

# Resolving taphonomic and preparation biases in silicified faunas through paired acid residues and X-ray microscopy

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Paired petrography and acid maceration has shown that preferential silicification of shelly faunas can bias recovery based on taxon and body size. Here, silicified fossils from the Upper Ordovician Edinburg Formation, Strasburg Junction, Virginia, USA, were analyzed using X-ray tomographic microscopy ( $\mu$ CT) in conjunction with recovered residues from acid maceration of the same materials to further examine sources of potential bias. Results reveal that very small ( $< \sim 1$  mm) fossils are poorly resolved in  $\mu$ CT when scanning at lower resolutions ( $\sim 30\mu\text{m}$ ), underestimating abundance of taxa including ostracods and bryozoans. Acid maceration, meanwhile, fails to recover poorly silicified fossils prone to disarticulation and/or fragmentation during digestion. Tests for patterns of breakage, however, indicate no significant size or taxonomic bias during extraction. Comparisons of 3-D fossil renders and extracted residues of individual fossils reveal taxon-specific patterns of fragmentation and allow biostratigraphic and preparational breakage to be differentiated. Paleoecological ordinations and cluster analyses of datasets based on  $\mu$ CT data and residues produce generally concordant results but indicate that true variation between taxonomic composition of our samples is compromised by the resolvability of small size classes in  $\mu$ CT imaging, limiting the utility of this method for addressing paleoecological questions in these specific samples. We suggest that comparability of results will depend strongly on the sample size, taphonomic history, textural, and compositional characteristics of the samples in question, as well as  $\mu$ CT scan parameters; applying similar methods to different deposits will help to generalize conclusions drawn on the relative strengths and weaknesses of the methods.

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## 40 **Abstract**

41 Paired petrography and acid maceration has shown that preferential silicification of  
42 shelly faunas can bias recovery based on taxon and body size. Here, silicified fossils  
43 from the Upper Ordovician Edinburg Formation, Strasburg Junction, Virginia, USA, were  
44 analyzed using X-ray tomographic microscopy ( $\mu$ CT) in conjunction with recovered  
45 residues from acid maceration of the same materials to further examine sources of  
46 potential bias. Results reveal that very small ( $< \sim 1$  mm) fossils are poorly resolved in  $\mu$ CT  
47 when scanning at lower resolutions ( $\sim 30\mu\text{m}$ ), underestimating abundance of taxa  
48 including ostracods and bryozoans. Acid maceration, meanwhile, fails to recover poorly  
49 silicified fossils prone to disarticulation and/or fragmentation during digestion. Tests for  
50 patterns of breakage, however, indicate no significant size or taxonomic bias during  
51 extraction. Comparisons of 3-D fossil renders and extracted residues of individual fossils  
52 reveal taxon-specific patterns of fragmentation and allow biostratigraphic and  
53 preparational breakage to be differentiated. Paleoeological ordinations and cluster  
54 analyses of datasets based on  $\mu$ CT data and residues produce generally concordant  
55 results but indicate that true variation between taxonomic composition of our samples is  
56 compromised by the resolvability of small size classes in  $\mu$ CT imaging, limiting the utility  
57 of this method for addressing paleoeological questions in these specific samples. We  
58 suggest that comparability of results will depend strongly on the sample size,  
59 taphonomic history, textural, and compositional characteristics of the samples in  
60 question, as well as  $\mu$ CT scan parameters; applying similar methods to different deposits  
61 will help to generalize conclusions drawn on the relative strengths and weaknesses of  
62 the methods.

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## 78 **Introduction**

79 Silicification is a common preservational mode in the fossil record, especially in the  
80 early Paleozoic (Butts, 2007; Butts & Briggs, 2010), involving the fabric-specific replacement of  
81 calcareous biota (calcite and aragonite) with silica, typically as chalcedony, opal, or  
82 sometimes as sparry quartz cement. While the fine-scale mechanisms of silicification in  
83 limestone remain unclear, the overall process is simple: pore fluids containing dissolved  
84 silica (whether volcanic, hydrothermal, or biogenic in origin) undergo transport and react  
85 with the host rock, causing contemporaneous dissolution of calcite and/or aragonite and  
86 the precipitation of silica through mechanisms involving both pH change and pressure  
87 dissolution (Maliva & Siever, 1988).

88 Due to their enhanced resistance to other fabric-destructive diagenetic process (Schubert,  
89 Kidder & Erwin, 1997), silicified fossils are typically preserved in great fidelity and  
90 abundance. Moreover, the replacement process frequently replicates fine morphological  
91 details if not internal microstructure, and the insolubility of silica to dissolution by acid  
92 allows such material to be prepared chemically rather than mechanically (Baars, 2008;  
93 Butts, 2014). Acid maceration is a broadly employed technique used to extract silicified  
94 fossils from limestone, relying on the high solubility of calcite in acidic solution and the  
95 comparatively low resistance of micritic matrix to alteration or dissolution (Donald W. St.  
96 Clair, 1935; Grant, 1989). However, such extraction methods can introduce potential  
97 biases that might skew perceptions of paleoecological studies investigating silicified  
98 specimens (Pruss, Payne & Westacott, 2015). Two primary sources of bias that can readily be  
99 identified include the differential likelihood that a given specimen will (1) silicify during  
100 diagenesis, and (2) once silicified, survive preparation and be recovered.

101           If biases in differential preservation and recovery methods go unaccounted for,  
102 certain types of analyses can be more severely affected. Systematic descriptions of  
103 organisms may be only mildly affected by the underrepresentation of some ontogenetic  
104 stages relative to others, but studies specific to the paleoecology or biofacies of an  
105 assemblage can suffer in quality due to their reliance on an accurate assessment of  
106 which biota are present in the deposit. For example, the near absence of a guild of  
107 predators or the preferential breakage of certain types of shell ornamentation might  
108 drastically alter the interpretation of trophic relationships and biotic interactions.  
109 Similarly, the loss of fossils in a certain size range during preparation can affect  
110 interpretations of biostratigraphic processes. Because these biases may skew  
111 paleoecological interpretations of fossil assemblages, quantifying or at least  
112 constraining them can improve confidence in the results of research relying on this  
113 common preparatory method.

114           Previous research has attempted to account for preservational and preparation  
115 bias in acid maceration of silicified fossils. Pruss et al. (2015) compared the relative  
116 abundances of fossil taxa in 11 hand samples of the Triassic Virgin Limestone Member  
117 of the Moenkopi Formation, from which the authors made a petrographic thin section  
118 and extracted fossil specimens via buffered acetic acid maceration. In general,  
119 echinoderms and gastropods were disproportionately over-represented in residues while  
120 bivalves were more likely to be identified in thin-section point counts (Pruss, Payne &  
121 Westacott, 2015).

122           Two questions not addressed by the aforementioned study are those of body  
123 size bias and breakage patterns. Since petrographic analysis can identify fossils by their

124 taxonomic group but offers minimal useful information with respect to their dimensions,  
125 detecting differences in body size distribution between thin section grains and fossils  
126 recovered in residues is not feasible. Similarly, it is impossible to tell by extraction  
127 methods whether fossils in residues suffered damage during digestion and washing or  
128 whether the breakage occurred prior to burial. Because thin sections and residues do  
129 not record the same fossils but rather disjoint subsets of the total fossils within the hand  
130 sample, no true pre and post comparison can be made. Indeed, since both methods  
131 applied in the Pruss et al. (Pruss, Payne & Westacott, 2015) study are destructive, a single fossil  
132 cannot be recorded by both. This inherent limitation can be avoided, however, by  
133 pairing a non-destructive whole-rock analytical method with acid maceration. X-ray  
134 tomographic microscopy ( $\mu$ CT) is well-suited to such an approach, allowing  
135 compositional and textural differences within a sample to be imaged as a grayscale  
136 representation of density variation in three dimensions, revealing fossil material within.  
137 By then subjecting the sample to acid maceration, individual fossils can be resolved and  
138 measured before and after chemical preparation, and any breakage or loss documented  
139 and quantified.

140         Herein, we employ a paired approach combining  $\mu$ CT-generated three-  
141 dimensional volume data of bulk rock samples with investigation of subsequently acid-  
142 extracted insoluble residues from the same scanned samples test the hypothesis that  
143 smaller fossils would be disproportionately less likely to survive preparation and be  
144 recovered from acid-digested residues and to further assess and quantify introduced  
145 biases and their implications for paleoecological studies. For this study, we used  
146 prepared cores of carbonate bulk rock samples from the Edinburg Formation of Virginia,

147 USA, which is well-known to host abundant and diverse silicified Ordovician marine  
148 fossils (Whittington & Evitt, 1953; KRAFT, 1962). The pervasive silicification of the fossils within  
149 this unit makes it amenable to delivering sufficient contrast between the skeletal  
150 material and the host matrix using  $\mu$ CT.

151

## 152 **Materials & Methods**

### 153 **Geological Setting**

154 The Edinburg Formation is an Upper Ordovician unit of massively bedded black  
155 limestone with occasional shaley interbeds, representing deep ramp to basinal  
156 deposition at the northwest of the Taconic foreland basin system after drowning of  
157 shallower ramp carbonate facies including the underlying Lincolnshire Limestone (Holland  
158 & Patzkowsky, 1996). Potassium bentonites derived from Taconic volcanism are sporadically  
159 present in the sediments of the foreland basin and have been sampled for radiometric  
160 ages, most notably the Millbrig bed, which lies up section of the Edinburg in the  
161 Martinsburg Formation and has a U-Pb date of 452.86 Ma (Mitchell et al., 2004; Sell, Ainsaar &  
162 Leslie, 2013).

