A previously overlooked, highly diverse early Pleistocene elasmobranch assemblage from southern Taiwan (#73574)

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A previously overlooked, highly diverse early Pleistocene elasmobranch assemblage from southern Taiwan

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Abstract

The Niubu fossil locality in Chiayi County, southern Taiwan is best known for its rich early Pleistocene marine fossils that provide insights into the poorly understood past diversity in the area. The elasmobranch teeth at this locality have been collected for decades by the locals, but are not formally described and have received little attention. Here, we described three museum collections of elasmobranch teeth (n = 697) from the Liuchungchi Formation (1.90–1.35 Ma) sampled at Niubu locality, with an aim of constructing a more comprehensive view of the past fish fauna in the subtropical West Pacific. The assemblage is composed of 20 taxa belonging to 9 families and is dominated by *Carcharhinus* and *Carcharodon*. The occurrence of †*Hemipristis serra* is of particular importance because it is the first Pleistocene record in the area. We highlight high numbers of large *Carcharodon carcharias* in our sample exceeding 4 m, along with the diverse fossil elasmobranchs, suggesting that a once rich and thriving marine ecosystem in an inshore to offshore shallow-water environment during the early Pleistocene in Taiwan.

Introduction

The Indo-West Pacific is regarded as one of the crucial marine biodiversity hotspots in the world (Myers et al., 2000; Bellwood & Meyer, 2009). Most of the species are concentrated in the coral triangle area that has its northern limit extending to southern Taiwan. A remarkable 181 chondrichthyan species have been recorded in the modern fish fauna of Taiwan (Ebert et al., 2013), approximating over 15% of the total number of global chondrichthyan species

(Weigmann, 2016). Such species diversity is regarded as one of the highest biodiversity hotspots for elasmobranchs when considering the size of Taiwan (Ebert et al., 2013). However, how this remarkable chondrichthyan fauna was formed and evolved in the past are not well understood, primarily because relevant fossil records are traditionally overlooked or unstudied, despite being well-represented in the marine deposits of Taiwan. Thus, comparisons for associated fossil fauna and past biogeographic distributions are limited, particularly in the tropical-subtropical Pacific. Lin et al. (2021) highlighted the need for paleontological data for understanding the historical context of fish fauna and further recommended research potentials in the region.

In the Western Foothills of Taiwan, numerous Neogene to Quaternary strata are known to be rich in marine fossils (e.g., Ribas-Deulofeu, Wang & Lin, 2021; Lin & Chien, 2022; Lin et al., 2022). In particular, the early Pleistocene Liuchungchi Formation in the Niubu area, Chiayi County, southwestern Taiwan is of particular research interest due to its abundance and diversity of marine fauna. The fauna includes mollusks (Hu, 1989; Xue, 2004), crabs (Hu, 1989; Hu & Tao, 1996, 2004; Xue, 2004), sea urchins (Hu, 1989; Xue, 2004), whale barnacles (Buckeridge, Chan & Lee, 2018), teleost (Tao, 1993) and otoliths (Lin et al., 2018), and elasmobranch teeth (Xue, 2004). Fossils from this area have been collected by the late W.-J. Xue during the 1980s-2000s, and currently this large and diverse collection (over 3,000 specimens) is mainly deposited in the Chiayi Municipal Museum, Chiayi City, Taiwan (CMM). There is a considerable number of elasmobranch teeth from Xue's collection that were partly reported by Xue (2004) in a form of photographic atlas without descriptions. Another collection donated by Prof. Hsi-Jen Tao (National Taiwan University) to the Biodiversity Research Museum, Academia Sinica, Taipei, Taiwan (BRMAS) is available. An additional small collection is also deposited in the National Taiwan Museum (NTM). The purpose of this present study is to properly document the occurrences of these elasmobranch fossils from the Liuchungchi Formation at the Niubu locality based on these two collections and few newly collected specimens. The diverse association of teeth provides opportunities for obtaining a more complete view of the Pleistocene elasmobranch fauna in the rarely explored subtropical West Pacific.

Geological setting

Since the late Miocene, the island of Taiwan was gradually uplifted by the Penglai orogeny—the collision between the Chinese continental margin and the Luzon Arc—and, subsequently, a series of subsiding foreland basins were formed in western Taiwan (Ho, 1976; Suppe, 1984; Lundberg et al., 1997; Lin & Watts, 2002; Nagel et al., 2013; Chen, 2016). These foreland basins gradually developed from north to south accumulating clastic sediments (Ho, 1967; Covey, 1984; Teng, 1990), and in the south, the basins have high deposition rates (700–900 m/Ma) due to a deeper depositional environment (Chen, Huang & Yang, 2011). Thus, the depositional sequences reflect sea-level changes during the Quaternary that followed the 100 ky orbit eccentricity cycles (Chen, Huang & Yang, 2011; Chen, 2016). Meanwhile, thick pre-orogenic and synorogenic sediments infilling the foreland basin were squeezed and uplifted, which formed

the 7–9 km Miocene to Pleistocene strata in the Western Foothills (Yu & Chou, 2001; Nagel et al., 2013).

The Liuchungchi Formation in the Niubu area, Chiayi County is exposed along the Bazhang River (Fig. 1B). Four successive formations from east to west, lower to upper can also be observed along this river: the Liuchungchi, Kanhsialiao, Erhchungchi, and Liushuang formations (Stach, 1957; Chou, 1975; Chen, Huang & Yang, 2011; Chen, 2016; Fig. 1B, C). The age of the Liuchungchi Formation is 1.90–1.35 Ma (Chen, 2016), with a deposition rate of about 700 m/Ma in the lower section and 1,100 m/Ma upsection, the maximum thickness of the formation is 760 m (Chen, Huang & Yang, 2011). The Liuchungchi Formation is composed of dozens of depositional sequences, each representing a 41 ky climate cycle (Chen, Huang & Yang, 2011; Chen, 2016). The depositional environment can be divided into two distinct sections, with the lower sequence composed of thick sandstone with cross bedding, parallel bedding, and strong bioturbation reflecting shoreface to the offshore transition zone, and the upper sequence composed of interbeds of sandstone and shale and storm deposits in the form of sandstone, indicating the inner offshore (Chen, Huang & Yang, 2011; Chen, 2016).

Materials & Methods

The fossil site is located in the Niubu area, Chiayi County, southwestern Taiwan, about 15 km east of Chiayi City (Fig. 1A). The layers containing fossils are exposed along the Bazhang River, just downstream of a dam near a high-voltage tower, where they are readily accessible during the winter and dry seasons when the water level is low (Supplemental Fig. S1). Fossil mollusks are very abundant in several of the condensed layers, as well as fragments of crabs, sea urchins, and teleost fish bones (Fig. 2A). Fossil shark teeth are rare based on both surface collecting and bulk sampling conducted during our several field trips in 2018–2022. Bulk sediment samples of over 830 kg (Sites 1–3 in Fig. 2B) were sieved (500-μm mesh) from the loosely cemented siltstone without any sedimentary structure, yielding a large number of otoliths (Lin et al., unpublished data), but only one shark tooth and two ray teeth. We note the discrepancy in the numbers of elasmobranch specimens between museum collections and our field surveys is present, and it is explained by the fact that the larger sample sizes in museum collections primarily reflect collecting based on fortuitous occurrences of shark teeth over the past 3-4 decades, compared to collecting based on our limited number of field surveys. Moreover, the initial purpose of the bulk sampling was for collecting teleost otoliths instead of elasmobranch teeth, and these selected sites contained more otoliths than teeth.

The upstream of Bazhang River contains strata older than the Liuchungchi Formation; these include the Neogene Tangenshan Sandstone and Yenshuikeng Shale, which are exposed approximately 1 km east to the weir (Figs. 1, 2), with an elevation of more than 300 m compared to the level of our sampling sites. Both stratigraphic units are composed of consolidated sandstones with some marine fossils such as mollusks, but these are different from the fine,

unconsolidated siltstones of the Liuchungchi Formation. After storms and rainy seasons, numerous blocks of sandstones from these older strata can be found along the riverbed of the Bazhang River (Supplemental Fig. S1), but these lithologically distinct rocks are confined to areas below our sampling sites. We thus consider that the mixing of older Neogene fossils with our Pleistocene specimens is improbable, and that our assemblage represents fossils derived entirely from the Liuchungchi Formation.

The BRMAS, CMM, and NTM collections analyzed here were collected from the surface exposure of the Niubu locality without bulk-sampling of sediments; however, the exact stratigraphic horizons and detailed lithologic character within the Liuchungchi Formation for each specimen are not known. Stacked images of teeth were taken and measurements of crown height (CH), mesial crown edge length (MCL), and basal crown width, (BCW) were noted wherever necessary. Specimens from the BRMAS are registered under ASIZF, CMM under CMM F, and those in the NTM are under NTM I. Because the Pleistocene is relatively close to modern times, the morphology of elasmobranch teeth has not changed much from that time to the present. Therefore, identifications of these fossil teeth were conducted by comparing them with teeth of extant taxa.

The diversity of our elasmobranch assemblage was compared with other Pleistocene assemblages to highlight its significance within the associated spatio-temporal context. Taxonomic composition and abundance data from the early Pleistocene temperate assemblages of Japan (Karasawa, 1989; Kawase & Nishimatsu, 2016; Tanaka & Taru, 2022) and tropical records from Java (Koumans, 1949; Yudha et al., 2018) and Sulawesi (Hooijer, 1954) were extracted. We calculated diversity indices, including species richness, Shannon's entropy, Simpson's diversity index, and Fisher's alpha for a general comparison.

Systematic Paleontology

A summary of taxa and their numeric abundance are listed in Table 1. The elasmobranch assemblage contains 697 teeth, consisting of 9 families and 20 taxa. The classification scheme follows that of Nelson, Grande & Wilson (2016), except for the family Galeocerdonidae, which we follow Fricke, Eschmeyer & Van der Laan (2022). General morphological terminology follows that of Compagno (1984, 2002), Purdy et al. (2001), Shimada (2002), Purdy (2006), Cappetta (2012), and Ebert et al. (2013). The synonymy list is limited to relevant records from Taiwan (Huang, 1965; Uyeno, 1978; Hu & Tao, 1993; Xue, 2004; Tao & Hu, 2008).

Class Chondrichthyes Huxley, 1880
Order Lamniformes Berg, 1958
Family Carchariidae Müller & Henle, 1838
Genus *Carcharias* Blainville, 1816
Carcharias taurus Rafinesque, 1810

(Fig. 3)

1978 *Odontaspis* sp.; Uyeno, pl. 1, fig. 5.

Referred specimens: n = 2: ASIZF0100320, CMM F0204.
Description: CH = 12.92–16.83 mm; MCL = 12.18–15.54 mm; BCW = 7.07–7.84 mm. The teeth are characterized by a slender, dagger-like main cusp and a single pair of small lateral cusplets. The crown exhibits no serrations. The lingual protuberance of the root is prominent.
Remarks: The teeth of *Carcharias taurus* are similar to those of *Odontaspis noronhai* and *O. ferox* by having a slender main cusp and lateral cusplet. However, the lateral cusplets of *Odontaspis* are more pronounced than those of *C. taurus*, including the fact that teeth of *O. ferox* typically exhibit multiple pairs of lateral cusplets.

Genus *Carcharodon* Smith, 1838
Carcharodon carcharias (Linnaeus, 1758)
(Fig. 4)

1978 *Carcharhinus* sp.; Uyeno, pl. 1, fig. 4, pl. 2, fig. 7.
1978 *Carcharodon carcharias*; Uyeno, pl. 3, figs. 12, 13.
2004 Elasmobranchii indet.; Xue, pl. 1, figs. 1–6, pl. 2, figs. 1–7, pl. 3, figs. 1, 2–7, pl. 7 fig. 2.

Referred specimens: n = 55: ASIZF0100322–0100346, 0100435, 0100465, 0100530. CMM F0001–F0005, F0007–F0010, F0012–F0022, F0210, F0212, F2824, F2825, F2830, NTM I01122, I01123.
Description: CH = 6.76–41.03 mm; MCL = 9.61–45.68 mm; BCW = 8.74–37.09 mm. The upper teeth (Fig. 4A–N) are broad and triangular. The cutting edge of both mesial and distal sides is almost straight with coarse serrations. The labial face of the crown is flat and the lingual face is convex, where the crown is erect and symmetric to slightly distally inclined depending on tooth positions. The root is slightly arched, and the nutritive foramina and transverse groove are not prominent or absent. The lower teeth (Fig. 4O–X) have a more robust but narrower serrated crown and bilobate roots with a rounded lingual face compared to the upper teeth.
Remarks: The genus *Carcharodon* is represented by three species: †*C. hastalis*, †*C. hubbelli*, and *C. carcharias*. Whereas *C. hastalis*, which was traditionally placed in the genus *Isurus* or †*Cosmopolitodus*, lived through the Miocene and early Pliocene, †*C. hubbelli* in the late Miocene and *C. carcharias* in the early Pliocene–Recent form a single lineage of chronospecies by developing serrations on their teeth (Ehret et al., 2012). The specimens described in this present paper exhibit well-developed serration consistent with teeth of *C. carcharias* (e.g., Hubbell, 1996), and not like the teeth of *C. hubbelli* with weak serrations (Ehret et al., 2012).

They include the largest dental remains among all the shark tooth specimens described in this paper.

Genus *Isurus* Rafinesque, 1810
Isurus oxyrinchus Rafinesque, 1810
(Fig. 5)

?1965 *Isurus hastalis*; Huang, pl. 22, figs. 12–14.
1993 *Isurus hastalis*; Hu & Tao, pl. 24, figs. 6, 8.
2004 Elasmobranchii indet.; Xue, pl. 5, fig. 3, pl. 8, fig. 4.
2008 *Isurus* sp.; Tao & Hu, pl. 2, figs. 1–2.

Referred specimens: n = 6: ASIZF0100317–0100319, 0100321, CMM F0242, NTM I01131_1.
Description: CH = 9.86–27.81 mm; MCL = 12.21–26.89 mm; BCW =9.31–9.96 mm. The mesial teeth have a slender, dagger-like, unserrated crown that is erect or lingually curved with an apical labial flexure (Fig. 5A–H). The root, if preserved, has two rather narrow lobes with a moderately tight basal concavity. The distal teeth have a flatter and broader, distally curved, unserrated crown with a short but mesiodistally wide root (Fig. 5I, J).
Remarks: Two extant species of *Isurus* are known: *I. oxyrinchus* and *I. paucus*. *Isurus oxyrinchus* has a more elongated and more labially curved crown than *I. paucus* (Whitenack & Gottfried, 2010). The teeth of *I. oxyrinchus* are also similar to those of *Carcharias taurus*, but the teeth of *C. taurus* have a pair of lateral cusplets that is absent in the teeth of *I. oxyrinchus* (Wilmers, Waldron & Bargmann, 2021). Huang (1965) reported a tooth of †*I. hastalis* (= *Carcharodon hastalis*; see above) from the Pleistocene Cholan Formation in Hsinchu, northern Taiwan; however, its species identification is questionable and the whereabouts of the specimens is unknown for verification.

Order Carcharhiniformes Compagno, 1973
Family Hemigaleidae Hasse, 1878
Genus *Hemipristis* Agassiz, 1835
†*Hemipristis serra* Agassiz, 1843
(Fig. 6)

1978 *Hemipristis serra*; Uyeno, pl. 1, fig. 2.
2004 *Hemipristis* sp.; Xue, pl. 5, figs. 1, 2, 5, 6, 7.
2004 Elasmobranchii indet.; Xue, pl. 5, fig. 5, pl. 7, figs. 3, 5, pl. 9, figs. 6, 7.
2008 *Hemipristis serra*; Tao & Hu, pl. 6, fig. 1.

Referred specimens: n = 7: ASIZF0100460–0100462, CMM F0232, F2826, F2827, NTM I01131_2.

