2} mures and metaconules that are submerged into the mures; and in having an M^2 metaloph that is connected both to the metaconule and the posteroloph. Differs from late Eocene *Talahphiomys libycus* from Libya in having a dP^4 metaloph that is connected to the metaconule, rather than distally oriented; in having a dP^4 mure; in lacking an M^{1-2} mesostyle and mesolophule; in having an M^2 metaloph that is connected both to the metaconule and the posteroloph; in having M^{1-2} mures; and in lacking posterior arms of the M^{1-2} paracones and anterior arms of the M^{1-2} metacones.

Description

Four crushed cranial elements of *Birkamys korai* (Figs. 3 and 4) together document much of the morphology of the rostrum, mid-cranium, and the complete upper dentition. The holotype rostrum CGM 66000 (Fig. 3) was subjected to severe post-mortem distortion that has led the specimen to be dorsoventrally flattened with numerous surface cracks and displacements; rather than attempt physical preparation of this tiny and very fragile specimen, we figure as much as is possible through volume rendering of the encasing block using high-resolution micro-CT scans (Fig. 3). CGM 66000 contains the premaxillae with two upper incisors, both maxillae with the entire dentition (dP^3 - M^3), and most of the frontal. The cranial parts in the hypodigm represent adult individuals, two of which preserve third and fourth deciduous premolars that are worn.

302 The premaxillae are preserved in the holotype, house the two upper incisors, and form most of
 303 the mediolaterally narrow and anteroposteriorly elongate rostrum and upper diastema (Fig. 3).
 304 The most striking feature of the rostrum is the capacious vacuity, referred to be some as an
 305 “anterior palatine fenestra” (e.g., Wood 1968), formed by the anteroposteriorly elongate and
 306 confluent incisive foramina, the anterior halves of which deeply excavate, and are formed by, the
 307 premaxillae. The posterior border of the incisive foramen extends posteriorly between dP³ and
 308 dP⁴. The suture between the premaxilla and the maxilla is well-preserved. *Birkamys* was
 309 hystricomorphous; that is, the infraorbital foramen was very large and presumably allowed for
 310 the passage of a greatly expanded medial masseter that inserted on the side of the rostrum,
 311 anterior to the zygomatic arch. The size and shape of the ventral margin of the infraorbital
 312 foramen is most clearly preserved on DPC 9276 (Fig. 4F). On the ventral surface of the maxilla,
 313 a small masseteric tubercle is situated immediately ventral to the infraorbital foramen and
 314 anterolateral to the alveolus of dP³, providing a point of origin (along with the zygomatic arch,
 315 which extends laterally at the level of the alveolus of dP³ and masseteric tubercle) for the
 316 superficial masseter. On the dorsal view of DPC 9276, the infraorbital fissure is relatively broad
 317 and deepens anteriorly, separating the orbital process from the alveolar portion (Fig. 4G). The
 318 alveolar foramen is oval in shape and lies within the medial wall of the infraorbital fissure, dorsal
 319 to the dP³ alveolus. The anterior portion of the maxilla protrudes roughly at the same level as the
 320 alveolar plane and preserves part of the articulation with the premaxilla. The palate is preserved
 321 in the holotype and is somewhat flat, slightly lower than the alveolar plane, and broad throughout
 322 its length. It houses the two major palatine foramina, which are relatively round and large, and
 323 which lie at the level of the first upper molar.

324 The upper deciduous third premolar (dP³) is generally a small peg-shaped tooth with a rounded
325 base, and abuts the mesial surface of dP⁴. The dP³, on both the left and right sides, is well-
326 preserved in the holotype and DPC 17457 (Figs. 3F and 4D). It has one large cusp that occupies
327 the distal portion of the crown and forms the apex of the tooth. There is a small depression on the
328 distolabial side of the former cusp.

329 The upper deciduous fourth premolar (dP⁴) is roughly quadrate in shape and bears four major
330 cusps (paracone, metacone, protocone and hypocone) as well as a small metaconule (Figs. 3F
331 and 4D). The paracone is of approximately the same size and height as the metacone, and the
332 hypocone is situated at the same level as the protocone. The occlusal configuration of the crown
333 is tetralophodont, with no mesolophule. The anteroloph that forms the anterior border of the
334 tooth is low and runs labially from the protocone, parallel to the protoloph, and fuses with a
335 weakly-developed parastyle just mesial to the base of the paracone. A well-developed and
336 transversely oriented protoloph joins the paracone and protocone. There is a small inflation at the
337 junction of the protocone and anteroloph that might ~~represent an incipient~~ anterostyle. The
338 metaloph takes a sinuous course, running lingually and mesially from the metacone, connecting
339 to the metaconule. There is a weak connection with the posteroloph at its lingual portion, which
340 together along with the hypocone and its anterior arm form a small fovea at the distolingual
341 corner of the crown. The posteroloph runs labially from the hypocone, courses around the
342 posterior margin of the tooth and connects to the distal base of the metacone. A weak and poorly
343 defined metaconule is centrally placed and connected to the hypocone via the latter cusp's
344 anterior arm, which is robust and oriented mesiolabially. The mure is complete, connecting the
345 metaconule with the protoloph. The labial wall bears a deep notch that extends to the level of the
346 central basin, and there is a very weakly developed knob that could be a remnant of the

mesostyle. The tooth lacks an endoloph, and the lingual sinus is wide, deep, and is not continuous with the central basin due to the presence of the mure.

The upper first molar (M^1) is the largest tooth of the upper dentition (Figs. 3F, 4D, 4L). Its occlusal pattern is nearly identical to that of dp^4 , but differs in having relatively tall lophs and cusps that are completely integrated into the four primary crests (anteroloph, protoloph, metaloph, and posteroloph) and in having a relatively tall paracone when compared to the metacone. The M^2 occlusal surface is similar to that of M^1 , differing only in being relatively shorter and broader. The hypocone is more labially situated with respect to the protocone.



The upper third molar (M^3) is smaller than M^{1-2} and has a relatively short lingual margin, leading to a somewhat oval outline (Figs. 3F, 4D, 4H). The tooth bears a reduced metacone and hypocone, which are relatively lingually and labially positioned, respectively, when compared with the same cusps on $M^{1,2}$. The metaloph is weakly developed and connects the metacone and the anterior arm of the hypocone. The central basin is closed by a weakly developed labial wall, and is closed lingually by a weak to well-developed mure that reaches the protoloph. On DPC 17457, the posterior part of the protocone extends distolabially toward the base of the metaconule, forming a high and continuous endoloph. Together with the mure, the protoloph and endoloph delimit a small fovea. On DPC 9276, there is a short and low anterior arm of the hypocone that is connected to the base of the metaconule. The latter is relatively well-developed and more distal in position when compared with the same cusp on M^2 . There is a small crest running longitudinally from the base of the posteroloph to connect with the metaconule distally. This crest contributes to a small fovea that is also bordered by the metaconule, the hypocone and its anterior arm, and the lingual part of the posteroloph. The posteroloph is relatively well-


developed with respect to that of M^2 , and courses around the posterior margin of the tooth, running labially from the hypocone to form a strong connection with the metacone.

The mandible (Fig. 5A-D) is slender, with a partially preserved ascending ramus and a well-preserved corpus. The angular process is placed lateral to the tooth row and the incisor, leaving a well-developed groove between the angular process and the incisor alveolus; the mandible is thus fully hystricognathous. On the lateral surface of the mandible, the ventral masseteric ridge extends laterally, arising near the midpoint of the horizontal ramus and continuing posteriorly toward the angular process; this ridge, which is an insertion site for the deep masseter muscle, extends anteriorly as a part of the masseteric fossa, and terminates beneath the posterior portion of the dP_4 at roughly the same level as the mental foramen. The dorsal masseteric ridge is less developed anteriorly and extends posteriorly with the coronoid process. The tip of the coronoid process is broken but it is apparently higher than the condylar process; it arises lateral to the third molar, leaving a deep fossa. The mental foramen is somewhat oval and can be seen in dorsal view, lying anterior to the anterior root of the dP_4 and ventral to the distal portion of the diastema. The diastema is deep and makes up about half the length of the horizontal ramus. The posterior part of the ascending ramus is extremely fragile and the tip of the angular process is not preserved. On the medial surface of the mandible, the mandibular foramen is situated in the area between the coronoid and condylar processes, on the dorsal margin of a strut that extends posteriorly from the rear part of the incisor alveolus. The condylar process is slightly higher than the tooth row and the tip of the incisor (which are roughly at the same level). The lower incisor is well-preserved; it is oval in cross-section, with somewhat flat medial and lateral margins. The anterior surface of the incisor is covered by smooth enamel that extends only to the labial surface, covering about one-third of the labial side of the incisor.

392 The lower deciduous premolar (dP₄) is slightly less worn than the lower molars (Fig. 5). A
 393 micro-CT scan revealed no hint of a developing p4 (Fig. 5F), suggesting that *Birkamys* likely did
 394 not replace dP₄ – a condition that occurs in later-occurring phiomorphs aside from *Phiomys*. The
 395 dP₄ is longer than it is wide, and has a somewhat rectangular outline with a wide talonid and a
 396 narrow trigonid. The tooth bears five major cusps (metaconid, entoconid, protoconid, hypoconid
 397 and hypoconulid). The lingual cusps are slightly smaller than the labial cusps, and the
 398 hypoconulid is the smallest cusp. The protoconid extends distal to the metaconid, and has a short
 399 crest running mesially from its mesiolingual portion that meets the anteroconid. A low and
 400 weakly developed anterocingulid extends around the mesial margin of the tooth. The middle
 401 portion of the crown is open mesially, due to the absence of the metalophulid I and II, and open
 402 lingually due to the absence of the anterior arm of the entoconid and the short posterior arm of
 403 the metaconid. The ectolophid is low and attaches to the anterior arm of the hypoconid near that
 404 crest's junction with the hypolophid. The entoconid is placed mesial to the hypoconid. The
 405 hypoconid is connected to a distinct hypoconulid by a well-developed posterolophid that runs
 406 across the distal border of the crown and does not reach the distal face of the entoconid, leaving
 407 the posterior basin open lingually. The labial sinusid is wide and shallow with no ectostylid.



408 The first lower molar (M₁) is somewhat rectangular in outline and relatively broad, and bears
 409 long and relatively well-developed lophs when compared with those of the dP₄ (Fig. 5). The
 410 mesial wall of the tooth is formed by a trenchant and concave metalophulid I that connects the
 411 protoconid and metaconid. A low anterocingulid is present mesial to the protoconid. The
 412 metalophulid II is represented only by an incipient bulb that protrudes lingually from the
 413 protoconid; thus the tooth bears only two major basins. The ectolophid is more lingually
 414 positioned than that on dP₄. The posterior arm of the metaconid tapers and ends near the midline

415 of the tooth, leaving the  anterior basin open lingually via a narrow valley. The  labial sinusid is
416 broad and deep.

417 The second lower molar (M_2) is the largest tooth of the lower dentition. The occlusal
418 configuration is identical to that of the first molar, but the tooth is relatively broad. The anterior
419 margin is straighter and the anterocingulid is less developed than on the M_1 and dP_4 . The
420 occlusal pattern on the third molar is also very similar to that of M_{1-2} , but differs in having a
421 relatively weak anterocingulid, a trigonid that is slightly broader than the talonid, a  relatively
422 well-developed metalophulid II, and a hypoconulid that is submerged into the posterolophid,
423 forming the rear lobe of the crown.

424

425 **Cf. *Birkamys* (Fig. 6)**

426 An almost complete hemi-mandible (DPC 17995) has been recovered from L-41 that preserves
427 the lower incisor and M_1 (Fig. 6). There are few differences in the morphology of the corpus,
428 ascending ramus, and in the position and development of the masseteric crests, but the anterior
429 border of the ascending ramus begins to curve dorsally just posterior to the distal aspect of M_1 ,
430 whereas on the mandible placed in the *Birkamys korai* hypodigm (DPC 22737), the dorsal
431 curvature begins at about the mid-point of M_3 ; therefore it seems likely that the ascending ramus
432 would have obscured M_3 in lateral view on DPC 17995. The M_1 preserved in DPC 17995 also
433 differs from that in DPC 22737, notably in being longer relative to width (1.26 times longer than
434 wide, as opposed to 1.10 times longer than wide); having a distinct, but short,  metalophulid II; in
435 having a lower crown height and relatively thin crests; in having a more broadly open
436  hypoflexid; and in having a metalophulid I that is mesially convex. We refrain from naming a

new taxon based on this material because only one diagnostic tooth is known, but the differences between DPC 17995 and DPC 22737 suggest that this might be an additional tiny new species that could even be distinct at the generic level.

***Mubhammys*, new genus (Figs. 7 and 8, Table 1)** urn:lsid:zoobank.org:act:[in process]

Type and only known species

Mubhammys vadumensis, new species urn:lsid:zoobank.org:act:[in process]

Etymology

Combination of *mubham*, Arabic word for enigmatic or mysterious, and *mys*, Greek for mouse.

Diagnosis

As for the type and only known species.

***Mubhammys vadumensis*, new species** urn:lsid:zoobank.org:act:[in process]

(Figs. 7 and 8, Table 1)

Etymology

From Greek *vadum* for shallow, in reference to the depositional environment of L-41.

Holotype

CGM 66001, a left maxilla with dP³⁻⁴ and M¹⁻³ (Fig. 7A-B, measurements in Table 1).

Referred specimens: DPC 14324, left maxilla with dP³-M¹ (Fig. 7C-H), DPC 13220, left mandibular fragment with dP₄-M₃ (Fig. 8E-H), DPC 14141, left mandibular fragment with dP₄-M₁ (Fig. 8A-D).

457 *Type locality*

458 Locality 41 (L-41).

459 *Age and Formation*

460 Latest part of late Eocene (latest Priabonian, ~34 Ma), lower sequence of Jebel Qatrani
461 Formation, northern Egypt.

462 *Diagnosis*


463 Relatively large basal phiomorph that shows no evidence for replacement of deciduous
464 premolars. Differs from *Birkamys korai* primarily in its larger size, but also in lacking a dP₄
465 anteroconid; in having metalophs on dP₄-M³ that are submerged into posterolophs; and in having
466 small M¹⁻² mesostyles. Differs from contemporaneous and sympatric *Acritophiomys bowni* in
467 showing no evidence for replacement of deciduous premolars; in lacking a distinct anteroconid,
468 mesostylid, mesolophid, metalophulid I, and metalophulid II on dP₄; in lacking a mesolophid,
469 metalophulid II, and incipient anteroconid on M₁₋₂; in having relatively large lower molar
470 protoconids when compared to metaconids; in having M^{3/3} relatively small when compared with
471 M^{2/2}; in lacking pericingula and mesolophules on dP₄-M²; in having relatively large M¹⁻²
472 metaconules; in having M¹⁻³ metalophs that are submerged into the posterolophs; and in lacking
473 enamel wrinkling and crenulation on molars. Differs from younger *Metaphiomys beadnelli*, also
474 ~~from the Fayum succession~~, in lacking a dP₄ anteroconid; in lacking a metalophulid II on dP₄; in
475 having a larger dP₄ hypoconulid; in having relatively large lower molar protoconids when
476 compared to metaconids; in having M^{3/3} relatively small when compared with M^{2/2}; in having a
477 relatively low dP₄ anterocingulum; in having dP₄-M³ metalophs that are submerged into
478 posterolophs; in lacking mesolophules on dP₄-M²; in having relatively large M¹⁻² metaconules;









and in having no posterior arm of the paracone on M^{1-2} . Differs from Oligocene *Turkanamys* *hexalophus* from Kenya in showing no evidence for replacement of deciduous premolars; in having a low anterocingulid on M_{1-2} ; in having small metastylids, no mesostylids, and ~~very weak~~ ~~metalophulid Hs~~ on M_{1-2} ; in having relatively large lower molar protoconids when compared to metaconids; in having no connection of the entoconid and hypoconid via the posterolophid on M_{1-2} ; in having $M^{3/3}$ relatively small when compared with $M^{2/2}$; in having small parastyles, weak anterostyles, and no mesolophules on M^{1-2} ; in having metalophs submerged into posterolophs on M^{1-3} ; and in lacking enamel wrinkling and crenulation on the molars.



Description


On the medial surface of the broken premaxilla of CGM 66001 (Fig. 7F-G) a complete right upper incisor is exposed; it is short and highly arched when compared with the lower incisor. It is oval in occlusal outline with a flat medial surface and curved dorsal and lateral surfaces. In lateral view, the occlusal surface is longer when compared with that of the lower incisor. The pulp cavity is short and slit-shaped, and placed at the middle of the occlusal surface. A smooth enamel layer covers the mesial surface of the upper incisor and extends labially to cover only one-third of the labial side.


The maxillary morphology of *Mubhammys* appears to be very similar to that of *Birkamys*, with a similar placement of the zygomatic process, and an anteroposteriorly extensive concave margin of the lateral border of the incisive foramen. As in *Birkamys*, the posterior margin of the incisive foramen would have extended posteriorly to approximately the point of DP^4 , and the anterior margin presumably extended far into the premaxilla. *Mubhammys* thus shares the remarkably enlarged incisive foramen morphology that is seen in *Birkamys*.

502 The upper deciduous third premolars (dP³) are preserved in the holotype and DPC 14324. It is a
 503 small peg-shaped tooth with a primary cusp and a rounded base, and abuts the mesial surface of
 504 dP⁴.  The dP³ is less worn than the upper molars (Fig. 7A-D, H).

505 The upper deciduous fourth premolar (dP⁴) has a ~~somewhat~~ trapezoidal outline, with a relatively
 506 short lingual margin and a broad labial margin (Fig. 7A-C, H). The occlusal surface has four
 507 enlarged cusps (paracone, protocone, metacone  and hypocone), all of which are about equal in
 508 size and height. The mesostyle is a distinct and isolated cusp  and is situated midway between the
 509 paracone and ~~the~~ metacone along the buccal margin of the tooth. An ~~inflation for the~~ anterostyle 
 510 is situated mesiolabial to the protocone, midway along the mesial margin of the tooth. A short
 511  and weakly developed crest runs longitudinally from the most mesiolabial part of the protocone
 512 to integrate into the anterostyle. The anteroloph is  low and short, runs labially from the
 513 anterostyle, and courses across the mesial margin of the tooth. It continues distally to end labially
 514 to the metacone, forming a cingulum around the mesiolabial corner of the crown. The protoloph
 515 is a very short transverse crest, runs labially from the protocone, and flares labially to meet the
 516 paracone, which is large and lacks a posterior arm. Due to the absence of the posterior arm of the
 517 paracone and the anterior arm of the metacone, the  central basin is open labially via a wide
 518 notch. The hypocone is well-developed and placed distal to the protocone. The anterior arm of
 519 the hypocone is robust and runs mesiolabially to end at the middle of the tooth. There is a very
 520 low and weakly developed mure, connecting the base of the anterior arm of the hypocone with
 521  the protoloph. The mesolophule is absent. The metaloph is robust but very short and its lingual
 522 end is submerged into the posteroloph, the latter of which runs labially from the hypocone and
 523 ends at the distal base of the metacone. The  posterior basin is shallow and very narrow when

524 compared with the  anterior basin. Distal to the anterostyle and mesial to the protocone, there is a
525 short and low anterocingulum. The  labial sinus is deep and narrow. There is no ectostyle.

526 The upper first molar (Fig. 7A-C, H) has a similar occlusal morphology to that of dp^4 , but it is
527 larger, and the lophs and cusps are relatively well-developed. The outline of the tooth is roughly
528 square. The anteroloph has a relatively straight course and is strongly connected lingually to a
529 well-developed protocone, unlike that of dp^4 . The metaloph is relatively long with respect to that
530 of dp^4 and turns backward where it is confluent with the posteroloph. The mesostyle is relatively
531 well-developed and extends toward the  middle basin via a short strut. In the holotype, the
532 metacone is broken.

