

A new Lower Triassic ichthyopterygian assemblage from Fossil Hill, Nevada

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We report a new ichthyopterygian assemblage from Lower Triassic horizons of the Prida Formation at Fossil Hill in central Nevada. Although fragmentary, the specimens collected so far document a diverse fauna. One partial jaw exhibits isodont dentition with blunt tipped, mesiodistally compressed crowns and striated enamel. These features are shared with the Early Triassic genus *Utatusaurus* known from coeval deposits in Japan and British Columbia. An additional specimen exhibits a different dentition characterized by relatively small, rounded posterior teeth resembling other Early Triassic ichthyopterygians particularly *Grippia*. This Nevada assemblage marks a southward latitudinal extension for Early Triassic ichthyopterygians along the eastern margin of Panthalassa and indicates repeated circum-oceanic dispersal events in Early Triassic ichthyopterygians.

1 A new Lower Triassic Ichthyopterygian assemblage from Fossil Hill, Nevada

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24 Abstract

25 We report a new ichthyopterygian assemblage from Lower Triassic horizons of the Prida
26 Formation at Fossil Hill in central Nevada. Although fragmentary, the specimens collected so far
27 document a diverse fauna. One partial jaw exhibits isodont dentition with blunt tipped,
28 mesiodistally compressed crowns and striated enamel. These features are shared with the Early
29 Triassic genus *Utatusaurus* known from coeval deposits in Japan and British Columbia. An
30 additional specimen exhibits a different dentition characterized by relatively small, rounded
31 posterior teeth resembling other Early Triassic ichthyopterygians, particularly *Grippia*. This
32 Nevada assemblage marks a southward latitudinal extension for Early Triassic ichthyopterygians
33 along the eastern margin of Panthalassa and indicates repeated trans-hemispheric dispersal
34 events in Early Triassic ichthyopterygians.

35

36 Main article text

37 Introduction

38 Ichthyosaurs were among the most enduring and successful Mesozoic marine reptile groups,
39 appearing in the Early Triassic and persisting some 150 million years until their extinction in the
40 Late Cretaceous (McGowan and Motani, 2003). The fossil record of early ichthyopterygians (the
41 clade comprising ichthyosaurs and close relatives) includes a variety of morphologically
42 disparate taxa from widespread localities in Asia, North America and the Arctic. Most of these
43 assemblages are broadly contemporaneous, all being late Spathian (late Early Triassic) in age.
44 Recent discoveries in China (Ji et al., 2014, Motani et al., 2015) have extended this record
45 earlier into the Spathian and have shed new light on the phylogenetic and biogeographic origins

46 of the clade. However, the rapid early diversification and trans-hemispheric dispersal history of
47 ichthyopterygians during the Early Triassic remains poorly understood.

48

49 Nevada has been an important source of Triassic marine reptile fossils since the 19th Century
50 producing abundant and well-preserved Middle Triassic (Leidy, 1868; Merriam, 1905, 1908,
51 1910; Sander et al., 1994, 1997; Fröbisch et al., 2006, 2013) and Late Triassic (Camp 1976,
52 1980) ichthyopterygian and sauropterygian fossils. Notably, Early 20th Century field work led by
53 John Merriam and Annie Alexander at the Fossil Hill locality in the Humboldt Range produced
54 several specimens of the ichthyosaur *Cymbospondylus* (Merriam 1908)—previously described by
55 Leidy (1868) on the basis of fragmentary remains—as well as the type specimens of
56 *Omphalosaurus nevadanus* (Merriam, 1906) and *Phalarodon fraasi* (Merriam 1910). Later work
57 by Camp (1976, 1980), Sander et al. (1994, 1997) and others illuminated rich Middle and Late
58 Triassic marine reptile assemblages preserved in Nevada.

59

60 In contrast, knowledge of Early Triassic marine reptile fossils in this region is scant. The only
61 published Early Triassic marine reptile occurrence from Nevada is based on a partial jaw referred
62 to the enigmatic genus *Omphalosaurus* and described as a second species, *O. nettarhynchus*
63 Mazin and Bucher, 1987. This specimen was collected from the Spathian-aged informally
64 designated “lower member” of the Prida Formation in the Humboldt Range, which sits
65 immediately below the well-known Fossil Hill Member of the Prida Formation, famous for its
66 rich marine reptile assemblage including the aforementioned *Cymbospondylus*, *Phalarodon*, and
67 *Omphalosaurus nevadanus*. Fragmentary, float-derived remains of Early Triassic
68 ichthyopterygians have been reported from Spathian horizons in southeastern Idaho (Massare

69 and Callaway, 1994; Scheyer et al., 2014) roughly 500 km to the northeast of the Fossil Hill
70 locality. Even further to the east, the early sauropterygian *Corosaurus alcovensis* Case, 1936 is
71 known from the Alcova limestone in Wyoming whose Early Triassic age was recently confirmed
72 (Lovelace and Doebbert, 2015).

73

74 Here, we report a new Early Triassic ichthyopterygian assemblage from the lower member of the
75 Prida Formation at the Fossil Hill Locality. These fossils are Spathian (Lower Triassic) in age
76 based on co-occurring conodont and ammonoid faunas and sit stratigraphically below the diverse
77 Middle Triassic marine reptile assemblage from the Fossil Hill Member of the Prida Formation.
78 These occurrences extend the southward latitudinal range of early ichthyopterygians in North
79 America and demonstrate that early in their evolutionary history, multiple ichthyopterygian taxa
80 quickly dispersed around or across wide expanses of ocean and ranged from sub-tropical to high
81 temperate waters on the eastern margin of northern Panthalassa.

82 *Institutional Abbreviations.* USNM, National Museum of Natural History, Smithsonian
83 Institution, Washington, D.C., U.S.A.