Description: CH = 5.21–30.81 mm; MCL = 8.73–41.38 mm; BCW = 6.50–36.59 mm. All collected specimens of this taxon represent upper teeth that are characterized by a distally inclined, broad triangular crown, and mesiodistally separated bilobate root. Coarse serrations are present along the distal cutting edge, whereas serrations along the meisal cutting edge are finer. The root has a prominent lingual protuberance with a deep nutritive groove, and has a notch-like shallow basal concavity. The crown overhangs the root, and their boundary, especially on the lingual face, is strongly arched.

Remarks: As presumed sister species, the teeth of extinct †*Hemipristis serra* and extant *H. elongata* are similar. However, compared to †*H. serra*, teeth of *H. elongata* possess a more gracile crown and a longer apex without serration, and a narrower root (Smith, 1957; Purdy et al., 2001). The Pleistocene records of †*H. serra* are rare globally compared to its Neogene records (Hooijer, 1954, 1958; Yabumoto & Uyeno, 1994; Carrillo-Briceño et al., 2015; Ebersole, Ebersole & Cicimurri, 2017; Boessenecker, Boessenecker & Geisler, 2018).

Family Carcharhinidae Jordan & Evermann, 1896

Genus *Carcharhinus* Blainville, 1816

Remarks: The identification based on teeth below the genus level is difficult for *Carcharhinus* (Compagno, 1984, 1988; Purdy et al., 2001; Naylor & Marcus, 1994; Marsili, 2006; Voigt & Weber, 2011; Ebert, Dando & Fowler, 2021). Most of the upper teeth are triangular with their crown inclining distally. In different species, the crown varies from narrow to broad, and has smooth to coarsely serrated cutting edges, different notch angles on distal cutting edges, and the straight to convex mesial cutting edge. At least nine species of *Carcharhinus* are recorded in the collections: *C. altimus*, *C. amboinensis*, *C. leucas*, *C. limbatus*, *C. longimanus*, *C. obscurus*, *C. plumbeus*, *C. sorrah*, and *C. tjutjot*. See remarks below for comparisons among other similar-looking species.

Carcharhinus altimus (Springer, 1950)

(Fig. 7)

Referred specimens: n = 17: ASIZF0100357, 0100359, 0100362, 0100363, 0100365, CMM F0080, F0101, F0113, F0134, F0214, F0224, F0293, F0304, F0322, F0363, TNM I01125, I01129_1.

Description: CH = 4.55–9.82 mm; MCL = 7.91–12.72 mm; BCW = 7.10–10.92 mm. The specimens examined in this study consist only of upper teeth. The crown of the upper teeth is finely serrated and varies in shape from a tall triangle to distally oblique. There is a notch on the distal cutting edge, whereas a slight constriction occurs on the lower part of the mesial cutting edge. The root is arched and has a nutritive groove. The roots of some specimens are not well-preserved (Fig. 7A, B, E, F, I–L), but where well-preserved (Fig. 7C, D, G, H), it is arched and exhibits a nutritive groove on the lingual face.

Remarks: Teeth of *Carcharhinus altimus* and *C. plumbeus* are similar. However, those of *C. altimus* exhibit a distally bent apex unlike those of *C. plumbeus* that show an apically directed apex (Figs. 7 vs. 13).

Carcharhinus amboinensis (Müller & Henle, 1839)
(Fig. 8)

Referred specimens: n = 5: ASIZF0100366, 0100368, 0100369, CMM F0209, F0229.
Description: CH = 6.88–8.95 mm; MCL = 9.28–14.74 mm; BCW = 9.16–16.86 mm. The triangular crown is broad and exhibits coarse serrations although the serrations become smaller towards the apex. A prominent tooth neck is present between the crown and root on the lingual face. There is a notch on the distal cutting edge, whereas the mesial cutting edge is nearly straight. The bilobed root is gently arched and has a nutritive groove on the lingual face.
Remarks: Teeth of *Carcharhinus amboinensis*, *C. leucas*, and *C. longimanus* are very similar (Marsili, 2006; Voigt & Weber, 2011). However, the angle of the notch on the distal cutting edge of *C. longimanus* is larger than *C. leucas* and *C. amboinensis*. Compared to the teeth of *C. leucas*, the upper teeth of *C. amboinensis* are somewhat broader, the crowns are generally lower and more distally curved, and their distal heel is more pronounced and is closer to the base of the crown (Kocsis et al., 2019).

Carcharhinus leucas (Valenciennes, 1839)
(Fig. 9)

?1965 *Carcharhinus gangeticus*; Huang, pl. 22, figs. 19, 20.
2004 Elasmobranchii indet.; Xue, pl. 3, fig. 2, pl. 4, fig. 3, pl. 7, fig. 7, pl. 9, fig. 4.

Referred specimens: n = 71: ASIZF0100390, 0100393–0100398, 0100400–0100404, 0100411, 0100419, 0100424, 0100425, 0100481, CMM F0154, F0155, F0157, F0159, F0162, F0163, F0165–F0168, F0170–F0175, F0180, F0183, F0186–F0188, F0190, F0192, F0198–F0201, F0205, F0206, F0221, F0222, F0227, F0231, F0240, F0244, F0246, F0249, F0288, F0290, F0297, F0299, F0301, F0317, F0319, F0321, F0328, F0332, F0334, F0341, F0342, F0348, F0354, F0362, NTM I01130_2.
Description: CH = 4.87–18.68 mm; MCL = 7.97–21.69 mm; BCW = 8.73–30.56 mm. The teeth of *Carcharhinus leucas* are generally robust. The crown of the upper teeth (Fig. 9A–P) is broad and triangular with a slight distal inclination. The middle of the distal cutting edge is concave, forming a weak notch, whereas the mesial cutting edge is straight to slightly convex. Both cutting edges are coarsely serrated, but the sizes of serrations are smaller at the base and apex of the crown than those in the middle. The boundary between the crown base and root on the lingual face displays a V-shape tooth neck. The bilobate root is arched and displays a weak

nutritive groove on the lingual face (Fig. 9A–H, K, L). The lower teeth (Fig. 9Q, R), that have fine serrations, are labiolingually thicker and mesiodistally narrower than the upper teeth.
Remarks: Marsili (2006) described the crown of *Carcharhinus longimanus* as larger, more elongate and possessing a straighter root margin compared to that of *C. leucas*. In addition, based on the images of *Carcharhinus* by Garrick (1982) and Voigt & Weber (2011), we find some other slight differences in tooth morphology between the two species. For example, the angle on the distal cutting edge of the upper teeth in *C. longimanus* is larger than that in *C. leucas*, making the crown of *C. leucas* incline more distally than that in *C. longimanus*. In addition, the tooth shape of *C. leucas* is close to a wide-bottom triangle, whereas that of *C. longimanus* forms a taller triangle. Furthermore, the lower teeth of *C. leucas* tend to exhibit a stronger demarcation between the main cusp and mesial and distal heels than those of *C. longimanus* with a smoother cusp-heel transition.

Carcharhinus limbatus (Valenciennes, 1839)
(Fig. 10)

1978 *Carcharhinus* sp.; Uyeno, pl. 3, fig. 14.
2004 Elasmobranchii indet.; Xue, pl. 8, fig. 5.

Referred specimens: n = 40: ASIZF0100467–0100480, 0100482, 0100483, CMM F0056, F0111, F0216, F0217, F0234, F0236–F0238, F0286, F0289, F0291, F0295, F0306, F0307, F0310, F0368–F0373, NTM I01127, I01133_2, I01134_2.
Description: CH = 7.70–9.31 mm; MCL = 10.02–13.26 mm; BCW = 10.26–14.70 mm. Our specimens consist only of upper teeth. The teeth of *C. limbatus* are serrated and are characterized by a narrow cusp that is erect to slightly oblique distally with a mesiodistally wide crown base. The serrations near the crown base are coarser than those towards the apex. The root is apicobasally shallow. Its base is straight to slightly arched with a prominent deep nutritive groove that forms a notch along the root base.
Remarks: Although similar, teeth of *Carcharhinus limbatus* can be distinguished from those of *C. amblyrhychoides*, *C. brachyurus* and *C. brevipinna*. Unlike the teeth of *C. limbatus*, the serrations on the cutting edges tend not to continue to the crown base in *C. amblyrhychoides*, and are absent or weak in *C. brevipinna*, and in *C. brachyurus*, the apex is more pointed and more distally directed than in *C. limbatus* (Garrick, 1982; Voigt & Weber, 2011). In addition, the crowns of *C. limbatus* have a narrow, erect cusp with a sharp transition to their broad crown base that is distinct from all other congeneric specimens in our material. The teeth of *C. limbatus* and *C. amblyrhynchoides* are, however, very difficult to distinguish. Kocsis et al. (2019) noted a narrower crown with finer serrations in *C. limbatus*, but this character is not clear in our specimens. Currently, no records of *C. amblyrhynchoides* have been reported in Taiwan (Ebert et al., 2013; Shao, 2022); therefore, we tentatively assign these specimens to *C. limbatus*.

Carcharhinus longimanus (Poey, 1861)
(Fig. 11)

1965 *Carcharhinus gangeticus*; Huang, pl. 22, figs. 21, 22.
2004 Elasmobranchii indet.; Xue, pl. 4, fig. 4, pl.7, fig. 6, pl. 9, fig. 1.

Referred specimens: n = 36: ASIZF0100370, 0100371, 0100373–0100382, 0100391, 0100392, 0100421, 0100422, 0100428, 0100466, CMM F0006, F0011, F0087, F0151, F0153, F0156, F0158, F0182, F0189, F0194, F0195, F0197, F0223, F0248, F0287, F0294, NTM I01128, NTM I01130.
Description: CH = 10.23–15.93 mm; MCL = 13.51–22.08 mm; BCW = 13.20–21.69 mm. The crowns of the upper teeth (Fig. 11A–P) are broad, triangular, and coarsely serrated. The distal cutting edge is weakly concave, whereas the mesial cutting edge is nearly straight. The crown base on the lingual side is deeply concave and is accompanied basally by a narrow tooth neck and a deep bilobate root with a shallow nutritive groove. The lower teeth (Fig. 11Q–T) are thicker and narrower than the upper teeth, they also have fine serrations on the cutting edges. The boundary between the crown base and root on the lingual side is also deeply concave with a V-shaped tooth neck.
Remarks: See remarks under *Carcharhinus leucas*.

Carcharhinus obscurus (Lesueur, 1818)
(Fig. 12)

2004 Elasmobranchii indet.; Xue, pl. 4, fig. 7.

Referred specimens: n = 25: ASIZF0100372, 0100383–0100389, 0100399, CMM F0123, F0143, F0148, F0160, F0164, F0176–F0179, F0181, F0184, F0196, F0208, F0338, F0353, NTM I1132_3.
Description: CH = 5.04–15.14 mm; MCL = 7.57–21.61 mm; BCW = 9.56–20.96 mm. The specimens in this study consist only of upper teeth. They are broad and triangular with coarse serrations, although the serrations tend to become finer apically. The mesial cutting edge is overall slightly convex with a marked distally directed apex. The distal cutting edge has a relatively deep notch, but the degree of the angle varies based on tooth positions within the dentition. The crown base on the lingual side is moderately concave and is accompanied by a prominent tooth neck and a relatively robust bilobed root that has a shallow nutritive groove.
Remarks: The crown of *Carcharhinus obscurus* is mesiodistally broad and typically exhibits coarse serrations along the middle section of both cutting edges, a feature for separating all other congeneric specimens in the present study.

Carcharhinus plumbeus (Nardo, 1827)

(Fig. 13)

Referred specimens: n = 51, ASIZF0100405–0100410, 0100412, 0100429, CMM F0074–F0077, F0079, F0086, F0088, F0091, F0096, F0100, F0106, F0115, F0124, F0144, F0146, F0169, F0225, F0228, F0292, F0302, F0316, F0318, F0320, F0325–F0327, F0330, F0331, F0333, F0335, F0337, F0346, F0347, F0349, F0350, F0352, F0356, F0360, F0361, F0364, F0365, F0367, NTM I01124.

Description: CH = 6.28–12.17 mm; MCL = 7.81–17.06 mm; BCW = 7.48–13.39 mm. The teeth that are referred to this species are all upper teeth. They are triangular with a slight distal inclination and with fine serrations. The mesial cutting edge is nearly straight, whereas the distal cutting edge tends to form a shallow notch close to the crown base. The root is bilobate and arched, and a shallow nutritive groove is present on the lingual face.

Remarks: The crown of *Carcharhinus. plumbeus* is narrower and elongate than that of *C. leucas*, *C. longimanus*, *C. obscurus*, and *C. amboinensis*, but it is wider than that of *C. altimus*.

Carcharhinus sorrah (Valenciennes, 1839)

(Fig. 14)

Referred specimens: n = 11: ASIZF0100418, CMM F0117, F0119, F0122, F0126, F0129, F0135, F0140, F0303, F0343, F0344.

Description: CH = 4.39–5.80 mm; MCL = 5.55–9.82 mm; BCW = 4.03–9.85 mm. All teeth identified to this species are represented by upper teeth. Their crown consists of a finely serrated triangular cusp that strongly inclines distally along with a coarsely serrated, relatively broad distal heel. The apex is narrow and may be slightly recurved apically (Fig. 14E–H). The serrations on the distal heel become smaller distally, where finer secondary serrations are observed on one or two of the mesial-most serrations. Well-preserved specimens exhibit a strong nutritive groove on the lingual face that forms a notch along the root base.

Remarks: According to Voigt & Weber (2011), the crown of the upper teeth in *Carcharhinus sorrah* is high, and its distal cutting edge is deeply notched. These features are seen in our specimens, but their descriptions for the serrations are slightly different. The serrations on the central part of the mesial cutting edges are coarser in Voigt & Weber (2011), whereas in our specimens, the coarsest serrations are on the basal part of the mesial cutting edges. Furthermore, the main cusp and coarse serrations in our specimens are farther apart than those figured by Voigt & Weber (2011). The teeth of *C. tjutjot* and *C. sorrah* are both characterized by a coarsely serrated distal heel, but the teeth of *C. tjutjot* differ from those of *C. sorrah* by having fewer but larger serrations forming a distal heel (Figs. 14 vs. 15).

Carcharhinus tjutjot (Bleeker, 1852)

(Fig. 15)

Referred specimens: n = 19: ASIZF0100413–0100417, CMM F0116, F0136–F0138, F0142, F0296, F0298, F0323, F0324, F0339, F0345, F0357, F0376, F0377.
Description: CH = 4.25–5.82 mm; MCL = 6.03–9.01 mm; BCW = 5.61–7.94 mm. The specimens of this species described here are all represented by upper teeth. They have a robust, distally inclined, triangular cusp, a small distal heel consisting of coarse serrations that rapidly diminish in size distally. The strongly inclined mesial cutting edge is relatively straight, where the apex may slightly recurve apically and its serrations become slightly coarser towards the base. Finer secondary serrations are observed on the first and possibly second mesial-most serrations on the distal heel. The root is weakly bilobate as the root base is nearly straight. Well-preserved specimens show a shallow nutritive groove on the lingual face of the root.
Remarks: The teeth of *Carcharhinus sealei*, *C. dussumieri*, *C. coatesi*, and *C. tjutjot* are very similar (White, 2012). The differences among them are the serrations on the cutting edges. The serrations of *C. sealei* are present only on the basal half of the mesial cutting edge, where the distal cutting edge, including the distal heel, is smooth (White, 2012). Both cutting edges of teeth in *C. coatesi* have fine to coarse serrations, but their distal heel is smooth. In *C. dussumieri*, both cutting edges of teeth, including the distal heel, have evenly-sized coarse serrations. The teeth of *C. tjutjot* also have evenly-sized serrated cutting edges, including the distal heel. *Carcharhinus dussumieri* and *C. tjutjot* have long been misidentified due to their similar appearance, but *C. dussumieri* is now considered a West Indian species distributed from the Persian Gulf to India, whereas *C. tjutjot* is distributed from Indonesia to Taiwan (White, 2012).