533 The upper second molar is only preserved in the holotype (Fig. 7A-B) and it is strongly worn,
534 but in general appears to be similar in morphology to M^1 . It differs in being larger and having a
535 reduced and more labially and lingually situated hypocone and metacone (broken), respectively.
536 The metaloph is relatively long with respect to that of M^1 . The  sinus is relatively wider when
537 compared with those of dp^4 and M^1 . The cingulum around the mesiolabial corner of the tooth is
538 absent.

539 The upper third molar (Fig. 7A-B) is heart-shaped and is smaller than all of the other upper teeth
540 aside from dp^3 . The anterior half of the tooth is similar to the corresponding part on M^{1-2} , but the
541 posterior part is very worn and the metacone is broken. The tooth has a large and more
542 crestiform protocone, occupying most of the lingual portion of the crown. The hypocone is very
543 small and is relatively labial in position with respect to the protocone. The course of the
544 metaloph is uncertain due to wear.

545 The mandible is slender and hystricognathous, with its angular process placed lateral to the plane
 546 of the incisor and the tooth row, leaving a well-developed groove between the angular process
 547 and this plane (Fig. 8B-D, F-H). DPC 13220 (Fig. 8E-H) represents a fragment of a left
 548 mandibular corpus with dp_4 - M_3 and the middle part of the incisor. On the labial surface of the
 549 specimen, the masseteric fossa is pronounced, ending beneath the first molar. The ventral
 550 masseteric ridge extends farther mesially below the dp_4 and is located lateral to the incisor and
 551 teeth row. The dorsal masseteric ridge is less developed when compared to the ventral ridge and
 552 extends distally along the base of the coronoid process. The groove between the tooth row and
 553 the base of the coronoid process is preserved in DPC 13220 (Fig. 8F, H). The lower incisor is
 554 oval with somewhat flat medial- and slightly convex lateral- margins. A smooth and thick enamel
 555 layer covers the ventral surface of the incisor and extends on both the mesial and distal sides.
 556 The pulp cavity is elongate in outline and sits in the middle of the dentine layer. The mental
 557 foramen is oval-shaped and relatively small and is located between the distal portion of the
 558 diastema, slightly above the level of the mesial end of the ventral masseteric ridge. The diastema
 559 is well-preserved in specimen DPC 14141 and is slightly deeper than the alveolar row and is as
 560 long as the length of the two first teeth. The tip of the angular process, coronoid process, and
 561 mandibular process are not preserved.

562 The two mandibular fragments of *Mubhammys vadumensis* (DPC 13220 & DPC 14141) preserve
 563 lower cheek teeth (Fig. 8). The dp_4 is longer than it is wide, and has a somewhat oval outline
 564 with a wide talonid and a narrow trigonid. The tooth bears a very weakly developed
 565 anterocingulid mesial to the protoconid. The ~~anterolophid~~ (the metalophulid I) and the posterior
 566 arm of the protoconid (~~the metalophulid II~~) are absent, leaving the ~~anterior basin~~ open mesially
 567 via a deep and narrow notch. The metaconid is larger than the protoconid and is more mesially

positioned. The hypoconid is placed slightly posterior to the entoconid. A short anterior arm of the hypoconid attaches to the hypolophid, which flares lingually along the apex of a large entoconid. The ectolophid is low relative to cusp height, and joins the protoconid at the junction of the anterior arm of the hypoconid and the hypolophid. The posterior arm of the metaconid slopes distally as a part of the lingual wall but terminates before the midline of the crown. The anterior arm of the entoconid is absent, leaving a wide notch along the lingual side of the ~~anterior basin~~. There is ~~no~~ mesoconid ~~or~~ mesolophid. The posterolophid is well-developed, running distolingually from the hypoconid, coursing around the posterior margin of the tooth to end distal to the entoconid, leaving a wide opening on the lingual margin of the ~~posterior basin~~. A well-developed hypoconulid occupies the middle portion of the posterolophid and forms the very distal tip of the tooth. A low, poorly developed postcingulid runs labially from the distal end of the hypoconulid. The ~~labial sinusid between the protoconid and hypoconid~~ is wide and deep. In DPC 14141, dP₄ has a well-developed cusp that abuts the metaconid distally, while the mesostylid is represented by a small cuspid.

The first lower molar is roughly rectangular in outline and bears relatively well-developed lophs and cusps; the mesial part of the tooth is slightly narrower than the distal part. An incipient and low anterocingulid runs parallel to the ~~anterolophid~~; it is relatively well-developed when compared with that on dP₄. The ~~anterolophid~~ is well-developed, forming the mesial border of the tooth, and runs from the anterior side of the protoconid toward the mesial side of the metaconid; it bears a shallow notch near its labial one-third. The posterior arm of the protoconid (~~the metalophid II~~) is represented by a very small knob on the mesial part of the ectolophid. The latter is well-developed and higher than that of the dP₄, and connects the protoconid to the junction of the hypolophid and the anterior arm of the hypoconid. The posterior arm of the

metaconid is relatively short when compared with that of the dP_4 , and only forms about one-third of the lingual wall between the metaconid and the entoconid. The anterior arm of the entoconid is absent, leaving the ~~anterior basin~~ open via a lingual notch that is narrower than that on dP_4 . The ~~labial sinus~~ is relatively narrow, and the postcingulid that runs labially from the distal end of the hypoconulid is relatively well-developed with respect to that on dP_4 .

The occlusal surface of M_2 (Fig. 8E-H) is very similar to that of M_1 . The M_2 differs in being relatively wider and shorter, and in having a slightly broader trigonid and a longer talonid. The notch of the ~~anterolophid~~ is relatively deep when compared with that on M_1 , but still higher than the anterior basin. The posterior arm of the protoconid is very short and is oriented toward the lingual wall of the tooth, but it is relatively long when compared with that on M_1 . The hypoconulid and its postcingulid are not as well-developed as they are on dP_4 - M_1 .

The mesial portion of M_3 (Fig. 8E-H) is somewhat similar to the corresponding part of M_1 and M_2 . The M_3 differs from the rest of the lower dentition in being relatively small, and in having a triangular outline. The trigonid is relatively large when compared with the talonid. The posterior arm of the protoconid is relatively long and is oriented toward the entoconid, which is reduced in size. The hypolophid is short and the hypoconid flares with the posterolophid distally, forming the posterior margin of the tooth. The hypoconulid is absent. The anterior arm of the entoconid is present, leading to a narrow and deep notch on the lingual wall; the ~~posterior basin~~ is closed lingually. The postcingulid is absent, and the anterior arm of the hypoconid is very short.





Comparisons



612 The contemporaneous and sympatric L-41 species of *Birkamys* and *Mubhammys* are quite
 613 different in tooth and mandible size (Fig. 9), but are very similar in the occlusal morphology. *B.*
 614 *korai* is the smallest known hystricognathous rodent from the Paleogene of Africa, and the first
 615 lower molar area of *M. vadumensis* is 3.5 times that of *B. korai*. The lower dentition of *Birkamys*
 616 is similar to that of *Mubhammys* in lacking metalophulid I and II on dp_4 , and in having an
 617 ~~incipient metalophulid II~~ on M_1 and a more robust ~~metalophulid II~~ on the last two molars that
 618 never reaches the lingual wall. Moreover the upper dentition of *Birkamys* is similar to that of
 619 *Mubhammys* in lacking the connection between the metacone and the anterior arm of the
 620 hypocone; instead the metaloph is directed distally and submerged into the posteroloph. The
 621 lower molars of *Birkamys* differ from those of *Mubhammys* in lacking the postcingulid, having a
 622 relatively robust metalophulid I on the lower molars, in having a relatively short M_1 . *Birkamys*
 623 also lack the M^2 mesostyle that is present in *Mubhammys*. Furthermore, when compared with
 624 that of *Mubhammys*, the dp_4 of *Birkamys* is relatively long, has a crest that runs mesially from
 625 the protoconid, has a well-developed anterocingulid, and a large protoconid when compared with
 626 the metaconid; *Mubhammys* has a relatively large metaconid. The M^3 of *Mubhammys* is heavily
 627 worn, which makes it difficult to compare it with that of *Birkamys*.

628 *Birkamys* and *Mubhammys* share a number of dental features with early Oligocene members of
 629 the genera *Phiomys* and *Neophiomys* (Coster et al., 2012b; Holroyd, 1994; Wood 1968) that were
 630 presumably present in the last common ancestor that these taxa shared with all later phiomorphs.
 631 The type species of *Phiomys* (*Phiomys andrewsi*) differs from *Birkamys* in being larger, having a
 632 well-developed metalophulid II, in replacing the dp_4 , and in having a relatively narrow anterior
 633 portion of dp_4 . When compared with *Mubhammys*, the lower teeth of *P. andrewsi* are smaller,
 634 with a relatively well-developed anterocingulid. The lower molars of *Birkamys* differ from those

635 of *Neophiomys* from the early Oligocene Fayum Quarry G (Coster et al., 2012b; Wood 1968) in
 636 their small size, in having a well-developed metalophulid I, a relatively weak anterocingulid, and
 637 a relatively long dP₄ that has the anteroconid connected to the protoconid. Furthermore, the M²
 638 of *Birkamys* differs from that of *Neophiomys* in lacking a mure, a mesolophule, and a metaloph
 639 that is connected with both the posteroloph and mesolophule. *Neophiomys* is similar to
 640 *Mubhammys* in having an M₁ that is the longest tooth of the lower dentition, in the development
 641 of the metalophulid II, and in having an anterocingulid, but differs in its small size, interrupted
 642 metalophulid I, and well-developed anterocingulids that are present on M₃. Furthermore, the dP₄
 643 of *Neophiomys* differs from that of *Mubhammys* in having an anteroconid and a complete
 644 metalophulid II. The upper dentition of *Neophiomys* differs from that of *Mubhammys* in having a
 645 double (mesial and distal) connection of metaloph on M¹, as well as a well-developed mure and
 646 mesostyle. When compared to *Birkamys*, the lower teeth of “*Phiomys* aff. *paraphiomyoides*”
 647 from the early Oligocene Fayum Quarry I are larger, and there is a small cusp on the
 648 anterocingulid of M₁₋₂. The dP₄ of “*Phiomys* aff. *paraphiomyoides*” also shows some characters
 649 that differ from those of *Birkamys*, such as being relatively short and in having a complete
 650 metalophulid II and a distinct anteroconid. The lower cheek teeth of *Mubhammys* show a great
 651 similarity to those of “*Phiomys* aff. *paraphiomyoides*”, differing only in having a dP₄ that lacks
 652 metalophulid II, and in having a relatively long M₁ with deep lingual notches, relatively well-
 653 developed postcingulids, and in lacking a small cusp on the weak anterocingulid.

654 “*Phiomys*” *lavocati* (Wood 1968) is roughly the same size as *Birkamys korai*, but has a more
 655 robust mandible that bears a more ventral position of the mental foramen, as well as a relatively
 656 short diastema. The lower teeth of “*Phiomys*” *lavocati* differ from those of *Birkamys* in being
 657 broader and shorter, having relatively weak metalophulid I and II, in lacking anterocingulids, and

in having  anterior and posterior basins that are closed lingually. The lower deciduous premolar of “*Phiomys*” *lavocati* is also relatively short, and has a relatively narrow trigonid and wide talonid, a relatively well-developed ectolophid, and an anteroconid. The dP₄ protoconid of “*Phiomys*” *lavocati* is more distally positioned with respect to the metaconid, and the anterocingulid is more labially placed. The upper second molar of “*Phiomys*” *lavocati* differs from that of *Birkamys* in having a relatively reduced and more lingually positioned metacone; a metaloph that is poorly developed and oriented toward the metaconule; a hypocone that is equal in size to the protocone and less labially positioned; an anteroloph that is relatively well-developed; a high labial wall that closes the ~~central basin~~  labially; a mesostyle that is integrated into the  labial wall, but still recognizable; an enterostyle that abuts the hypocone, and a mure extending from the metaconule  to contact the posterior portion of the protoconid.

The lower molars of “*Phiomys*” cf. *lavocati* from the early Oligocene of Oman (Thomas et al. 1989) are similar to those of *Birkamys* in having an anterocingulid and short posterior arm of the protoconid, and a lingually closed  anterior basin on M₃. But the only upper molar from that locality shows significant differences when compared with that of *Birkamys*, such as the presence of a complete mure, an interrupted mesolophule, a metaloph that is connected to both the metaconule and posteroloph, a labial wall that closes the ~~central basin~~ , and a mesostyle. Furthermore the first lower molar of “*Phiomys*” cf. *lavocati* has both a well-developed postcingulid and a lingual wall, and the posterior arm of the protoconid protrudes from the middle of the ectolophid, rather than from the protoconid as in *Birkamys*.

The diminutive species *Phiocricetomys minutus*, from Fayum Quarry I, differs from *Birkamys* in exhibiting strong distal reduction of the cheek teeth (i.e., loss of M₃), having relatively rounded cusps, no metalophulid II, and an anterior end of the masseteric fossa that extends as far forward

681 as the anterior end of the premolars. The mental foramen is placed beneath this ridge in
 682 *Phiocricetomys*, whereas in *Birkamys* the anterior end of the masseteric fossa lies just below the
 683 mental foramen. *Phiocricetomys* also has a cingulid that courses all the way around the labial
 684 surface of M₁, and a strong anterocingulid that bears an anteroconid and some small accessory
 685 cusps. On the dP₄ of *Phiocricetomys*, the metaconid and its posterior arm are more mesially
 686 positioned with respect to the protoconid, the anteroconid is well-developed, the anterolophid is
 687 interrupted labially, the hypolophid is absent, the hypoconulid is well-developed and relatively
 688 large, with a posterior cingulid, and the posterolophid is very weak. The M₂ of *Phiocricetomys*
 689 has a reduced entoconid, and a protoconid that is concave mesially and convex distally. The
 690 hypolophid is absent and the posterolophid is incipient, with no hypoconulid.

691 Contemporaneous *Acritophiomys bowni* (Sallam et al. 2012) is the same size as *Mubhammys*
 692 *vadumensis*, but it differs in replacing its deciduous premolars, and in having well-developed
 693 mesolophules, double connections of the metaloph, complete mures, and relatively long lophs on
 694 the upper molars, while the lower molars bear well-developed metalophulid IIs. Moreover, the
 695 dP₄ of *A. bowni* differs than that of *M. vadumensis* in having a well developed anterocingulid, a
 696 complete posterior arm of protoconid and a well-developed mesolophid, mesostylid, and
 697 ectostylid.

698 *Birkamys* and *Mubhammys* differ from the primitive hystricognaths *Protophiomys* and *Waslamys*
 699 from the earliest late Eocene (~37 Ma) of the Fayum area (Sallam et al, 2009) in retaining
 700 deciduous premolars, having no mesolophid or mesoconid on dP₄, in having lingually open
 701 anterior basins on the lower molars, and, on the upper molars, lacking endolophs, mesolophules,
 702 and labial walls on the upper molars; *Birkamys* and *Mubhammys* also have M² metalophs that are
 703 oriented distally, meeting the posterolophs.

In addition to its small size, *Birkamys korai* shares some dental features with *Kahawamys mbeyaensis* from the late Oligocene of the Rukwa Rift Basin in Tanzania (Stevens et al. 2009), such as a more centrally positioned ectolophid, a relatively mesial position of the entoconid with respect to the hypoconid, and a ~~crest~~ extending distally from the dP₄ anteroconid to connect with the protoconid. The lower molars of *Birkamys* differ from those of *Kahawamys* in lacking an anterior arm of the entoconid, having relatively short and lingually open **posterior basins**, relatively distinct ~~metalophulid II crests~~ that increase in length distally, anterocingulids on all molars, and a dP₄ that is relatively wide, with a taller trigonid.

The extant cane rat *Thryonomys* resembles *Birkamys* **in retaining deciduous premolars** throughout its life, but in addition to being much larger in size, differs from *Birkamys* and *Mubhammys* in having relatively short and hypsodont crowns, ~~no~~ anterocingulids ~~or metalophulid IIs~~ on lower teeth, and in having a complete metalophulid I and a mesolophid in dP₄. Furthermore, the M² of *Thryonomys* differs from that of *Birkamys* and *Mubhammys* in having a complete mure and in lacking a metaconule. The M³ of *Thryonomys* differs from that of *Birkamys* in lacking an ~~endoloph~~, a metaloph and a well developed metaconule.

Phylogenetic analysis

Parsimony analysis

Parsimony analysis in PAUP 4.0b10 returned 10 equally parsimonious trees of length 907, the strict consensus of which is shown in Fig. 10. In all trees, *Birkamys* and *Mubhammys* are nested deep within **Phiomorpha** as the sister taxon of extant *Thryonomys*, with the sister group of the *Birkamys*-*Mubhammys*-*Thryonomys* clade containing early Oligocene “*Paraphiomys*” *simonsi*

726 and Miocene *Paraphiomys* and *Paraulacodus*. With the placement of *Canaanimys* (late middle
 727 Eocene of Peru) taken to indicate the divergence of Caviomorpha from Phiomorpha, there is a
 728 pectinate sequence of basal phiomorphs that includes “*Phiomys*” *hammadai* and *Turkanamys* as
 729 its most basal members, followed sequentially by *Acritophiomys*, *Prepomonomys*, *Phiomys*,
 730 *Neophiomys*, *Metaphiomys*, and *Diamantomys*. *Birkamys* and *Mubhammys* are placed as sister
 731 taxa with strong support (bootstrap support (BS) = 82). Among derived hystricognaths, the only
 732 other clade that was supported by BS >50 was *Canaanimys* + *Gaudeamus* (late Eocene and early
 733 Oligocene of Africa) (BS=62), *Gaudeamus* (BS=99), and a *Gaudeamus* clade that excludes late
 734 Eocene *Gaudeamus aslius* (BS=83). Elsewhere in the tree, the oldest African hystricognath,
 735 “*Protophiomys*” *tunisiensis*, is placed as the sister species of the phiocricetomyines
 736 *Talahphiomys*, “*Phiomys*” *lavocati*, and *Phiocricetomys*, while the next-oldest species from
 737 northern Africa, from Locality BQ-2 (*Protophiomys aegyptensis* and *Waslamys*) and Bir el-Ater
 738 (*Protophiomys algeriensis*) are placed outside of the Phiomorpha-Caviomorpha clade, with *P.*
 739 *algeriensis* being the sister taxon of derived Asian “baluchimyines”, *Waslamys* forming a clade
 740 with “*Protophiomys*” *durattalahensis*, and *Protophiomys aegyptensis* intervening along the
 741 phiomorph-caviomorph stem between the divergences of the “*P.*” *tunisiensis*-phiocricetomyine
 742 clade and the *Waslamys*-“*P.*” *durattalahensis* clade.

743

744 *Bayesian phylogenetic analysis*

745 As in the parsimony analysis, the “allcompat” (majority-rule plus compatible groups) consensus
 746 derived from the Bayesian analysis (Fig. 11) recovered a *Birkamys*-*Mubhammys* clade, with
 747 moderate support (posterior probability (PP) = 0.68), but this clade was not so deeply nested
 748 within Phiomorpha, instead being the sister group of a well-supported (PP = 0.91) clade

containing Oligocene (*Metaphiomys*, “*Paraphiomys*” *simonsi*) and Miocene (*Diamantomys*, *Paraphiomys*, *Paraulacodus*) species, as well as a *Thryonomys*-*Prepomonomys* clade. Also as in the parsimony analysis, the most basal phiomorphs are “*Phiomys*” *hammudai* and *Turkanamys*, followed by the sequentially more crownward genera *Acritophiomys*, *Phiomys*, and *Neophiomys*. The arrangement of species along the pectinate stem at the base of Phiomorpha in the allcompat tree is supported by PPs in the range of 0.49-0.56. *Canaanimys* and *Gaudeamus* form a well-supported clade (PP = 0.99), and “*P.*” *tunisiensis* is again placed as the sister taxon of phiocricetomyines, but with very weak support (PP = 0.36). There is strong support (PP = 0.91) for a clade containing African species, *Canaanimys*, and advanced “baluchimyines” to the exclusion of basal “baluchimyines” (*Baluchimys ganeshaper*, *Baluchimys krabiense*, *Confinniumys*, *Lindsaya*, *Ottomania*), but relationships among the basal members of the clade (*Protophiomys*, *Waslamys*) in the allcompat tree are only weakly supported (PPs = 0.25-0.5).