163 All samples in this study were collected from the Liberty Hall facies of the  
164 Edinburg Formation exposed at Strasburg Junction rail cut, a well-studied site in the  
165 Shenandoah Valley of Virginia (Fig. 1, Table 1; (Cooper & Cooper, 1946; Whittington & Evitt, 1953;  
166 Read, 1980; Jacobs & Carlucci, 2019). Hand sample observations revealed two starkly different  
167 lithologies (Table 1). Rocks from Horizon 1, approximately 0.5 m above the contact  
168 between the Lincolnshire Limestone and Edinburg Formation, were coarse-grained, with  
169 a sparry or dismicritic texture likely indicative of recrystallization during early diagenesis.  
170 Sparse fossils and possible intraclasts visible on fresh surfaces were weakly aligned

171 with the direction of original bedding, suggesting directional sorting during deposition.  
172 Horizons 3, 4, 5, and 6 cropped out over 30 m up section (Table 1) and were finely  
173 grained with only occasional sparry or hematitic grains. Fossils, where visible on fresh  
174 or weathered surfaces, were oriented randomly to original bedding and comprised  
175 arthropods, brachiopods, and occasional crinoid ossicles. Rocks from these horizons  
176 also displayed heavy rinds on weathered surfaces, frequently stained rusty orange to  
177 pale yellow with iron oxides/oxyhydroxides.

### 178 **Fossil Extraction**

179 From the five horizons with fossils visible on weathered surfaces, 15 samples were  
180 prepared as cylindrical cores (Table 1), approximately 2 cm in diameter and ranging  
181 from 1 to 4 cm in height, normal to bedding. Each core was imaged via X-ray  
182 tomographic microscopy using a Zeiss Xradia 510 Versa X-ray microscope with an  
183 isotropic voxel size of approximately 30  $\mu\text{m}$ , producing three-dimensional renders of the  
184 core interiors. All  $\mu\text{CT}$  scans were processed using Dragonfly software Build 941–  
185 v.4.2.2 for Windows, Object Research Systems (ORS) Inc, Montreal, Canada, 2018  
186 (<http://www.theobjects.com/dragonfly>). Post-image processing within the software was  
187 conducted to reduce imaging artifacts using the Ring Removal and Median filters.  
188 Regions of high brightness (corresponding to high-density ferrous material) and low  
189 brightness (corresponding to low-density siliceous material) relative to the local  
190 background were thresholded and manually segmented, excluding regions at the very  
191 top and bottom of cores due to persistent boundary artifacts. Due to radial and  
192 longitudinal variation in background brightness across the cores, multiple brightness  
193 thresholds were used to segment siliceous material in some cores; in such cases, the

194 total siliceous volume is the union of the volumes from segmentation by the various  
195 thresholds. Segmented volumes for ferrous and siliceous material were refined by  
196 removing small (<100 voxel total volume) islands and were then partitioned into distinct  
197 regions-of-interest (ROIs), using 6-connectivity (voxels considered to be connected if  
198 sharing faces rather than only edges or vertices) for both purposes. Individual fossils  
199 and other objects were identified among these ROIs by visual inspection in 3-D  
200 representations of voxels and against individual 2-D projections of  $\mu$ CT imagery. In  
201 cases where a fragmentary or incompletely resolved fossil fell into multiple non-  
202 connected ROIs, those ROIs were merged into one; in cases where a single ROI  
203 contained multiple contiguous objects, it was manually partitioned.

204         After each core had been imaged with  $\mu$ CT, they were macerated in 10% acetic  
205 acid to dissolve calcareous material until fully disaggregated and no longer visibly  
206 evolving gas bubbles. Insoluble residues were washed with water and sonicated in 30-  
207 second intervals, iterating until the supernatant ran clear. This washing process was  
208 then repeated using Calgon solution (0.052 M  $\text{Na}_6(\text{PO}_3)_6$ , 0.286 M  $\text{NaHCO}_3$ ) as a  
209 deflocculant, sonicating as before, until all loose clay was removed. Cleaned residues  
210 were then washed over a 250- $\mu\text{m}$  sieve, oven dried, and picked for identifiable fossils.

## 211 **Statistical Analyses**

212 Following scanning and segmentation, each object resolved in the  $\mu$ CT data was  
213 measured for minimum, maximum, and mean three-dimensional Feret diameter. Fossils  
214 that could be confidently identified to at least the phylum level were scored for additional  
215 taxon-specific features and anatomical measurements (Table 2).

216 Fossils recovered from residues were photographed using a GIGAmacro  
217 Magnify2 Robotic Imaging System with Canon EOS Rebel T8i DSLR and Nikon T1 1×  
218 and 3× objectives. For large specimens, additional photographs were taken under a  
219 reflected light microscope (Nikon SMZ1500 tethered to a Nikon D600 DSLR) to record  
220 features not visible in top-down view. Photographs were processed and analyzed with  
221 FIJI/ImageJ software, using the Trainable Weka Segmentation plugin to isolate fossils  
222 against the image background (Andreola et al., 2004a,b; Schindelin et al., 2012). Each  
223 fossil was measured for minimum and maximum two-dimensional Feret diameter as  
224 well as taxon-specific measurements corresponding to those taken from  $\mu$ CT data  
225 (Table 2).

226 For each core, both the segmented  $\mu$ CT render and acid residues were counted  
227 for the total abundance of fossil taxa. Trilobites were identified to the family level based  
228 on their general geometry, furrow pattern, and, in residues, their prosopon. Trilobite  
229 material not reliably assignable to a single family was treated as a separate category.  
230 Prosopon and other textural features were not easily resolvable in the  $\mu$ CT dataset, and  
231 so were not generally considered for taxonomic assignment in counts based on  $\mu$ CT  
232 results. Bryozoans were classified by growth form either as thin-branching, thick-  
233 branching, or fenestrate. Ostracods, gastropods, and bivalves were not further classified  
234 due their generally small sample size and coarse silicification, precluding the reliable  
235 identification of taxonomically relevant features.

236 Using R statistical software, datasets were subjected to NMDS in three  
237 dimensions based on the Bray-Curtis dissimilarity index, using the *metaMDS* function  
238 provided in the *vegan* package (Dixon, 2003; R Core Team, 2021). Sites with no

239 counted fossils were excluded from the analysis. Further paleoecological analysis of  
240 taxon abundances within and between samples was performed using PAST statistical  
241 software, calculating ecological dominance within samples (Simpson's *D*) and  
242 assessing compositional similarity between samples using the Bray-Curtis index (Hammer  
243 & Harper, 2001, 2022).

244

## 245 **Results**

246 A total of 582 objects were resolved via  $\mu$ CT, 460 as a low-opacity siliceous phase and  
247 122 as a high-opacity ferrous phase; of these, 241 siliceous objects and 14 ferrous  
248 objects were identifiable to at least a coarse taxonomic level, with the remaining 219  
249 and 108 respectively left unidentified (Fig. 2). Siliceous fossils were dominated by  
250 trilobites ( $n = 225$ ), of which slightly less than half ( $n = 99$ ) could be confidently assigned  
251 to a family classification; including Asaphidae (c.f. *Isotelus*), Cheiruridae (*Ceraurus*),  
252 Metagnostidae (*Trinodus*), Pterygometopidae (*Calyptaulax*), Raphiophoridae (*Ampyx*,  
253 *Lonchodomas*), and Remopleurididae (*Remopleurides*), all of which are previously  
254 known from this site (Whittington & Evitt, 1953; Evitt, 1961). Non-trilobite siliceous material  
255 consisted of thin-branching and fenestrate bryozoans and a single valve from an  
256 ostracod. Ferrous fossils were mostly gastropods, with two bivalves, one possible thin-  
257 branching bryozoan, and one infilling of a raphiophorid cranidium (Table 3). This  
258 cranidium was counted towards siliceous and ferrous  $\mu$ CT object totals due to being  
259 preserved as silicified cuticle filled in by moldic pyrite but was counted as one individual  
260 in taxon totals (Fig. 2D).

261 Residues yielded 1,349 recognizable objects, 1,222 preserved as silica, 125 as  
262 iron minerals (pyrite, hematite, and/or limonite), and the remaining two as other

263 materials (one of them appears to be a crystal of muscovite; neither was identifiable as  
264 a fossil). Nearly all siliceous fossils were identifiable to at least a coarse taxonomic  
265 level, with only 48 left unidentified. Conversely, 113 of the iron-preserved objects could  
266 not be identified. Siliceous objects in residue were made up largely of trilobites,  
267 fragments of thin-branching bryozoans, and ostracods. All trilobite genera represented  
268 in  $\mu$ CT were found also in residues, along with trilobites from the family Odontopleuridae  
269 (c.f. *Ceratocephala*). The remainder consisted of thick-branching and fenestrate  
270 bryozoans along with occasional brachiopod fragments. Ferrous fossils were exclusively  
271 represented by gastropods, with the exception of a single bivalve (Table 3).

## 272 **Tests of Preparation Bias**

273 A contingency table was constructed containing  $\mu$ CT identified specimens  
274 counts, broken down by taxonomic grouping and by presence or absence in residues.  
275 The relationship between taxon and recovery was investigated using Pearson's  $\chi^2$  test  
276 (Table 4). The only taxonomic group which deviated significantly ( $p < 0.05$ ) from overall  
277 likelihood of recovery was Unidentified Trilobita, material clearly from trilobites but  
278 lacking anatomical features in  $\mu$ CT sufficient to assign it to a family; such fossils  
279 identified in  $\mu$ CT are disproportionately unlikely to be recovered in residues.

280 Application of Pearson's  $\chi^2$  test to total counts of taxa identified in  $\mu$ CT and in  
281 residue found strong support ( $p < 0.000001$ ) for different proportional abundances in the  
282 two preparation types (Table 5). Raphiophorids, remopleuridids, unidentified trilobites,  
283 bivalves, and gastropods are significantly ( $p < 0.05$ ) more abundant in  $\mu$ CT, while  
284 branching bryozoans (thin and thick) and ostracods are significantly more abundant in  
285 residues.

286 The presence of size bias in the likelihood of recovery was evaluated using the  
287 Mann-Whitney U-test with a null hypothesis of no difference in medians and a standard  
288 threshold of significance of  $\alpha = 0.05$  (Mann & Whitney, 1947). In comparisons between  
289 fossils identified in both  $\mu$ CT and residue and those identified in  $\mu$ CT but *not* in residue,  
290 there was no statistically significant difference of maximum ( $p = 0.544$ ), minimum ( $p =$   
291  $0.088$ ), or mean ( $p = 0.244$ ) Feret diameters, nor of elongation factor ( $p = 0.159$ ) defined  
292 as the ratio of maximum to minimum Feret diameters. However, when testing overall  
293 distributions of size between all fossils identified in  $\mu$ CT and all those found in residue,  
294 the U-test found a dramatic difference in maximum Feret diameter ( $p = 1.75 \times 10^{-79}$ )  
295 between the medians of the two groups.