Genus *Rhizoprionodon* Whitley, 1929
Rhizoprionodon acutus (Rüppell, 1837)
(Fig. 16)

Referred specimens: n = 8: ASIZF0100463, 0100464, CMM F0110, F0120, F0121, F0130, F0131, F0218.
Description: CH = 3.97–5.35 mm; MCL = 6.22–10.82 mm; BCW = 7.68–10.69 mm. The upper teeth of this species have a crown that is strongly inclined distally and is accompanied by a low distal heel (Fig. 16A–H). Both cutting edges, including the distal heel, are smooth or exhibit fine irregular serrations. The mesial cutting edge is overall straight, whereas the junction between the cusp and distal heel is deeply notched. A deep nutritive groove is present on the lingual side of the root that continues to the root base. The root is low with little to no basal concavity. ASIZF0100464 (Fig. 16I, J) is a lower tooth, where its crown that is unserrated is more gracile than the upper teeth with a concave mesial cutting edge. The root morphology is similar to that of lower teeth.
Remarks: The teeth of *Rhizoprionodon acutus* are serrated in adults (Compagno, 1984). In our specimens, the serrations are absent, indicating immature individuals. Distinguishing between the teeth of *R. acutus* and *R. oligolinx* is difficult, where both have very fine irregular serrations.

However, due to the questionable distribution of *R. oligolinxi* in Taiwan (Ebert et al., 2013; Froese & Pauly, 2022), we tentatively assign these specimens to *R. acutus*.

Family Galeocerdonidae *sensu* Ebersole, Cicimurri & Stringer, 2019
Genus *Galeocerdo* Müller & Henle, 1837
Galeocerdo cuvier (Péron & Lesueur, 1822)
(Fig. 17A–H)

?1965 *Galeocerdo aduncus*; Huang, pl. 22, figs. 10, 11.
?1978 *Galeocerdo aduncus*; Uyeno, pl. 1, fig. 3.
2004 Elasmobranchii indet.; Xue, pl. 6, figs. 1–7, pl. 8, fig. 6.

Referred specimens: n = 7: ASIZF0100459, CMM F0213, F0215, F0245, F2823, F2829, NTM I01121.
Description: CH = 12.27–17.72 mm; MCL = 17.37–26.88 mm; BCW = 18.10–28.05 mm. The teeth of *G. cuvier* are characterized by a coarsely serrated crown with a cusp that strongly curves distally and a prominent distal heel demarcated by a deep notch at approximately 90 degrees angle along the distal cutting edge. Fine secondary serrations are present on the coarser primary serrations (Fig. 17E, H). The serrations on the distal heel in ASIZF0100459 (Fig. 17F,G) are weak, the width to crown height ratio suggests this tooth represents a posterior position. CMM F0245 and CMM F0215 are anterior teeth with well-marked serration (Fig. 17A–D).
Remarks: Five extinct species and one extant species of *Galeocerdo* are considered valid: the Eocene †*G. clarkensis* and †*G. eaglesomei*, Oligocene–late Miocene †*G. aduncus*, Miocene †*G. mayumbensis*, Pliocene †*G. capellini*, and the Pleistocene–Recent *G. cuvier* (Purdy et al., 2001; Türtscher et al., 2021). The specimens described here are identified as *G. cuvier*, particularly because of the presence of secondary serrations (Cigala-Fulgosi & Mori, 1979; Türtscher et al., 2021). Huang (1965) reported a questionable occurrence of †*G. aduncus* from the Pleistocene Cholan Formation in Hsinchu, northern Taiwan, however we consider the specimen lost. Uyeno (1978) reported another occurrence of †*G. aduncus* from the poorly constrained Plio-Pleistocene strata along the Tsailiao River in Tainan, southwestern Taiwan (as Miocene to Pleistocene in Uyeno, 1978). Although Uyeno's collection was deposited in the NTM, we were not able to locate the specimen of †*G. aduncus* in the collection. Nevertheless, although the whereabouts of the specimen is uncertain, it is interpreted here to have also belonged to *G. cuvier*.

Family Sphyrnidae Bonaparte, 1840
Genus *Sphyrna* Rafinesque, 1810
Sphyrna lewini (Griffith & Smith, 1834)
(Fig. 17I–L)

?1978 *Sphyrna* sp.; Uyeno, pl. 2, fig. 8

Referred specimens: n = 2: CMM F0235, F0312.
Description: CH = 4.64–6.85 mm; MCL = 7.51–11.17 mm; BCW = 7.19–10.53 mm. The tooth crown of *S. lewini* is characterized by a slender distally inclined cusp with a narrow, mesially extended base separated by a slight concavity along the mesial cutting edge and a low distal heel demarcated by a deep notch. Both cutting edges are smooth without serrations. The root is low, and its base is straight. It has a deep nutritive groove on the lingual side and extends to the root base.
Remarks: The teeth of *Sphyrna lewini* are most similar to *S. macrorhynchos* and *Loxodon macrorhinus*, but a slight concavity is present on the base of the mesial cutting edge in *S. lewini*, whereas the edge is almost straight in the latter two species (Ebert et al., 2013).

Order Myliobatiformes Compagno, 1973
Family Dasyatidae Jordan & Gilbert, 1879
Dasyatidae indet.
(Fig. 18)

Referred specimens: n = 2: ASIZF0100590, 0100591.
Description: The specimens are roughly hexagonal with a globular, thick crown and a well-divided bilobed root that is smaller than the crown and extends ventrally. The apex of the crown in both specimens is flat, but the specimen ASIZF0100590 (Fig. 18A–D) has blunt, rounded corners compared to ASIZF0100591 (Fig. 18E–H).
Remarks: The teeth referred to this taxon may belong to the genus *Dasystis* or *Himantura*, but because teeth of dasyatid taxa are highly variable in morphology, including sexual dimorphism and differences in ornamentation pattern (Taniuchi & Shimizu, 1993; Kajiura & Tricas, 1996; Herman, Hovestadt-Euler & Hovestadt, 1998, 1999, 2000), we refer our materials simply to Dasyatidae indet. Uyeno (1978) reported teeth of *Dasyatis* sp. from the Miocene to Pleistocene of Taiwan. However, whether our specimens are conspecific with Uyeno's (1978) specimens cannot be ascertained.

Family Aetobatidae Agassiz, 1858
Genus *Aetobatus* Blainville, 1816
Aetobatus sp.
(Fig. 19)

Referred specimens: n = 58: ASIZF0100549–0100580, CMM F0380, F0382, F0388, F0395, F0399–0409, F0412, F2848–F2850, F2852–F2854, NTM I01116, I01117, I01119, I01120.
Description: Teeth of *Aetobatus* are characterized by strongly extended roots on the lingual (posterior) side and the arcuate crown in apical view with a flat occlusal surface. The crown overhangs the root on the labial (anterior) side and the root is more prominent than the crown on

the lingual side. Both lingual and labial crown faces have fine vertical grooves as ornamentation. The root is polyaulocorhizous, consisting of anteroposteriorly oriented, densely packed, vertical lamellar plates.

Remarks: Five species of Myliobatidae (one *Aetobatu*, three *Aetomylaeus*, and one *Myliobatis*) are known from Taiwan (Ebert et al., 2013). All of which have grinding-type dental plates but each with different shapes and forms. The upper medial teeth of *Aetobatus ocellatus* are straight and elongate but slightly distally deflected towards the lingual side; its lower teeth are strongly arched towards the labial side. Considerable ontogenetic morphological change in dental plate is known in *Aetomylaeus* (Hovestadt & Hovestadt-Euler, 2013). Both upper and lower dental plates of adult *Aetomylaeus* are similar to the upper teeth of *Aetobatus*. Unlike adult individuals that have a single row of medial teeth, juveniles of *Aetomylaeus* have one medial, two lateral, and one posterior tooth rows (Hovestadt & Hovestadt-Euler, 2013). The hexagon shape of medial teeth is very similar to those of juvenile *Myliobatis* (Hovestadt & Hovestadt-Euler, 2013). Teeth of *Aetobatus* have weak ornaments on the labial and lingual crown, but in *Aetomylaeus*, beaded ridges with reticulated and pitting patterns are observed (Ebersole, Cicimurri & Stringer, 2019). Moreover, *Aetobatus* has teeth with a strong arched appearance than other myliobatid genera (see also remarks under *Myliobatis* sp.).

Family Myliobatidae Bonaparte, 1835

Genus *Myliobatis* Cuvier, 1816

Myliobatis sp.

(Fig. 20)

Referred specimens: n = 30: ASIZF0100581–0100589, CMM F0378, F0379, F0381, F0383–F0387, F0389–F0393, F0394, F0396–F0398, F0410, F0411, F2855, NTM I01118.

Description: Each tooth of *Myliobatis* has a flat occlusal surface and is laterally elongated and hexagonal that may be straight or slightly arched. The root is polyaulocorhizous with well-defined anteroposteriorly oriented, vertical lamellar plates separated by deep grooves, where the crown overhangs the root on the labial (anterior) face. The lingual and labial faces are ornamented with a network of fine reticulated ridges that grade into longitudinal ridges in the apical and become finer and anastomotic.

Remarks: The tooth plates of *Myliobatis* are similar to those of *Aetomylaeus* and *Aetobatus*, but the lateral angle of the hexagonal tooth plates in *Aetomylaeus* is more oblique than that of *Myliobatis* (Ebersole, Cicimurri & Stringer, 2019). The vertical lamellar plates of the root in *Myliobatis* are coarser than *Aetobatus*. Teeth of *Myliobatis* lack the tuberculated enameloid on the occlusal surface, whereas teeth of *Aetomylaeus* are reticulated on the labial and lingual faces (Ebersole, Cicimurri & Stringer, 2019). Because the total morphological variation range of teeth in many of the aetobatid and myliobatid (Myliobatinae) species is unknown (e.g., see Hovestadt & Hovestadt-Euler, 2013), we refrain from assigning the *Aetobatus* (see above) and *Myliobatis* teeth described here to the species level.

Discussion

Previous works on fossil elasmobranchs in Taiwan are very scarce, where they were limited in scope, often lacked formal description, and were mostly not based on specimens in museum repositories (Lin et al., 2021). Huang (1965) reported three shark taxa while describing a fossil whale tympanic bone from the early Pleistocene Cholan Formation in northern Taiwan (as early Pliocene in Huang, 1965). Although the whereabouts of the specimens is unknown, it is one of the earliest accounts reporting fossil shark teeth in Taiwan. Uyeno (1978) listed nine elasmobranch taxa from the Pleistocene Chochen–Tsailiao area with images of the specimens but without descriptions, and these materials are reviewed in this present study.

Perhaps the most complete description on a single fossil shark assemblage in Taiwan is the one by Tao & Hu (2008) from the late Miocene Tangenshan Sandstone Formation in Chiahsien County, Kaohsiung. They described five taxa common in late Miocene marine deposits (*Otodus megalodon*, *Odontaspis* sp., †*Isurus hastalis*, †*Hemipristis serra*, and *Carcharhinus* sp.) as well as a new extinct species of *Hemipristis*, *H. liui* Tao & Hu, 2008. The specimen of *H. liui* is an upper tooth and is characterized by asymmetric serrations on the distal and mesial cutting edges. The occurrences of *Otodus megalodon* are sparsely recorded from Taiwan (Hu & Tao, 1993; Tao & Hu, 2008) and are mostly present in private collections, which are potentially one of the directions for future research efforts (Haug et al., 2020; Lin et al., 2021).

The materials were mainly based on surface collecting that spanned over more than three decades, and we note that our bulk sediment samples (830 kg, see methods) only yielded three specimens (ASIZF0100548, ASIZF0100590, and ASIZF0100591). Surface collecting likely results in sampling bias towards larger specimens underrepresenting smaller specimens (Welton & Farish, 1993; Perez, 2022). Nevertheless, 697 elasmobranch teeth from the Liuchungchi Formation in Niubu described in this study reveal the presence of at least 20 elasmobranch taxa (Table 1). The excellent overall preservation allowed species-level taxonomic identification for most of the specimens, which in turn, permitted the elucidation of the diverse elasmobranch community. In fact, the assemblage represents the most diverse elasmobranch paleofauna from Taiwan reported to date. Moreover, the species richness and diversity indices suggest that our assemblage is highly diverse even with respect to other contemporaneous assemblages from temperate and tropical West Pacific (Table 2, Supplemental Table S1). Together, the high diversity captured in our study is significant in the spatio-temporal context.

The abundant and large teeth of *Carcharodon carcharias* are remarkable. *Carcharodon carcharias* is distributed along southern, eastern, and northeastern Taiwan today, but not on the west coast where the fossils are found (Teng, 1958; Shen, 1993; Ebert et al., 2013; Shao, 2022). According to the Fisheries Agency, Council of Agriculture, Taiwan (Taiwan Fisheries Agency,

2021), a total of 39 individuals of *C. carcharias* were caught between 2012 and 2021, with the majority of landings being in northeastern Taiwan. However, at our fossil sites, teeth of *C. carcharias* are the second most abundant remains (n = 55) among specimens identified to the species level in this study, next to teeth of *Carcharhinus leucas* (n = 71, Table 1). Of the 55 isolated teeth that are interpreted to have most certainly come from 55 different individuals, 44 of them are well-preserved, offering their tooth position identifications and crown heights (CH). Based on the linear regression equation between the CH and fish total length (TL) for each tooth position in extant *Carcharodon carcharias* presented by Shimada (2003), the CH of each of the 44 teeth was used to estimate the TL of each fossil individual (Supplemental Table S2). Our specimens are normally distributed (Shapiro–Wilk test = 0.853, p = 0.08) and range in fish length from 1.9 to 5.6 m, with a mean of 3.5 m (Supplemental Table S2), suggesting presence of many mature, large individuals (Ebert et al., 2013).

One of the most noteworthy occurrences reported in this study is that of the extinct species †*Hemipristis serra*. The species is known worldwide, but most of the documented occurrences are from the Miocene and Pliocene deposits (e.g., Yabumoto & Uyeno, 1994; Sánchez-Villagra et al., 2000; Marsili et al., 2007; Portell et al., 2008; Visaggi & Godfrey, 2010; Carrillo-Briceño et al., 2015; Kocsis et al., 2019). The fossil record indicates that the fossil species preferred warm neritic environments (Cappetta, 2004, 2012). Although most previous studies suggest its last appearance at the end of the Pliocene, new evidence indicates that †*H. serra* persisted into the Pleistocene in North America (Ebersole, Ebersole & Cicimurri, 2017; Boessenecker, Boessenecker & Geisler, 2018; Perez, 2022). Teeth of *Hemipristis* that may belong to *H. serra* have been reported from Pleistocene and 'Plio-Pleistocene' deposits in Sulawesi and Java, Indonesia (*H.* cf. *serra* by Hooijer, 1954, 1958; simply "*Hemipristis*" by Yudha et al., 2018). Previous records of †*H. serra* from Taiwan were reported by Uyeno (1978) from an uncertain stratigraphic horizon along Tsailiao River, and that by Tao & Hu (2008) from the Miocene Kueichulin Formation in southern Taiwan. The †*H. serra* specimens described here are the first confirmed Pleistocene record in Taiwan, and along with the putative Indonesian records (Hooiher, 1954, 1958; Yudha et al., 2018), the geologically youngest records of the extinct species in the Northwest Pacific, meaning that the North American occurrences were not isolated. It should be noted that the preservation of all specimens of †*H. serra* described in this study are excellent (Fig. 6), and thus they most certainly do not represent fossils derived from an older horizon.

The assemblage is dominated by two genera, *Carcharhinus* (Carcharhinidae, n = 462) and *Carcharodon* (Lamnidae, n = 52), comprising more than 77.5% of the total specimen count and about half of the taxa identified (11 out of 20). From a paleoecological perspective, the composition is roughly similar to the one found in modern western Taiwan (Ebert et al., 2013; Shao, 2022). For example, the most abundant species of *Carcharhinus* in this study, *C. leucas*, presently lives close to the coastal area of tropical and subtropical riverine and lacustrine

(Compagno, 1984). The second-most abundant species in this study, *Carcharodon carcharias*, inhabits inshore shallow water to open ocean and, as a top predator, feeds on larger marine mammals and fishes (Ebert et al., 2013; Compagno, 2002). While pelagic sharks *Carcharhinus plumbeus* and *C. longimanus* are also frequently represented in the Pleistocene assemblage, the occurrences of *C. altimus*, *Aetobatus* sp., and *Myliobatis* sp. may suggest the possible presence of deeper sandy, flat bottoms (Compagno, 1984). The abundant associated marine vertebrate fossils, including teleost bones (Tao, 1993), otoliths (Lin et al., 2018), and whale bones (Xue, 2004), indicate a rich, thriving marine ecosystem in the area. The sedimentary environment of the Liuchungchi Formation further points to shoreface to inner offshore setting, with several transgressive and regressive cycles (Chen, 2016). Taken together, the coastal areas in southwest Taiwan during the early Pleistocene can be interpreted as an inshore to offshore shallow-water environment with sandy bottoms.