Tip-dating analysis with the fossilized birth-death prior

The “allcompat” consensus summarizing the 50,000 post-burn-in trees from the tip-dating analysis of the 118-character matrix with broad uniform priors on tip ages (i.e., analysis TD1, see Fig. S1) provided tip estimates that were averaged on a locality-by-locality basis for analysis TD2. These averages allowed for the ordering of localities from oldest to youngest as follows (Fig. 12, see Table 2 for 95% HPD): Subathu “Zone VIII” (India), 45.8 Ma; Khaychin II-III-IV (Mongolia), 43.5 Ma; Rencun Member (China), 41.0 Ma; Djebel el Kébar (Tunisia), 39.2 Ma; Birket Qarun Locality 2 (Egypt), 37.3 Ma; Bir el-Ater (Nementcha, Algeria), 37.3 Ma; Dur at-Talah DT1 (Libya), 36.2 Ma; Dur at-Talah DT2 (Libya), 35.9 Ma; Fayum Locality 41 (Egypt), 34.4 Ma; Fayum Quarries A and B (Egypt), 33.8 Ma; Krabi Bang Mark Pit (Thailand), 33.5 Ma;

772 Lokone (Kenya), 31.9 Ma; Hsanda Gol (Mongolia), 32.0 Ma; Y-GSP 417 (Pakistan), 31.6 Ma;
 773 Süngülü (Turkey), 32.1 Ma; Fayum Quarry E (Egypt), 32.0 Ma; Fayum Quarry G (Egypt), 30.8
 774 Ma; Paali Nala C2 (Pakistan), 29.9 Ma; Fayum Quarries I and M (Egypt), 29.6 Ma; and Silica
 775 North (Namibia), 25.7 Ma.

776 The resulting “allcompat” tree from analysis TD2, with tips fixed to the mean dates above (Fig.
 777 13; see Fig. S2 and Dataset S5 for absolute median rates for each branch), is effectively the same
 778 as that from TD1 [Fig. S1; the sole difference being that *Acritophiomys* joins “*Phiomys*”
 779 *hammadai*, with very low probability (PP=0.30)], but shows some notable differences from that
 780 based on the standard Bayesian analysis; few of the different placements are supported by high
 781 posterior probabilities, however. In contrast to the Bayesian allcompat consensus, the tip-dating
 782 allcompat 1) places *Birkamys* and *Mubhammys* in a slightly more basal position in phiomorph
 783 phylogeny, being the sister group of all younger phiomorphs aside from *Phiomys*; 2) shows
 784 much stronger support for the monophyly of progressively more nested phiomorph clades that
 785 include *Turkanamys*, “*Phiomys*” *hammadai*, and *Acritophiomys* (TD1 PP=0.90, TD2 PP=0.92,
 786 standard Bayesian PP=0.55), *Phiomys* (TD1 PP=0.90, TD2 PP=0.91, standard Bayesian
 787 PP=0.49), and *Birkamys-Mubhammys* (TD1 PP=0.94, TD2 PP=0.96, standard Bayesian
 788 PP=0.50); 2) the sole undoubted caviomorph in the matrix (*Canaanimys*) in a particularly basal
 789 position, and with no special relationship to *Gaudeamus*, in strong contrast to the Bayesian
 790 analysis that supported a *Canaanimys-Gaudeamus* clade with a posterior probability of 0.99; 3)
 791 *Tsaganomys* and “basal baluchimyines” form a well-supported clade (TD1 PP=0.87, TD2
 792 PP=0.92) rather than a paraphyletic stem with respect to African and South American
 793 hystricognaths (as found in both the standard Bayesian and parsimony analyses). Despite the
 794 differences in overall branching sequence between the tip-dating allcompat consensus and the

standard Bayesian allcompat consensus, in the former the only higher-level clades that are supported by posterior probabilities >0.5 are Phiocricetomyinae (TD1 PP=0.93, TD2 PP=0.94), a clade of “advanced” baluchimyines (TD1 PP=0.94, TD2 PP=0.98), Gaudeamuridae (TD1 and TD2 PP=1.0), and the clade including “phiomyids” and derived phiomorphs (TD1 PP=0.90, TD2 PP=0.92). As such, the interrelationships among those clades, *Canaanimys*, *Waslamys* and the various species assigned to the genus *Protophiomys* are not well-resolved, though support for a “*Protophiomys*” *tunisiensis*-Phiocricetomyinae clade increases to PP=0.52 in analysis TD2.


Evolutionary rates (calculated for each branch as number of changes per site per Ma by multiplying the median rate for the branch in the allcompat consensus by the median estimate for the base clockrate that is output in the MrBayes *.pstat file; see Beck and Lee, 2014) are fairly consistent across most nodes in the tree, with the most striking accelerations being along the branches leading to *Canaanimys* and to the *Birkamys*-*Mubhammys* clade. The split between Caviomorpha and Phiomorpha (in this case including basal baluchimyines and *Tsaganomys*) is estimated to have occurred at either 43.5 Ma (TD1) or 43.2 (TD2), and the largely African clade containing phiocricetomyines, gaudeamurids, and undoubted phiomorphs is estimated to have appeared at either 41.8 Ma (TD1) or 41.5 Ma (TD2). The advanced phiomorph clade that includes all species showing suppression of $P^{4/4}$ eruption (i.e., the clade including *Birkamys* and *Mubhammys* but excluding *Phiomys*) is estimated to have appeared at 35.1 Ma in both TD1 and TD2.

Evolution of M_1 size among early hystricognaths

With $\ln M_1$ area treated as a continuous variable evolving on the pruned tip-dating “allcompat” tree from TD2 (Fig. S3), the directional and random walk models (both with a lambda scaling parameter) returned roughly equal likelihoods, with no basis for preferring one model over the other. In Figure 14 we present the results of runs based on both models, with mean estimates for ancestral nodes from each model represented by single points and the intervening space infilled to reflect uncertainty (see Dataset S6 for means and upper and lower 95% HPD for each node reconstructed).

While the two models leave considerable uncertainty about the mean M_1 area estimate along the stem leading to the African hystricognath radiation, the estimate for that group’s common ancestor is well-constrained, with the random walk model recovering a mean estimate of 2.86 mm² and the directional model a mean estimate of 2.69 mm² — i.e., about the size of *Protophiomys aegyptensis* from Locality BQ-2, and a little bit larger than the oldest Afro-Arabian hystricognath, “*Protophiomys*” *tunisiensis* (Marivaux et al., 2014). From this point of origin for African hystricognaths, there is an immediate size-related divergence between phiocricetomyines (here including “*Protophiomys*” *tunisiensis*) and the lineage leading to derived phiomorphs. Phiocricetomyines decrease in size at a fairly constant rate through time, finally terminating with the diminutive early Oligocene species *Phiocricetomys minutus*. There is little change leading to the initial divergence of *Protophiomys sensu stricto*, with slight increases in M_1 area along the lineages leading to gaudeamurids, “*Phiomys*” *hammudai*, and *Turkanamys*, but in the early Priabonian a dwarfing event is implied, paralleling the trend seen in phiocricetomyines, along a trajectory that ultimately leads to the tiny species *Phiomys andrewsi*.

There is a reconstructed reversal of that trend in the Priabonian, close to the point of origin of the *Birkamys-Mubhammys* clade and the reconstructed acquisition of $P^{4/4}$ suppression, followed by

gradual increases in size through the early Oligocene with the evolution of more derived and deeply nested phiomorphs (e.g., *Metaphiomys*). Given the trends reconstructed here, the small size of *Neophiomys* is, like *Phiomys*, also reconstructed as having been due to dwarfing, rather than retention of ancestral small size. Most remarkable in these size trends is the rapid divergence of *Birkamys* and *Mubhammys* from a common ancestor into dramatically different size categories, along trajectories that are almost horizontal on the size change  versus time plot, showing that major change in size has occurred over only a very short period of time (note, though, that the implied change along the *Birkamys* branch appears somewhat exaggerated due to the use of logarithmically transformed variables). The upper and lower 95% HPDs for most ancestral estimates are quite broad (see Dataset S6), but the 95% HPD for the common ancestor of all Afro-Arabian hystricognaths (random walk model, 0.3205-1.8077; directional model, 0.2243-1.7649) nevertheless clearly excludes values in the range of *Birkamys korai*.

Discussion


Origin and evolution of the Birkamys-Mubhammys clade

The placement of *Birkamys* and *Mubhammys* as the exclusive sister taxa of extant *Thryonomys* in the maximum parsimony analysis must be considered highly unlikely given the numerous extensive ghost lineages that the topology implies throughout phiomorph phylogeny (Fig. 15). When tips are scaled to the age estimates provided by the TD1 tip-dating analysis, and internodes are arbitrarily separated by 1 Ma, the strict consensus derived from the parsimony analysis requires a total of 422.7 Ma along all branches (Fig. 15A, calculated in Mesquite 2.75 (“sum of branch lengths”); Maddison and Maddison, 2011), while the standard Bayesian allcompat tree, which places *Birkamys* and *Mubhammys* much more basally in phiomorph phylogeny, provides a

significant reduction in overall time required (Fig. 15B, 357.4 Ma). When compared to the standard Bayesian allcompat tree, the tip-dating allcompat tree from TD2 requires only 45% of the total time accumulated across all branches (Fig. 15D, 159.4 Ma total). This tip-dating tree has several zones with very rapid divergences and short internodes, but even if branch lengths of terminal taxa and internodes are set to 0 in the parsimony tree (Fig. 15C, it still requires 125% more time (198.7 Ma versus 159.4 Ma) than the TD2 tip-dating tree. In light of the parsimony tree's poor fit to the fossil record, and the fact that denser sampling of Neogene thryomyids (which was outside the scope of this analysis) has provided strong evidence for the nesting of extant *Thryonomys* among middle Miocene *Paraulacodus*, late Miocene *Protohumus*, and extinct late Miocene and Pliocene species of *Thryonomys* (Kraatz et al., 2013), we focus our discussion on the results from the Bayesian analyses, which congruently place *Birkamys* and *Mubhammys* as basal phiomorphs.

Taken together, the phylogenetic and morphometric analyses presented here suggest that *Birkamys* and *Mubhammys* are members of a previously unrecorded late Eocene African lineage of early phiomorph rodents that diverged dramatically in size following a rapid change in tooth morphology in the later Eocene. The tip-dating analyses recovered particularly high evolutionary rates along the stem leading to the *Birkamys*-*Mubhammys* clade (Fig. 12), but ancestral reconstructions of M_1 area indicate only a slight size decrease along this branch; instead the most dramatic size change is the rapid dwarfing event that is reconstructed along the branch leading to *Birkamys*. This pattern suggests that a relatively rapid change in dental morphology (overall simplification, involving the loss or reduction of transverse crests) might have facilitated, or driven, *Birkamys*' expansion into a new niche space that was either unoccupied during the latest Eocene in this particular part of Africa, or that overlapped with the niche spaces of

885 phiocricetomyines (small members of which have been recovered at L-41; descriptions currently
886 in preparation).

887 The dwarfing event implied for the *Birkamys* lineage also suggests that  the complete suppression
888 of $P^{4/4}$ development and eruption likely first occurred at a body size larger than that of the tiny
889 species *B. korai*, perhaps in populations with individuals that were about the size of
890 *Protophiomys aegyptensis* or *Pr. algeriensis*. Given the ancestral reconstructions depicted in Fig.
891 14, initial suppression, however, would have evolved following an earlier, less dramatic, late
892 Eocene dwarfing event from somewhat larger ancestors. After the Eocene-Oligocene boundary,
893 members of the clade that evolved suppression subsequently undergo a steady increase in M_1
894 size. Wood (1968) suggested that in early phiomorphs the $P^{4/4}$ might have “lagged in the race to
895 become molariform” (p. 84), perhaps due to the longer retention of $dP^{4/4}$ in basal stem
896 phiomorphs, and that life-long retention of $dP^{4/4}$ would have been advantageous if selection
897 favored individuals with increased capacity for grinding across the post-diastemal dentition. The
898 dwarfing event reconstructed prior to the evolution of $P^{4/4}$ suppression opens up the possibility
899 that $dP^{4/4}$ retention might have first evolved in populations that were somewhat neotenuous
900 relative to their ancestors -- i.e, containing individuals whose growth trajectories (including the
901 timing of dental eruption) had been truncated, so that replacement of $dP^{4/4}$ ultimately never
902 occurred. Only after $P^{4/4}$ suppression had been effectively “fixed” might there have then been
903 selection for the increased hypsodonty of all unreplaced teeth in various Oligocene lineages.
904 Given the presence of what might be yet another tiny phiomorph species at L-41 (cf. *Birkamys*,
905 Fig. 6), however, it is also possible that the M_1 size change scenario presented in Fig. 14 is
906 overly simplistic; the hypothesis presented here will have to be tested with the recovery of
907 additional species from late Eocene sites throughout Africa and Arabia.

908

909 *Broader implications of tip-dating topologies for early hystricognath evolution*


910 The parsimony, standard Bayesian, and tip-dating analyses all congruently supported a
 911 phiomorph clade that includes as its most basal members late Eocene *Acritophiomys* and
 912 “*Phiomys*” *hammudai* and Oligocene *Turkanamys hexalophus*. In the parsimony and standard
 913 Bayesian analyses, this clade was found to be the sister taxon of Caviomorpha — i.e., either a
 914 Gaudeamuridae+Caviomorpha clade (parsimony) or a
 915 Gaudeamuridae+Caviomorpha+*Waslamys*+“*Protophiomys*” *durattalahensis* clade (standard
 916 Bayesian). In the tip-dating analysis, the caviomorph *Canaanimys* was not placed with
 917 gaudeamurids, and in fact was placed as the most basal of all hystricognaths, though with very
 918 weak support for its exclusion from more nested positions. One of the more remarkable
 919 topological rearrangements in the tip-dating analyses was the placement of primitive
 920 “baluchimyines” (*Baluchimys*, *Confiniummys*, *Lindsaya*, *Ottomania*) in a well-supported clade
 921 with early Oligocene *Tsaganomys*. This result is not entirely surprising, because the evidence for
 922 the paraphyly of the group with respect to derived hystricognaths is weak in the parsimony and
 923 standard Bayesian analyses (Figs. 10 and 11), and furthermore the paraphyly of the group
 924 requires extensive ghost lineages (Fig. 15) and presumably very low rates of evolution along
 925 terminal branches. However it is surprising that the monophyly of the group could be strongly
 926 supported (PP=0.92 in the TD2 analysis) given these conditions. “Advanced” baluchimyines
 927 (*Bugtimys*, *Hodsahibia*, *Lophibaluchia*) are nested within the African hystricognath radiation
 928 across all analyses, implying an African origin for that clade and a dispersal to Asia, which is
 929 estimated to have occurred in the middle or late Eocene by the tip-dating analyses. However

while this topological result is strongly supported by the standard Bayesian analysis, it is not well-supported in either the parsimony or tip-dating analyses.

Telling time with rodent teeth: implications of tip estimates for the chronology of Paleogene hystricognath-bearing sites

In the absence of radioisotopically datable rocks, terrestrial mammal faunas from spatially and/or temporally isolated horizons can be ~~extraordinarily~~ difficult to date. The magnitude of the difficulty is proportional to the intensity of sampling of that temporal interval elsewhere on a landmass; if many other faunas of different ages bracket a fauna of interest, it is more likely that the same species will be recovered from multiple localities, and, in such cases, standard biochronology based on first/last appearances can be used to order localities, using for instance Appearance Event Ordination (Alroy, 1994). In the Paleogene of Afro-Arabia, sampling of the terrestrial mammal record has been so limited, and so patchy in space and time, that it is very rare for localities to show overlapping species (Seiffert, 2006, 2010; Coster et al., 2012). The taxonomy that workers choose to use can further obfuscate the situation; i.e., if newly discovered fossils are uncritically assigned to new species, there will be no species overlap for biochronological analysis.

These conditions have led to an unfortunate situation in which ordering of sites in the Paleogene of Africa is more often than not based on assumption-laden “stage of evolution” arguments that compare species from two localities and determine that one is older than the other because species A is “more primitive” than species B, or species B is of a more “advanced evolutionary stage” than species A. For instance, in attempting to determine the age of the Dur at-Talah

952 localities in Libya that yielded some of the species included in our analysis, Jaeger et al. (2010a)
 953 argued that “*Protophiomys* is a primitive representative of the phiomyid African radiation and it
 954 is represented in Dur At-Talah by a slightly more derived species (*Pr. durattalahensis*) than that
 955 of Nementcha (*Pr. algeriensis*), thereby suggesting a younger age for the Dur At-Talah deposits”
 956 and that “the L41 rodent assemblage (Holroyd, 1994) contains more derived species than that of
 957 Dur at-Talah” (p. 211), leading them to argue that Dur at-Talah is also older than L-41. Sallam et
 958 al. (2012, p. 297) argued that ““*Phiomys*” *hammudai* from Dur at-Talah is perhaps slightly more
 959 primitive than *Acritophiomys* from L-41,  but nevertheless is clearly more derived than the
 960 hystricognaths from BQ-2 (*Waslamys attiai* and *Protophiomys aegyptensis*)” in suggesting that
 961 Dur at-Talah is probably intermediate in age between L-41 and BQ-2.

962 Though such assessments might be based on compelling background information, the evidence
 963 and assumptions underlying the arguments are rarely explicit. On the broadest level, in the
 964 absence of a phylogenetic analysis, the comparison of the characters of species A with those of
 965 species B assumes that the two species are closely related and that the characters in question are
 966 homologous; it further assumes the evolutionary trajectory of the features, and that the presence
 967 of a presumed apomorphy in species B indicates that that feature has appeared more recently in
 968 time than the presumed plesiomorphic state in species A. These arguments thus also make
 969 assumptions about rates of evolution — they assume that a (presumed) plesiomorphic taxon
 970 species A is likely to be temporally older than a (presumed) apomorphic taxon species B because
 971 the reverse arrangement would imply relatively slow evolutionary rates in species A (i.e., stasis)
 972 and fast evolutionary rates in species B. It might even be assumed that the presence of a
 973 presumed apomorphy in species B implies that a certain amount of speciation must have
 974 occurred along the lineage leading to species B to account for that amount of change.

Bayesian tip-dating with the fossilized birth-death prior takes into account the important background information that must underlie these “stage of evolution” arguments — ages of related species, phylogenetic relationships among those species, rates of evolution, and patterns in speciation and fossilization — but in a much more explicit, objective, and replicable manner. Here we suggest that Bayesian tip-dating analysis with the fossilized birth-death prior is thus not only of use for dating internal nodes, but can also reasonably be “turned on its head” to provide age estimates for temporally poorly-constrained *tips* that have been assigned broad uniform priors on tip age — i.e., taking into account multiple biochronologically-relevant parameters to determine both phylogenetic position *and* the most likely point in time at which a morphological pattern would likely be present, given its broad uniform prior on age, its phylogenetic position, and the base clock rate of the tree.