#### 296 **Exploratory Paleoecological Ordinations**

297 NMDS performed separately on the  $\mu$ CT (Table 6) and residue (Table 7)  
298 datasets resulted in ordinations with stress scores of 0.06 and 0.08, respectively (Fig.  
299 4). Taxon loadings in both ordinations consistently grouped bivalves and gastropods  
300 nearby each other, which may reflect either a shared infaunal environment or instead  
301 early burial conditions conducive to pyritization in those deposits. However, the low  
302 abundance of bivalves (2 in  $\mu$ CT, 1 recovered from residues, always alongside multiple  
303 gastropods) makes this apparent association tenuous. It is noteworthy that  
304 disseminated pyritization in the form of non-fossil granules was present in several  
305 samples (e.g., 05-C4) that contained no mollusks, suggesting that the presence or  
306 absence of mollusks in a sample may reflect their abundance in the paleoenvironment  
307 and is not solely controlled by whether early burial chemistry allowed pyritization to  
308 proceed.

309           Ordinations of both  $\mu$ CT and residues tended to group medium-sized benthic  
310 trilobites (raphiophorids, pterygometopids, and cheirurids) together along with  
311 unidentified trilobite material (most of which likely derived from one of those families),  
312 with pelagic trilobites (remopleuridids) plotting closer to gastropods and bivalves.  
313 Asaphids, likely represented here by the extremely large benthic trilobite *Isotelus*,  
314 consistently fall near the latter cluster but are represented only by fragments of cuticle.  
315 The ordination position of agnostids, whose life habit remains controversial (Fortey &  
316 Owens, 1999), is inconsistent between  $\mu$ CT and residues, falling near the benthic and  
317 pelagic trilobite clusters in those analyses, respectively. In both cases, they are closely  
318 accompanied by ostracods (which can occupy pelagic or benthic niches and whose life  
319 habit was not interpreted in this study). Due to the extreme disparity between  
320 identification in  $\mu$ CT and recovery in residues of ostracods, their location within the  
321 ordination for  $\mu$ CT data is likely not informative.

322           Thin-branching and fenestrate bryozoans plotted relatively close together,  
323 forming an isolated cluster. Thick-branching bryozoans were not closely associated with  
324 the other two morphotypes in the residue ordination and were not detected in  $\mu$ CT.

325           Samples from individual horizons tended to fall near each other in loose  
326 association. Two main clusters are apparent in both residue and  $\mu$ CT ordinations: one  
327 characterized by low NMDS1 scores and dominated by bryozoans, and one with higher  
328 NMDS1 scores and dominated by arthropods (trilobites and ostracods).  
329 For each sample, dominance was computed for both  $\mu$ CT (Table 6) and residue (Table  
330 7) taxon totals, with 95% confidence intervals based on 9999 bootstrap replicates (Fig.  
331 5). The results of this analysis differed notably between the two;  $\mu$ CT dominance values

332 fell in a tight band between 0.25 and 0.40 (excepting two outliers with  $n < 3$ ), while  
333 residue dominance values had a bimodal distribution with one group of values between  
334 0.40 and 0.65 and another falling between 0.15 and 0.30. (Fig. 5) Rarefaction curves  
335 were calculated for both  $\mu$ CT (Table 6) and residue (Table 7) counts of each sample  
336 using PAST (Hammer & Harper, 2022; Fig. 6). Curves were visually inspected for the  
337 presence of an inflection point as a rough qualitative assessment of sampling  
338 completeness; a sample's curve "leveling off" (sudden decrease in slope) is considered  
339 informal evidence that taxa in the true population are well-represented in the sample  
340 (Raup, 1975; Sanders, 1968). This inflection point was observed in residue counts from  
341 most samples (prominently in 03-C1, 03-C3, 04-C1, and 06-C3), but not in  $\mu$ CT counts.  
342 This suggests that the smaller total counts of individuals in  $\mu$ CT data are leading to the  
343 non-recovery of rarer taxa. Such under-sampling is a liability of ecological analyses  
344 based on those samples.

### 345 **Cluster Analysis**

346 Multivariate cluster analyses were performed in PAST to provide alternate  
347 metrics of similarity in the paleoecological composition of the samples, using the Bray-  
348 Curtis index as before (Bray & Curtis, 1957; Hammer & Harper, 2022). Dendrograms  
349 with branch length scaled to similarity were constructed for  $\mu$ CT counts, residue counts,  
350 and a combined dataset treating  $\mu$ CT and residue counts for each sample as two  
351 separate sites. (Figs. 7 and 8) Cluster analysis of  $\mu$ CT counts grouped samples from  
352 Horizon 04, and treated Horizon 05 similarly, but scattered samples from Horizons 03  
353 and 06 across the tree. Using residue data, Horizon 05 formed a cluster as before but  
354 Horizon 04 split, with cores 04-C2 and 04-C3 remaining close but 04-C1 further

355 removed; Horizons 03 and 06 were dispersed as in the  $\mu$ CT tree. In both trees, samples  
356 from Horizon 01 were far removed from other samples, tending to form a cluster basal  
357 to the rest of the samples.

358         The combined tree retains the general close groupings of the two individual trees  
359 due to sharing the same dissimilarity metric (Fig. 8). Notably, even for Horizons (01, 04,  
360 05) whose samples tended to cluster in both the  $\mu$ CT and residue trees, those clusters  
361 are not closely related in the overall analysis; there are few clear patterns in this tree,  
362 but  $\mu$ CT samples tend to cluster more closely with other  $\mu$ CT samples, and residue  
363 samples with residue samples, than do  $\mu$ CT and residue counts of the same sample or  
364 even of samples from the same horizon.

365

## 366 **Discussion**

### 367 **Size and Taxon Effects on Recovery**

368 We hypothesized at the outset of this study that body size would be negatively  
369 associated with likelihood of recovery in residues; larger fossils may be more  
370 susceptible to breakage, and larger organisms with thicker skeletal elements may not  
371 fully silicify, producing a brittle outer husk. However, the results of our analysis do not  
372 support this possible effect. While fossils recovered in residue have a smaller median  
373 body size than those resolved in  $\mu$ CT, this appears to be due not to preferential  
374 breakage of larger fossils during washing but rather disproportionate non-resolution of  
375 smaller fossils in  $\mu$ CT datasets. Rather than  $\mu$ CT acting as a baseline to test the biases  
376 of acid maceration based on body size, the results of this study suggest the opposite.

377         Revealing that the number of unidentified trilobite fossils recognized in  $\mu$ CT is  
378 significantly less likely to be recovered in residue is to be expected. Fragmentation and

379 incomplete preservation are common causes for trilobite material to lack identifying  
380 characteristics, and both are likely to promote further breakage and degradation during  
381 washing by compromising the structural integrity of the sclerite. This may also influence  
382 likelihood of recovery by impeding identification of the fossil in residue; a fragment of  
383 cuticle lacking identifying characteristics is less likely to be recognized as corresponding  
384 to an object observed in  $\mu$ CT, and minor breakage is likely to disrupt recognizable  
385 aspects of its outline and other key features.

386         Overrepresentation with respect to the abundance of pyritized taxa (gastropods  
387 and bivalves) in  $\mu$ CT may be the result of the high contrast between iron-bearing  
388 phases and the matrix, making such fossils easier to resolve in  $\mu$ CT than otherwise-  
389 comparable siliceous fossils. Alternatively, this may represent preferential breakage of  
390 pyritized material during washing due to differences in replacement texture. Ostracods,  
391 meanwhile, are disproportionately overrepresented in residue due to their small size  
392 making them difficult to resolve in  $\mu$ CT at the resolution used in this study. Branching  
393 bryozoans similarly also suffer from this, but their numbers in residue are likely inflated  
394 by fragmentation of large individuals into many smaller ones.

395

### 396 **Nonrecovery Bias**

397 Diagenetic bias arises from a variety of sources and is heavily dependent on shell  
398 microstructure, organic matter, and availability of reactive and replacive ions to the  
399 shell. It is well-documented that certain taxa are more susceptible to silicification than  
400 others, and even within taxa textural differences can make certain skeletal elements  
401 more likely to be preserved than others (Roberta L. Daley (\*), 1996; Cherns et al., 2011;

402 Butts, 2014). Organic matter content within a shell also influences the likelihood of  
403 silicification by providing nucleation sites for the deposition of silica, though the  
404 chemistry of this is complex (Wallace, DeYoreo & Dove, 2009; Butts, 2014). Similarly, replacement  
405 via pyrite tends to initiate at sites of organic matrix in the shell but is also dependent  
406 upon reducing microenvironments and/or microbial zonation within the sediment (FISHER,  
407 1986; Canfield & Raiswell, 1991; Schiffbauer et al., 2014).

408 Further biases can arise from the sample extraction and preparation process  
409 itself. During washing, larger fossils may be tumbled against other grains or container  
410 walls, leading to abrasion or breakage, while smaller fossils may be crushed beneath  
411 larger objects as grains settle. Elongated grains may be more prone to breakage than  
412 spheroidal grains of the same volume due to a lower minimum cross-sectional area,  
413 which can lead to taxon or sclerite bias against fossils with rod-like geometry, such as  
414 branching bryozoans and the spines of some trilobites. Breakage during preparation  
415 can cause taxa to be undercounted due to indiscriminate destruction of individuals or  
416 damage sufficient to remove recognizable features. Paradoxically, it can also inflate  
417 taxon counts by turning one fossil into many still-recognizable fragments. These  
418 overlapping effects can produce a wide variety of biases in size and shape, and the  
419 cumulative effect on measurements of abundance cannot reasonably be predicted *a*  
420 *priori*.

421 Composition of the host rock can further impose preparation-related biases. For  
422 instance, well-cemented rocks frequently require longer maceration periods than poorly  
423 consolidated ones, while argillaceous rocks may need more thorough washing to drive  
424 off insoluble clays, often including sonication. The longer sediment is washed and

425 manipulated, the more breakage tends to occur, making matrix texture an important  
426 factor of preparation bias. While the acid maceration processes can be accelerated by  
427 using alternative acids, such as hydrochloric, the violent effervescence caused by the  
428 intensity of the reaction can also have a deleterious effect on the extraction of delicate  
429 forms.