Conclusions

Fossil elasmobranch fauna from the tropical-subtropical West Pacific is poorly known compared to its modern analog, impeding our understanding of the formation of this current marine biodiversity hotspot. Using elasmobranch fossils from an early Pleistocene locality in southern Taiwan, we report a previously unreported but highly diverse shark and ray fauna from the region. The taxonomic composition of the assemblage reveals a nearshore shallow-water paleoenvironment and supports the sedimentary interpretation. In addition, the presence of †*Hemipristis serra* and large specimens of *Carcharodon carcharias* highlight the potential for similar studies from other strata and localities, which would allow a more comprehensive picture of the evolutionary history and biogeographic distribution of the species. The present study can be regarded as the most extensive on elasmobranch fossils from Taiwan.

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814 Dasyatinae – Genera: *Amphotistius*, *Dasyatis*, *Himantura*, *Pastinachus*, *Pteroplatytrygon*,
815 *Taeniura*, *Urogymnus* and *Urolophoides* (incl. supraspecific taxa of uncertain status and
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817 *Gymnura*, Superfamily Plesiobatoidea – Family Hexatrygonidae – Genus: *Hexatrygon*.
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822 Suborder Myliobatoidei – Superfamily Dasyatoidea – Family Dasyatidae – Subfamily
823 Dasyatinae – Genera: *Taeniura*, *Urogymnus*, *Urolophoides* – Subfamily
824 Potamotrygoninae – Genera: *Disceus*, *Plesiotrygon*, and *Potamotrygon* (incl. supraspecific
825 taxa of uncertain status and validity), Family Urolophidae – Genera: *Trygonoptem*,
826 *Urolophus* and *Urotrygon* – Superfamily Myliobatidea – Family: Gymnuridae – Genus:
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833 Dasyatinae – Genus: *Urobatis*, Subfamily Potamotrygoninae – Genus: *Paratrygon*,
834 Superfamily Plesiobatoidea – Family Plesiobatidae – Genus: *Plesiobatis*, Superfamily
835 Myliobatoidea – Family Myliobatidae - Subfamily Myliobatinae – Genera : *Aetobatus*,
836 *Aetomylaeus*, *Myliobatis* and *Pteromylaeus*, Subfamily Rhinopterinae – Genus: *Rhinoptera*
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Supplemental file

Supplemental Figure S1. Aerial photographs of the Niubu locality showing the closest sampling site to the weir (red square, Site 1 in Fig. 2B). A, general view; B, C, closer views of the sampling site. Note the sandstones in the river bed transported from the upstream older strata.

Supplemental Table S1. List of taxa and their abundance extracted from the Pleistocene elasmobranch assemblages from the West Pacific. Only teeth are included, i.e., spines, vertebrae and other remains are not considered. Nomenclatures remain as those reported in the literature.

Supplemental Table S2. Measurement, tooth position, and estimated fish total length of 44 better-preserved specimens of *Carcharodon carcharias* specimens form the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. * = estimated crown height values for specimens with slightly worn apex.

Figure captions

Figure 1. Summary of the sampling sites. A, overview of geological map of Taiwan (modified after Chen, 2016). B, geological map of Nuibu area, Chiayi (map extracted from National Geological Data Warehouse, Central Geological Survey, MOEA). Yellow stars = sampling sites (see Fig. 2B for details). C, stratigraphic correlation of the Western Foothills (modified after Chen, 2016). The examined Liuchungchi Formation is indicated in red.

Figure 2. A, stratigraphic column (modified after Huang, 2010). B, details of the sampling sites. GPS coordinates: Site 1 = 23°26'23.4"N, 120°35'35.5"E; Site 2 = 23°26'22.6"N, 120°35'32.7"E; Site 3 = 23°26'23.5"N, 120°35'29.8"E.

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Figure 5. Teeth of *Isurus oxyrinchus* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, C, ASIZF0100317; D, E, F, ASIZF0100318; G, H, ASIZF0100321; I, J, CMM F0242. A, D, G, I = lingual views; B, E, H, J = labial views; C, F = lateral views. Scale bar = 1 cm.

Figure 6. Teeth of †*Hemipristis serra* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, CMM F0232; C, D, ASIZF0100460; E, F, ASIZF0100461; G, H, ASIZF0100462. A, C, E, G = lingual views; B, D, F, H = labial views. Scale bars = 1 cm.

Figure 7. Teeth of *Carcharhinus altimus* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, ASIZF0100357; C, D, ASIZF0100359; E, F, CMM F0363; G, H, CMM F0293; I, J, CMM F0322; K, L, ASIZF 0100365. A, C, E, G, I, K = lingual views; B, D, F, H, J, L = labial views. Scale bar = 1 cm.

Figure 8. Teeth of *Carcharhinus amboinensis* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, CMM F0209; C, D, ASIZF0100368; E, F, ASIZF0100366; G, H, ASIZF0100369. A, C, E, G = lingual views; B, D, F, H = labial views. Scale bar = 1 cm.

Figure 9. Teeth of *Carcharhinus leucas* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, ASIZF0100398; C, D, ASIZF0100397; E, F, ASIZF0100394; G, H, ASIZF0100411; I, J, ASIZF0100396; K, L, ASIZF 0100395; M, N, ASIZF0100400; O, P, ASIZF0100402; Q, R, ASIZF0100390. A, C, E, G, I, K, M, O, Q = lingual views; B, D, F, H, J, L, N, P, R = labial views. Scale bar = 1 cm.

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Figure 11. Teeth of *Carcharhinus longimanus* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, ASIZF 0100371; C, D, ASIZF0100376; E, F, ASIZF0100370; G, H, ASIZF0100377; I, J, ASIZF0100375; K, L, ASIZF0100378; M, N, ASIZF0100374; O, P, ASIZF0100373; Q, R, ASIZF0100392; S, T, ASIZF0100391. A, C, E, G, I, K, M, O, Q, S = lingual views; B, D, F, H, J, L, N, P, R, T = labial views. Scale bar = 1 cm.

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Figure 13. Teeth of *Carcharhinus plumbeus* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, ASIZF0100412; C, D, ASIZF0100406; E, F, ASIZF0100405; G, H, ASIZF0100410; I, J, ASIZF0100409; K, L, ASIZF0100408; M, N, ASIZF0100407. A, C, E, G, I, K, M = lingual views; B, D, F, H, J, L, N = labial views. Scale bar = 1 cm.

Figure 14. Teeth of *Carcharhinus sorrah* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, CMM F0129; C, D, CMM F0119; E, F, ASIZF0100418; G, H, CMM F0126; I, J, CMM F0135; K, L, CMM F0122; M, N, CMM F0140. A, C, E, G, I, K, M = lingual views; B, D, F, H, J, L, N = labial views. Scale bar = 1 cm.

Figure 15. Teeth of *Carcharhinus tjutjot* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, ASIZF0100415; C, D, ASIZF0100414; E, F, CMM F0116; G, H, ASIZF0100413; I, J, ASIZF0100417; K, L, ASIZF0100416; M, N, CMM F0323; O, P, CMM F0324. A, C, E, G, I, K, M, O = lingual views; B, D, F, H, J, L, N, P = labial views. Scale bar = 1 cm.

Figure 16. Teeth of *Rhizoprionodon acutus* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, ASIZF0100463; C, D, CMM F0120; E, F, CMM F0121; G, H, CMM F0131; I, J, ASIZF0100464. A, C, E, G, I = lingual views; B, D, F, H, J = labial views. Scale bar = 1 cm.

Figure 17. Teeth of *Galeocerdo cuvier* and *Sphyrna lewini* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A–H, *Galeocerdo cuvier*; A, B, CMM F0245; C, D, E, CMM F0215; F, G, H, ASIZF0100459. I–L, *Sphyrna lewini*; I, J, CMM F0235; K, L, CMM F0312. A, C, F, I, K = lingual views; B, D, G, J, L = labial views; E, H = details of secondary serrations. Scale bars = 1 cm.

Figure 18. Teeth of Dasyatidae indet. from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, C, D, ASIZF0100590; E, F, G, H, ASIZF0100591. A, E = labial views; B, F = basal views; C, G = occlusal views; D, H = lateral views. Scale bar = 5 mm.

Figure 19. Teeth of *Aetobatus* sp. from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, CMM F2854; C, D, CMM F2850; E, F, CMM F0408; G, H, ASIZF0100549. A, C, E, G = occlusal views; B, D, F, H = basal views. Scale bar = 1 cm.

Figure 20. Teeth of *Myliobatis* sp. from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan. A, B, C, D, ASIZF0100582; E, F, G, H, ASIZF0100587; I, J, K, L, ASIZF0100586; M, N, CMM F0395; O, P, CMM F2855; Q, R, CMM F0393; S, T, CMM

F0398. A, E, I, M, O, Q, S = occlusal views; B, F, J, N, P, R, T = basal views; C, G, K = lingual views; D, H, L = labial views. Scale bars = 1 cm.

Table captions

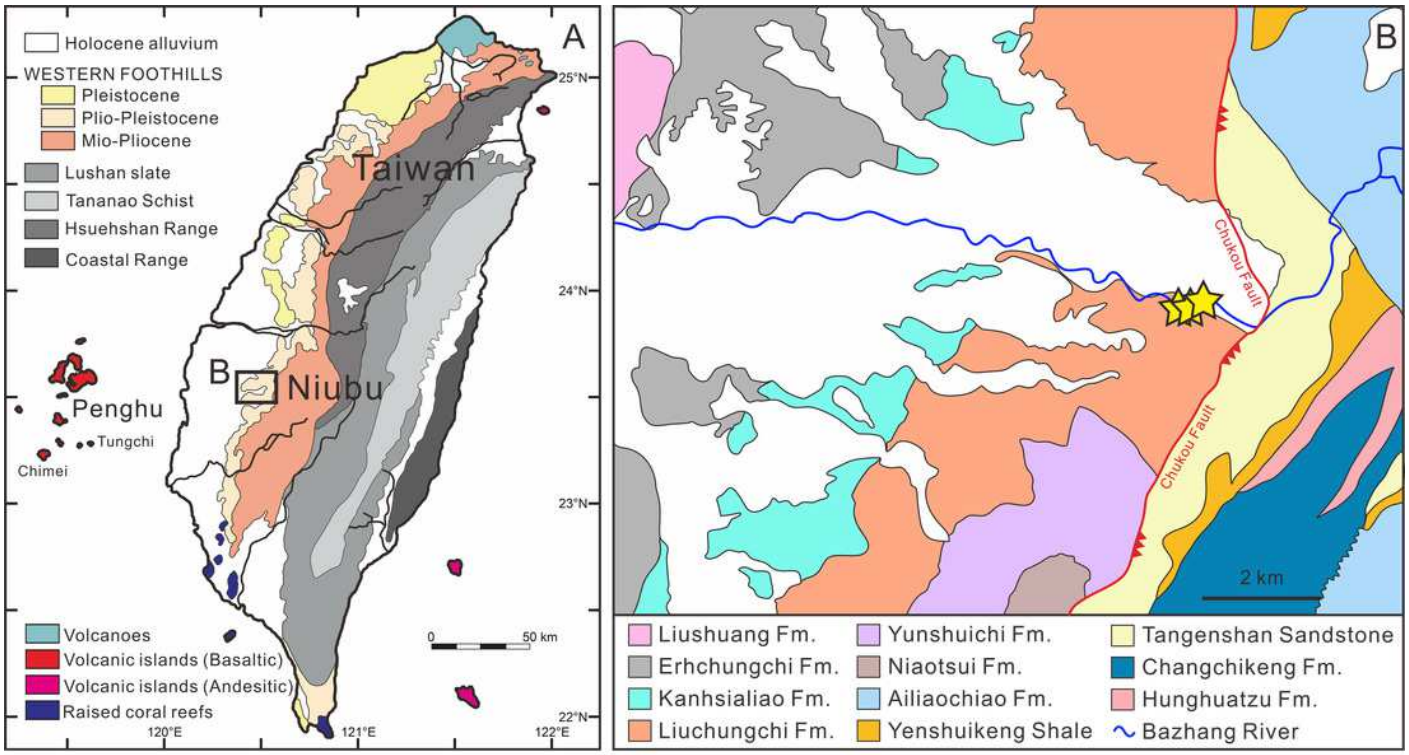
Table 1. Elasmobranchs from the early Pleistocene Liuchungchi Formation of Niubu, southern Taiwan.

Table 2. Various diversity indices from the Pleistocene West Pacific elasmobranch assemblages showing high diversity of the present material. See Supplemental Table S1 for details of the data.

Figure 1

Summary of the sampling sites.

A, overview of geological map of Taiwan (modified after Chen, 2016). B, geological map of Nuibu area, Chiayi (map extracted from National Geological Data Warehouse, Central Geological Survey, MOEA). Yellow stars = sampling sites (see Fig. 2B for details). C, stratigraphic correlation of the Western Foothills (modified after Chen, 2016). The examined Liuchungchi Formation is indicated in yellow.



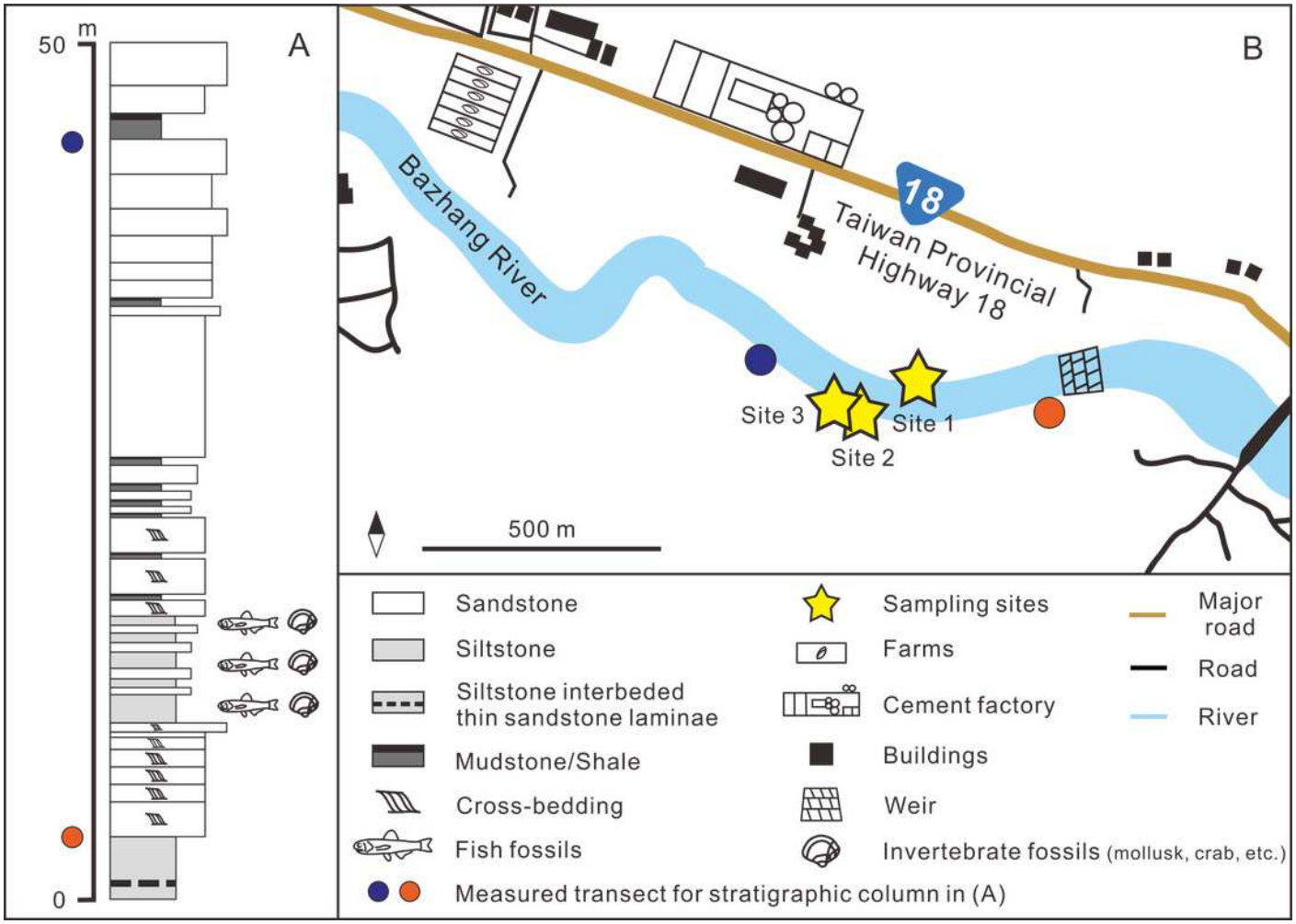
C

AGE			N — Western Foothills — S			
			Keelung / Taoyuan	Hsinchu / Taichung	Nantou	**Chiayi**
Quaternary	Holocene	0.0117 Ma				Tainan Fm.
	Pleistocene "Tarantian"		Chungli Fm.	Lateritic Conglomerates		? Liushuang Fm.
	Chibanian	0.129 Ma – 0.774 Ma			Toukoshan Fm.	Erhchungchi Fm.
	Calabrian	1.80 Ma	Linkou Fm.	Toukoshan Fm.		Kanhsialiao Fm.
			Cholan Fm.	Cholan Fm.	Cholan Fm.	Liuchungchi Fm.
	Gelasian	2.58 Ma				Yunshuichi Fm.