One possible concern with this approach is that age estimates for species will simply fall near the middle of their uniform age priors. Our results clearly indicate that this is not the case. A particularly striking example is provided by *Prepomonomys bogenfelsi*, a species of contested age from the Sperrgebiet area of Namibia (Pickford et al., 2008). *P. bogenfelsi* and the other mammals from the Sperrgebiet Silica North locality were first described as Lutetian (early middle Eocene) in age (Pickford et al., 2008) and later as Bartonian (late middle Eocene) (Pickford et al., 2014), but multiple authors (Coster et al., 2012; Marivaux et al., 2014), including Pickford et al. (2008) themselves, have noted that the species from Silica North are similar to Miocene species from east Africa. Pickford et al. (2008) even placed some of the Silica North species in, or close to, otherwise exclusively Miocene phiomorph genera (*Apodecter*, cf. *Bathyergoides*). If the Silica North specimens are Bartonian in age as Pickford et al. (2014) suggest, their taxonomic identifications as *Apodecter* and cf. *Bathyergoides* would require ~17-

998 21-million-year-old temporal extensions for these genera far back into the Eocene, into time
 999 intervals when crown hystricognaths were (given a direct reading of the fossil record) only just
 1000 starting to diversify. However with our very conservative and broad 20-47 Ma uniform age prior
 1001 on *Prepomonomys*, the TD1 analysis favored an age of 25.7 Ma (late Oligocene), i.e., far into the
 1002 youngest part of the age prior, with a 95% HPD interval (20.0-30.41 Ma) that excludes most of
 1003 the early Oligocene, and the Eocene entirely. Based solely on the fossils that have been described
 1004 from Silica North and Silica South thus far (Pickford et al., 2008), and the strongly conflicting
 1005 information provided by middle and late Eocene hystricognaths in northern Africa (Jaeger et al.,
 1006 1985, 2010a; Marivaux et al., 2014; Sallam et al., 2009, 2011, 2012), a late Oligocene age would
 1007 appear to us to make better sense of the rodent fauna from these localities, and could also explain
 1008 why Silica North would have a relatively derived rodent fauna without any of the immigrant
 1009 mammals that are thought to have arrived in Africa near the Oligocene-Miocene boundary
 1010 (Rasmussen and Gutiérrez, 2009). The recently described mammals from the Eocliff and Eoridge
 1011 sites (e.g., Pickford, 2015a, b), also in the Sperrgebiet area and also considered to be of
 1012 Bartonian age by Pickford et al. (2014), are in our opinion also consistent with a late Oligocene
 1013 age; the Eocliff tenrecoids (Pickford, 2015a) are morphologically intermediate between early
 1014 Miocene species and those known from the late Eocene and Oligocene of Egypt (Seiffert and
 1015 Simons, 2000; Seiffert et al., 2007; Seiffert, 2010), while the presence of a derived
 1016 anthracotheriid artiodactyl (Pickford, 2015b) — an immigrant clade that has never been found at
 1017 any of the earliest Priabonian Birket Qarun localities in the Fayum, and first appear in the Dir
 1018 Abu Lifa Member of the stratigraphically younger Qasr el-Sagha Formation — strongly supports
 1019 a maximum age of latest Priabonian for Eocliff, and more clearly an Oligocene age.

Similarly, Asian “baluchimyines” and *Turkanamys* have relatively old (early Oligocene) mean estimates within the broad (Oligocene-wide) uniform limits of their age priors. “Baluchimyine” species from the lower part (Bugti Member) of the Chitarwata Formation of Pakistan were initially thought to be early Miocene in age (Flynn, 1986), but there is now agreement that this unit is Oligocene, though estimates for its maximum age range from the early Oligocene into the late Oligocene (Lindsay et al., 2005; Metais et al., 2013). In the case of the five Chitarwata Formation “baluchimyine” species (of 12 known) sampled in this analysis, the resulting mean estimates (31.57 Ma for Y-GSP 417 and 29.9 Ma for Paali Nala C2) are consistent with the early Oligocene estimates of Welcomme et al. (2001) and Metais et al. (2013) based on biochronological interpretation of multiple vertebrate lineages. The mean age estimates for three species from Y-GSP 417 (as output in the MrBayes “.vstat” file) are tightly constrained at 31.7 Ma (28.1-33.9 Ma 95% HPD) for *Baluchimys ganeshapher*, 31.7 Ma (28.2-33.9 Ma 95% HPD) for *Lindsaya derabugtiensis*, and 31.3 Ma (27.1-33.9 Ma 95% HPD) for *Lophibaluchia pilbeami*, while *Bugtimys zafarullahi* and *Hodsahibia gracilis* from Paali Nala C2 were estimated to be 30.2 Ma (25.6-33.9 Ma 95% HPD) and 29.6 Ma (24.9-33.9 Ma 95% HPD), respectively. Lindsay et al. (2005) suggested that the baluchimyine-bearing base of the Chitarwata Formation is likely to be either ~29.8 Ma or ~25.8, and of these options our data support the former interpretation. Two other “baluchimyines” in the analysis, *Confiniummys* and *Ottomania* from Süngülü in Turkey, were thought to be close in age to the Eocene-Oligocene boundary (de Bruijn et al., 2003), but here were estimated to be about two million years younger (32.1 Ma), despite broad uniform priors extending from the end of the Oligocene (23 Ma) all the way back to the beginning of the late Eocene (37.8 Ma). Given the age estimates provided by the tip-dating analysis, *Baluchimys krabiense* from the Bang Mark Pit in Krabi, Thailand (Marivaux

et al., 2000) would be the oldest “baluchimyine” from Asia at 33.5 Ma (31.6-35.0 Ma 95% HPD). Finally, the 31.9 Ma (28.3-33.9 Ma 95% HPD) estimate that *Turkanamys hexalophus* provides for the Lokone Hill sites in the Turkana Basin of Kenya is consistent with the broad late early to late Oligocene age that was first considered by Ducrocq et al. (2010), though more recently a late Oligocene age has been favored by Marivaux et al. (2012). Seiffert (2012) suggested that the Lokone Hill sites were close in age to the boundary between the early and late Oligocene (~28 Ma).

Finally, our analysis has implications for the age of the Dur at-Talah faunas DT1 and DT2, which have been described by Jaeger et al. (2010a, b) on “stage of evolution” grounds as being late middle Eocene (Bartonian) in age, but were instead considered to be late Eocene (Priabonian) by Sallam et al. (2012) and Antoine et al. (2014). Both of these faunas present interesting problems in that they preserve remains of primitive species (“*Protophiomys*” *durattalahensis* at DT-1 and “*Protophiomys*” aff. *durattalahensis* at DT-2) that resemble *Waslamys* from BQ-2, combined with phiocricetomyines [*Talahphiomys libycus* (DT-1) and *Talahphiomys lavocati* (DT-2)] that have no relatives at BQ-2 and that more closely resemble species from the Oligocene levels of the Jebel Qatrani Formation. “*Phiomys*” *hammudai* from DT-1 shares derived features with *Acritophiomys* from the latest Eocene L-41 locality, as well as more derived phiomorphs from younger Fayum levels, and is not known from BQ-2. It is perhaps not surprising, then, that the tip estimates for the species from these localities are the most disparate of all in the analysis: the “*Protophiomys*” species are assigned mean estimates of 37.6 Ma (*durattalahensis*) and 38.2 Ma (aff. *durattalahensis*), which are in line with Jaeger et al.’s (2010a) late middle Eocene estimates, while the others Dur at-Talah species are estimated to be considerably younger: “*Phiomys*” *hammudai* is assigned a mean estimate of 35.6 Ma (1.7 Ma

younger than the overall mean estimate for BQ-2), and *Talahphiomys* species are estimated to be even younger (35.3 Ma for *T. libycus* and 33.6 Ma for *T. lavocati*). Ultimately, the mean estimates based on all species are 36.2 Ma for DT-1 (3 species) and 35.9 Ma for DT-2 (2 species) — i.e., intermediate in age between BQ-2 and L-41, as was argued by Sallam et al. (2012) and Seiffert (2012). The mean estimates for the ages of DT-1 and DT-2 do not require particularly fast or slow rates and are broadly consistent with adjacent branches (Fig. 13).

We would not argue that this method should be used in place of traditional biochronological methods that can be employed on landmasses with better sampling and sufficient species overlap, but it is certainly a more rigorous and repeatable approach than the data-free and assumption-laden “stage of evolution” arguments that have otherwise been applied to the ordering of terrestrial mammal sites in the very sparsely sampled Afro-Arabian Paleogene. We would expect this method to converge on increasingly robust age estimates as evidence is brought to bear from multiple biochronologically useful clades, and to be most useful when there are long and relatively well-dated reference sections available — for instance, in the case of Paleogene hystricognaths, the ability to integrate relatively tight age priors for species from the long Fayum succession undoubtedly helps to constrain several key parameters that in turn constrain estimates for poorly-constrained tips.

Incisive foramina of early phiomorphs

The most striking feature of the rostrum of *Birkamys* is the great enlargement of the apparently confluent incisive foramina, most clearly seen on the holotype specimen (CGM 66000, Fig. 3B) but also evident from the anteroposteriorly elongate, concave, and smooth lateral borders of the

foramen on isolated maxillae [DPC 9276 (Fig. 4E) and DPC 15625 (Fig. 4B)]. The same pattern holds for the one maxilla of *Mubhammys* that preserves this area (Fig. 7C). An anteroposteriorly enlarged incisive foramen was also identified in *Metaphiomys* by Wood (1968, his Fig. 6) and referred to as an “anterior palatine fenestra”; of this, he noted (pp. 51-52) that “ventrally, the anterior palatine fenestrae are large and not sunk into a palatal depression as much as in *Petromus* or *Thryonomys*, although there is a shallow depression lateral to the fenestra (fig. 6D), within which the fenestrae lie, which suggests the initial stages of a *Petromus* type of modification. The palatine fenestra is more rounded anteriorly than in the recent genera. No suggestion of an interpremaxillary foramen is present in any of the genera, although there is a paired foramen at the anterior end of the anterior palatine fenestra in *Petromus* and *Thryonomys* not seen in *Metaphiomys*. Posteriorly, there is a broad depression extending as far back as the anterior end of dP⁴ (Fig. 6A, 16), which seems to be identical to the deep fossa that contains the anterior palatine fenestrae in *Petromus* and *Thryonomys*.” With the recovery of fairly complete, but crushed, crania of *Acritophiomys* and *Gaudeamus* from L-41 (Sallam et al., 2011, 2012), it is now clear that enlarged incisive foramina are also present in those genera. The only known maxilla of *Waslamys* is not well-preserved, but appears to have a smooth margin anterior to the P⁴ that is similar to that of *Birkamys* (Fig. 16A). Among early Miocene taxa, similarly enlarged foramina are seen in *Lavocatmys aequatorialis* (Fig. 16F), *Simonimys genovefae* (Fig. 16H) and *Paraphiomys stromeri*. *Diamantomys leuderitzi* appears to have confluent foramina, but they are mediolaterally narrow and bordered by ventrally protruding flanges (Lavocat, 1973), perhaps correlated with the anteroposterior elongation of this region of the cranium. *Kenyamys mariae* also has more restricted foramina that are separated by a midline bony partition (Lavocat, 1973). In strong contrast, the incisive foramina of the Miocene bathyergoid *Renefossor songhorensis* are

1111 very small (Fig. 16C), and are either absent or tiny in *Proheliophobius leakeyi* (Fig. 16G).

1112 Among early Miocene bathyergoids, the foramina of *Efeldomys loliae* from Namibia (Mein &
1113 Pickford, 2008) appear to be the largest relative to tooth size (unless their apparent size is due to
1114 breakage), but are still much smaller than those of the L-41 species.

1115 From these observations the question arises as to whether the very enlarged foramina of
1116 *Acritophiomys*, *Birkamys*, *Gaudeamus*, and *Mubhammys* are primitive within Phiomorpha (or
1117 possibly even at some more inclusive level within Hystricognathi), or instead a synapomorphy of
1118 Thryonomyoidea (*Petromus-Thryonomys*) relative to Bathyergoidea, in which case *Efeldomys*
1119 and *Renefossor* retain the ancestral condition within Phiomorpha. The molecular divergence
1120 estimates of Patterson and Upham (2014) place the thryonomyoid-bathyergoid divergence at
1121 36.3 Ma, but the divergence of gaudeamurids from other hystricognaths in the matrix is
1122 estimated by the tip-dating analyses to have occurred ~39 Ma; assuming homology, this suggests
1123 an origin for such enlarged foramina well into the middle Eocene. There is no clear fossil record
1124 of the hystricid lineage before the Miocene, but, as in extant species, the late Miocene form
1125 “*Hystrix*” *gansuensis* has restricted foramina (Wang & Qiu, 2002). Early Oligocene *Tsaganomys*,
1126 which is aligned with “baluchimyines” in the tip-dating analysis, has restricted foramina as well
1127 (Bryant & McKenna, 1995). Among early caviomorphs, *Incamys bolivianus* has enlarged
1128 foramina, while *Branisamys luribayensis* does not (Patterson & Wood, 1983; their Figs. 14 and
1129 23). Unfortunately the origin of this distinctive feature cannot be adequately addressed without a
1130 matrix that samples more comprehensively from living and extinct ctenohystricans, but the
1131 recognition of the early ubiquity of this feature provides an interesting and easily identifiable
1132 new piece of evidence that will no doubt be of great importance for ongoing efforts to unravel
1133 early hystricognath phylogeny.

The broader question of the functional and behavioral implications of such variation will require much more research into the morphology of extant hystricognaths. The morphology seen in early phiomorphs might relate in some way to both the peculiar position of the vomeronasal organ of rodents, which opens anterior to the incisive foramina (Wöhrmann-Repenning, 1982; Giere et al., 1999), and the unique transformations of the rostral nasal skeleton and soft tissue structures of hystricognaths (Mess, 1999). For instance, Mess (1999) found that extant non-bathyergoid hystricognaths lacked fusion of the processus **lateralis ventralis and the lamina transversalis** anterior, leading to a non-continuous rostral nasal floor, and that the nasal septum forms a ventrally protruding keel onto which facial musculature inserts; she suggested that this morphology could lead to increased mobility of the rostral nasal skeleton in these forms. In contrast, bathyergoids have a continuous rostral nasal floor and much smaller incisive foramina; of interest in this regard is the observation that among the bathyergoid phiomorphs, the naked mole-rat *Heterocephalus* has a very small vomeronasal organ, perhaps relating to eusociality (Smith et al., 2007).

Conclusions

The late Eocene genera described here, *Birkamys* and *Mubhammys*, further expand the morphological diversity observable in the terminal Eocene radiation of Afro-Arabian hystricognaths, and provide the first compelling evidence for a key phiomorph synapomorphy — **suppression of P⁴/₄ eruption** — having evolved by the latest Eocene. In their lower molar morphology, *Birkamys* and *Mubhammys* show considerable similarity to members of the genus *Phiomys*; this shared morphology likely would have characterized the late Eocene-aged last common ancestor of the phiomorph clade that contains *Birkamys*, *Mubhammys*, and *Phiomys*.

The combined evidence from phylogenetic analysis and estimation of ancestral sizes of the first lower molar across early hystricognaths imply that *Birkamys* and *Mubhammys* underwent exceedingly rapid divergence in tooth size in the latest Eocene; the reasons for such a dramatic change are unclear. Cranial evidence from *Birkamys*, *Mubhammys*, and several Fayum and early Miocene has unexpectedly revealed that greatly enlarged incisive foramina are likely to be an ancient feature of phiomorph rodents, and possibly synapomorphic at an even more inclusive level within Hystricognathi.

Institutional abbreviations

CGM	Egyptian Geological Museum, Egypt
DPC	Duke Lemur Center Division of Fossil Primates
MUVP	Mansoura University Vertebrate Paleontology Center

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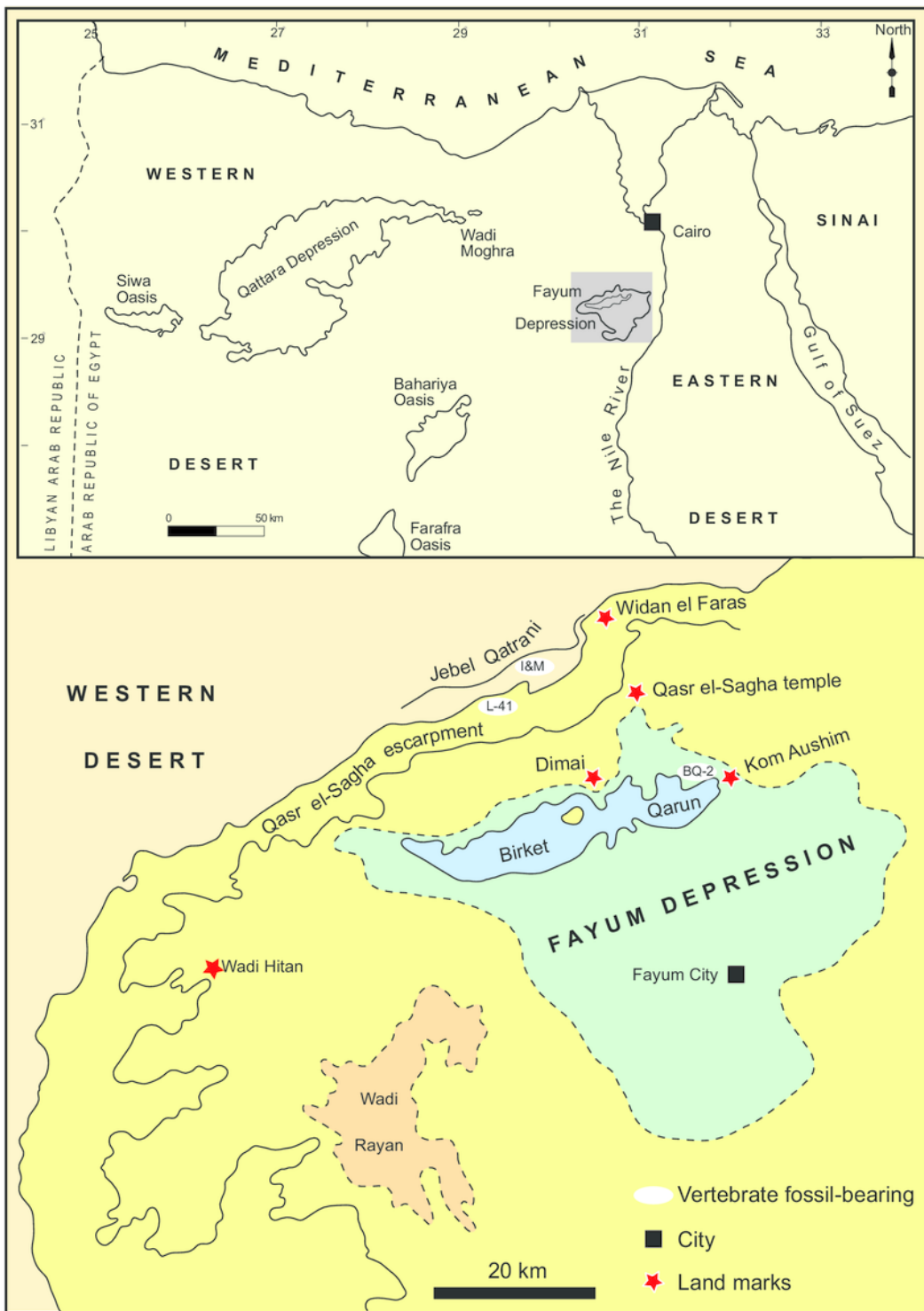
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1

Major fossil localities in the Fayum Depression of northern Egypt

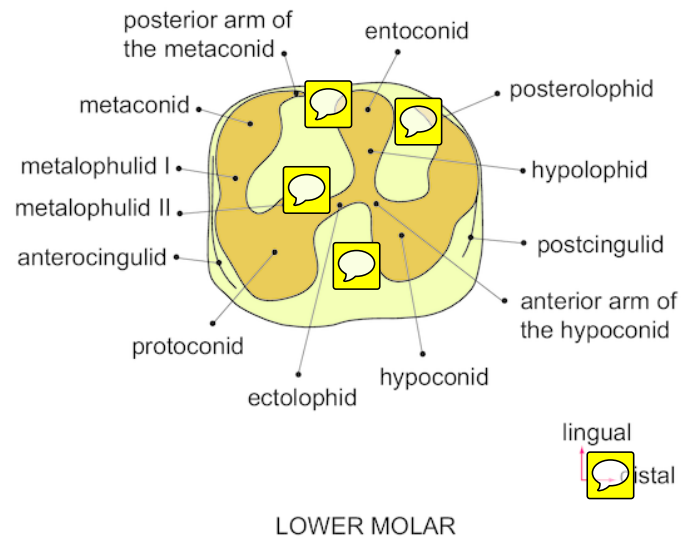
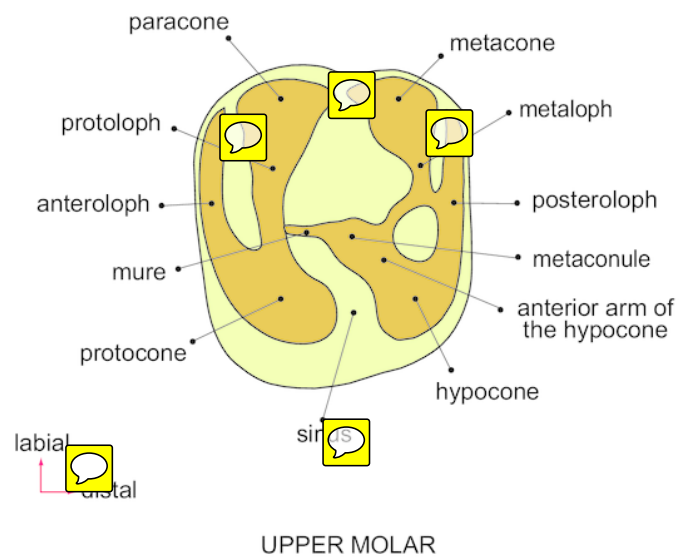
Left, stratigraphic positions and age estimates for major mammal-bearing fossil localities in the Fayum succession, and approximate position of Eocene-Oligocene boundary, following Seiffert (2006). Upper right, map of northern Egypt showing common landmarks and localities near the Fayum Depression. Lower right, map of the Fayum area showing the approximate position of Locality 41 (L-41), which occurs near the middle of the section approximately midway between the oldest rodent-bearing fossil locality (BQ-2) and the youngest rodent-bearing fossil localities (I&M). Modified after Bown and Kraus (1988).