84

85 Materials and Methods

86 **Geological and Stratigraphic Setting**

87 The new fossils reported here were collected from multiple horizons within the unnamed lower
88 member of the Prida Formation of the Star Peak Group at Fossil Hill, on the eastern flank of the
89 Humboldt Range in Pershing County, Nevada (Fig. 1). The Star Peak Group consists of a
90 sequence of syndepositionally deformed carbonate-dominated units deposited on what was then
91 the western shelf of North America (Nicholls & Silberling, 1977; Wyld, 2000). In the study area,

92 the lower member of the Prida Formation forms the base of the Star Peak Group and sits
93 unconformably atop the Permian/Lower Triassic aged Koipato Group volcanics (Wyld, 2000).
94
95 The lower member of the Prida Formation transitions from siliciclastic sand and conglomerate
96 layers near the contact with the underlying Koipato Group to dark-grey limestone above with
97 intermittent microbialite, conglomerate and chert-dominated beds. The presence of
98 conglomerates and microbialites indicate relatively shallow conditions with a general trend
99 towards deeper water facies characteristic of the overlying Fossil Hill Member (Wyld, 2000).
100 Gastropods and bivalves are abundant in lower layers whereas conodonts and ammonoids are
101 found locally within middle and upper layers of the lower member. Scattered vertebrate fossils
102 occur in multiple horizons within the lower member (Fig. 1), but are most abundant in the middle
103 carbonate layers where they are associated with the conodont *Triassospathodus symmetricus*
104 (Orchard, 1995) and the ammonoid *Prohungarites gutstadt* (Guex et al., 2010) indicating a late
105 Spathian age (Subcolumbites ammonoid biozone). These fossils were collected by the landowner
106 and co-author Patrick Embree, who donated the material to the Smithsonian National Museum of
107 Natural History.

108

109 Results

110 **Systematic Paleontology**

111 Ichthyopterygia Owen, 1840

112 cf. *Utatusaurus* Shikama, Kamei and Murata, 1978

113 **Diagnosis.** Teeth smaller than marginal dentition present on pterygoid; squamosal not entirely
114 eliminated from supratemporal fenestra by supratemporal; interclavicle cruciform; dorsal margin

115 of external naris formed entirely by nasal; prefrontal shelf prominent; transverse flange of the
116 pterygoid well defined and anterolaterally projecting; supratemporal terrace present; tooth
117 implantation subthecodont, with both dental groove and shallow socket; tooth crowns of middle
118 to posteriorly placed teeth distomesially compressed; humerus as wide proximally as distally;
119 humerus longer than wide; ulnar facet of humerus as wide as radial facet; no more than five
120 phalanges in any digit; posterior dorsal vertebrae cylindrical in outline (from Cuthbertson et al.,
121 2013a; after Motani, 1999; McGowan and Motani, 2003).

122

123 **Referred specimen.** USNM 559349 Partial mandible including teeth. (Fig. 2)

124 **Locality.** Fossil Hill, Humboldt Range, Pershing County, Nevada.

125 **Horizon and age.** Found as surface float within an outcrop of Lower Triassic (upper Spathian)
126 lower member of Prida Formation, Star Peak Group. Based on location and matrix lithology this
127 jaw is inferred to derive from horizon FH1-7 (Fig. 1), which is Spathian based on the occurrence
128 of the ammonoid *Prohungarites gutstadi* (i.e. Subcolumbites Zone of Guex et al., 2010) and
129 conodonts *Triassospathodus symmetricus* (Orchard, 1995) and *Neostrachanognathus* sp.
130 extracted from the matrix.

131

132 **Description.** USNM 559349 is a partial mandible measuring 82 mm long. The jaw fragment
133 preserves portions of the dentary, surangular and splenial. The surfaces of the dentary and
134 surangular are heavily striated and the orientation of these striations differs between the bones.
135 The suture between the dentary and surangular is long and straight, extending across the entire
136 preserved portion of the jaw. In places this suture is indistinct but can be traced by the
137 contrasting surface striation patterns of the dentary and surangular. The splenial can be observed

138 at the broken anterior edge of the fossil where it comprises the medial and ventral portion of the
139 jaw where a thin projection wraps underneath the surangular. A row of irregular weathered
140 depressions follows the approximate course of the suture between the surangular and dentary but
141 it is not clear if these represent natural foramina or are simply artifacts of weathering. Judging
142 from the arrangement of the bones, the fragment likely represents a central-posterior portion of
143 the left mandibular ramus anterior to the coronoid process.

144

145 Thirteen lower teeth are present, along with an additional poorly preserved isolated tooth
146 between the tenth and eleventh in-place teeth, which may be either a disarticulated upper or
147 lower tooth. The teeth are set within alveoli along a continuous groove. No distinct bony septa
148 between alveoli are visible but may be present at the bottom of the dental groove, being
149 concealed in matrix that is very difficult to remove through mechanical preparation. The roots of
150 some teeth are clearly expanded at the base and exhibit plications that are coarser than crown
151 striation. The most anterior tooth is completely exposed anteriorly, revealing its root structure
152 inside the dental groove. It is seen there that the root ceases its expansion once inside the
153 groove, and teeth are embedded to both the labial wall and the base of the groove. A narrow gap
154 emerges between the lingual wall and the root toward the dentigerous margin. Tooth
155 implantation is likely subthecodont (*sensu* Motani, 1997a), although histological study is
156 necessary to firmly establish this. The root cross-section is much wider than long, as reported for
157 *Utatsusaurus hataii* (Motani, 1996).

158

159 Teeth are isodont and conical with striated crowns. Tooth roots are extensively exposed above
160 the alveolar margin such that they account for half or more of the exposed height of each tooth.