Figure 2

A, stratigraphic column (modified after Huang, 2010). B, details of the sampling sites.

GPS coordinates: Site 1 = 23°26'23.4''N, 120°35'35.5''E; Site 2 = 23°26'22.6''N, 120°35'32.7''E; Site 3 = 23°26'23.5''N, 120°35'29.8''E.



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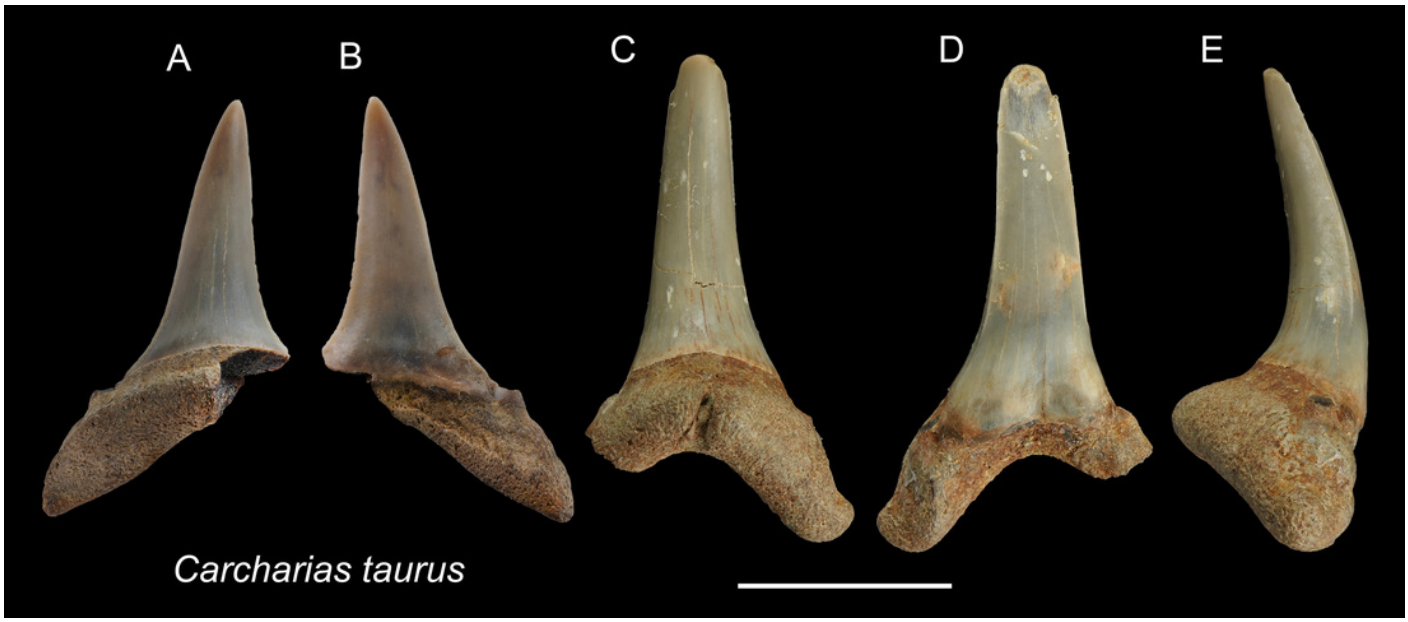
A, B, ASIZF0100320; C, D, E, CMM F0204. A, C = lingual views; B, D = labial views; E = lateral view. Scale bar = 1 cm.



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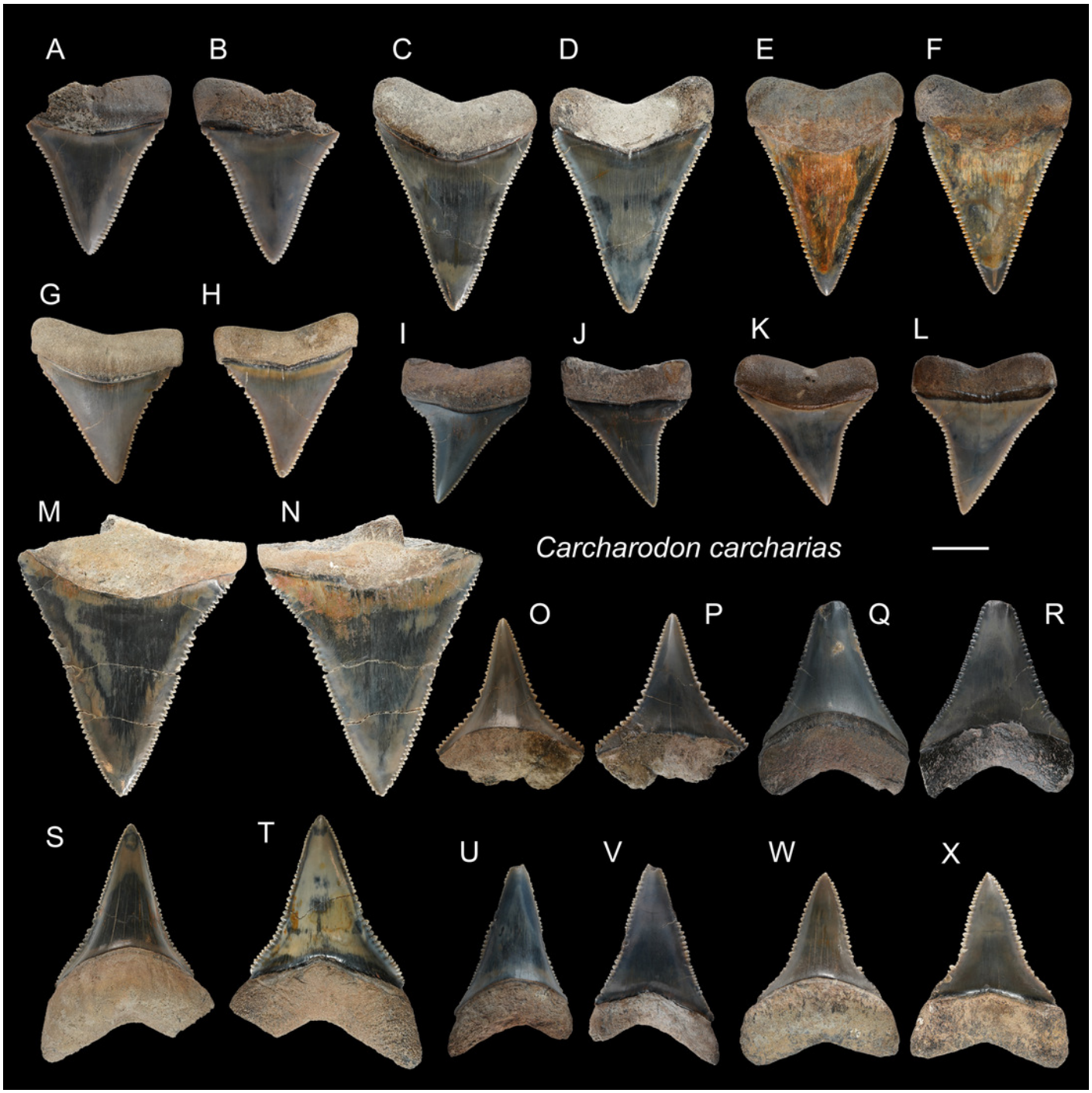
A
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Carcharodon carcharias
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U
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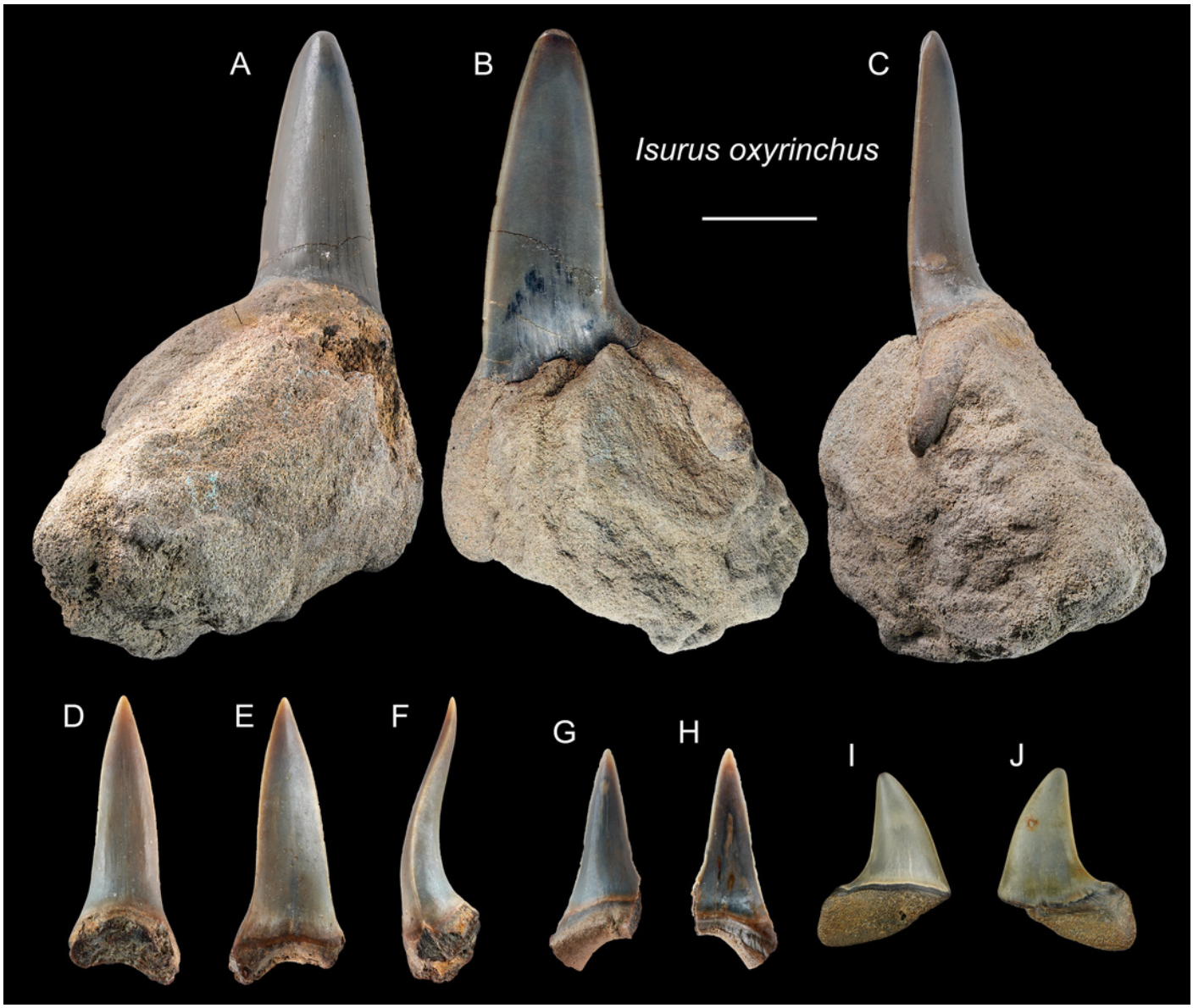
W
X

Figure 5

Teeth of *Isurus oxyrinchus* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan.

A, B, C, ASIZF0100317; D, E, F, ASIZF0100318; G, H, ASIZF0100321; I, J, CMM F0242. A, D, G, I = lingual views; B, E, H, J = labial views; C, F = lateral views. Scale bar = 1 cm.



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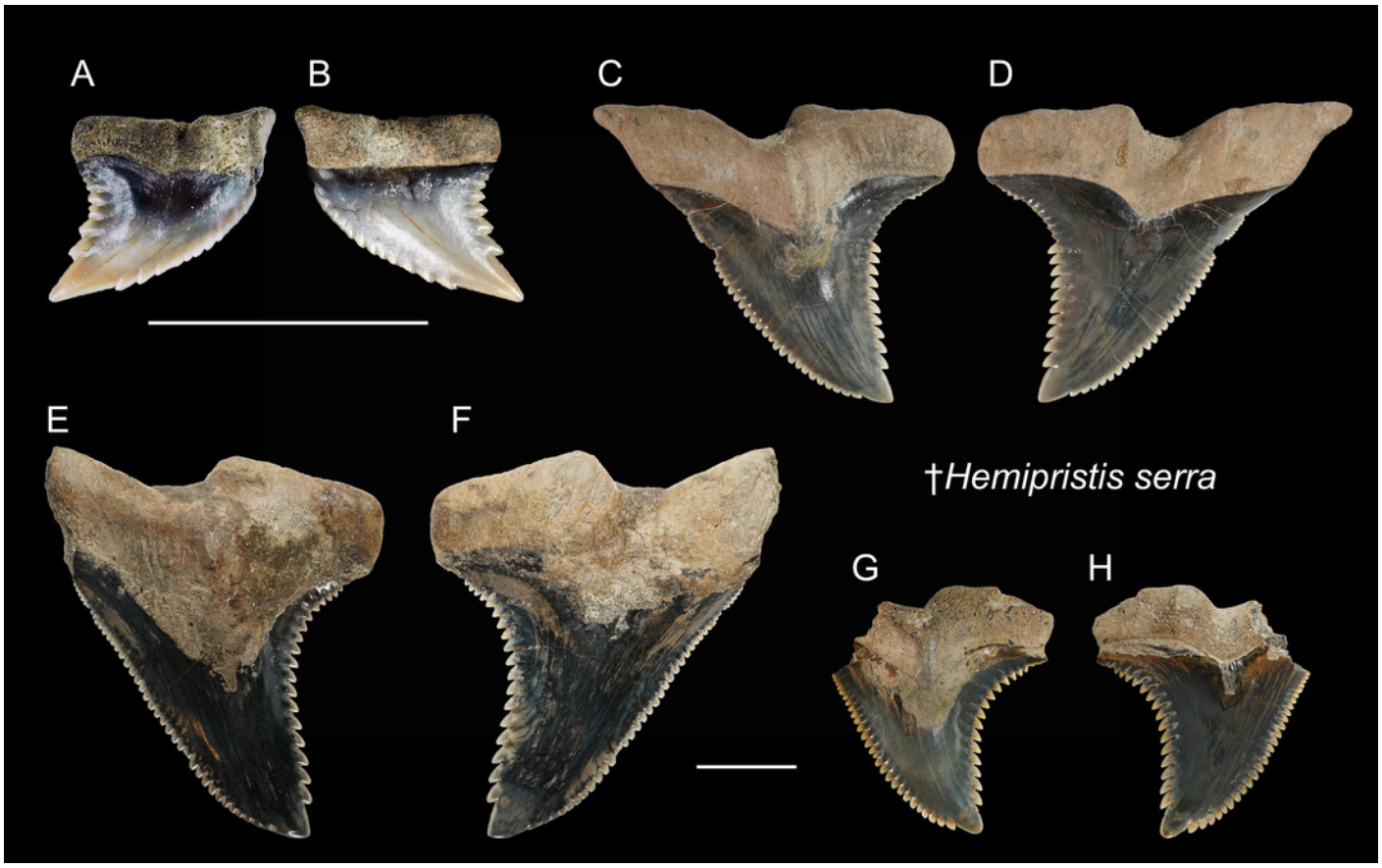