### 430 **Artifacts and Limitations of $\mu$ CT**

431 While a powerful tool for visualization,  $\mu$ CT does present some unavoidable  
432 sources of potential error: artifacts arising from specimen capture and processing,  
433 particularly the “hardening” of the X-ray beam through preferential absorption of lower-  
434 energy photons by the sample. Beam hardening is a ubiquitous issue in  $\mu$ CT, but not all  
435 hardening artifacts are of equal impact. Since the beam is slightly hardened by passing  
436 through the outer surface of the sample, the interior is always slightly darkened relative  
437 to the outermost layer; these “cupping artifacts” are produced predictably based on  
438 sample geometry and can be corrected relatively easily (Schladitz, 2011; Jung et al., 2011; Abel,  
439 Laurini & Richter, 2012).

440 Heterogeneous samples suffer from more vexing artifacts. Regions of higher-  
441 density material (e.g., iron minerals in sediments, bones in biological samples) can  
442 drastically harden beams passing through them, causing dark blotches to appear  
443 around bright features within a sample and especially in the spaces between multiple  
444 bright features. These artifacts are far more difficult to correct due to their irregular  
445 shape and remain the subject of ongoing research in economic geology and other  
446 materials science fields (Remeysen & Swennen, 2006; Park, Chung & Seo, 2015; Bam et al., 2019). Given  
447 the abundant pyrite in the material examined in this study, this stands as a caveat to the

448 results reported here; similar methods applied to rocks with lower density variation may  
449 yield clearer  $\mu$ CT data and therefore more complete identification of fossils within the  
450 samples. Filtering during pre-processing, or correction applications in post-processing  
451 software packages, can reliably improve these artifacts, but often at the cost of  
452 introducing noise along the axis of sample rotation (Kyriakou, Prell & Kalender, 2009; Yousuf &  
453 Asaduzzaman, 2009).

454         The 3-D nature of  $\mu$ CT data can lead to more conceptual issues when drawing  
455 comparisons to features in corresponding 2-D imagery. One such caveat relates to  
456 resolution, the smallest possible distance between features that can be resolved by an  
457 imaging method. Not only is this distance always larger (coarser) than the side length of  
458 pixels or voxels making up the image produced, but it depends on the nature of the  
459 sample, with more internally complex and heterogeneous materials having larger  
460 effective resolution than homogeneous ones (Elkhoury, Shankar & Ramakrishnan, 2019).  
461 Resolution can also be larger and more challenging to measure in 3-D imaging than in  
462 2-D. Determining by observation whether adjacent features are connected or distinct in  
463 a 2-D plane is relatively easy, but due to limitations of how displays and human vision  
464 project three dimensions down to two, the equivalent task in 3-D volume is more  
465 challenging. In practice, this means that features near the theoretical resolution limit of  
466 3-D imaging methods are likely to be lost, especially when interpreted by human  
467 observers. When assessing body size effects on recovery in 3-D  $\mu$ CT imagery versus 2-  
468 D residue photographs, this disparity in effective resolution should be taken into  
469 account.

470 **Relevance for Paleoecological Interpretation**

471 Broad categories of biofacies are generally consistent between residue- and  
472  $\mu$ CT-derived taxon abundances as interpreted using NMDS. While total counts for  $\mu$ CT  
473 are much lower than those of residues, relative abundances in the former largely  
474 recapitulate the latter with some notable exceptions; ostracods are almost absent from  
475  $\mu$ CT data, due to resolution constraints imposed from the diameter of the cores, even  
476 when highly abundant in residues, and bryozoan counts tend to be much higher in  
477 residues than in  $\mu$ CT due to breakage transforming one large fragment into many small  
478 fragments. For this reason, the application of  $\mu$ CT data alone to describe  
479 paleoecological structure should be undertaken with caution. Ordinations used in  
480 paleoecology are sufficiently complex that such biases in recovery can, by introducing  
481 error and uncertainty upstream, irreparably taint conclusions drawn downstream.

482 Overall, the use of  $\mu$ CT counts on their own to characterize silicified remains in  
483 these limestone biofacies is of uncertain value. Sample volume constraints make bulk  
484 samples unfeasible, and the cost of instrument time can be a roadblock to replication  
485 and large-scale sampling. However,  $\mu$ CT may be best suited when chemical preparation  
486 is dangerous or impractical due to matrix or fossil mineralogy, or when studying  
487 sponges, bryozoans, corals, or other modular organisms for which fragmentation during  
488 washing can obscure or inflate the number of individuals present. Samples best suited  
489 to the approach outlined in this study will be densely fossiliferous (mitigating the  
490 limitation of sample volume), bear taxa with morphological features conducive to  
491 identification in  $\mu$ CT (related to shape rather than texture) and contain grains more  
492 susceptible to dissolution than the matrix or cement joining them together. Such rocks  
493 should also have a strong contrast in density (and therefore X-ray attenuation) between

494 grains and matrix, but with relatively little influence of beam hardening artifacts.  
495 Conodonts and other phosphatic fossils, for example, are occasionally known from  
496 silicified sediments, and characterizing taxonomic assemblages using  $\mu$ CT techniques  
497 may constitute a welcome alternative to digestion with the notoriously toxic and  
498 hazardous hydrofluoric acid normally used to liberate fossils from siliceous cements  
499 (Green, 2001).

500

## 501 **Conclusions**

502 The main goal of this study was to quantify the effects of taxonomic affinity and body  
503 size on the likelihood of fossil recovery, but it is the qualitative results that may have the  
504 broadest relevance to further work. Breakage remains an unavoidable concern with acid  
505 maceration, but the degree to which breakage rates depend on body size and  
506 taxonomic factors likely varies substantially between deposits. This variation can stem  
507 not only from easily observable sedimentological features such as grain size and  
508 composition but also from redox chemistry and solute profiles of pore fluids during burial  
509 and early diagenesis, which can be challenging to infer from samples without more in-  
510 depth geochemical analyses. Original shell composition, both in terms of organic  
511 content and aragonitic versus calcitic (high- or low-magnesium) mineralogy is likely  
512 relevant due to its influence on the spontaneity and kinetics of silicification chemistry.  
513 Information on the breadth of variation is limited but suggests a wide range of possible  
514 taxonomic outcomes based on sample lithology.  
515 While  $\mu$ CT imaging of silicified fossils in limestone matrix can resolve morphological  
516 features of interest, the issues faced in this specific study may limit this technique's  
517 ability to answer broader paleoecological questions. Analyzing original rock contents as

518 a means of detecting breakage during preparation may be valuable as a control on  
519 fragmentary abundance counts, but limitations of sample size and the potential for  
520 taxonomic bias to affect ordinations present serious pitfalls to analysis or  
521 characterization of biofacies. It can, however, be useful in establishing broad categories  
522 if not finer gradations between related assemblages. Since the drawbacks resulting  
523 from beam hardening and other  $\mu$ CT artifacts are heavily dependent on the properties of  
524 the sample, these methods may prove more effective when applied to rocks containing  
525 different fossil taxa or with different lithologic compositions from the materials studied  
526 here.

527

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531

532

533

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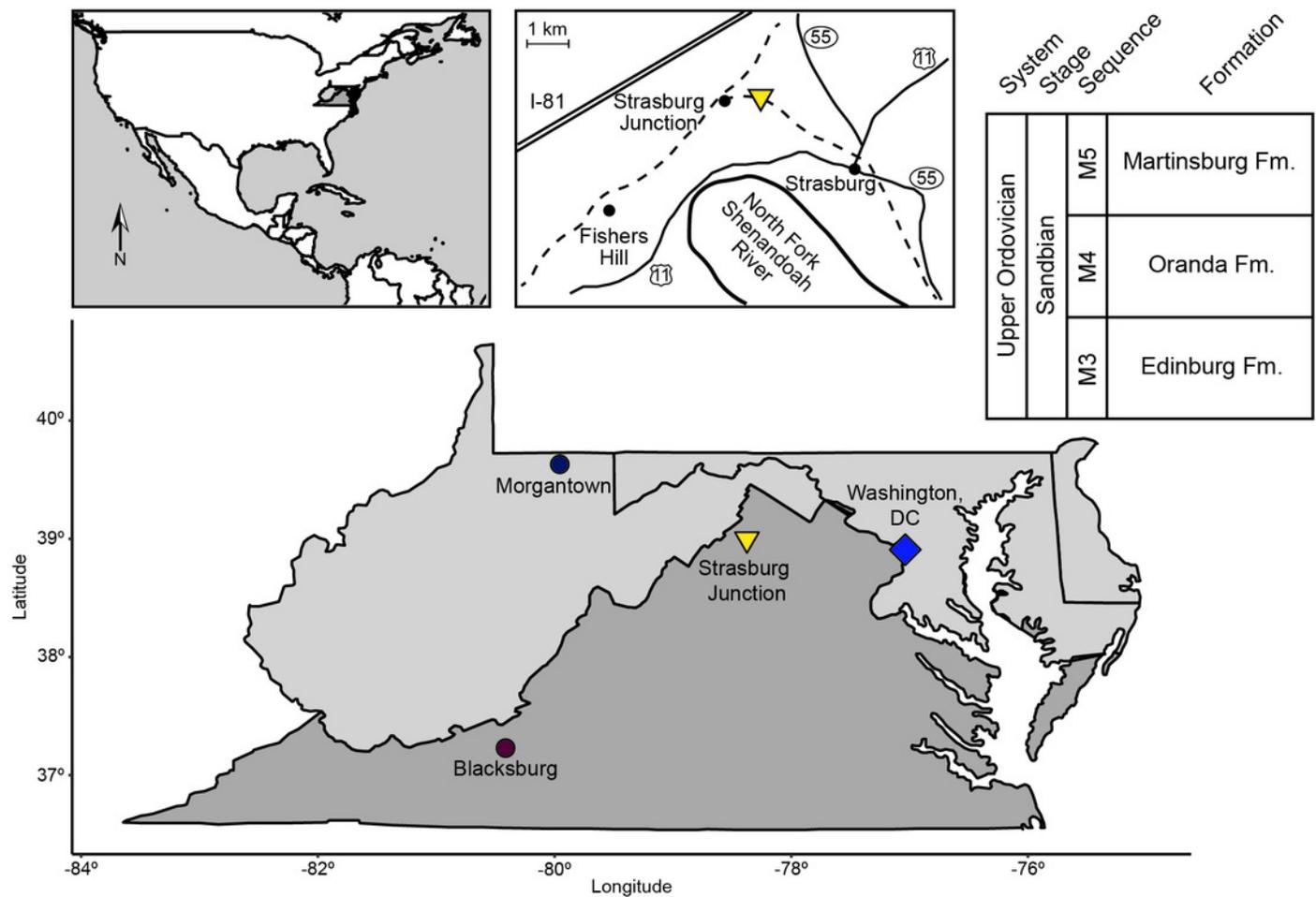
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# Figure 1