Age	Rock unit	Thickness (m)	Quarries	Million years
Basalt (Widan El-Faras)		340		
early Oligocene	Jebel Qatrani Formation	Upper Sequence		
		250	M	30
		246	I	
		180	G & L-12	
		96	V	
late Eocene	Qasr el-Sagha Formation	Lower Sequence		
		65	E	34
		65	A & B	
		46	L-41	
		0		
late Eocene	Birket Qarun Formation	Temple Member		
		-171	BQ-2	37
		Um Rigi Member		



2

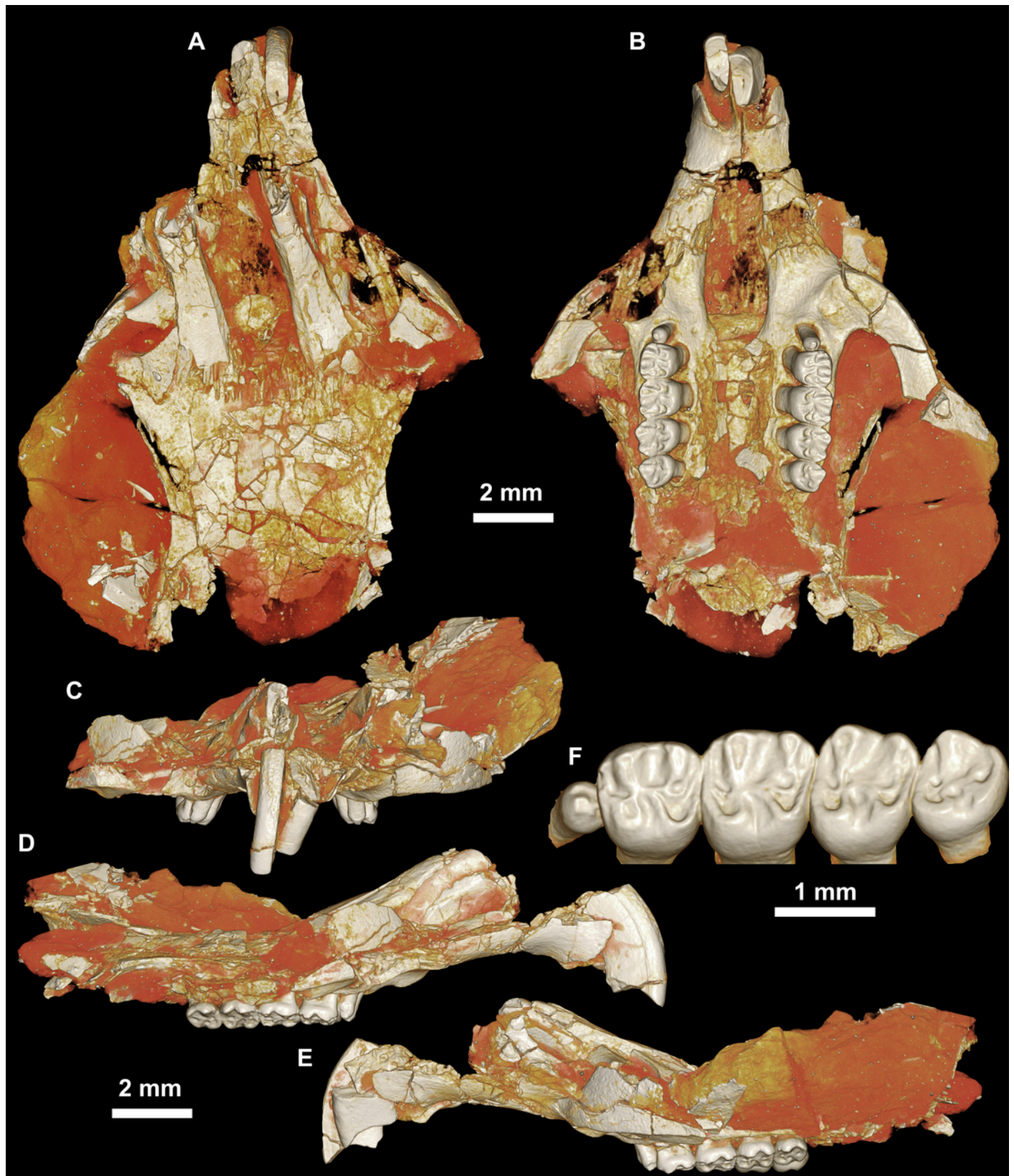
Dental terminology, based on Wood and Wilson (1936) and Marivaux *et al.* (2004).



3

CGM 66000, holotype cranium of *Birkamys korai*, new genus and species, from the latest Eocene Locality L-41, Jebel Qatrani Formation, Fayum Depression, northern Egypt.

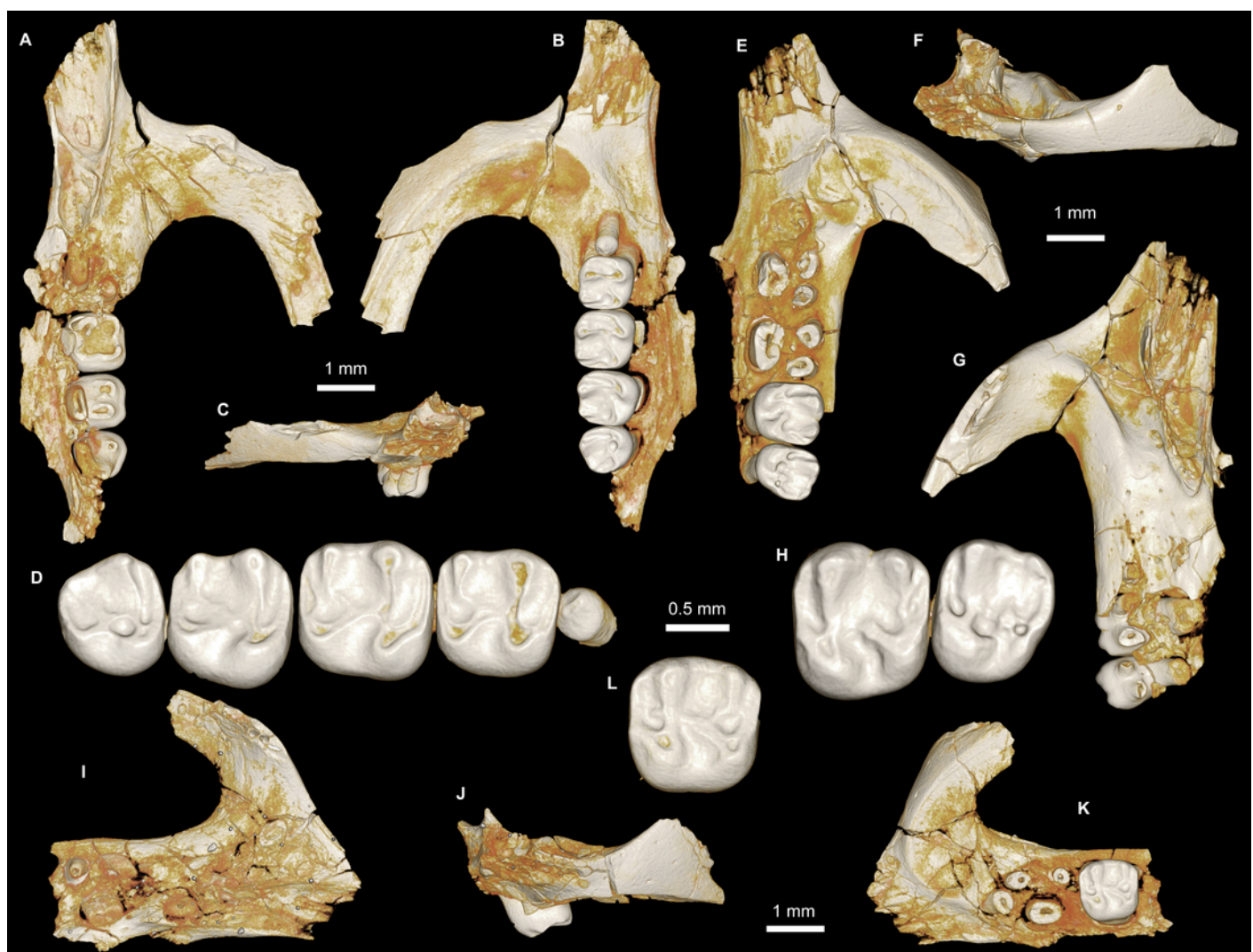
Volume renderings of high-resolution micro-CT scans of CGM 66000, in (A) dorsal view, (B) ventral, (C) anterior, (D & E) lateral views. Occlusal surface of the upper teeth is shown in (F).



4

Maxillae and upper dentition of *Birkamys korai*, new genus and species, from Quarry L-41.

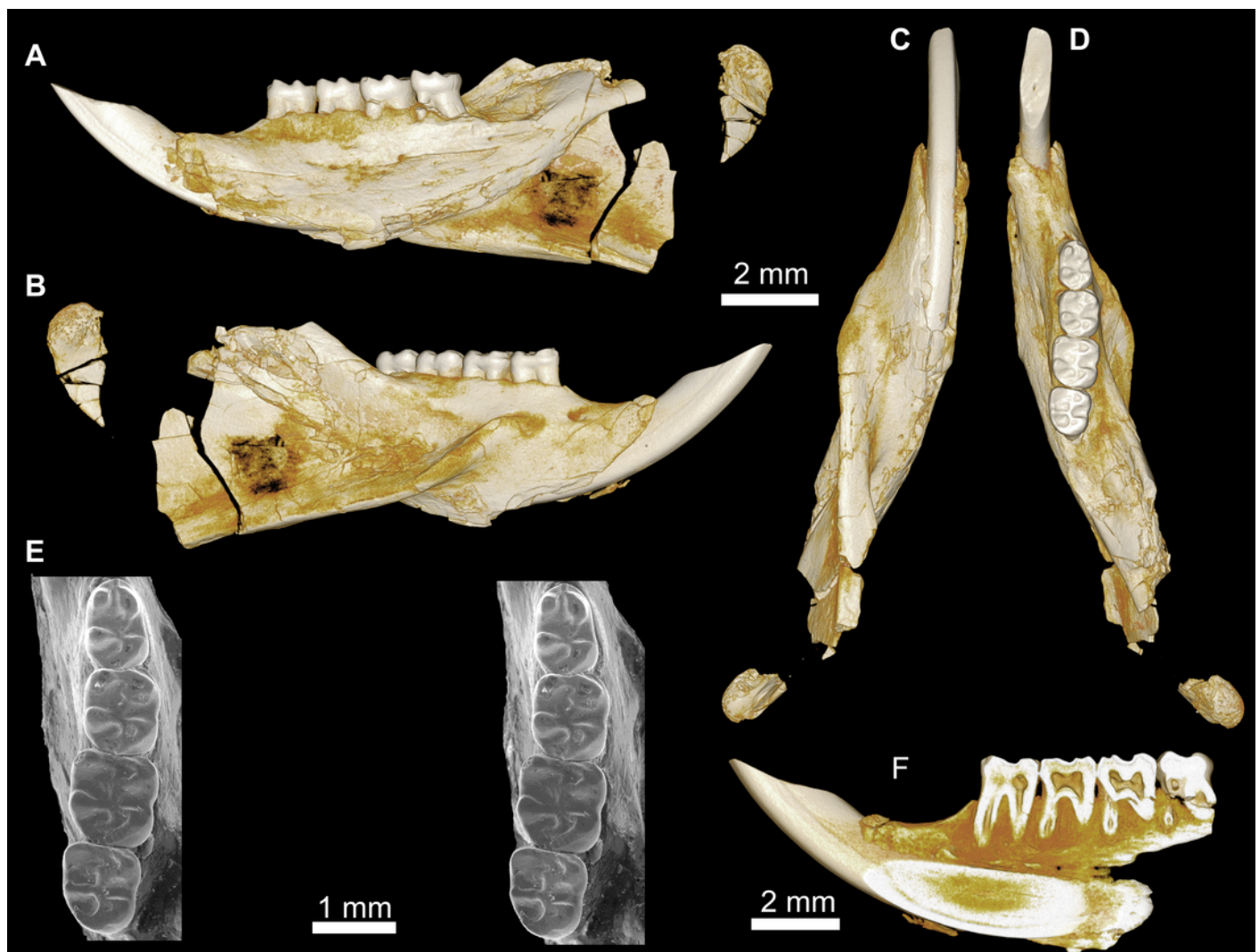
A-D) right partial maxilla of DPC 17457 with dP³-M³, in (A) dorsal, (B) ventral, (C) anterior views and (D) occlusal surface; E-H) DPC 9276, left partial maxilla with M²⁻³ and alveoli for dP³⁻⁴ and M¹, in (E) ventral, (F) anterior, (G) dorsal views and (H) occlusal surface; DPC 15625, left partial maxilla with M¹ and alveoli for dP³⁻⁴, in (I) dorsal, (J) anterior, (K) ventral views and (L) occlusal surface.



5

Mandible and lower dentition of *Birkamys korai*, new genus and species, from Quarry L-41.

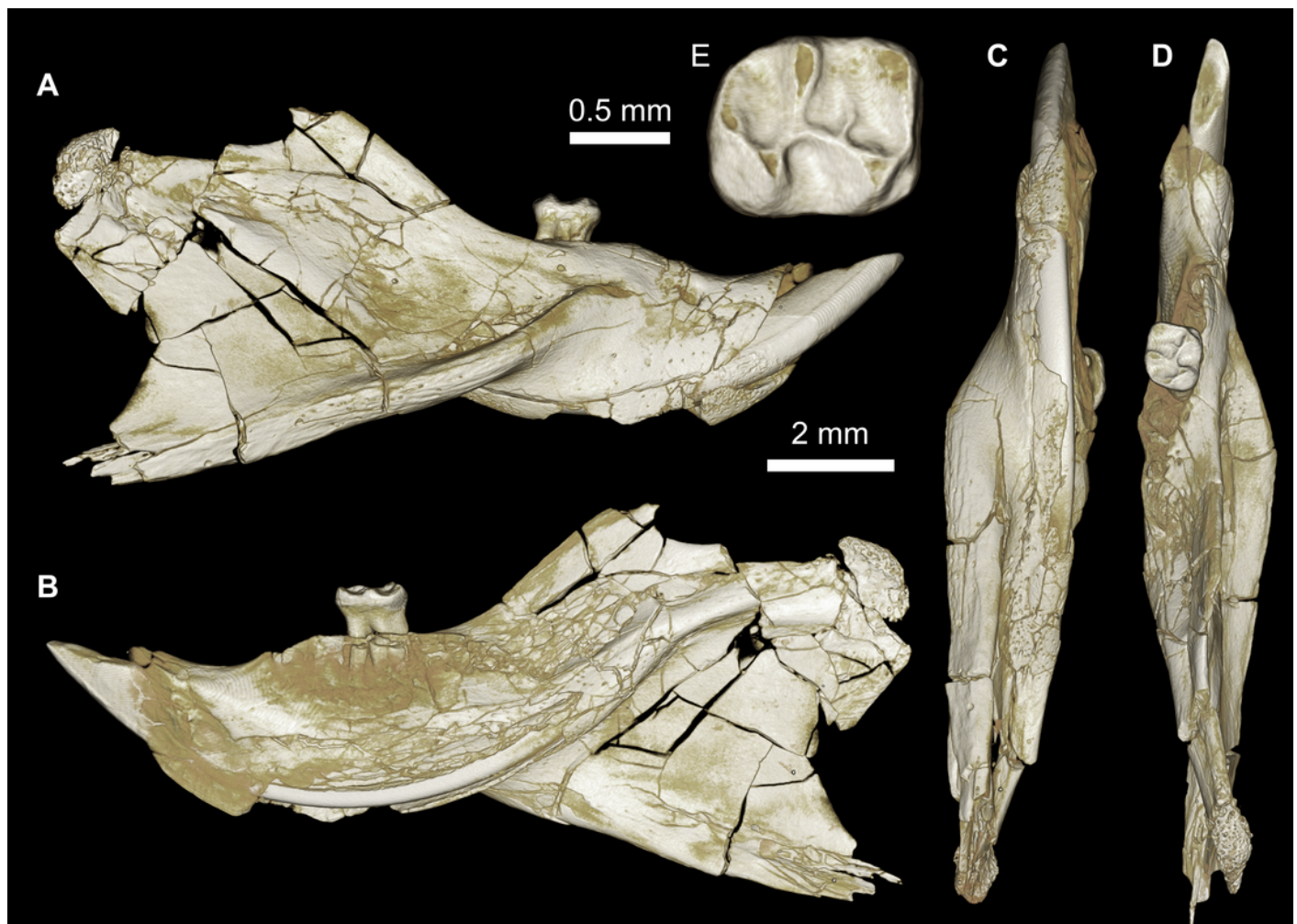
A-F) DPC 22737, left mandible with dP₄-M₃, in (A) medial, (B) lateral, (C) ventral and (D) dorsal views; (E) scanning electron stereopair illustrating the occlusal surfaces of dP₄-M₃; (F) micro-CT scans, illustrating the deep roots of the dP₄ in cross-section.



6

Mandible of cf. *Birkamys* from Quarry L-41.