161 Tooth crowns are distinguished from these exposed roots by a distinct margin, with most crowns
162 slightly constricted at their base. Some crowns exhibit slightly higher convexity of their anterior
163 surface relative to the posterior surface given them a slightly recurved appearance. The teeth are
164 also recurved lingually toward the tip, as clearly seen in the most anterior tooth (Fig 2B). This
165 curvature closely resembles what was described for *Utatusaurus hataii* (Motani, 1996:Fig. 3;
166 Cuthbertson et al., 2013a:Fig 7C, D). The tips of the teeth are relatively blunt. Tooth crowns are
167 approximately 3.6 mm tall and 2 mm in mesiodistal diameter. Spacing between teeth ranges from
168 3 mm to 6 mm, more widely spaced teeth may have replacement teeth between them. One 10
169 mm gap along the tooth row likely represents at least one missing tooth. Several teeth are broken,
170 either at the root or the crown, revealing a pulp cavity without evidence of infolding of the
171 dentine.

172

173 **Remarks.** The tooth morphology observed in this specimen closely resembles that described for
174 *Utatusaurus hataii* from the Lower Triassic Osawa Formation of Kitakami, Japan (Shikama et
175 al., 1978). Most notably, the teeth curve lingually and slightly posteriorly toward the tip, which is
176 a feature that is uniquely known in *Utatusaurus* among basal ichthyopterygians. Other shared
177 features include: isodont dentition, tall exposed roots, blunt conical striated crowns, a slightly
178 constricted base of some crowns, and an absence of infolding in the pulp cavity. Each of these
179 features is observed in the holotype of *Utatusaurus hataii* (IGPS 95941) (Motani, 1996) and
180 most are reported in a referred specimen (UHR 30691) (Cuthbertson et al., 2013a). No other
181 Early Triassic or later ichthyopterygian exhibits this suite of dental characters.

182

183 The teeth of this specimen do diverge from *Utatsusaurus hataii* in their much larger size. The
184 maximum tooth exposed height in this specimen is 11 mm, whereas the maximum crown height
185 and width are 4.5 mm and 2.3 mm respectively (Table 1), compared with 3.3 mm, 1.7 mm and
186 0.9 mm for the same measurements of teeth in the holotype of *Utatsusaurus hataii* (Motani,
187 1996). However, the holotype represents a juvenile (Motani, 1997c), so the size difference may
188 partly be explained as ontogenetic variation. A referred specimen (UHR 30691) is somewhat
189 larger than the holotype, however, the teeth of this specimen are still considerably smaller than
190 those of USNM 559349. Despite this difference in size, the overall shape of the teeth in each of
191 these specimens is very similar. Motani (1996) reported “crown shape index” values—calculated
192 as crown height divided by average basal diameter of the tooth crown (Massare, 1987)—ranging
193 between 0.9 and 3.1 with an average value of 1.9 in the type specimen of *Utatsusaurus hataii*,
194 while these values range from 1.3 to 2.3 with an average of 1.8 in USNM 559349. Likewise, the
195 “crown ratio”—calculated as crown height divided by total exposed tooth height—averages 0.51 in
196 IGPS 95941 and 0.43 in USNM 559349.

197

198 The overall arrangement of bones in the fragmentary mandible is consistent with that observed in
199 the *Utatsusaurus* holotype (IGPS 94941) and the referred specimen (UHR 30691), notably the
200 long, straight contact between the dentary and surangular. The suture between the surangular and
201 dentary is well developed in IGPS 94941 but indistinct in the larger UHR 30691, possibly a
202 consequence of differing ontogenetic stages, or differential weathering, between these two
203 specimens. USNM 559349 approaches UHR 30691 in that the suture between the dentary and
204 surangular is somewhat obscured, but its approximate course can be traced by a contrast in bone
205 texture and an irregular row of weathered pits. This might suggest a relatively mature individual,

206 as has been proposed for IGPS 9494, but this is highly speculative given the incomplete nature of
207 the material.

208

209 A partial skull from the Lower Triassic Vega Phroso Member of the Sulphur Mountain
210 Formation from British Columbia, Canada was referred to *Utatusaurus* sp. by Nicholls and
211 Brinkman (1993), based largely on the presence of the same dental features detailed above. The
212 teeth of the British Columbia specimen are similar in size to those of the Nevada specimen
213 described here (Nicholls & Brinkman, 1993) and distinctly larger than those found in the
214 holotype of *Utatusaurus hataii* from Japan (Motani, 1996). It is therefore possible that these
215 larger-toothed specimens from the eastern margin of Panthalassa (Nevada, British Columbia)
216 represent a form allied with but distinct from *Utatusaurus hataii*; however more complete
217 material is needed before this can be confirmed. Recently, Cuthbertson et al. (2014) described
218 another partial skull from the Vega Phroso Member, which they also referred to *Utatusaurus*
219 sp., although they concluded that the material originally referred to this genus by Nicholls and
220 Brinkman (1993) was non-diagnostic at the genus level. Unfortunately this recently described
221 material lacks a lower jaw or teeth and cannot be compared to USNM 559349.

222

223 Ichthyopterygia Owen, 1840

224 cf. Grippiidae Wiman, 1929

225 **Definition.** The last common ancestor of *Grippia longirostris* and *Gulosaurus helmi*, and all its
226 descendants (Ji et al., 2015).

227 **Diagnosis.** Maxilla with multiple tooth rows; posterior tooth crown rounded; supratemporal-
228 postorbital contact present; proximal manual phalanges not closely packed proximo-distally
229 (from Ji et al., 2015).

230

231 **Referred specimen.** USNM 559350 Partial maxilla including teeth. (Fig. 3)

232 **Locality.** Fossil Hill, Humboldt Range, Pershing County, Nevada.

233 **Horizon and Age.** Collected from FH1-5 (Fig. 1) horizon which is Spathian in age based on the
234 occurrence of the ammonoid *Prohungarites gutstadi* (Guex et al., 2010) and conodonts including
235 *Triassospathodus* and *Neostrachanognathus*. This horizon is also characterized by distinctive
236 spherical structures originally interpreted as microbial (i.e. ‘oncoids’) but more recently
237 suggested to represent sponge “reefs” (Brayard et al., 2011; see further discussion below).