Figure 1

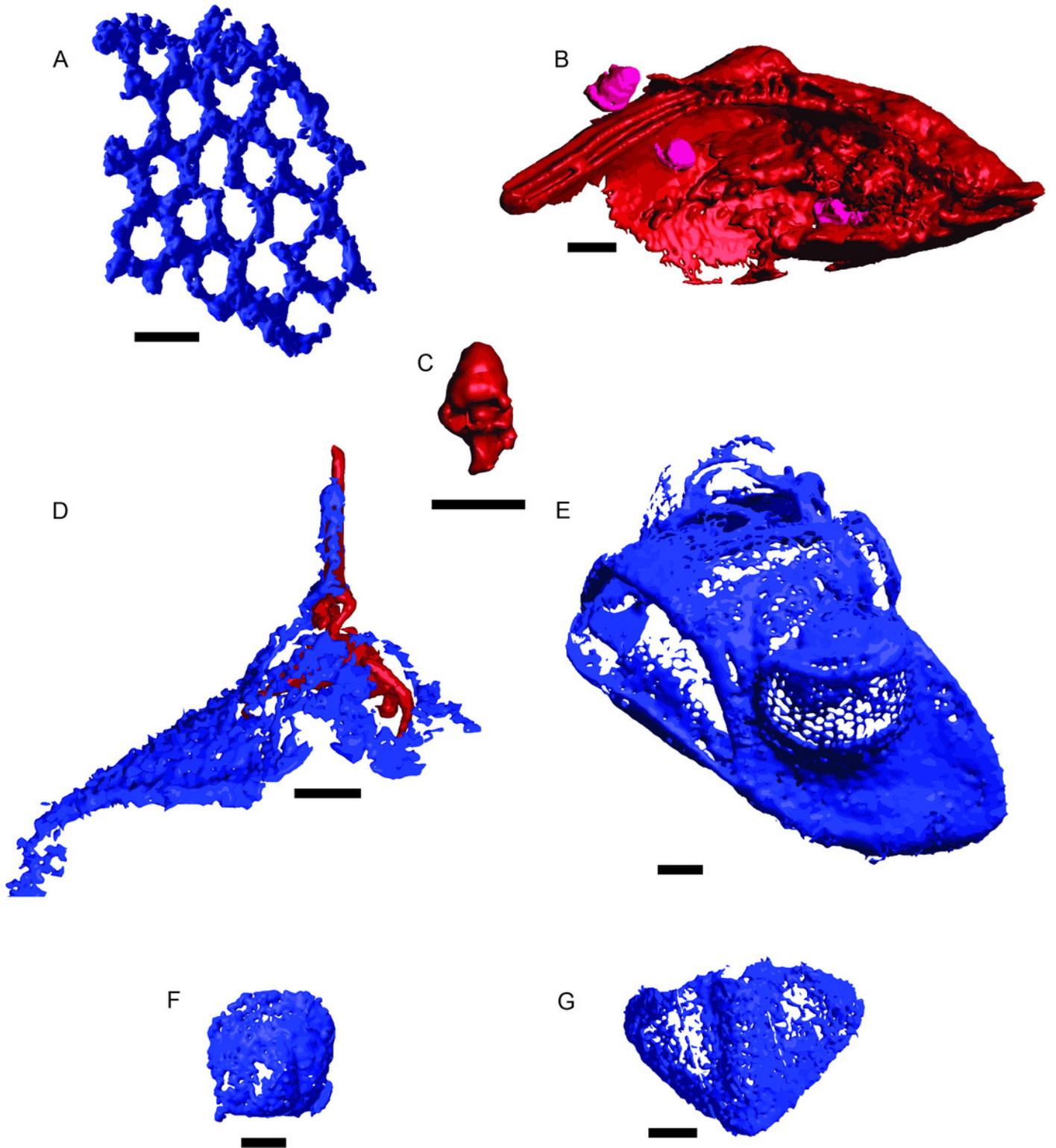
Figure 1. Location map and generalized stratigraphy of samples analyzed in this study. Adapted from Jacobs and Carlucci (2019).



## Figure 2

### Figure 2

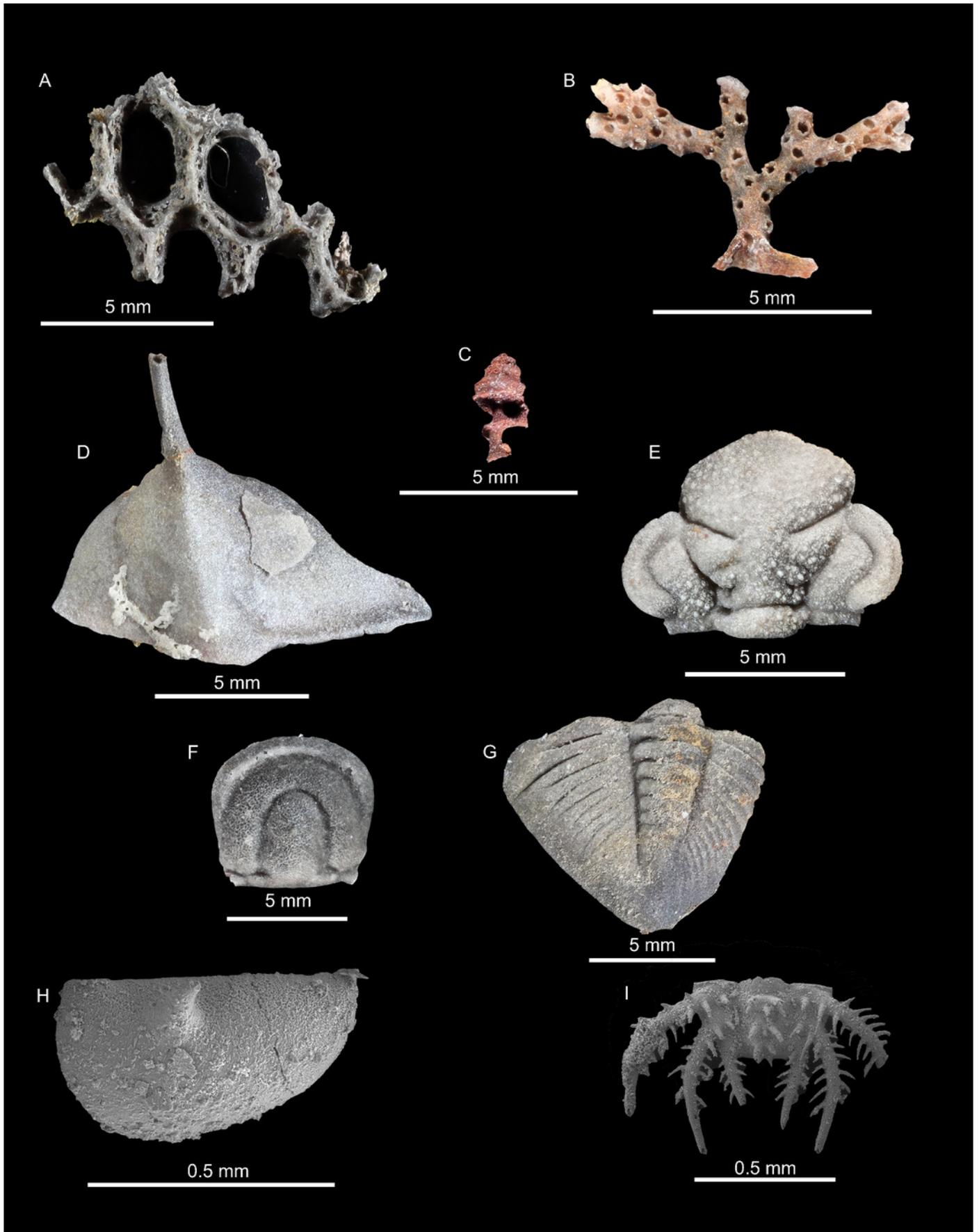
Figure 2. CT renderings of fossils identified in cores of the Edinburg Formation. Blue indicates silica and red indicates Fe-rich mineralization. A) Fenestrate bryozoan (01-C4-001). B) Nuculanid bivalve (03-C3-064) with three small gastropods (03-C3-060, 03-C3-062, 03-C3-065) in pink. C) Closer view of gastropod (03-C3-060). D) Raphiophorid cranidium preserved in silica (05-C4-005) with pyrite infilling (05-C4-047). E) Pterygometopid cephalon (04-C2-028). Note the compound eyes. F) Agnostid fragment (06-C4-027). G) Pterygometopid pygidium (05-C1-019). All scale bars approximate 1 mm.