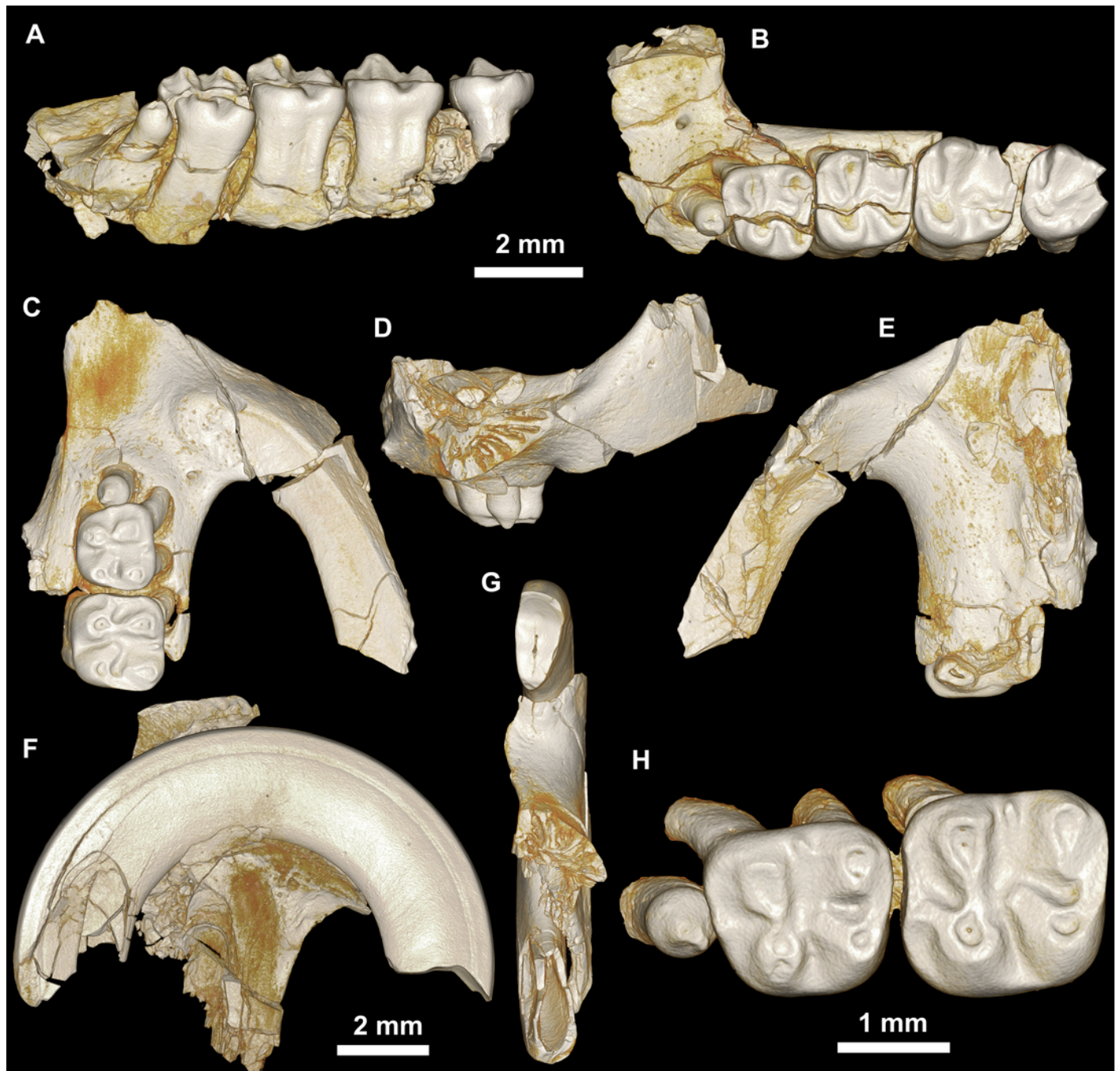
A-E) DPC 17995, almost complete right mandible with I and M₁, in (A) lateral, (B) medial, (C) ventral and (D) dorsal views; (E) occlusal surface.



7

Maxillae and upper dentition of *Mubhammys vadumensis*, new genus and species, from the latest Eocene Quarry L-41, Jebel Qatrani Formation, Fayum Depression, northern Egypt.

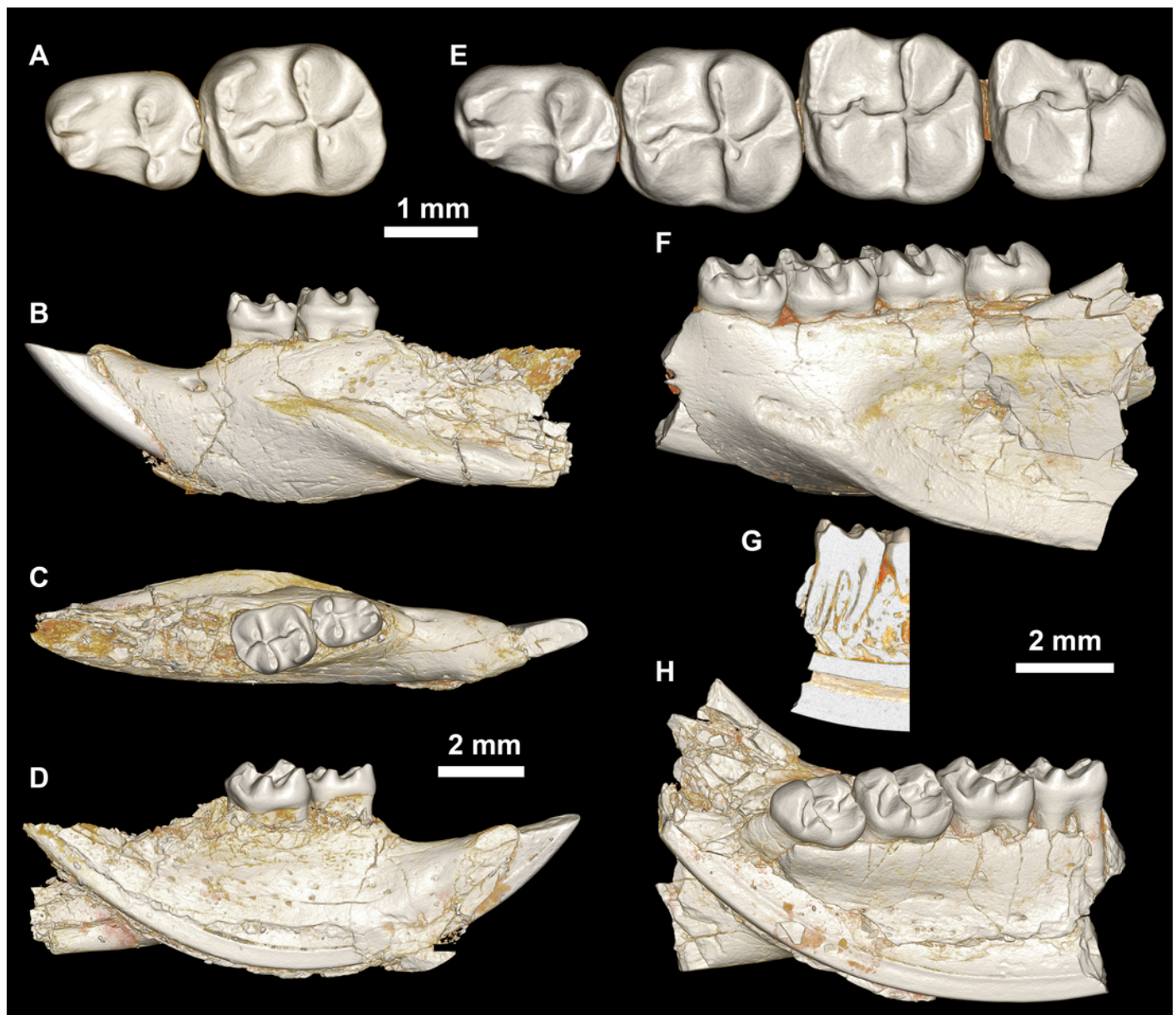
A-B) CGM 66001, holotype left partial maxilla (upper left incisor, dP^{3-4} and M^{1-3}) in (A) medial and (B) occlusal views; C-H) DPC 14324, left partial maxilla and upper dentition (I, dP^{3-4} and M^1) in (C) ventral, (D) anterior and (E) dorsal views; (F and G) medial and ventral surface of upper left incisor; (H) occlusal surface.



8

Lower dentition of *Mubhammys vadumensis*, new genus and species.

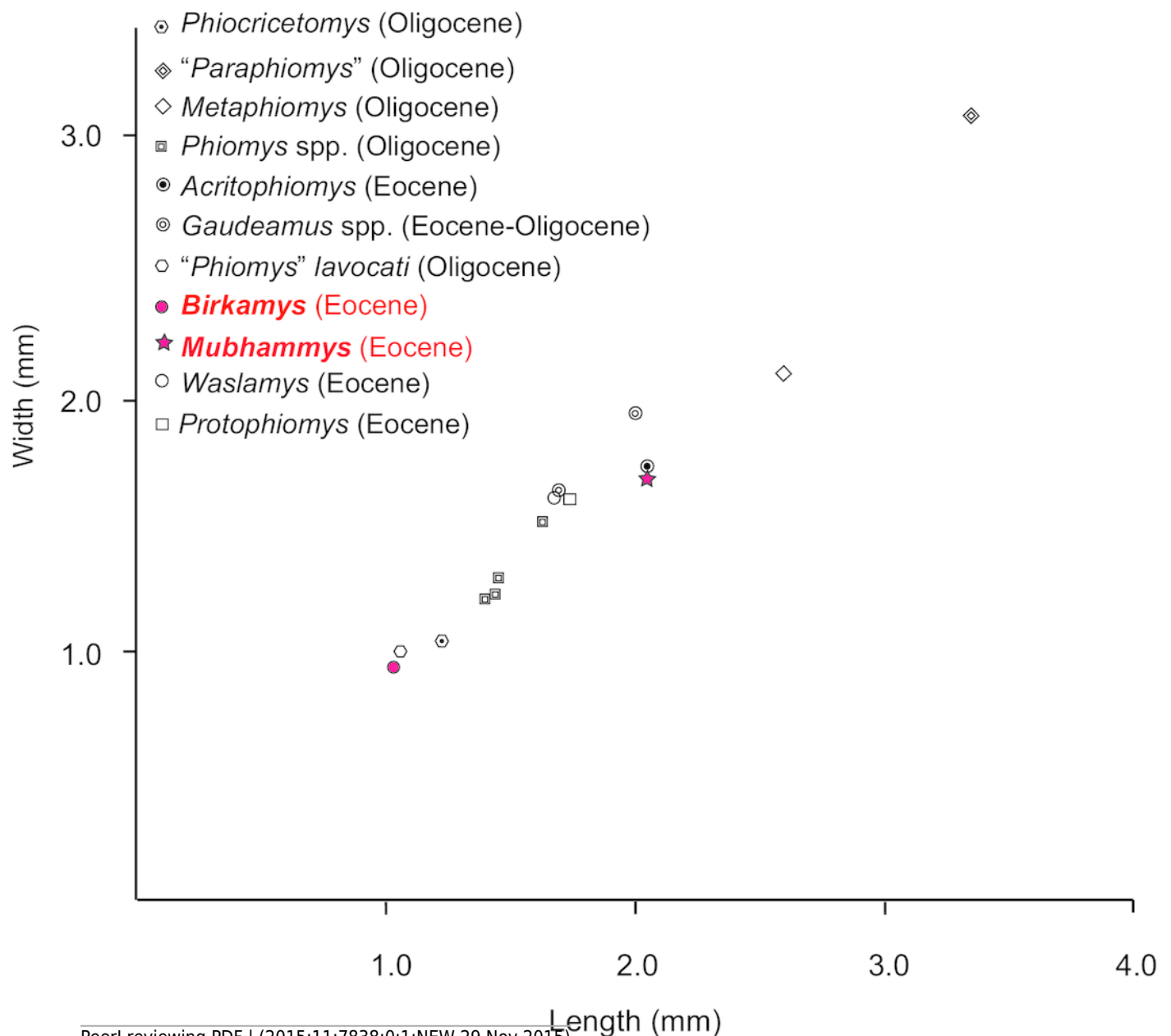
A-D) left mandibular fragment and lower dentition DPC 14141 (dP₄-M₁), (A) occlusal surface, (B) lateral, (C) dorsal and (D) medial views; E-H) left mandibular fragment and lower dentition DPC 13220 (dP₄-M₃), (E) occlusal surface, (F) lateral, and (H) medial views and (G) micro-CT scans, illustrating the deep roots of the dP₄ in cross-section.



9

Plot of length versus width of M_1 comparing *Mubhammys vadumensis* and *Birkamys korai* with other hystricognaths from the Fayum Depression.

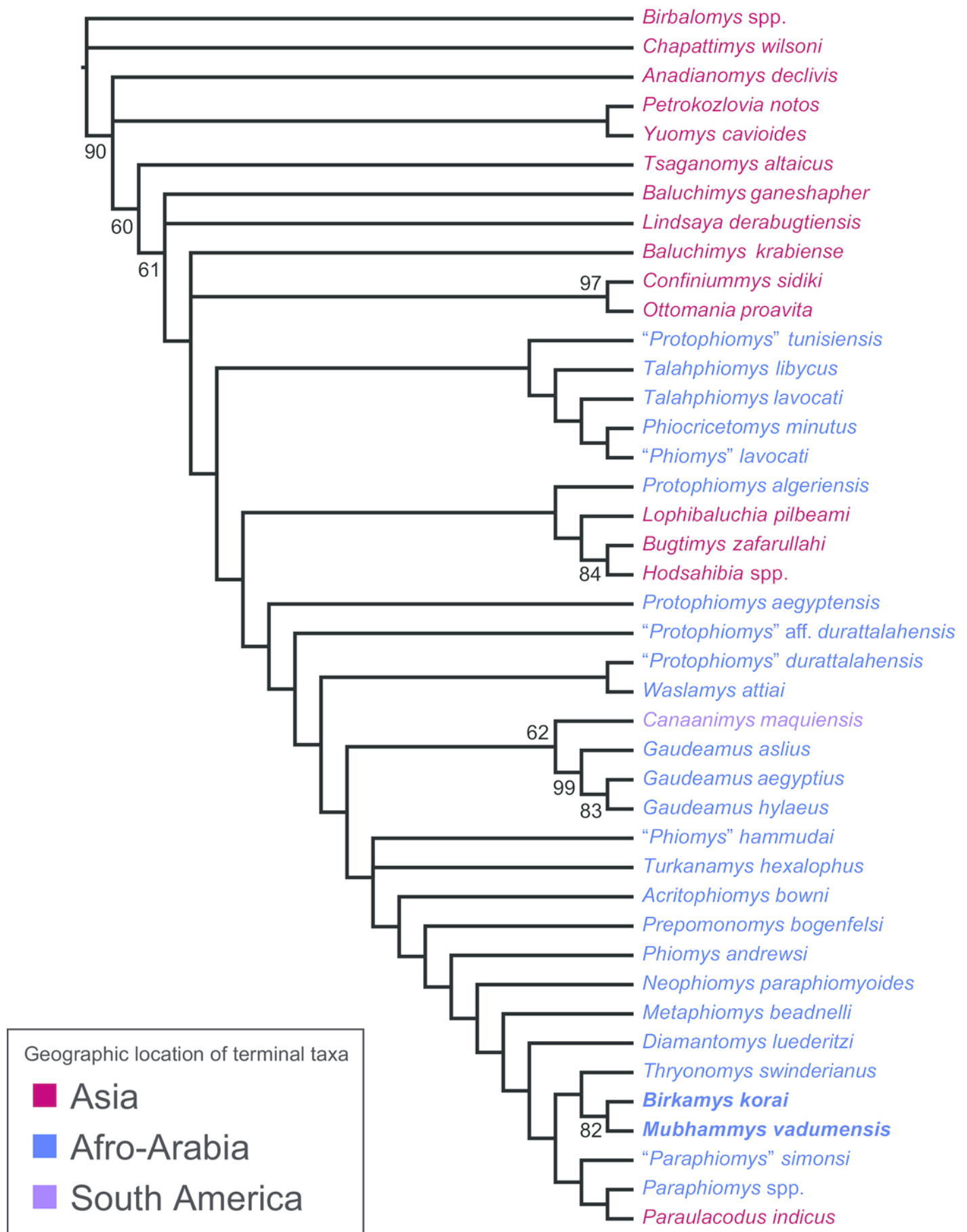
Measurements for "*Paraphiomys*", *Metaphiomys*, *Phiomys*, and "*Phiomys*" *lavocati* are from Wood (1968) and Holroyd (1994); *Gaudeamus* spp. from Wood (1968) and Sallam et al. (2011); *Acritophiomys* from Sallam et al. (2012); *Waslamys* and *Protophiomys* are from Sallam et al. (2009).



10

Parsimony analysis of living and extinct hystricognathous rodents, based on 118 morphological characters, largely from the dentition, 77 of which were treated as ordered and unweighted.

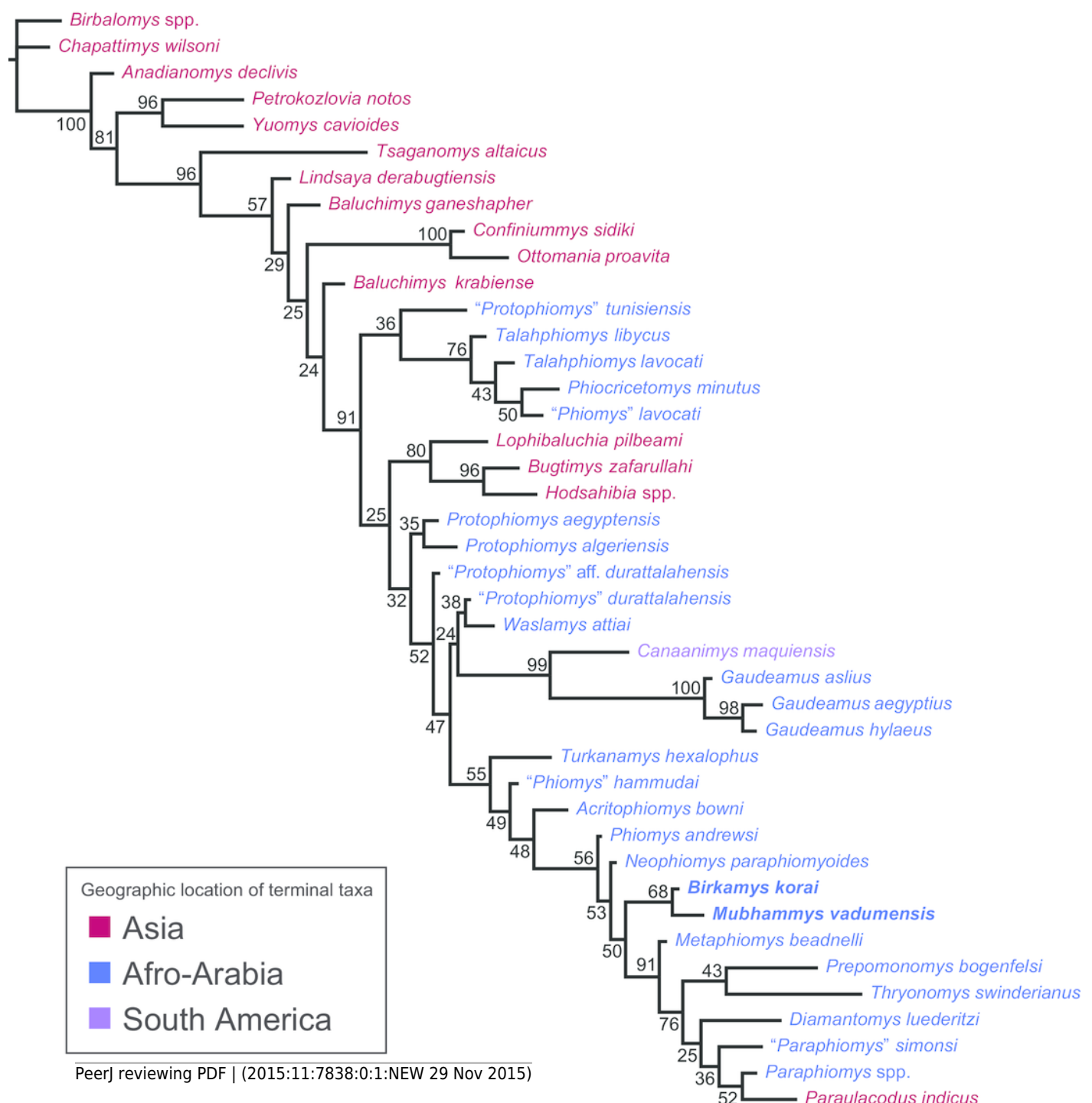
Strict consensus of 10 MPTs; tree length (TL) = 907; consistency index excluding uninformative characters = 0.2892; retention index = 0.5821; rescaled consistency index = 0.1720. Taxon names are colored according to their continental geographic location. Bootstrap support values, based on 1000 pseudoreplicates, are found above or below branches that are supported at a level >50%.