238 **Description.** USNM 559350 (Fig. 3) is a partial maxilla measuring 20 mm and bearing five teeth
239 exposed in medial view. The teeth are robust cones exhibiting a trend of posteriorly increasing
240 basal diameter, whereas the crown height remains constant giving the posteriormost preserved
241 tooth a distinctly rounded shape. There is a distinct constriction below the crown separating it
242 from the root below, however the constriction is very slight in the anteriormost tooth. The
243 anteriormost crown height and width are 1.7 mm and 1.2 mm respectively; in the second
244 posteriormost tooth, which is better preserved than the posteriormost tooth, crown height and
245 width are 1.3 mm and 2.1 mm. The tooth enamel appears smooth and polished with little
246 indication of striation; however, this could be attributed to tooth wear. Faint plication is visible
247 on some roots. Although some teeth are abraded, none expose the pulp cavity clearly enough to
248 determine presence or absence of infolded dentine.

249

250 In medial view the teeth are attached to the lingual wall of the maxilla, representing pleurodont
251 tooth attachment. An expanded bone of attachment conceals the bases of the two posteriormost
252 teeth, suggesting subleurodont attachment, a modified form of pleurodont attachment (Motani,
253 1997a), in at least the posterior region of the maxillary tooth row. While only a single row of
254 teeth is observed, a shallow depression on the lingual margin of the tooth row immediately
255 anterior to the second posteriormost tooth could represent the attachment facet of a missing
256 tooth. If this were the case it might possibly represent a second row immediately lingual to the
257 preserved teeth. Wide spacing between the four anteriormost teeth would easily accommodate an
258 additional offset tooth row as observed in the maxillary dentition of *Grippia* (Motani, 1997b) and
259 *Gulosaurus* (Cuthbertson et al., 2013b).

260

261 **Remarks.** Among Early Triassic ichthyopterygians, small, robust teeth, similar to those reported
262 here, are typical of the posterior dentition of *Grippia* (Motani, 1997b). Rounded teeth are also
263 observed in the Early Triassic genus *Chaohusaurus* (Motani & You, 1998) and, to a lesser
264 extent, *Gulosaurus* (Cuthbertson et al. 2013b). *Grippia* was previously reported from the Lower
265 Triassic Vega Phroso Member of the Sulphur Mountain Formation in British Columbia
266 (Brinkman et al., 1992). This specimen was later redescribed as a distinct taxon, *Gulosaurus*
267 *helmi* (Cuthbertson et al., 2013b) and found to be sister taxon to *Grippia longirostris*. Similarly,
268 recent work by Ji et al. (2015), established a new clade Grippioidea including *Grippia*,
269 *Gulosaurus*, *Utatusaurus*, and *Parvinator*, although the precise relationships among these taxa
270 varied somewhat depending on taxon and character inclusion. The enigmatic marine reptile
271 *Omphalosaurus nettarhynchus* (Mazin & Bucher, 1987), previously reported from Spathian
272 lower member of the Prida Formation, also possesses rounded dentition, but is distinct from this

273 specimen by its much larger size and in exhibiting a broad pavement of rounded teeth on the
274 mandible.

275

276 Alternatively, this specimen may have some affinity with *Chaohusaurus*, an Early Triassic
277 ichthyopterygian from China in which some specimens also show distinctly rounded posterior
278 dentition (Motani & You, 1998). However the posterior teeth of *Chaohusaurus* are generally
279 smaller and more tightly packed than in USNM 559350, averaging approximately ten teeth over
280 20 mm rather than the five teeth observed over the same distance in this specimen.

281 *Chaohusaurus* was previously regarded as a grippidian partly on the basis of possessing multiple
282 maxillary tooth rows and rounded posterior dentition (Motani, 1999). However, more recent
283 analyses (Cuthbertson et al. 2013b, Ji et al., 2015) do not support this placement.

284

285 The crown shape index in USNM 559350 ranges from 1.0-2.1 and averages 1.4 (Table 1),
286 similar to the average shape index reported for the maxillary dentition of *Grippia*, 1.4 (Motani,
287 1997b) and *Gulosaurus* 1.64 (Cuthbertson et al. 2013b). Thus, the rounded dentition observed in
288 USNM 559350 is strongly suggestive of an affinity with some other Early Triassic
289 ichthyopterygians but more precise placement will require more complete skeletal material.

290

291 Discussion

292 Despite the fragmentary nature of the remains described here, their resemblance with the
293 distinctive dentitions of other Early Triassic ichthyopterygians allows tentative interpretations to
294 be made. The presence of *Utatusaurus*-like and *Grippia* or *Chaohusaurus* -like forms suggests
295 similarity with the Lower Triassic Vega-Phroso assemblage from the Wapiti Lake region of

296 British Columbia, from which *Utatusaurus* (Nicholls & Brinkman, 1993; Cuthbertson et al.,
297 2014) and grippidians (Brinkman et al., 1992; Cuthbertson et al., 2013b) have also been reported.
298 However, the type locality of *Utatusaurus* is in the Osawa Formation of Japan (Shikama et al.,
299 1978), whereas the type localities of *Grippia* and *Chaohusaurus* are in the Vikinghøgda
300 Formation (= “Sticky Keep Formation” of older references) of Spitsbergen (Wiman, 1929, 1933;
301 Hounslow et al., 2008) and the Nanlinghu Formation of Anhui Province, China (Young & Dong,
302 1972), respectively. Thus, Early Triassic ichthyopterygian taxa were widely distributed around
303 the margins of northern Panthalassa (Cuthbertson et al., 2013b).