## Figure 3

### Figure 3

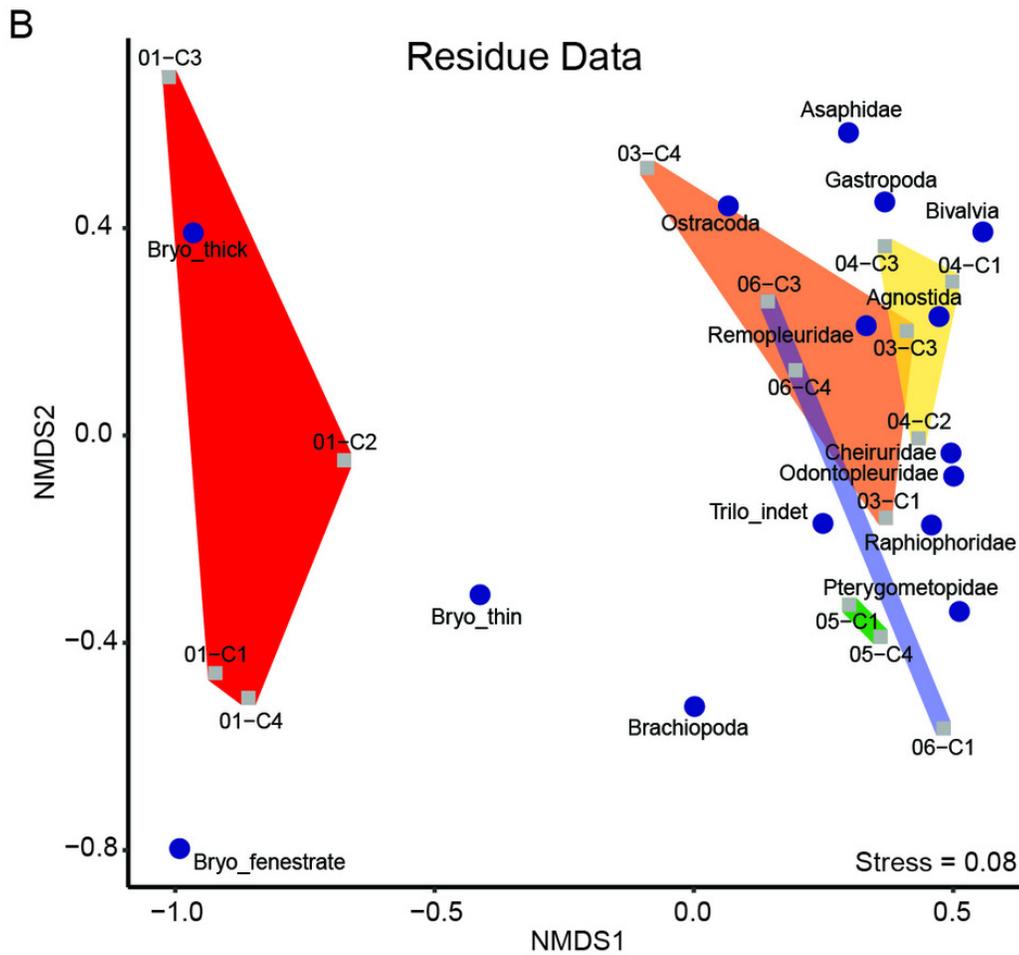
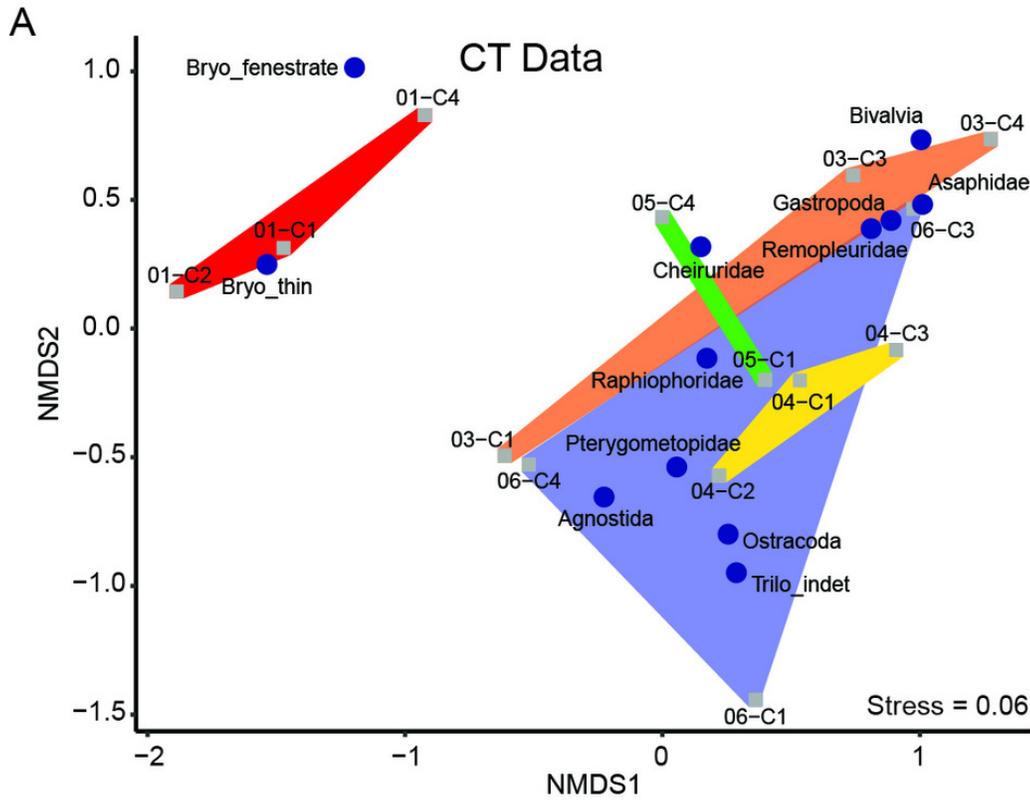
Figure 3. Photomicrographs and SEM images of fossils recovered in macerate residues. A) Silicified fenestrate bryozoan (01-C1). B) Silicified erect branching bryozoan (03-C3). C) Pyritized gastropod (03-C3). D) Silicified raphiophorid cranidium (03-C1). E) Silicified *Calyptaulax* cranidium (06-C4). F) Silicified agnostid fragment (06-C4). G) Silicified *Calyptaulax* pygidium (03-C1). H) SEM-image of silicified ostracod valve (04-C1). I) SEM-image of silicified trilobite pygidium (04-C1).



## Figure 4

Fig 4

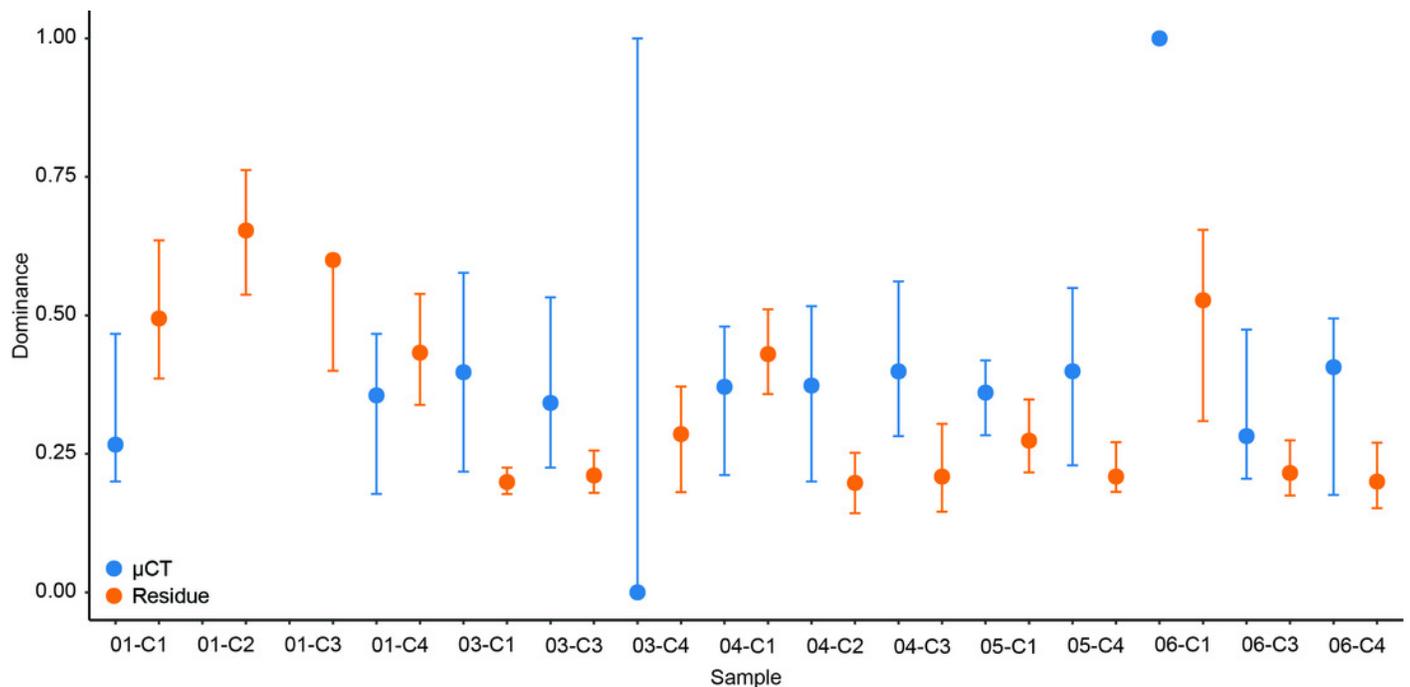
Figure 4. Results of NMDS of taxon abundance data for A)  $\mu$ CT data (stress = 0.06) and B) acid maceration residue data (stress = 0.08). Samples are plotted as grey squares. Taxon scores are indicated by blue circles. Convex hulls contain samples from each sampling horizon (Table 1). Red: EB18-01. Orange: EB18-03. Yellow: EB18-04. Green: EB18-05. Blue: EB18-06.