11

“Allcompat” consensus (majority-rule plus compatible groups) of 50,000 post-burn-in trees retained by Bayesian analysis of the 118-character matrix in MrBayes 3.2.5.

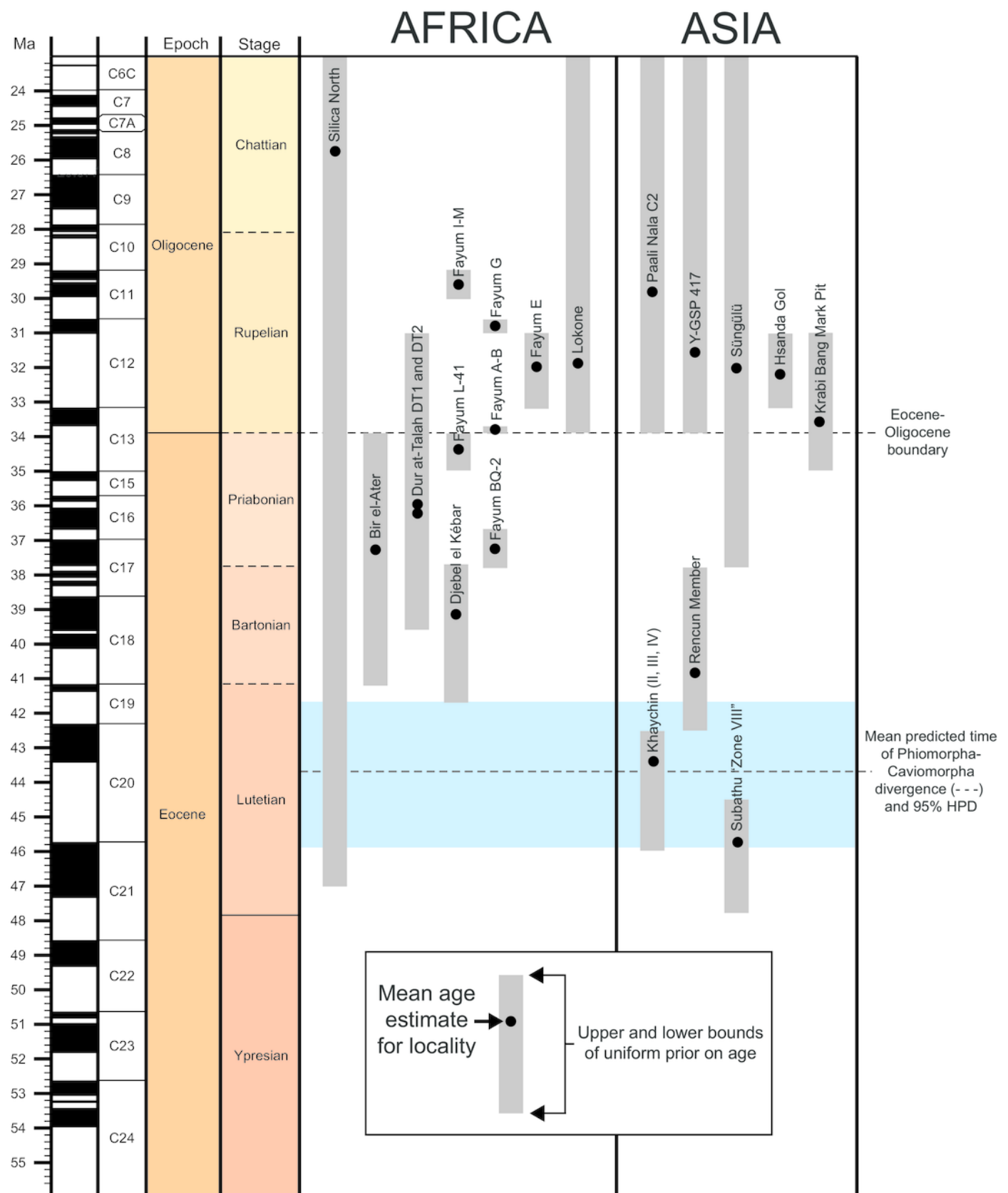
Taxon names are colored according to their continental geographic location. Numbers above or below branches represent posterior probabilities (x 100).



12

Uniform age priors for localities from which species in this analysis were sampled, with the sole or mean estimates provided by the tip-dating analysis (as output in the MrBayes “*.vstat” file.

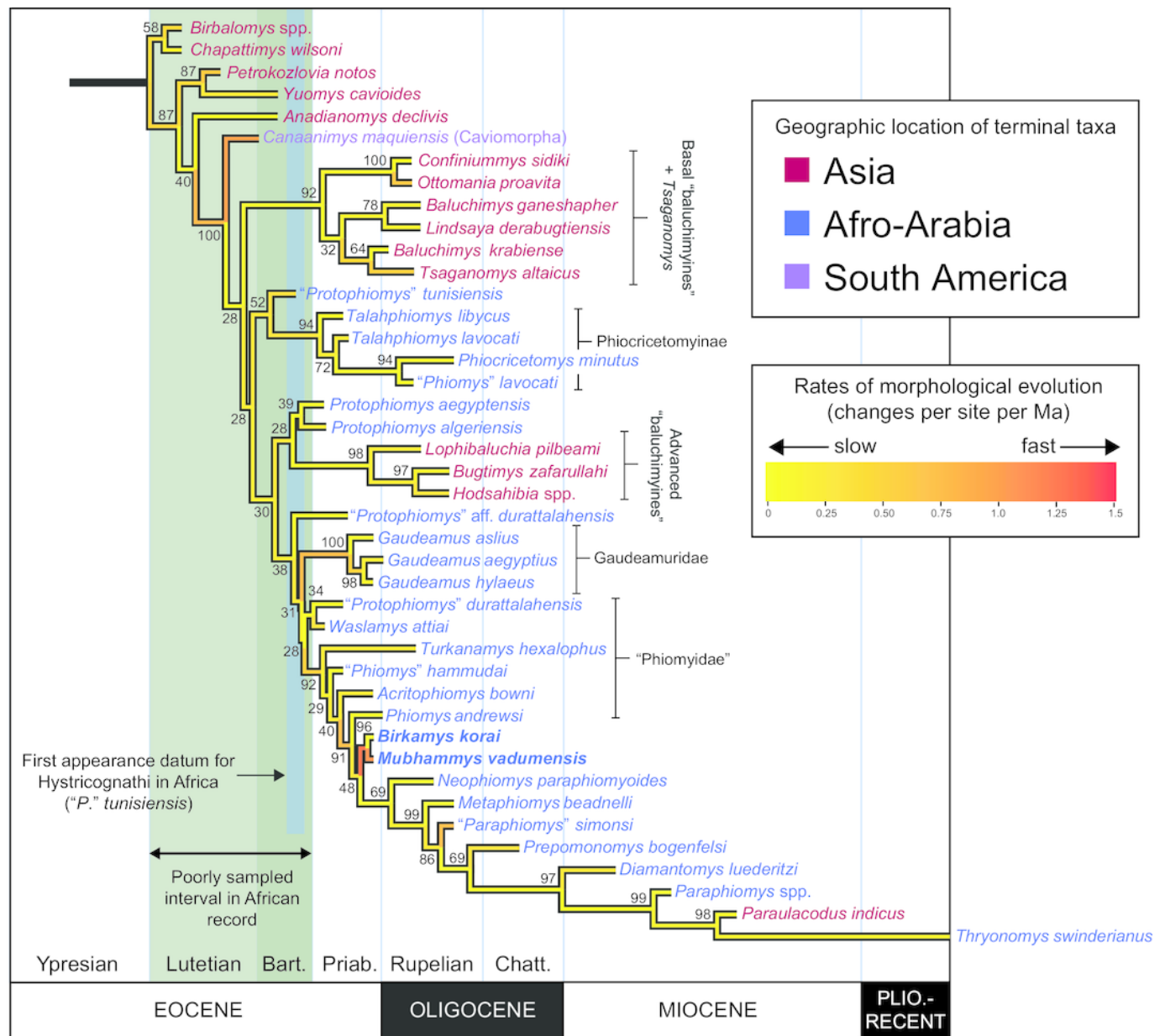
Geological timescale was created using the program TSCreator 6.4 (<http://www.tscreator.org>).



13

“Allcompat” consensus (majority-rule plus compatible groups) of 50,000 post-burn-in trees retained by tip-dating analysis (“TD2”) of the 118-character matrix in MrBayes 3.2.5 with locality ages fixed based on sole or mean estimates calculated by analysis

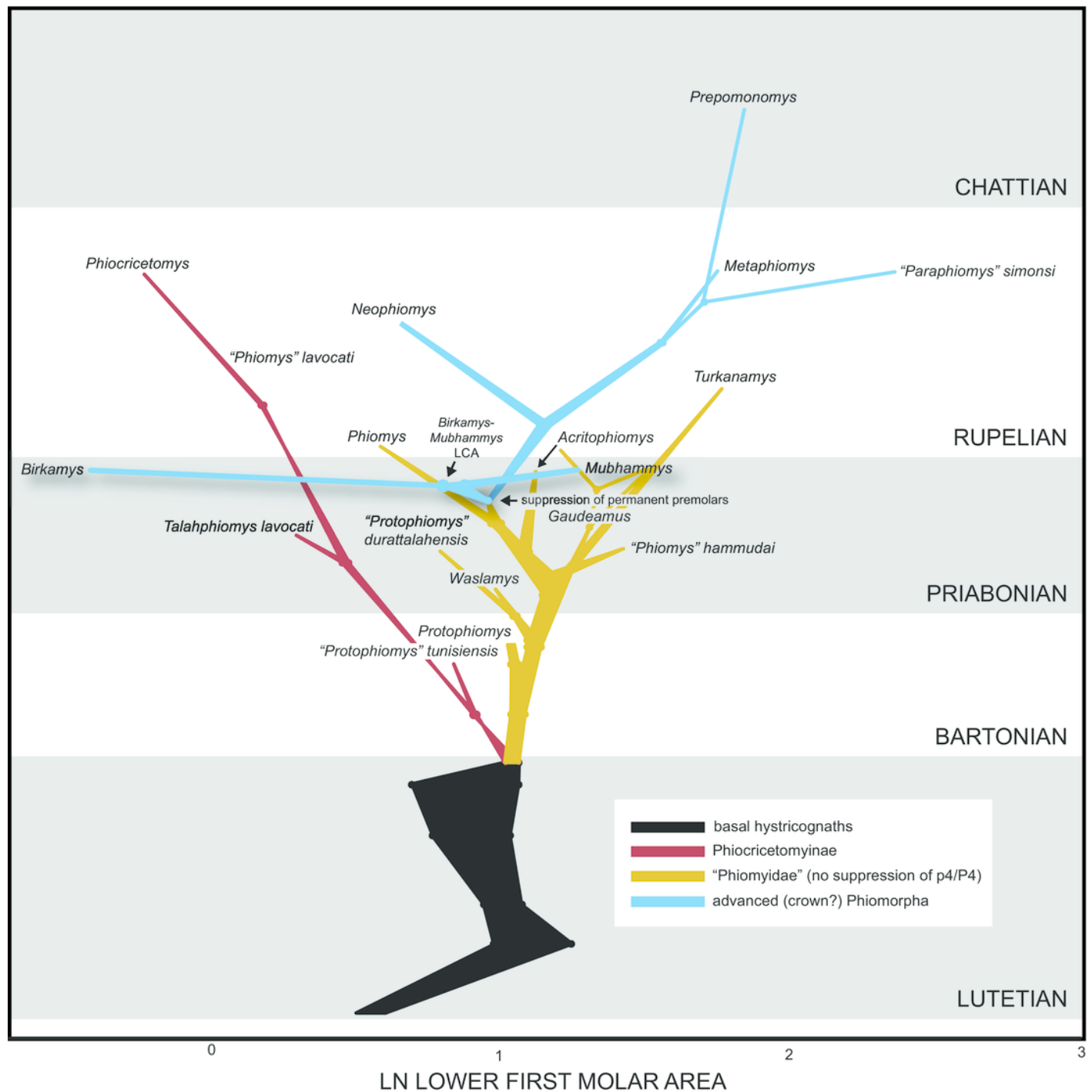
Branch lengths are colored according to rates of morphological evolution (calculated by multiplying the median rate for each branch, by the median clock rate for the entire tree), with the adjacent heat map showing the range of variation in the dataset. Taxon names are colored according to their continental geographic location. Numbers above or below branches represent posterior probabilities (x 100).



14

Evolution of lower first molar area (on a natural log scale) (x- axis) through time (y-axis), based on ancestral values calculated in BayesTraits, using both directional and random walk models, each with a lambda scaling parameter.

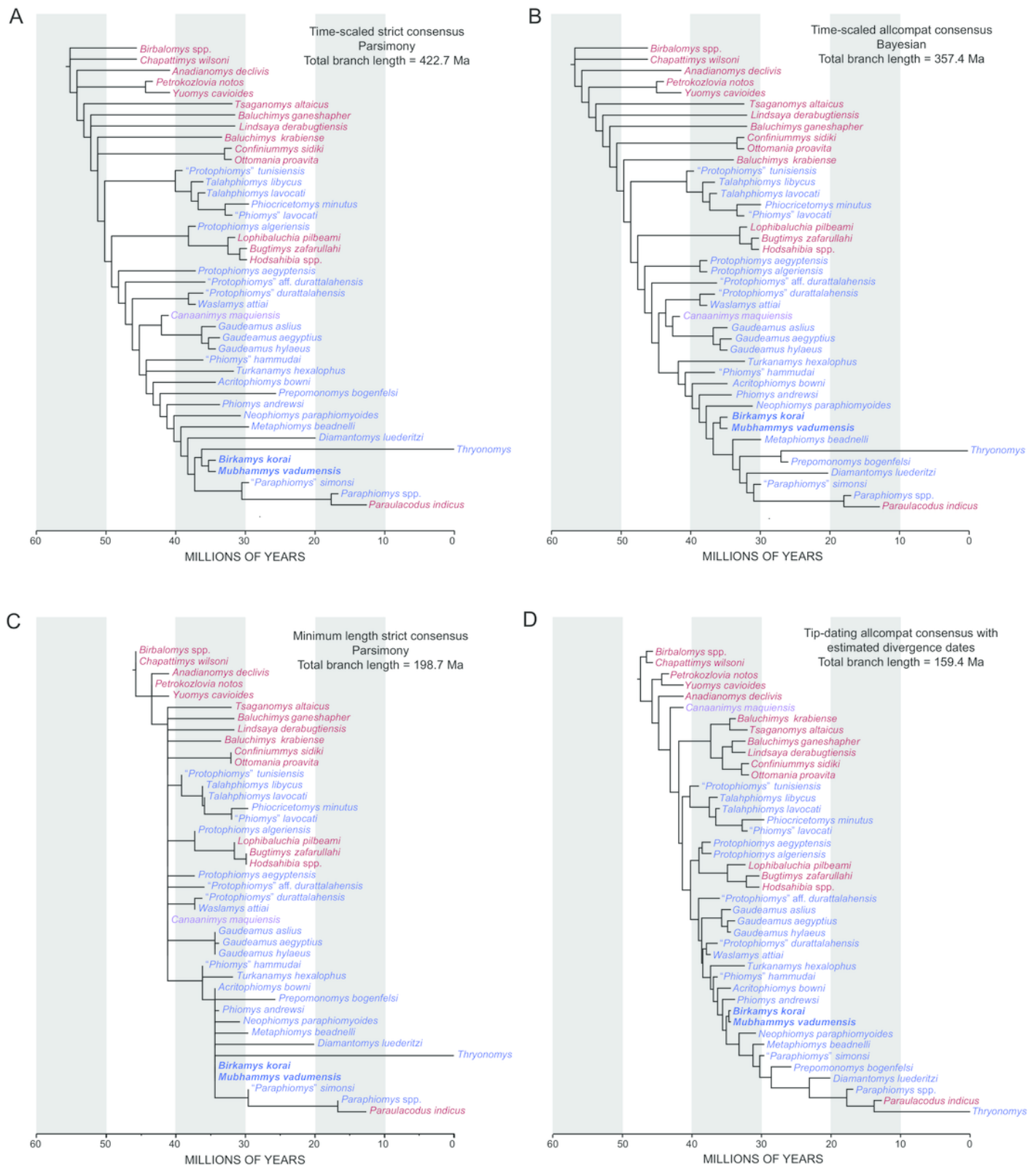
Ancestral values from each model are represented by points, and the intervening space at each node is infilled to reflect uncertainty between the mean estimates.



15

"Sum of branch lengths" calculation across time-scaled topologies.

Tip-dating topology (D) compared to time-scaled parsimony (A) and Bayesian allcompat (B) topologies (species durations and internodes are arbitrarily scaled to 1 Ma) and (C) minimum (0) length parsimony topology.



16

Comparison of the incisive foramina in late Eocene and early Miocene phiomorphs.

A) CGM 66000, late Eocene *Birkamys korai* from Quarry L-41; B) DPC 21311, late Eocene *Acritophiomys bowni* from Quarry L-41; C) KNM-SO 710, early Miocene *Renefossor songhorensis* from Songhor, Kenya; D) CGM 66006, late Eocene *Gaudeamus aslius* from Quarry L-41; E) mirror-imaged maxilla of CGM 83690, late Eocene *Waslamys attiai* from Locality BQ-2; F) KNM-SO 884, early Miocene *Lavocatomys aequatorialis* from Songhor, Kenya; G) KNM-RU 2318, early Miocene *Proheliophobius leakeyi* from Rusinga Island, Kenya; H) KNM-LG 834, early Miocene *Simonimys genovefae* from Legetet, Kenya.