304

305 This broad distribution early in their evolutionary history, from numerous Late Spathian
306 (Subcolumbites Zone) localities of broadly coeval age (Scheyer et al., 2014), has made it
307 difficult to pinpoint the biogeographic origins of the group. However, recent work in China has
308 extended the biostratigraphic range of ichthyopterygians to the underlying Procolumnbites Zone
309 (Motani et al., 2014; Ji et al., 2014). Furthermore, the occurrence of diverse and endemic
310 hupehsuchians, widely regarded as the ichthyopterygian sister-group, and the plesiomorphic
311 ichthyosauromorph *Cartorhynchus* (Motani et al., 2015) are consistent with an origin of
312 ichthyopterygians near the south China block in equatorial western Panthalassa.

313

314 The inferred nearshore lifestyle of most Early Triassic ichthyopterygians has led others to
315 propose that these early marine reptiles dispersed along coastlines or across transient
316 epicontinental corridors (Cuthbertson et al., 2013b). However, there is little geological evidence
317 for such corridors in the Early Triassic, which was a time of relatively low global sea level
318 (Miller et al., 2005). Furthermore the absence of Early Triassic ichthyopterygian fossils in

319 Western Tethys is surprising under this scenario. Conversely, the biogeographic histories of
320 other aquatic—and even terrestrial—reptile groups are marked by occasional transoceanic dispersal
321 events (Rocha et al., 2006; Velez-Juarbe et al., 2007), and such events could explain the
322 distribution of Early Triassic ichthyopterygians on opposite shores of Panthalassa.

323 Brayard et al. (2009) identified trans-Panthalassan distribution patterns in Spathian
324 ammonoids, identifying similar ammonoid faunas in Nevada, Kitakami and British Columbia,
325 which they attributed to oceanographic currents. The occurrence of some Early Triassic marine
326 reptile taxa (e.g. *Utatusaurus*) on both the eastern and western margins of of Panthalassa might
327 reflect sporadic crossing of deep ocean basins by these lineages, potentially facilitated by the
328 same ocean currents that mediated transoceanic dispersals of contemporaneous marine
329 invertebrates (Fig. 4). The wide distribution of Early Triassic sauropterygians, including the
330 South China Block (Jiang et al., 2014) and western margin of North America (Storrs, 1991;
331 Lovelace & Doebbert, 2015) on opposite shores of Panthalassa indicates to a similar dispersal
332 history in the early members of that marine reptile clade. Isolated terranes such as South
333 Kitakami, South Primoyre and Chulitna could have served as stepping-stones for shallow marine
334 taxa. Dispersal along coastlines around the northern margins of Panthalassa remains an
335 alternative scenario that could explain the broad distribution of some Early Triassic
336 ichthyopterygians, with pronounced global warmth in the Early Triassic mediating limiting
337 climatic conditions at high latitudes (Sun et al., 2012). However, the apparent absence of
338 ichthyopterygian fossils from high latitudes on the western margin of northern Panthalassa
339 remains a puzzle under this scenario.

340

341 Intriguingly, the oldest marine reptile bearing horizons at Fossil Hill are associated with a
342 prominent limestone marker bed bearing distinctive spherical structures ~1–2 cm in diameter
343 (Fig. 5). We initially interpreted these structures as microbial ‘oncoids.’ Widespread
344 microbialite-dominated facies are characteristic of Lower Triassic strata globally, including in
345 the western United States (Pruss & Bottjer, 2004; Baud et al., 2007), and are interpreted as a
346 byproduct of the end-Permian mass extinction and subsequent delayed biotic recovery of
347 metazoan reefs (Pruss et al., 2004). A similar association between the basal sauropterygian
348 *Corosaurus* and stromatolites in the Lower Triassic Alcova limestone has been reported
349 previously (Storrs, 1991). More recently, similar spheroidal structures from the Humboldt Range
350 and other localities in western North America have been interpreted as ‘transient sponge reefs’
351 (Brayard et al., 2011). Thus, the diversification and dispersal of Early Triassic marine reptiles
352 was apparently well underway at the end of the Early Triassic (Scheyer et al., 2014) despite some
353 lingering signs of continued environmental stress preserved in the same strata. Future work at
354 this new locality, and elsewhere, may help to clarify the role that large-scale environmental
355 changes played in shaping the early evolutionary history of Mesozoic marine reptiles.