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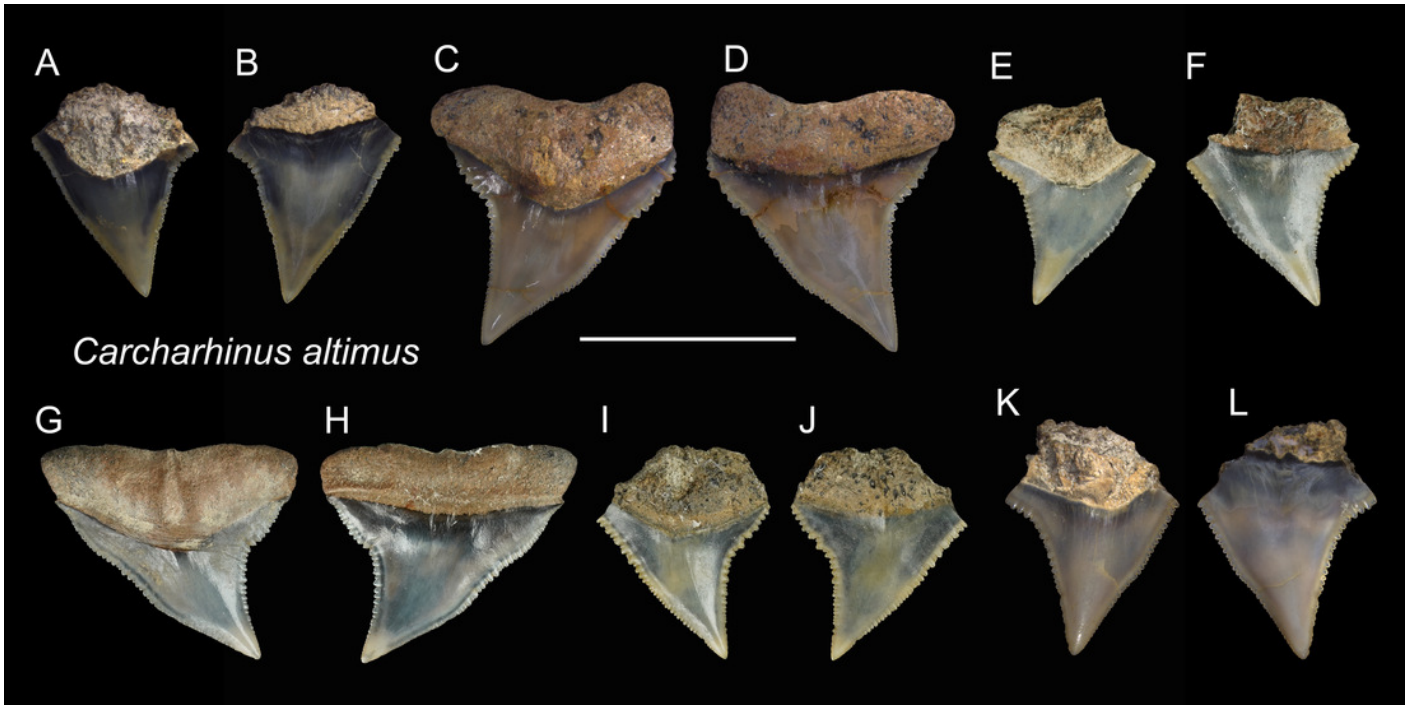

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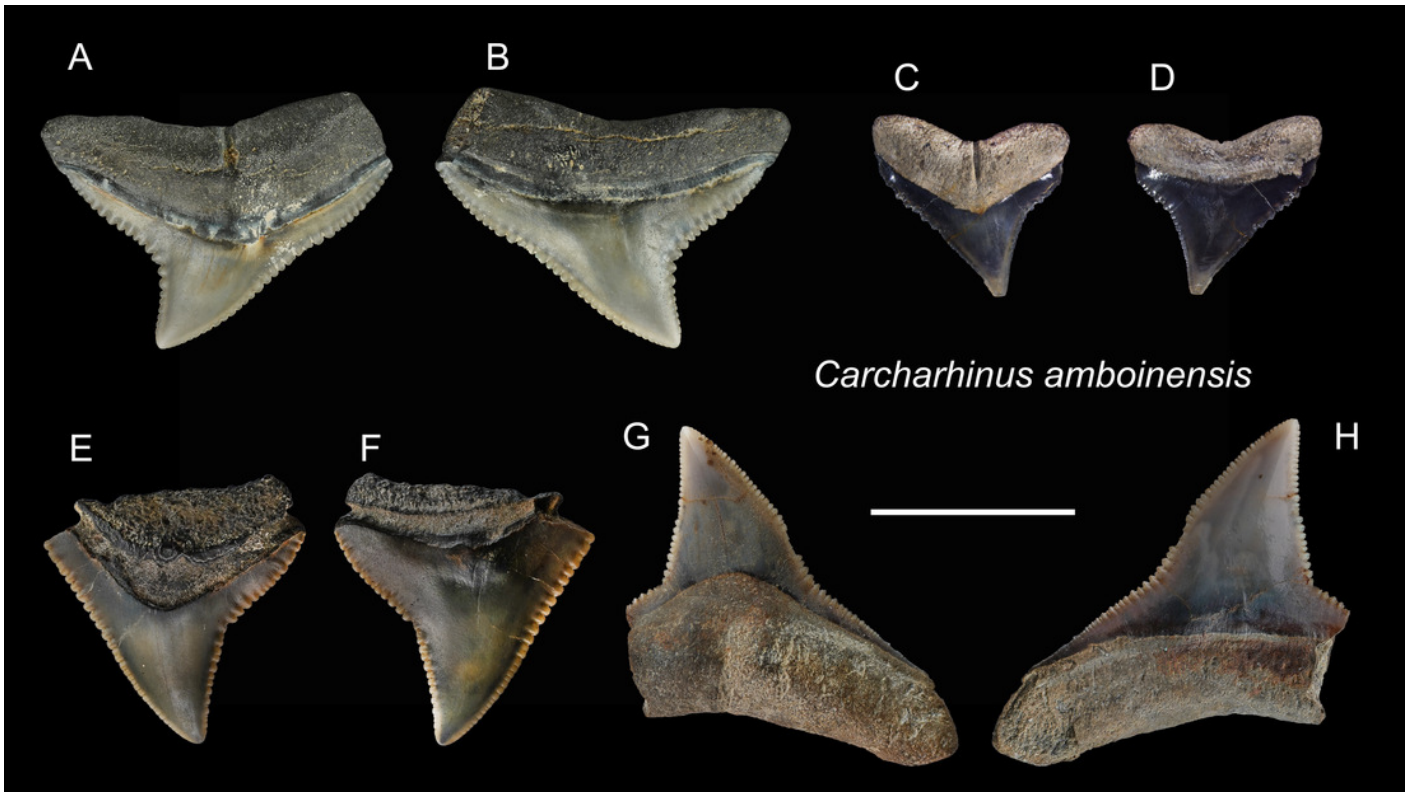

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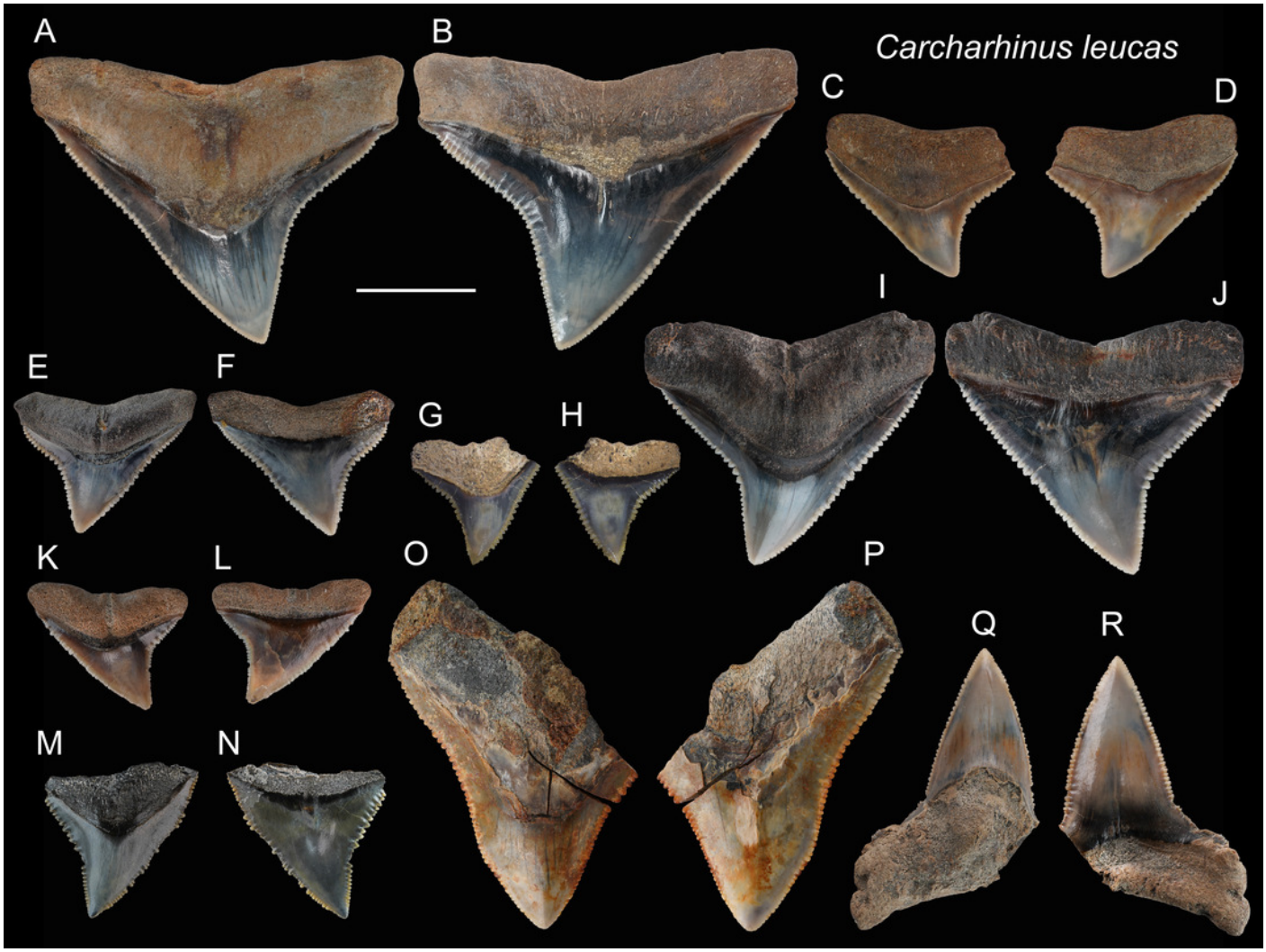
A, B, ASIZF0100398; C, D, ASIZF0100397; E, F, ASIZF0100394; G, H, ASIZF0100411; I, J, ASIZF0100396; K, L, ASIZF 0100395; M, N, ASIZF0100400; O, P, ASIZF0100402; Q, R, ASIZF0100390. A, C, E, G, I, K, M, O, Q = lingual views; B, D, F, H, J, L, N, P, R = labial views. Scale bar = 1 cm.



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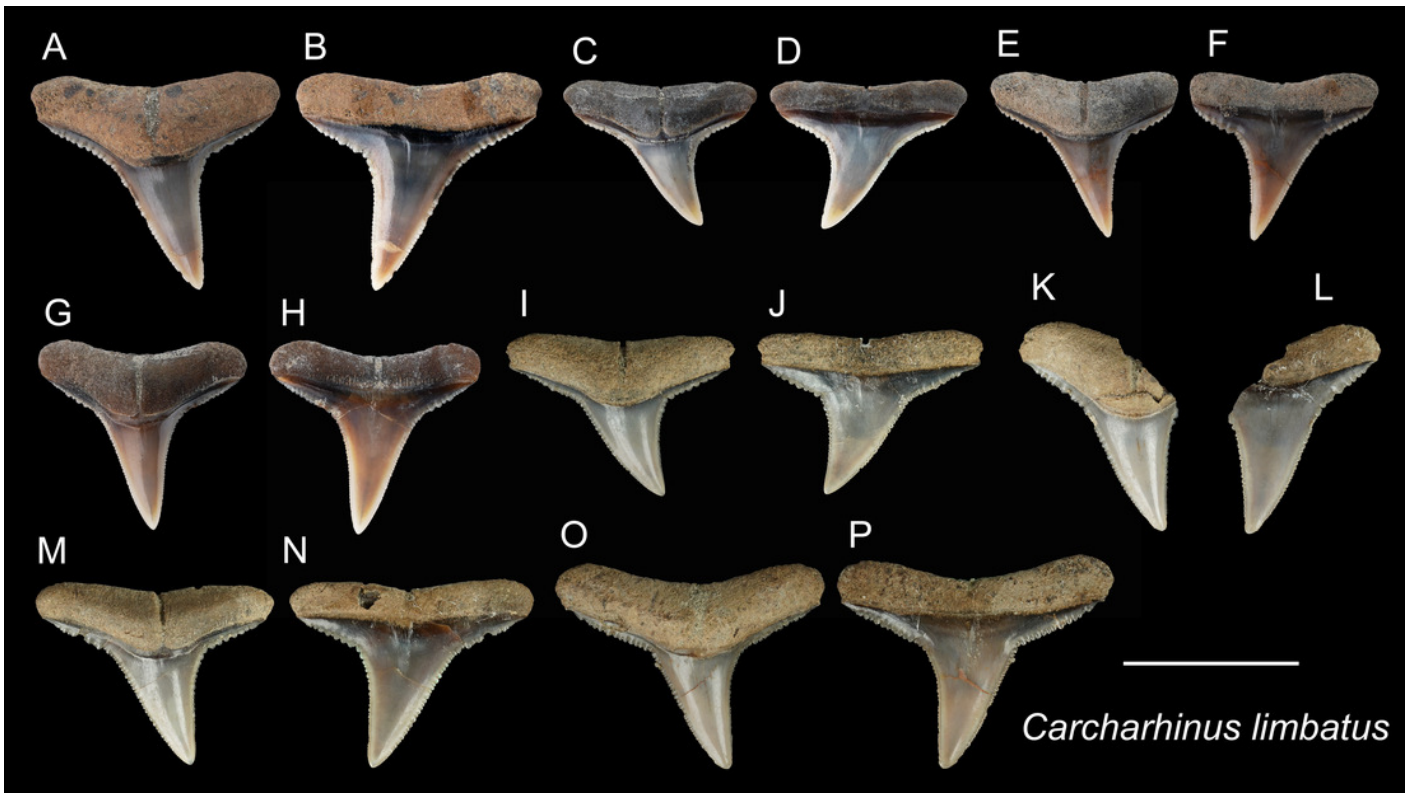
A, B, ASIZF0100470; C, D, ASIZF0100476; E, F, ASIZF0100469; G, H, ASIZF0100468; I, J, CMM F0236; K, L, CMM F0111; M, N, CMM F0237; O, P, CMM F0238. A, C, E, G, I, K, M, O = lingual views; B, D, F, H, J, L, N, P = labial views. Scale bar = 1 cm.