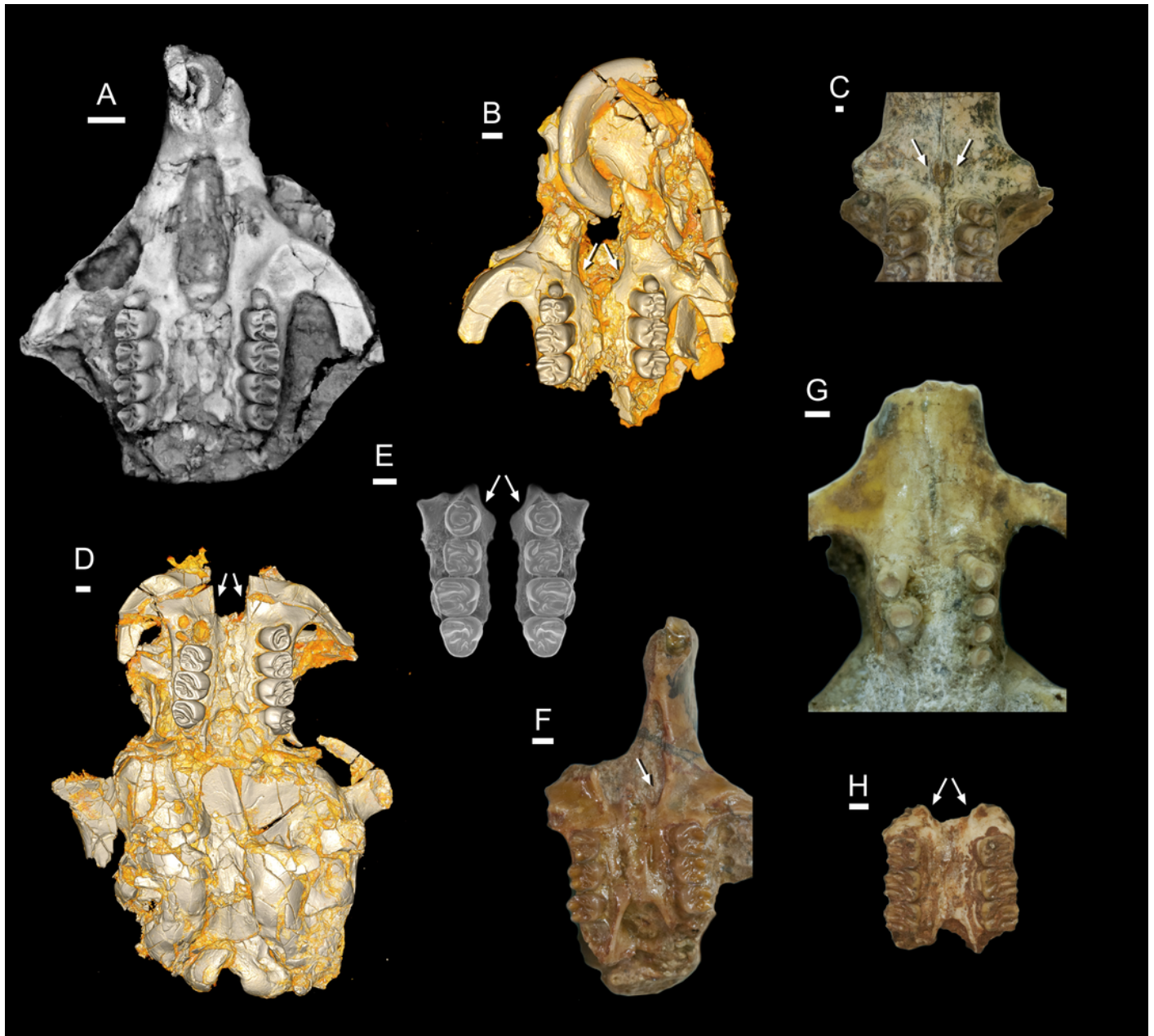


Table 1(on next page)

Table 1. Length and width of teeth (in millimeters) in the hypodigm of *Birkamys korai*, cf. *Birkamys*, and *Mubhammys vadumensis*.

- 1 **Table 1.** Length and width of teeth (in millimeters) in the hypodigm of *Birkamys korai*, cf.
- 2 *Birkamys*, and *Mubhammys vadumensis* gen. et sp. nov.

Birkamys korai

Specimen No.	Side	dP ³		dP ⁴		M ¹		M ²		M ³	
Upper teeth		L	W	L	W	L	W	L	W	L	W
CGM 66000 (Holotype)	Left	0.36	0.33	1.03	1.03	1.07	1.13	1.00	1.23	0.87	0.80
	Right	0.27	0.30	1.07	1.07	1.07	1.13	1.00	1.23	0.83	1.07
DPC 17457	Right	0.30	0.30	0.97	1.03	1.06	1.10	0.93	1.13	0.83	0.97
DPC 15625	Left	--	--	--	--	1.06	1.03	--	--	--	--
DPC 9276	Left	--	--	--	--	--	--	1.08	1.25	0.95	1.10
Lower teeth				dP ₄		M ₁		M ₂		M ₃	
DPC 22737	Right			1.10	0.76	1.03	0.93	1.10	1.06	1.03	0.96

Cf. *Birkamys korai*

Specimen No.	Side	M ₁	
DPC 17995	Right	1.10	0.87

Mubhammys vadumensis

Upper teeth		dP ³		dP ⁴		M ¹		M ²		M ³	
CGM 66001 (Holotype)	Left	0.68	0.76	1.8	1.9	1.90	--	1.95	2.10	--	1.80

DPC 14324	Left	0.58	0.68	1.70	1.60	1.85	1.80	--	--	--	--
<hr/>											
Lower teeth				dP ₄		M ₁		M ₂		M ₃	
DPC 13220	Left			1.83	1.43	2.00	1.75	1.95	1.80	1.83	1.70
DPC 14141	Left			1.66	1.29	1.90	1.67	--	--	--	--

Table 2(on next page)

Estimated mean ages and upper and lower HPD intervals for species with relatively broad uniform age priors, derived from the tip-dating analysis 1 (TD1) with a fossilized birth-death prior.

Table 2. Estimated mean ages and upper and lower HPD intervals for species with relatively broad uniform age priors, derived from the tip-dating analysis with a fossilized birth-death prior.

Taxon	Locality	Mean	L 95% HPD	U 95% HPD
<i>Baluchimys ganeshaper</i>	Y-GSP 417	31.66	28.11	33.90
<i>Baluchimys krabiense</i>	Krabi, Bang Mark Pit	33.54	31.59	35.00
<i>Bugtimys zafarullahi</i>	Paali Nala C2/Y-GSP 417(?)	30.21	25.64	33.90
<i>Confiniummys sidiki</i>	Süngülü	32.76	28.68	37.20
<i>Hodsahibia gracilis</i>	Paali Nala C2	29.64	24.93	33.90
<i>Lindsaya derabugtiensis</i>	Y-GSP 417	31.72	28.16	33.90
<i>Lophibaluchia pilbeami</i>	Y-GSP 417	31.34	27.09	33.90
<i>Ottomania proavita</i>	Süngülü	31.40	28.10	35.58
<i>“Phiomys” hammudai</i>	Dur at-Talah DT1	35.56	32.31	38.56
<i>Prepomonomys bogenfelsi</i>	Silica North	25.71	20.00	30.41
<i>Protophiomys algeriensis</i>	Bir el-Ater	37.27	33.90	40.45
<i>Protophiomys durattalahensis</i>	Dur at-Talah DT1	37.60	35.49	39.60
<i>Protophiomys aff. durattalahensis</i>	Dur at-Talah DT2	38.22	35.84	39.60
<i>Protophiomys tunisiensis</i>	Djebel el Kébar	39.17	37.70	41.18
<i>Talahphiomys lavocati</i>	Dur at-Talah DT2	33.65	31.00	36.60
<i>Talahphiomys libycus</i>	Dur at-Talah DT1	35.31	31.20	38.87
<i>Turkanamys hexalophus</i>	Lokone	31.88	28.30	33.90

Supporting Appendix 1: Justification for priors set on ages of included taxa. Note that upper and lower limits of magnetochrons is based on the recently revised GPTS presented by Ogg (2012).

Non-hystriognathous “ctenodactyloid” outgroups

Chapattimys wilsoni and *Birbalomys* spp. (primarily based on *Birbalomys sondaari*) are from the upper part (“faunal zone VIII”) of the Subathu Formation in northwestern India (Kumar et al., 1997; Gupta & Kumar, 2015). This level is considered to be “early” Lutetian in age by Gupta & Kumar (2015), but no specific upper boundary is provided for this estimate; however this interpretation gains support from the fact that older *Birbalomys* specimens (*Birbalomys* cf. *sondaari*) have recently been discovered in the later Ypresian of India (Gupta & Kumar 2015). Rather than assign included species a prior that covers the entire Lutetian, which could recover unrealistically young age estimates, we divided the Lutetian in half based on its currently recognized boundaries of 41.2–47.8 Ma on the Geological Time Scale 2012, leading to a uniform age prior of 44.5–47.8 Ma for a uniform “early” Lutetian prior.

Anadianomys declivis is from the Rencun Member in the lower part of the Heti Formation exposed in Henan, China (Tong, 1997), and is considered to fall within the Sharamurunian Asian Land Mammal Age based on its fauna (Russell & Zhai, 1987; Holroyd & Ciochon, 1994), an assessment that is supported by Appearance Event Ordination (Tsubamoto et al., 2004). Wang et al. (2007) consider the end of the Shuramurunian to be coincident with the end of the Bartonian (late middle Eocene), which is currently placed at 37.8 Ma on the Geological Time Scale 2012, but place the boundary between the Irдинmanhan and Shuramurunian Asian Land Mammal Ages at ~42.5, earlier than the currently recognized boundary between the Lutetian and Bartonian stages (41.2). Based on this combined evidence, we assigned *A. declivis* a uniform prior of 37.8–42.5 Ma.

Yuomys cavioides has been found in the Rencun Member in the lower part of the Heti Formation, as well as the Shara Murun and Jiyuan faunas (Russell and Zhai, 1987; Meng and McKenna, 1998; Bowen et al., 2002; Tsubamoto et al., 2004). All of these faunas are considered to fall within the Shuramurunian Asian Land Mammal Age; as with *Anadianomys declivis*, we accordingly restrict *Y. cavioides* to a uniform age prior of 37.8–42.5 Ma.

Petrokozlovina notos has been found in the Khaychin (II, III, IV) faunas on the Mongolian Plateau (Meng & McKenna, 1998; Bowen et al., 2002), which fall within the Irдинmanhan Asian Land Mammal Age (Tsubamoto et al., 2004). Wang et al. (2007) correlated the Arshantan-Irдинmanhan boundary with the Bridgerian-Uintan boundary of North America (~46 Ma), while the Irдинmanhan-Shuramurunian boundary was placed at ~42.5 Ma; accordingly we restrict *P. notos* to a uniform age prior of 42.5–46 Ma.

Asian “baluchimyines”

Confinniumys sidiki and *Ottomania proavita*, from Süngülü, Turkey (de Bruijn et al., 2003), have been interpreted as being of late Eocene or early Oligocene age based on comparison of these and other species (dipodids and murids) with roughly contemporaneous African and Asian species, but are poorly constrained due to the highly endemic nature of the fauna. We conservatively

place these species within the late Eocene-early Oligocene interval based on the current boundaries in the Geological Time Scale 2012 with a wide uniform prior (i.e., 28.1-37.8 Ma).

“Baluchimyines” from the Bugti Hills and the Zinda Pir Dome in Pakistan that were scored in this analysis — i.e., *Baluchimys*, *Bugtimys*, *Hodsahibia*, *Lindsaya*, and *Lophibaluchia* (Flynn et al., 1986; Marivaux et al., 2002; Marivaux & Welcomme, 2003) — all come from the lower part of the Chitarwata Formation (Welcomme et al., 2001; Lindsay et al., 2005; ~~Metais-Métails~~ et al., 2009) and are of contentious age. They were originally described as being of early Miocene age (Flynn et al., 1986), but are now universally considered to be of Oligocene age (Welcomme et al., 2001; Lindsay et al., 2005; ~~Metais-Métails~~ et al., 2009), though there is still debate about whether they are early Oligocene or late Oligocene in age. Two different interpretations of the magnetostratigraphic evidence proposed by Lindsay et al. (2005) suggest that the base of the formation could fall within Chron 11n.1r, or Chron 7Ar; i.e. either in the early Oligocene or late Oligocene. With this uncertainty in mind, we applied a wide uniform prior for the entirety of the Oligocene for these species (23-33.9 Ma).

Baluchimys krabiense, from the Bang Mark pit in the Krabi coal mine of Thailand (Marivaux et al., 2000), has been correlated with either Chron 12r or 13r (Benammi et al., 2001), the current boundaries of which form our uniform prior for this species (31 to 35 Ma).

Though the species has been found at several localities, the best-figured and described material of *Tsaganomys altaicus* from the Hsanda Gol Formation in Mongolia (Bryant & McKenna, 1995; Wang, 2001) is considered to be of early Oligocene age and was recovered from below a basalt that has been dated to ~31.5 Ma. We follow Kraatz & Geisler (2010) in considering these lower levels of the Hsanda Gol Formation to correlate with Chron 12r of the GPTS (31.03-33.16), and we use this range as the uniform prior on the appearance of *T. altaicus*’ diagnostic morphology as scored in our matrix.

Paleogene African and South American hystricognaths

Canaanimys maquiensis was recovered from the top of the ~~Yahuarango~~ Pozo Formation at the CTA-27 locality at Contamana, Loreto, ~~in~~ Peruvian Amazonia (Antoine et al., 2012), and the authors employed ⁴⁰Ar/³⁹Ar dating and mammalian biochronology to narrow the age estimate for the locality to 41.6-40.94 Ma, which we employ as the bounds of our uniform prior for this species.

“*Protophiomys*” *tunisiensis* has been found at the Djebel el Kébar locality in central Tunisia. Marivaux et al. (2014) reported that glauconite grains on the fossils from Djebel el Kébar returned K-Ar ages ranging in age from 38.7±1.0 and 40.7±1.0. We place a broad uniform prior on this species that takes into account the uncertainty in these K-Ar dates, ranging from 37.7 to 41.7 Ma.

Protophiomys algeriensis has been found at the Bir el-Ater or Nementcha locality in northern Algeria (Coiffait et al., 1984; Jaeger et al., 1985), which is dated entirely on the basis of its mammalian fossils. There is now a general consensus that Bir el-Ater is probably either earliest late Eocene (Priabonian) or latest middle Eocene (Bartonian) in age (Seiffert, 2010; Coster et al.,

Commentaire [LM1]: This debate concerns exclusively Lindsay.

The current debate is as to whether the Paali locality could be rather late Eocene (by some authorities).

Commentaire [LM2]: In the Bugti Hills, there is a long section above the level of the Paali nala locality. This long section, which is fossiliferous, is late Oligocene in age. So, the question as to whether the Paali Nala C2 is early or late Oligocene is “out of date”.

Commentaire [LM3]: Antoine et al., 2015 (updated version of the formation names).

2012; Marivaux et al., 2014), and the upper and lower boundaries of these stages delimit the uniform prior for this species (33.9–41.2 Ma).

Multiple species from two localities (DT-1 [*Phiomys* *hammudai*, *Protophiomys* *durattalahensis*, and *Talahphiomys libycus*] and DT-2 [*Protophiomys* aff. *durattalahensis*, *Talahphiomys lavocati*]) have been found along the Dur at-Talah Escarpment in central Libya (Jaeger et al., 2010), in beds that have been called the Idam Unit (Wight, 1980) or the Bioturbated Unit (Jaeger et al., 2010). Prior to the description of the species from DT-1 and -2, the Idam Unit mammals were most recently interpreted as possibly being intermediate in age between Quarries L-41 and A/B in the Fayum succession, putting them very close to the Eocene-Oligocene boundary (Seiffert, 2010). Jaeger et al. (2010) presented new magnetostratigraphic evidence, which showed that the entire sampled zone of the Dur at-Talah Escarpment is of normal polarity, but they favored a correlation of this normal polarity zone escarpment with C18n.1n, which would put those sites in the late middle Eocene (Bartonian). Sallam et al. (2012) and Antoine et al. (2012) have since favored an age intermediate between the Fayum localities BQ-2 and L-41, a zone that is, like the Dur at-Talah Escarpment, also of entirely normal polarity (Seiffert et al., 2005; Seiffert, 2006). To take into account the great uncertainty in the age of DT-1 and DT-2 reflected in these various interpretations, we set a wide uniform prior that spanned from the oldest possible age given the magnetostratigraphic correlation proposed by Jaeger et al. (2010) (39.6 Ma) to the younger bound of the Chron in which the only rodent species that Jaeger et al. (2010) consider to be shared with the Fayum succession (*Talahphiomys lavocati*) occurs (i.e., Fayum Quarry E), given Seiffert's (2006) preferred magnetostratigraphic correlation (31 Ma).

Commentaire [LM4]: A unique normal polarity is "evidence of nothing"....

Turkanamys hexalophus has been found solely at the LOK 13 locality at Lokone Hill, in the Lokone Sandstone Formation that is exposed in the Turkana Basin of northern Kenya (Ducrocq et al., 2010; Marivaux et al., 2012). The age of the site is constrained entirely by mammalian biochronology; based on available evidence Ducrocq et al. (2010) suggested an early late Oligocene age, and this estimate was followed by Marivaux et al. (2012). Given the ambiguity of the evidence that has been presented thus far, however, we considered it preferable to assign a broad uniform prior that encompassed the entire Oligocene (i.e., 23–33.9 Ma).

The Fayum succession provides important temporal control on this analysis because it includes several rodent faunas of various ages (Wood, 1968; Holroyd, 1994; Sallam et al., 2009, 2011, 2012), all of which have been tied into a single magnetostratigraphic column (Kappelman et al., 1992; Seiffert et al., 2005, 2008; Seiffert, 2006). Here we employ the magnetostratigraphic correlation that was proposed by Seiffert (2006), and use the upper and lower bounds of each chron to delimit the uniform priors for each species. Species from Locality BQ-2 (*Protophiomys aegyptensis* and *Waslamys attiai*) fall into a zone of normal polarity that Seiffert et al. (2005) and Seiffert (2006) correlated with Chron 17n.1n, the boundaries of which are now 36.7–37.8 Ma. As noted by Seiffert (2006), correlation of the zone of reversed polarity sampled at Quarry L-41 with Chron 13r of the GPTS would not rule out the possibility of an earliest Oligocene age (because Chron 13r is largely late Eocene but does cross the Eocene-Oligocene boundary), but he argued that a major unconformity just above the locality was likely due to near-coastal erosion associated with the major marine regression that occurred near the Eocene-Oligocene boundary. For this reason we have set a uniform prior for L-41 species (*Acritophiomys bowni*, *Birkamys korai*, *Gaudeamus aslius*, *Gaudeamus hylaeus*, *Mubhamys vadumensis*) that extends from the

base of Chron 13r to the Eocene-Oligocene boundary (35 Ma to 33.9 Ma), while *Gaudeamus aegyptius* and *Phiomys andrewsi* from Quarry A and Quarry B respectively (probably from the same zone of reversed polarity as that documented at L-41, as there are no intervening samples of normal polarity) are assigned a uniform prior that is post-Eocene (33.9) but before the termination of Chron 13r (33.7 Ma). Both *Gaudeamus aegyptius* and *Phiomys andrewsi* have also been identified at Quarry E (Holroyd, 1994), but note that their extensions to this younger level do not reflect age uncertainty but rather range extension, and in this case we are most concerned about the time by which the diagnostic morphology of *G. aegyptius* and *P. andrewsi* (as scored in this matrix) had appeared, not how long it persisted. Regarding other species, *Neophiomys paraphiomysoides* from Quarry G was placed in Chron 12n (30.6-31 Ma); “*Paraphiomys*” *simonsi* is only known from Quarries I and M, *Phiocricetomys minutus* is only found at Quarry I, and the definitive *Metaphiomys beadnelli* specimens that were scored for this matrix are based on specimens from I and M — both Quarries I and M fall within a zone of normal polarity that Seiffert (2006) correlated with Chron 11r (29.2-30 Ma); and “*Phiomys*” *lavocati* makes its first definitive appearance in the Fayum succession at Quarry E, which falls within Chron 12r (31-33.2 Ma) given Seiffert’s preferred correlation.

Neogene African hystricognaths

Diamantomys luederitzi has been found at several sites in east Africa and Namibia (e.g., Lavocat, 1973; Mein & Pickford, 2008), with the more complete specimens on which scoring is based having been found at Songhor in western Kenya. The fossils from Songhor are likely no older than 20 Ma (Pickford & Andrews, 1986; Cote et al., 2014), but fossils from Namibia are not as well constrained (Mein & Pickford 2008). We place a uniform prior on age of 19-21 Ma to partially account for this uncertainty.

Scorings for *Paraphiomys* spp. are based largely on specimens of *Paraphiomys pigotti* from Rusinga Island in western Kenya (Lavocat, 1973). This species has been found in both the Hiwegi and Kulu Formations on Rusinga (Lavocat, 1973; Peppe et al., 2009), based on available evidence these sites likely range in age from 15-17.8 Ma (Peppe et al., 2009), which we use as our uniform prior.

The age range of middle Miocene *Paraulacodus indicus*, from the Siwaliks of Pakistan, has been estimated by Flynn and Winkler (1994) as being 12.5-12.9 Ma, and we use these bounds as the limits of our uniform age prior for this species.

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