356

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362

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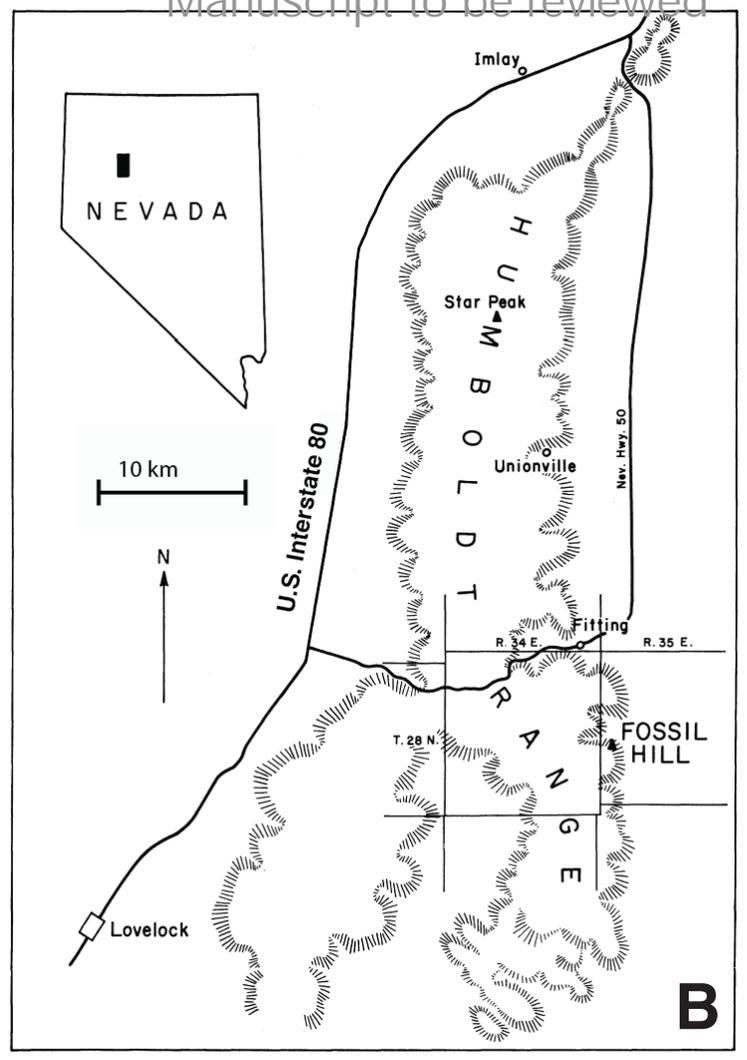
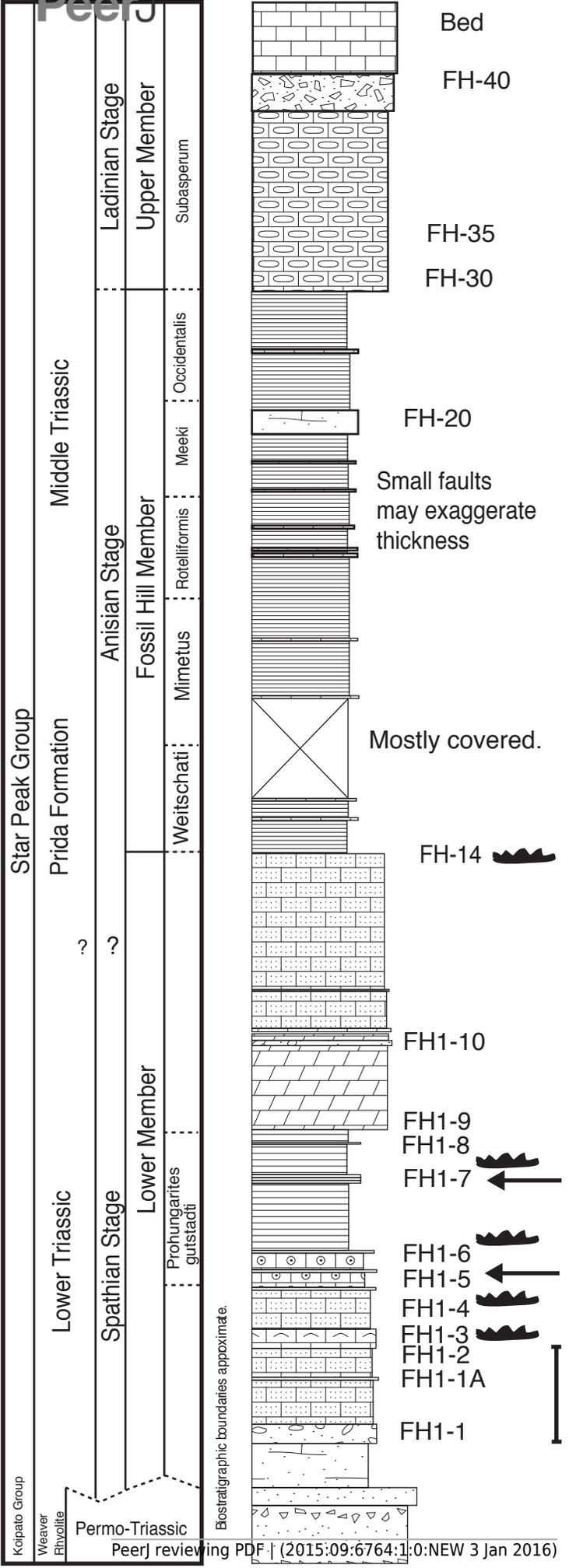
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Figure 1 (on next page)

Summarized stratigraphy and regional map

- (A) Stratigraphy of the Triassic Prida Formation near Fossil Hill in the Humboldt Range, Nevada indicating horizons of specimens USNM 559349 and 559350 and conodont samples.
- (B) Regional map, modified from Silberling (1962).



Location of conodont samples

A

2

Specimen USNM 559349, partial ichthyopterygian jaw cf. *Utatsusaurus*.

(A) Complete specimen, in labial view, anterior to the left. Squares on scale bar equal 5 mm.

(B) Magnified view of anterior dentition, squares on scale bar equal 1 mm.



3

USNM 559350 Partial ichthyopterygian maxilla cf. *Grippidia*.

A) Partial maxilla in lingual view, anterior to the left. Squares on scale bar equal 1 mm. White arrow indicates possible attachment facet for tooth in second lingual tooth row. B) Magnified view of dentition.