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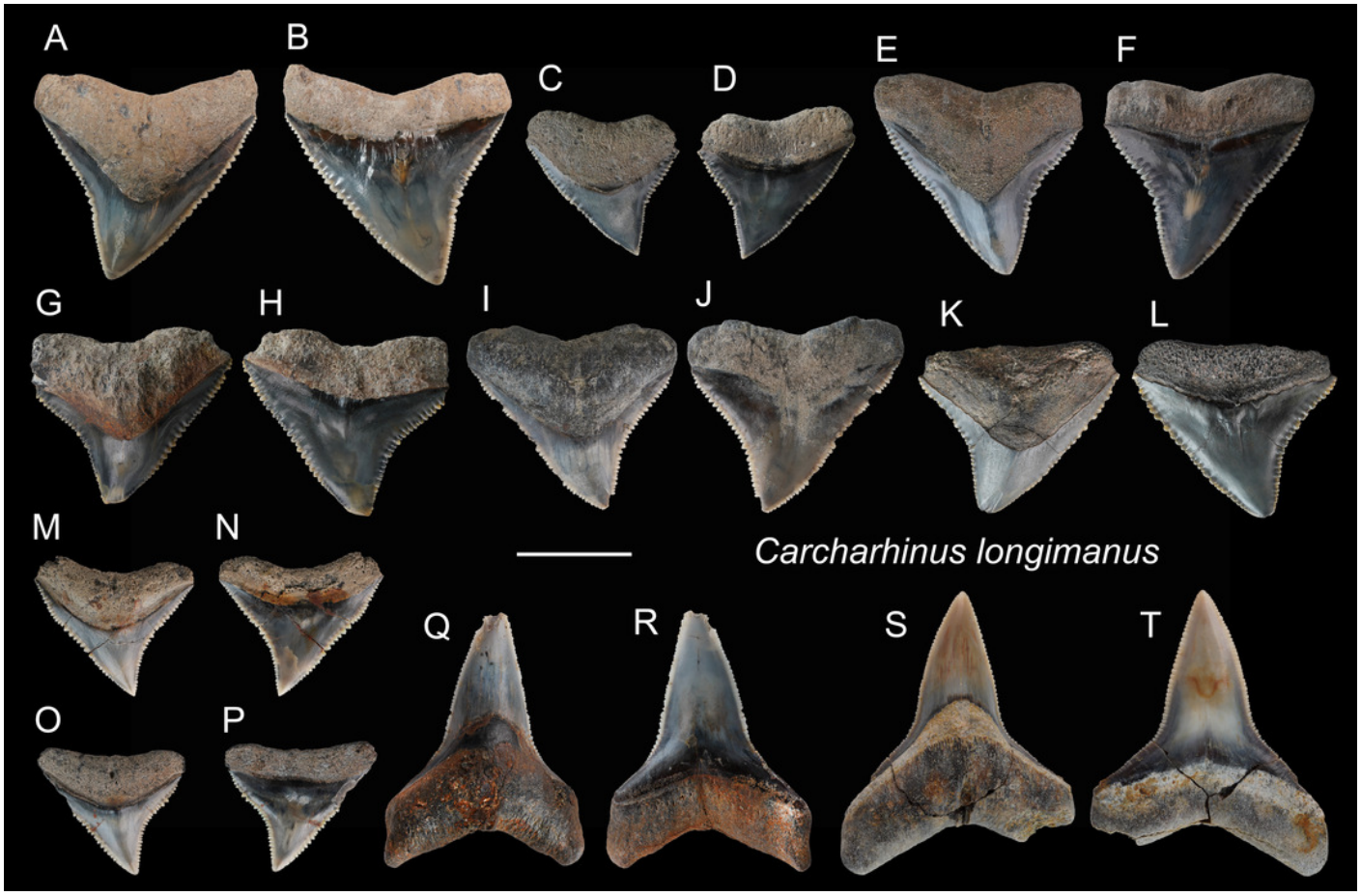
Teeth of *Carcharhinus longimanus* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan.

A, B, ASIZF 0100371; C, D, ASIZF0100376; E, F, ASIZF0100370; G, H, ASIZF0100377; I, J, ASIZF0100375; K, L, ASIZF0100378; M, N, ASIZF0100374; O, P, ASIZF0100373; Q, R, ASIZF0100392; S, T, ASIZF0100391. A, C, E, G, I, K, M, O, Q, S = lingual views; B, D, F, H, J, L, N, P, R, T = labial views. Scale bar = 1 cm.



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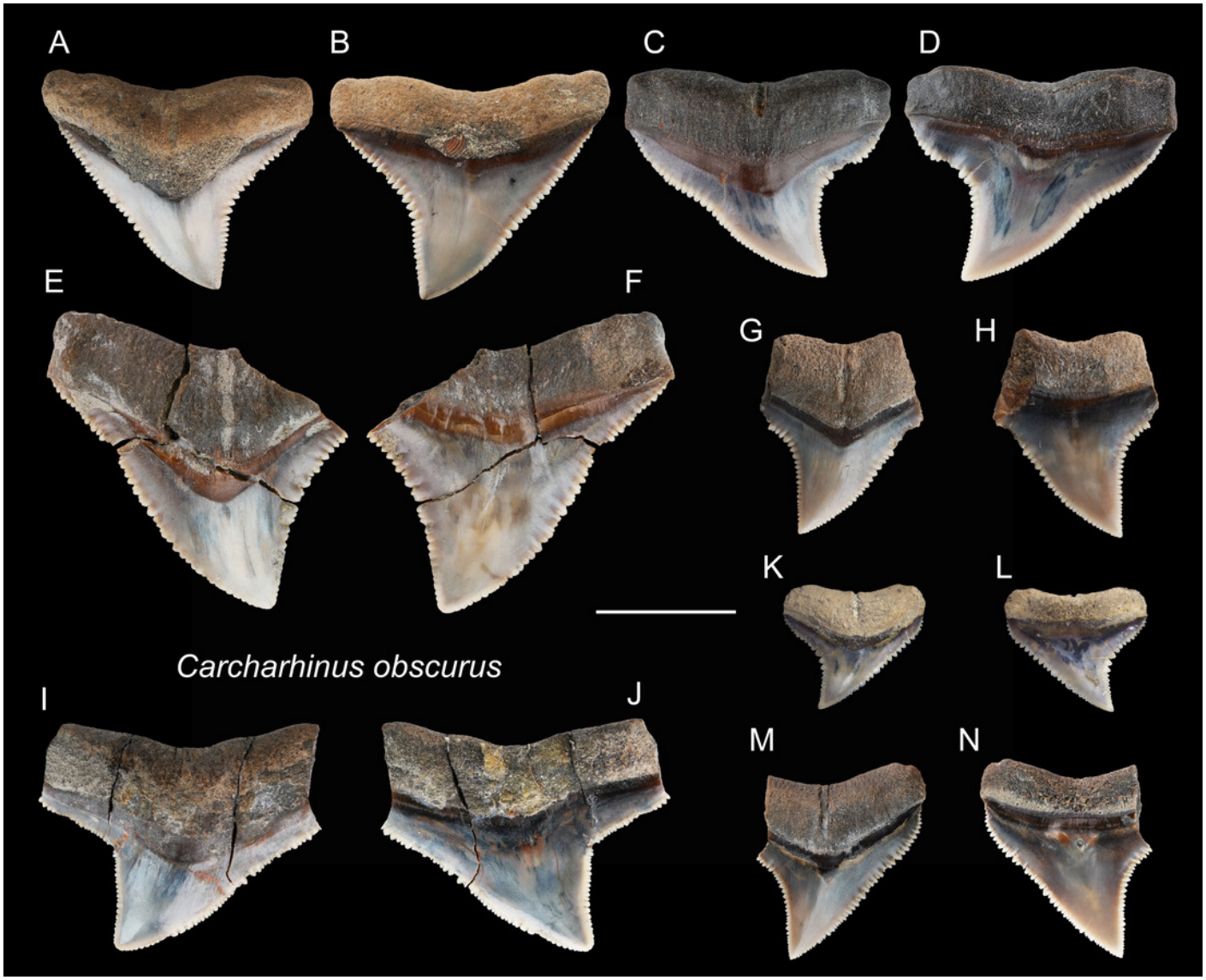
A, B, ASIZF0100372; C, D, ASIZF0100384; E, F, ASIZF0100385; G, H, ASIZF0100386; I, J, ASIZF0100388; K, L, ASIZF0100387; M. N, ASIZF0100383. A, C, E, G, I, K, M = lingual views; B, D, F, H, J, L, N = labial views. Scale bar = 1 cm.



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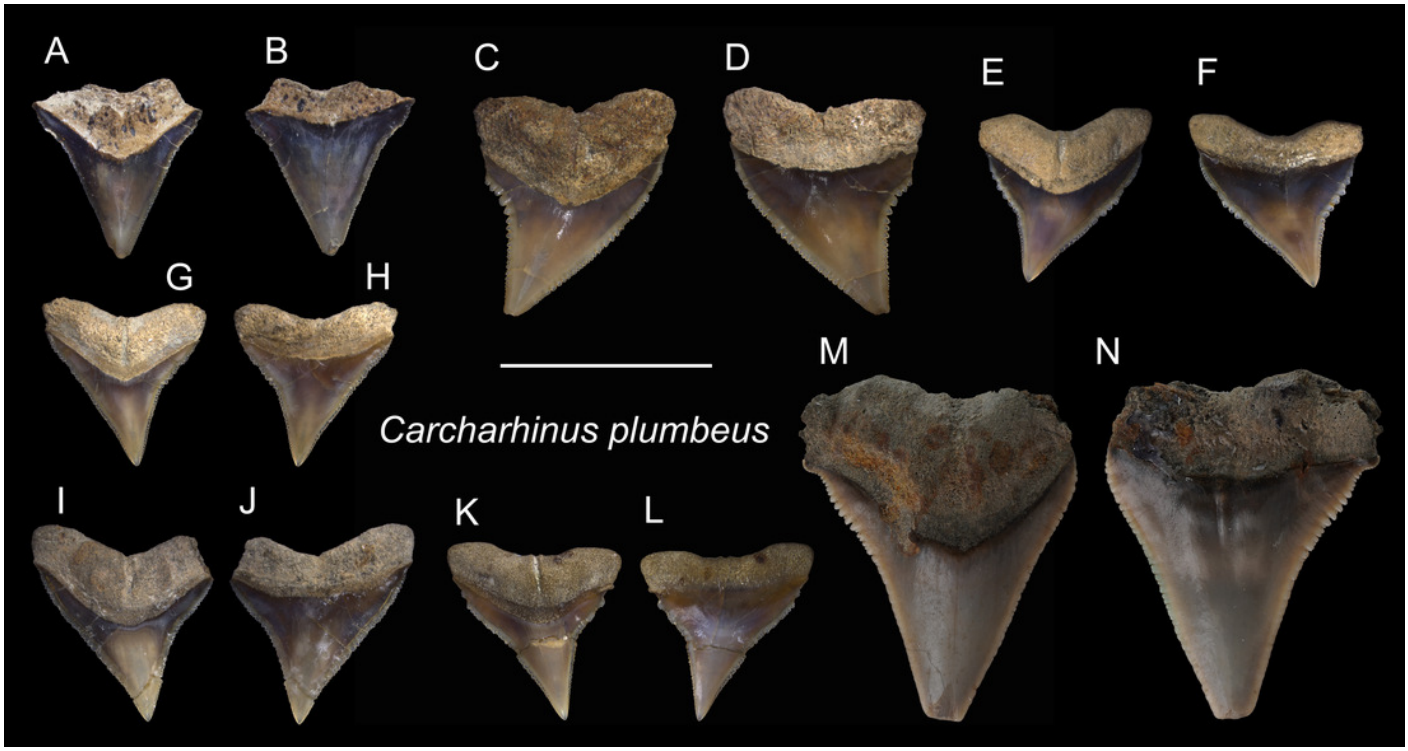

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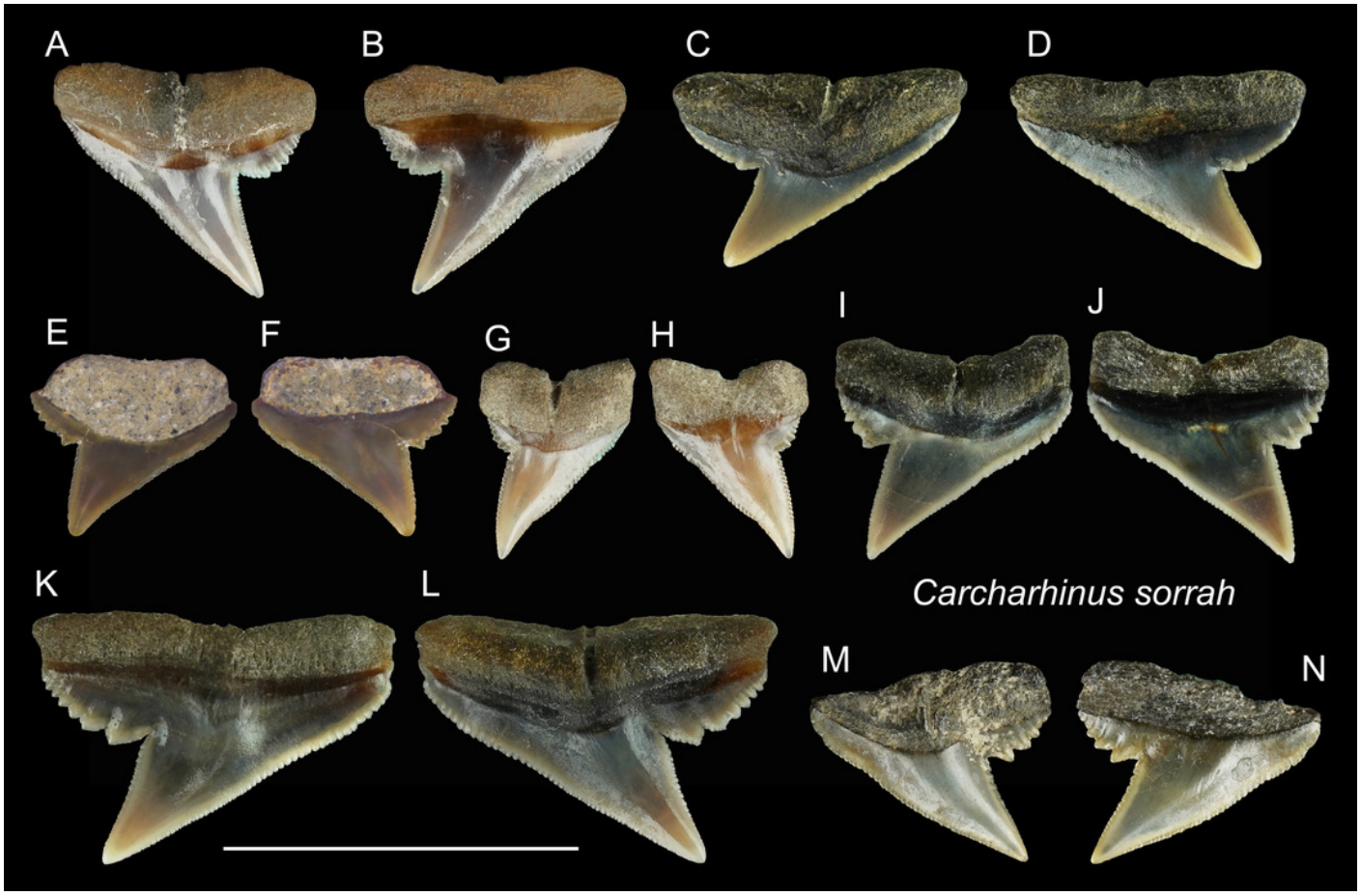