# Figure 5

Fig 5

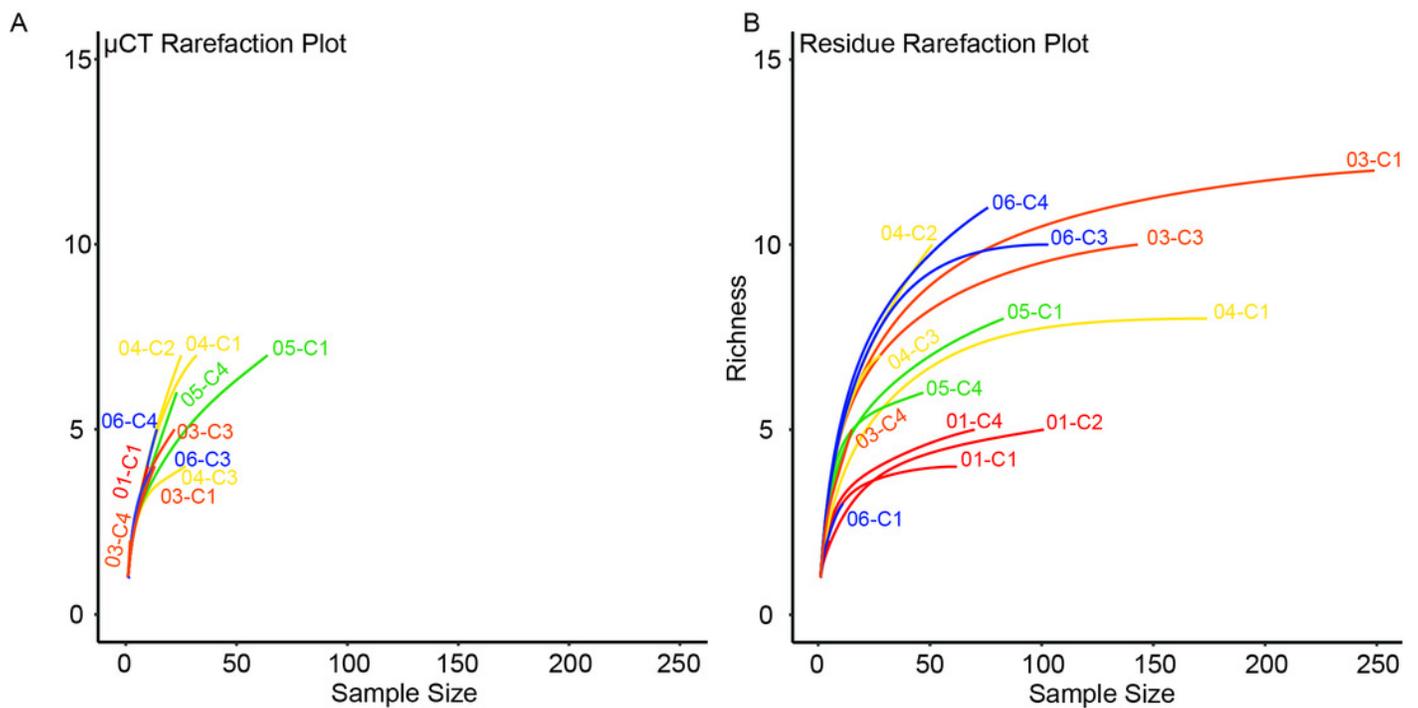
Figure 5. Simpson's Dominance (D) and bootstrapped 95% confidence intervals calculated for each sample by data type. data are in blue and acid maceration residue data are in orange.



# Figure 6

Fig 6

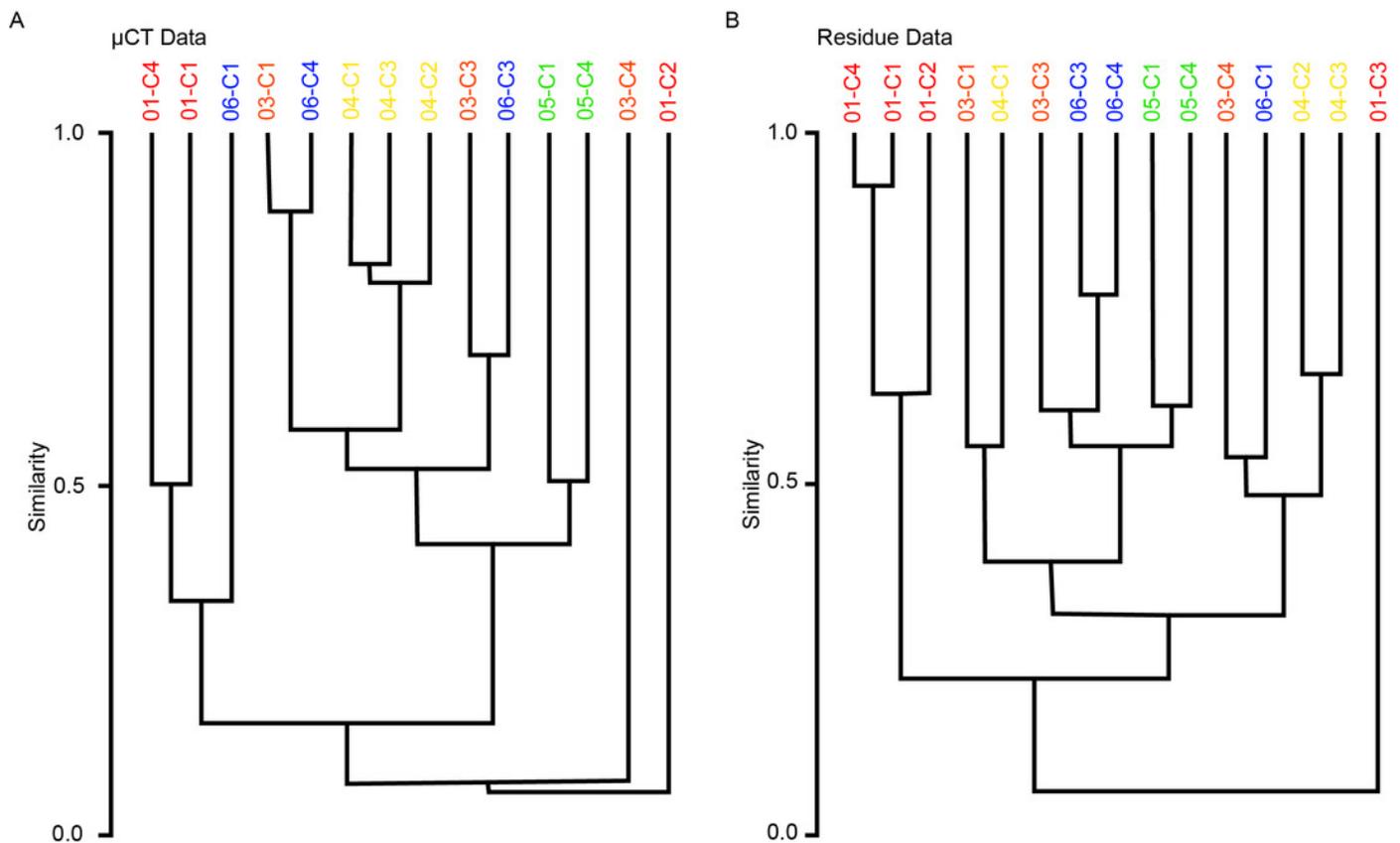
Figure 6. Rarefaction analysis of taxonomic richness for A)  $\mu$ CT samples (Table 6) and B) acid maceration residue samples (Table 7).



# Figure 7

Fig 7

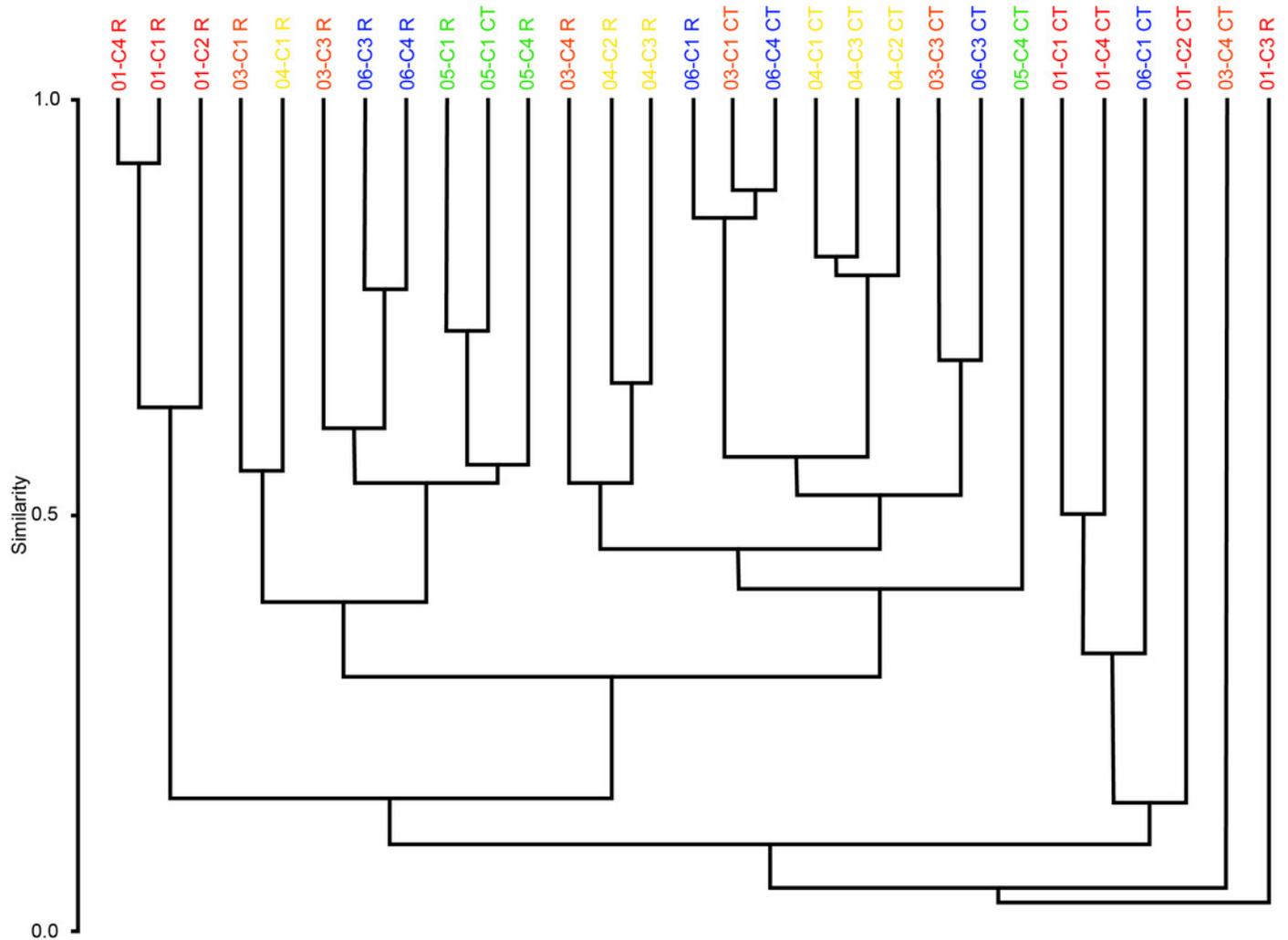
Figure 7. Bray-Curtis similarity dendrograms of taxon abundance values in A)  $\mu$ CT samples (Table 6) and B) acid maceration residue samples (Table 7) calculated separately.