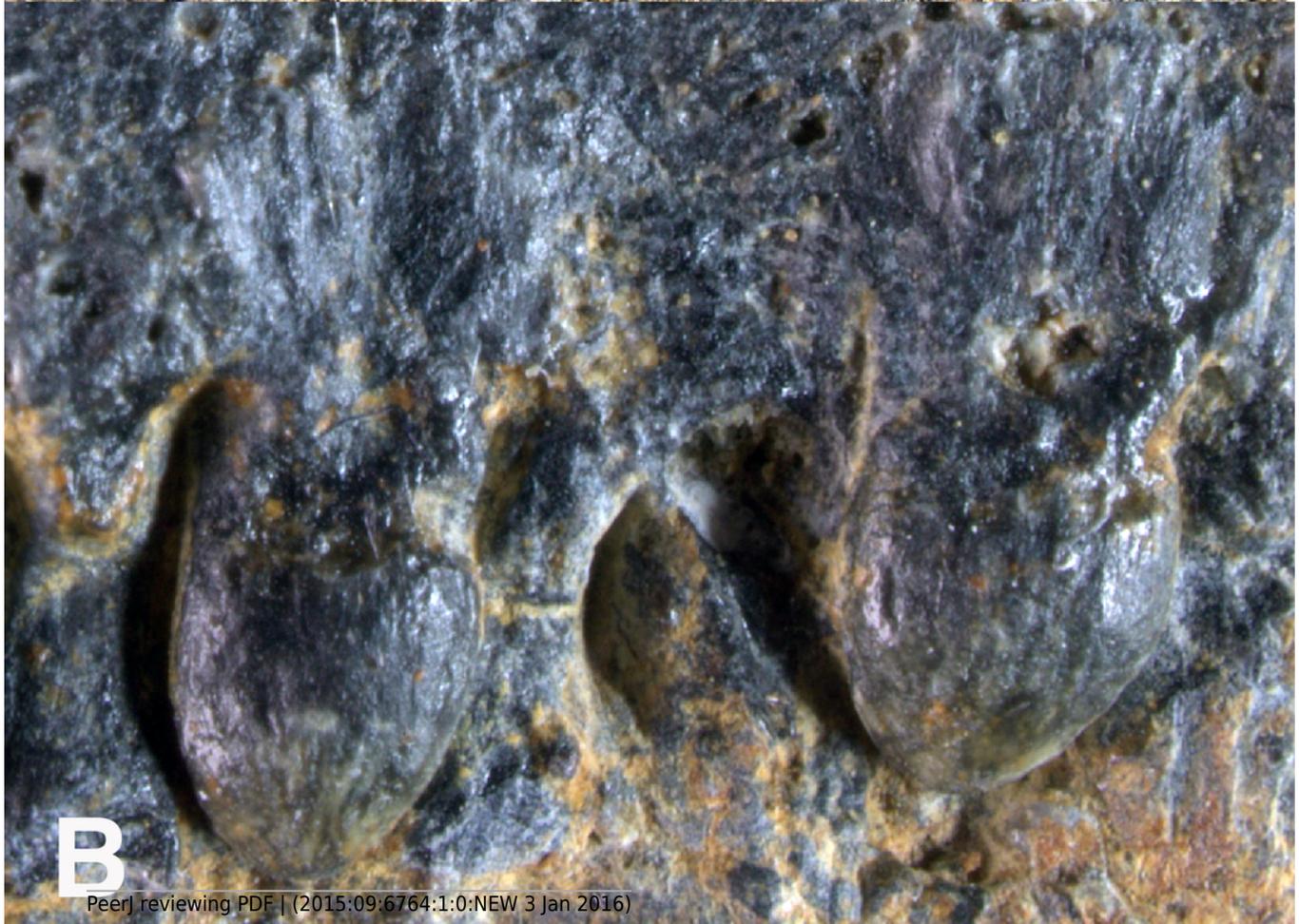
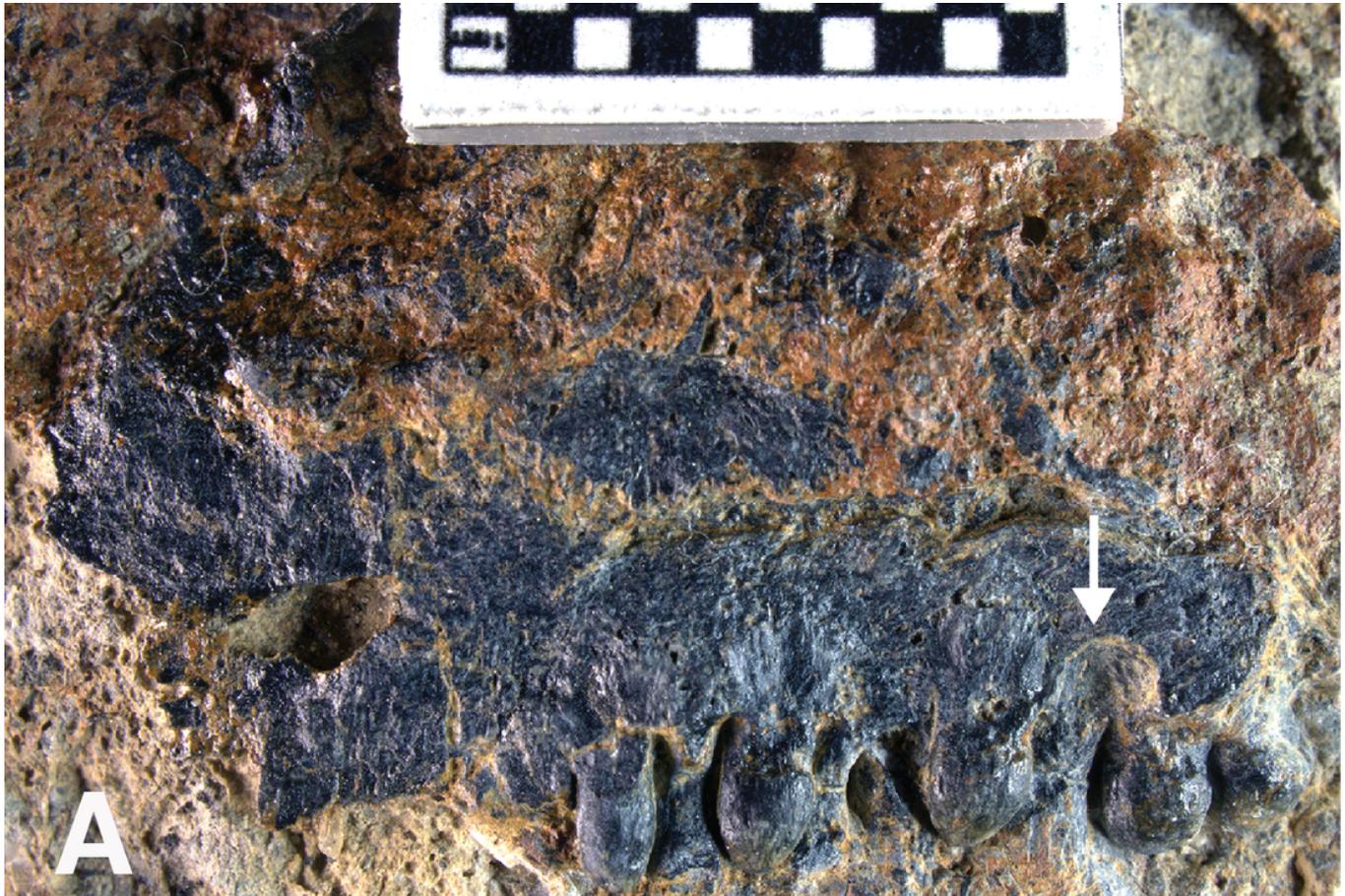
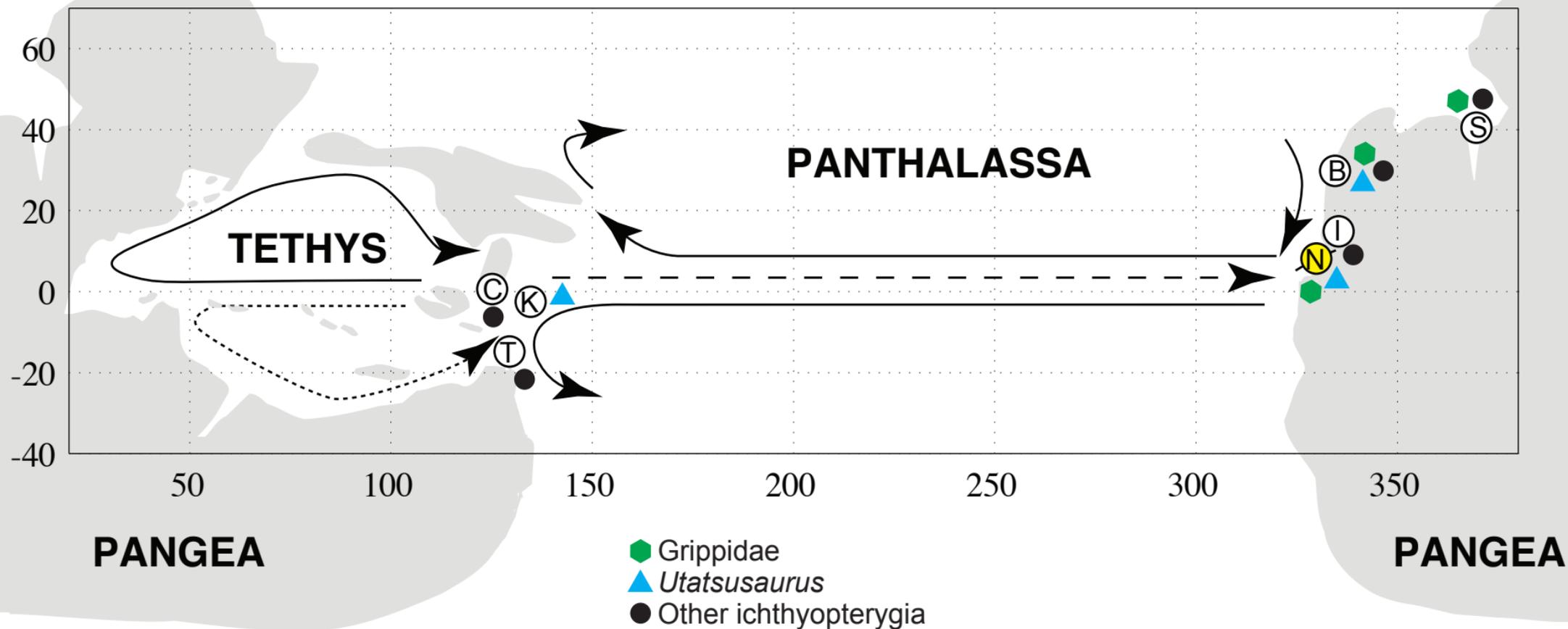