# Figure 8

Fig 8

Figure 8. Bray-Curtis similarity dendrograms of pooled taxon abundance values in  $\mu$ CT samples (Table 6) and acid maceration residue samples (Table 7).



**Table 1** (on next page)

Tables 1-8

This file contains all eight tables each with the appropriate legend

1  
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## Tables

Horizon	Latitude	Longitude	Notes
EB18-01	38.9969°N	78.3748°W	~0.5 m above contact between Lincolnshire LS and Edinburg Fm
EB18-03	38.9968°N	78.3744°W	~20-30 m up section from EB18-01
EB18-04	—	—	0.5 m up section from EB18-03
EB18-05	—	—	1.0 m up section from EB18-04
EB18-06	—	—	~5 m up section from EB18-05

4 Table 1. Summary of horizons and samples used in this study.

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Taxon		Measurements and Classifications
Trilobita	Cranidium	Maximum width (tr.) of occipital lobe*, maximum width (tr.) between eyes**, total length (sag.)
	Thoracic segment	Maximum width (tr.) of axial ring, width (tr.) between fulcra of left and right pleurae, length (sag.) of axial ring
	Pygidium	Maximum width (tr.) of first axial ring, total length (sag.)
Bryozoa		Growth form: dendroid or fenestrate
Gastropoda		Presence/absence of outer walls of whorls, total height of shell, height of last complete whorl, width of last complete whorl, half-angle of teleoconch***
Ostracoda		Body length (a.-p.), body height (d.-v.), articulation/disarticulation, presence/absence of lateral spine

11 Table 2. Anatomical measurements and classifications by taxon and sclerite. Measurements  
12 taken from CT and residue imagery for each of several common taxa. The main sclerites of  
13 trilobites (excluding hypostomes and librigenae) were assigned different sets of measurements  
14 due to fundamental differences in structure. Lengths, widths, and heights are measured in mm;  
15 angles are measured in degrees; other category variables were recorded as Booleans.

16 \* Raphiophorids have an effaced prosopon with indistinct furrows, making identification of the  
17 occipital lobe difficult in CT especially. For this family of trilobites, occipital width was instead  
18 measured as the full width (tr.) of the sclerite from gena to gena at the occiput.

19 \*\* For eyeless trilobites of the families Metagnostidae and Raphiophoridae, interocular distance  
20 was excluded.

21 \*\*\* This is the angle formed at the apex between the axis of coiling and a line tangent to the  
22 outer walls of the body whorls.

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Taxon		$\mu$ CT	Residues
Total Trilobita		225	594
Trilobita	Asaphidae	6	20
	Cheiruridae	2	3
	Metagnostidae	6	14
	Odontopleuridae	0	18
	Pterygometopidae	6	41
	Raphiophoridae	57	149
	Remopleurididae	22	59
	Unknown Trilobita	126	290
Bivalvia		2	1
Brachiopoda		0	5
Bryozoa	Thin branching	9	304
	Thick branching	0	49
	Fenestrate	7	22
Gastropoda		10	11
Ostracoda		1	232
Unknown		327	161
<b>Total</b>		<b>581</b>	<b>1379</b>

32 Table 3. Fossils Recovered by Preparation Method and Taxon. Total counts of all objects  
 33 identified in CT and residue imagery, grouped by taxon. Trilobites are further broken down to  
 34 the family level. Trilobite material lacking sufficient anatomical features to be confidently  
 35 assigned to a family was given its own category.

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Taxon	Not Recovered	Recovered	Recovered residual	p-value
Asaphidae	3	2	0.213	0.84
Cheiruridae	1	1	0.285	0.67
Metagnostidae	2	4	1.855	0.11
Pterygometopidae	4	1	-0.787	0.46
Raphiophoridae	37	20	-0.376	0.90
Remopleurididae	13	8	0.493	0.81
Unknown Trilobita	71	28	-7.389	0.04
Bivalvia	1	1	0.285	0.67
Bryozoa (thin)	4	4	1.140	0.39
Bryozoa (fenestrate)	2	4	1.855	0.11
Gastropoda	4	6	2.425	0.10

39 Table 4. Taxonomic Effects on Recovery. Contingency table containing counts of fossils with  
 40 distinctive geometry identified in CT imagery and either recovered or not recovered in residue.  
 41 Residuals of recovered counts are standardized to the expected number of recovered specimens  
 42 for a taxon; positive residuals indicate disproportionately high likelihood of recovery in residues,  
 43 while negative values indicate lower-than-average likelihood of recovery.  
 44 Statistically significant results ( $p < 0.05$ ) are shaded.

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Taxon	CT	Residue	CT residual	Residue residual	p-value
Asaphidae	6	20	0.715	-0.326	0.43
Cheiruridae	2	3	1.224	-0.559	0.18
Metagnostidae	6	14	1.372	-0.627	0.13
Odontopleuridae	0	18	-1.762	0.805	0.05
Pterygometopidae	6	41	-0.741	0.338	0.41
Raphiophoridae	57	149	3.598	-1.643	2.0 E-05
Remopleurididae	22	59	2.146	-0.980	0.02
Unknown Trilobita	126	290	6.399	-2.922	9.9 E-17
Bivalvia	2	1	2.060	-0.941	0.02
Brachiopoda	0	5	-0.929	0.424	0.31
Bryozoa (thin)	9	304	-6.125	2.797	3.3 E-14
Bryozoa (thick)	0	49	-2.908	1.328	1.2 E-3
Bryozoa (fenestrate)	7	22	0.892	-0.407	0.32
Gastropoda	10	11	3.350	-1.530	2.1 E-4
Ostracoda	1	232	-6.183	2.824	1.3 E-13

50 Table 5. Comparison of Taxonomic Abundances. Contingency table containing total counts of  
 51 taxa identified in CT and in residue imagery. Residuals of recovered counts are standardized to  
 52 the expected number of recovered specimens for the combination of taxon and preparation;  
 53 positive CT residuals indicate overrepresentation in CT for that taxon, while positive residue  
 54 residuals indicate overrepresentation in residues (relative to average proportional abundance for  
 55 all fossils). Statistically significant results ( $p < 0.05$ ) are shaded.

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	01-C1	01-C2	01-C4	03-C1	03-C3	03-C4	04-C1	04-C2	04-C3	05-C1	05-C4	06-C1	06-C3	06-C4	Sum
Agnostida	0	0	0	1	0	0	1	2	0	1	0	0	0	1	6
Asaphidae	0	0	0	0	0	1	2	0	1	1	1	0	0	0	6
Cheiruridae	0	0	0	0	0	0	0	0	0	1	1	0	0	0	2
Odontopleuridae	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
Pterygometopidae	0	0	0	0	0	0	1	1	0	3	0	0	0	1	6
Raphiophoridae	0	0	0	3	1	0	5	4	5	23	14	0	0	2	57
Remopleuridae	0	0	1	0	5	1	2	1	5	4	1	0	2	0	22
Trilo indet	2	0	1	8	12	0	19	15	16	31	5	2	6	9	126
Bivalvia	0	0	0	0	1	0	0	0	0	0	0	0	1	0	2
Brachiopoda	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
Bryo thin	3	1	2	1	0	0	0	0	0	0	1	0	0	1	9
Bryo thick	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
Bryo fenestrate	1	0	6	0	0	0	0	0	0	0	0	0	0	0	7
Gastropoda	0	0	0	0	3	0	2	1	0	0	0	0	4	0	10
Ostracoda	0	0	0	0	0	0	0	1	0	0	0	0	0	0	1
<b>SUM</b>	<b>6</b>	<b>1</b>	<b>10</b>	<b>13</b>	<b>22</b>	<b>2</b>	<b>32</b>	<b>25</b>	<b>27</b>	<b>64</b>	<b>23</b>	<b>2</b>	<b>13</b>	<b>14</b>	

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Table 6. Abundance of taxonomic groups identified in CT data used in NMDS.

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	01-C1	01-C2	01-C3	01-C4	03-C1	03-C3	03-C4	04-C1	04-C2	04-C3	05-C1	05-C4	06-C1	06-C3	06-C4	Sum
Agnostida	0	0	0	0	6	2	0	2	1	0	0	0	0	2	1	14
Asaphidae	0	0	0	0	0	2	1	4	1	2	0	0	0	4	6	20
Cheiruridae	0	0	0	0	1	0	0	0	0	1	0	1	0	0	0	3
Odontopleuridae	0	0	0	0	9	0	0	4	1	0	3	0	0	0	1	18
Pterygometopidae	0	0	0	0	12	4	0	5	2	0	1	5	1	4	7	41
Raphiophoridae	0	0	0	0	59	8	1	21	7	2	19	12	2	10	8	149
Remopleuridae	0	1	0	0	8	30	0	6	2	2	2	0	0	4	4	59
Trilo indet	2	8	0	2	72	51	6	22	17	8	37	12	8	29	16	290
Bivalvia	0	0	0	0	0	1	0	0	0	0	0	0	0	0	0	1
Brachiopoda	0	0	0	1	2	0	0	0	1	0	1	0	0	0	0	5
Bryo thin	42	81	0	43	17	24	0	0	6	3	11	13	0	36	28	304
Bryo thick	8	8	4	15	4	0	1	0	0	0	0	0	0	7	2	49
Bryo fenestrate	10	0	0	9	2	0	0	0	0	0	0	0	0	0	1	22
Gastropoda	0	0	0	0	0	6	0	0	0	0	0	0	0	5	0	11
Ostracoda	0	3	1	0	57	15	6	110	13	10	9	4	0	2	2	232
<b>SUM</b>	<b>62</b>	<b>101</b>	<b>5</b>	<b>70</b>	<b>249</b>	<b>143</b>	<b>15</b>	<b>174</b>	<b>51</b>	<b>28</b>	<b>83</b>	<b>47</b>	<b>11</b>	<b>103</b>	<b>76</b>	

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Table 7. Abundance of taxonomic groups recovered in acid maceration residues used in NMDS.