Figure 4(on next page)

Distribution of Early Triassic ichthyopterygians.

Paleogeographic distribution of Early Triassic ichthyopterygians, map modified from Brayard et al. (2009). Locality abbreviations as follows: (B) British Columbia; (C) South China; (I) Idaho; (K) South Kitakami; (N) Nevada (present study, highlighted in yellow); (S) Spitsbergen; (T) Timor. Arrows indicate inferred ammonoid dispersal routes (Brayard et al., 2009)



5

Distinctive sedimentary structures associated with horizon of USNM 559350

Spherical structures in FH1-5 that may represent microbial structures or sponges. This appears to be a widespread and distinctive regional Lower Triassic facies associated with recovery from the end-Permian mass-extinction. Vertebrate fossils also occur in this horizon including USNM 559350 described here. Hammer for scale is approximately 30 cm in length.



Table 1 (on next page)

Summarized tooth measurements from USNM 559349 and USNM 559350.

All measurements in mm except for shape index and crown ratio, which are ratios.

Specimen		Proximal width (mm) ^a	Exposed height (mm) ^b	Crown width (mm) ^c	Crown height (mm) ^d	Crown shape index ^e	Crown ratio ^f
USNM 559349	Max.	4.2	11.1	2.3	4.5	2.3	0.51
	Min.	2.5	6.4	1.9	2.7	1.3	0.39
	Mean	3.2	8.5	2.1	3.6	1.8	0.43
USNM 559350	Max.	1.6	2.5	2.1	1.7	2.1	0.88
	Min.	1.0	1.8	1.2	1.1	1.0	0.61
	Mean	1.4	2.2	1.6	1.5	1.4	0.69

1

^a Measured as mesio-distal width of the root at the jawline, following Motani (1996).

^b Measured as distance from tip of crown to jaw margin.

^c Measured as mesio-distal width of the crown at its widest point.

^d Base of crown is distinctive in USNM 559349 due to crown ornamentation; base of crown in USNM 559350 is less distinct but can be approximated by slight basal constriction.

^e Calculated as crown height/crown width, following Massare (1987).

^f Calculated as crown height/exposed tooth height, following Motani (1996).