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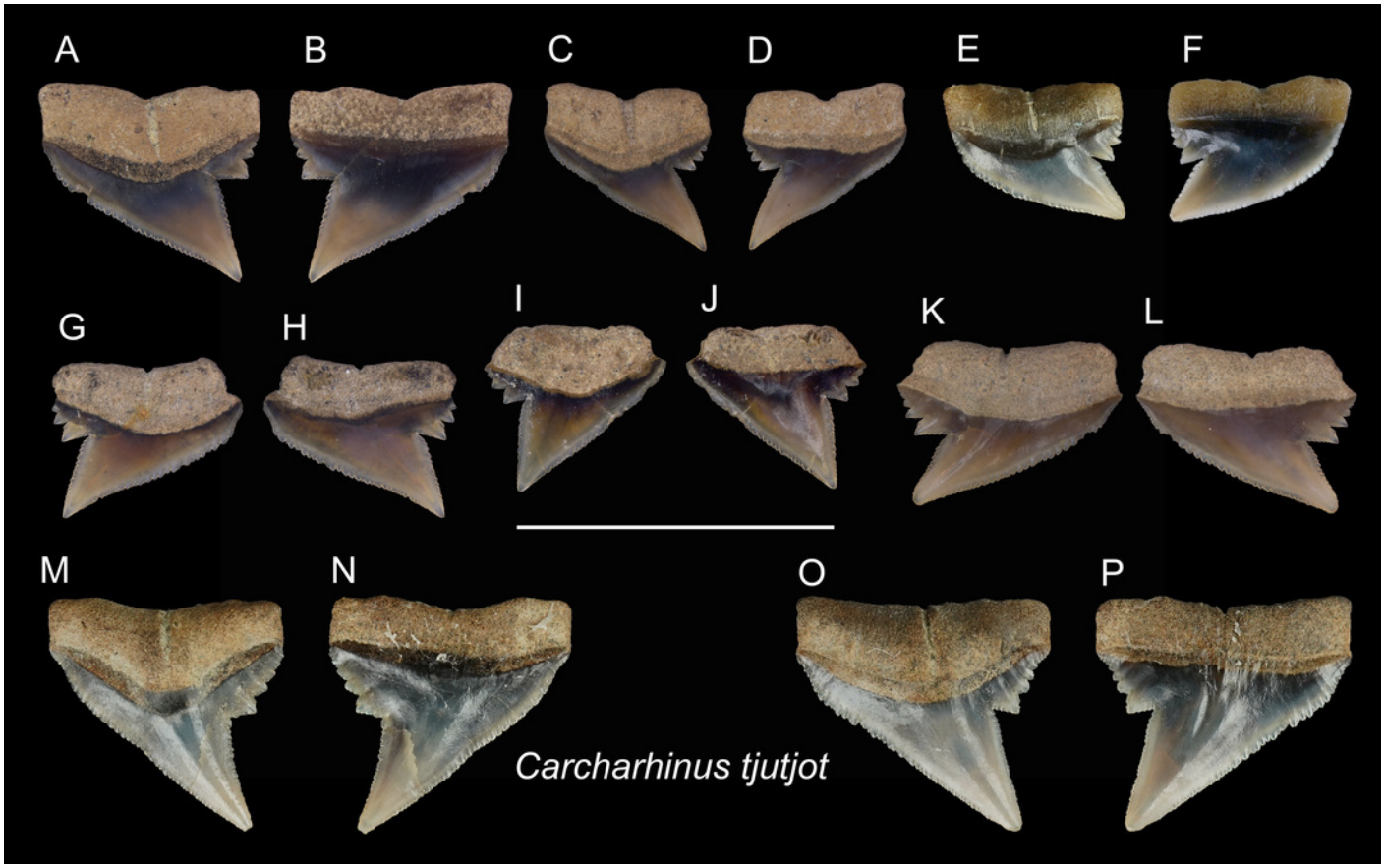

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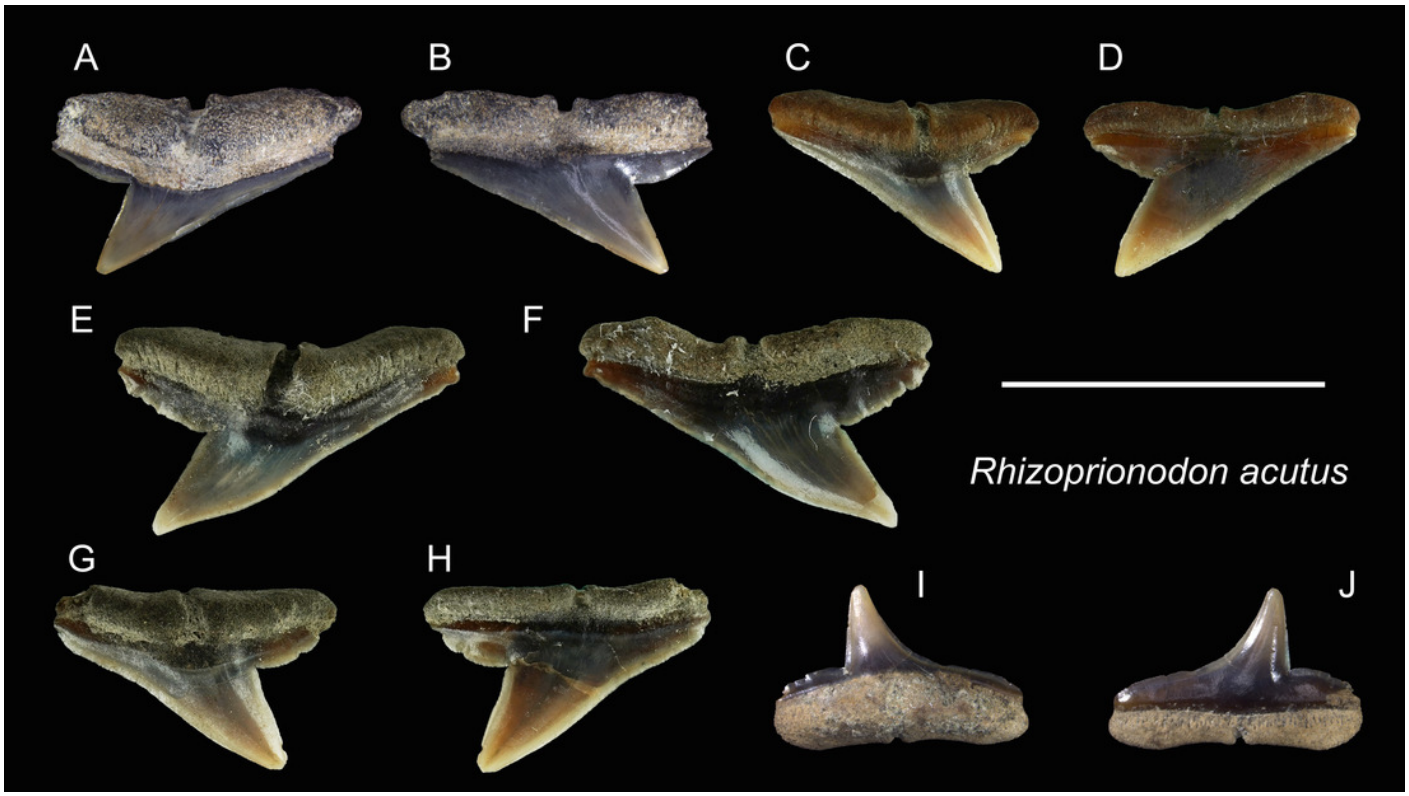

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Teeth of *Galeocerdo cuvier* and *Sphyrna lewini* from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan.

A–H, *Galeocerdo cuvier*; A, B, CMM F0245; C, D, E, CMM F0215; F, G, H, ASIZF0100459. I–L, *Sphyrna lewini*; I, J, CMM F0235; K, L, CMM F0312. A, C, F, I, K = lingual views; B, D, G, J, L = labial views; E, H = details of secondary serrations. Scale bars = 1 cm.

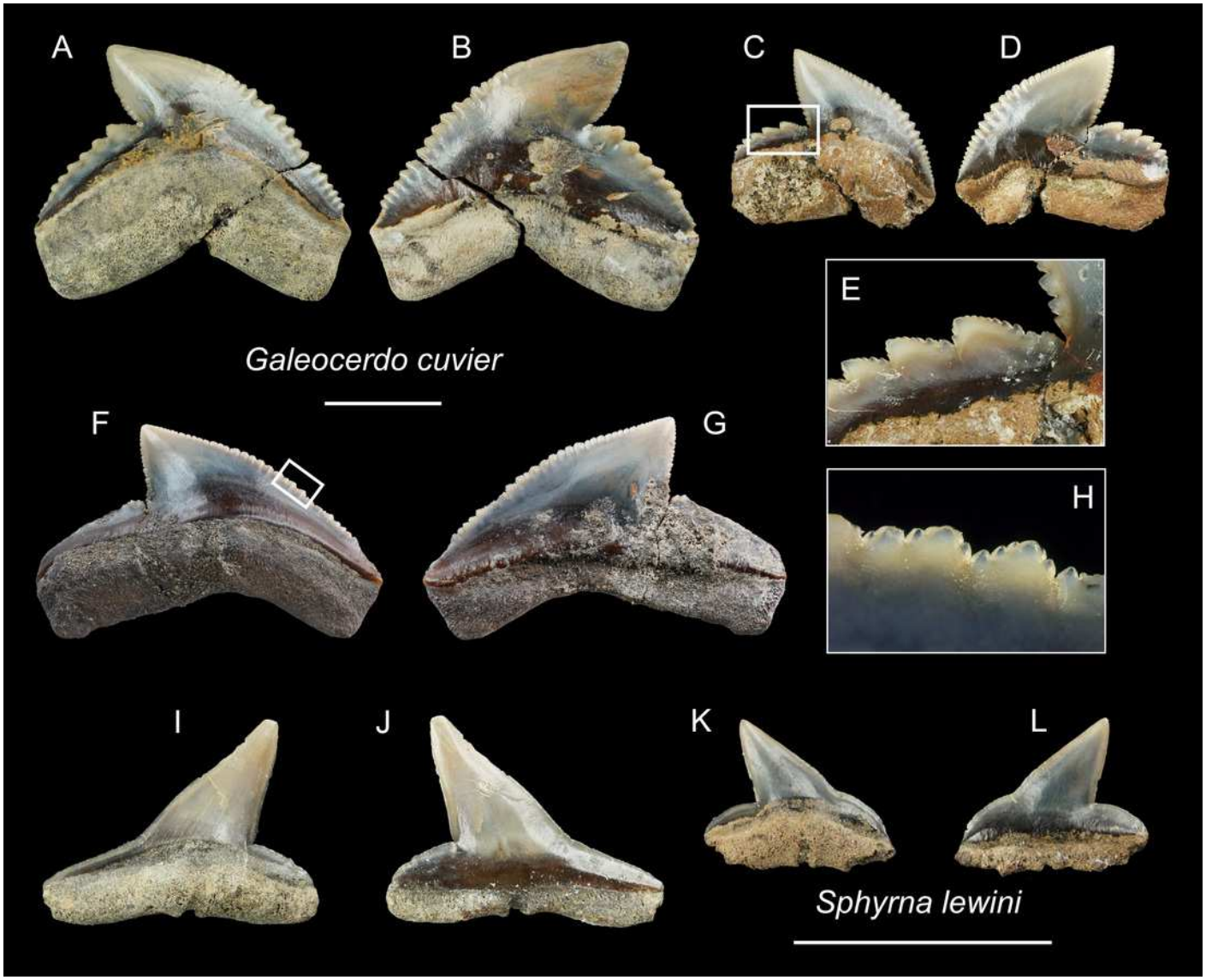

Figure 18

Teeth of Dasyatidae indet. from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan.

A, B, C, D, ASIZF0100590; E, F, G, H, ASIZF0100591. A, E = labial views; B, F = basal views; C, G = occlusal views; D, H = lateral views. Scale bar = 5 mm.

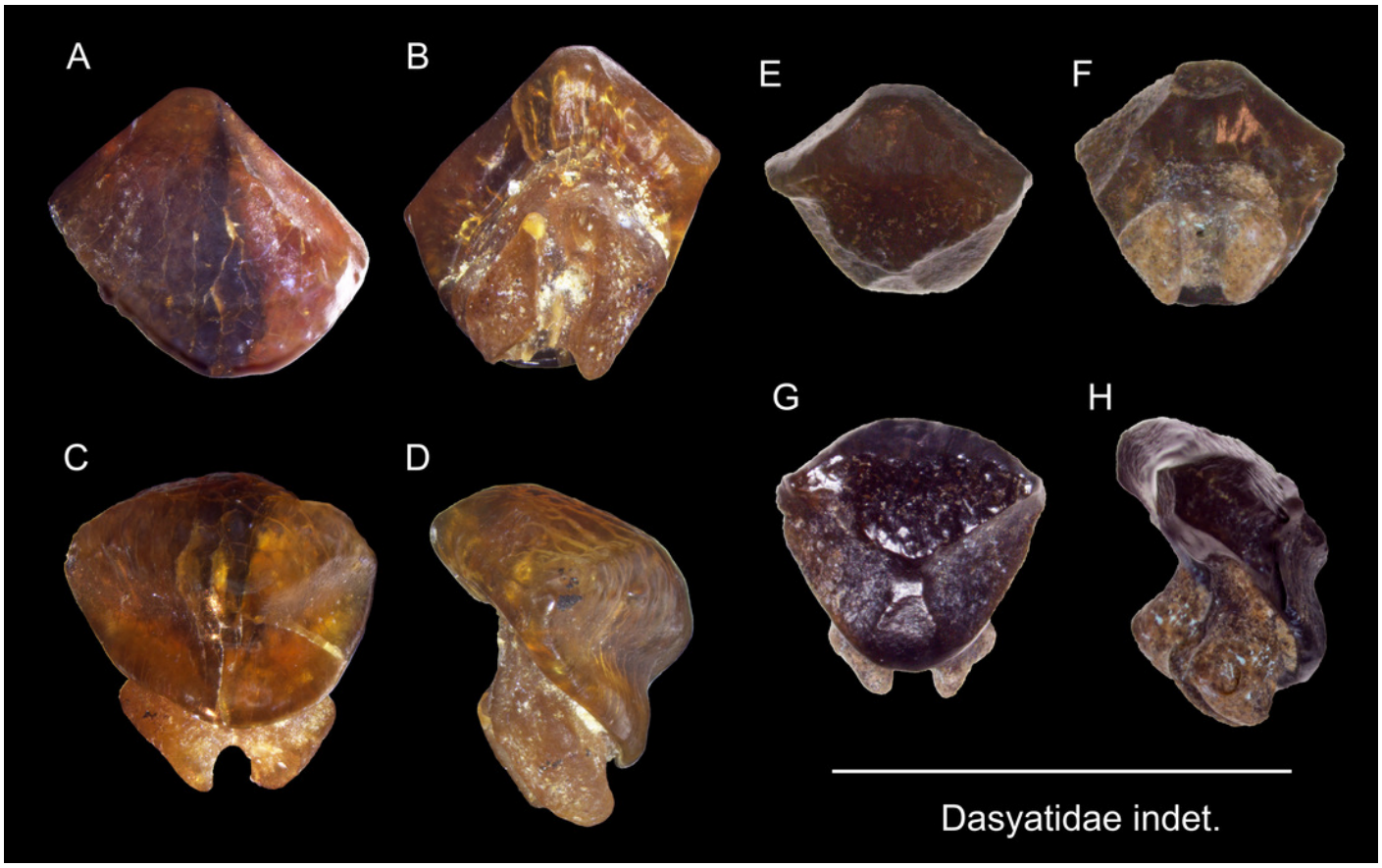

Figure 19

Teeth of *Aetobatus* sp. from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan.

A, B, CMM F2854; C, D, CMM F2850; E, F, CMM F0408; G, H, ASIZF0100549. A, C, E, G = occlusal views; B, D, F, H = basal views. Scale bar = 1 cm.

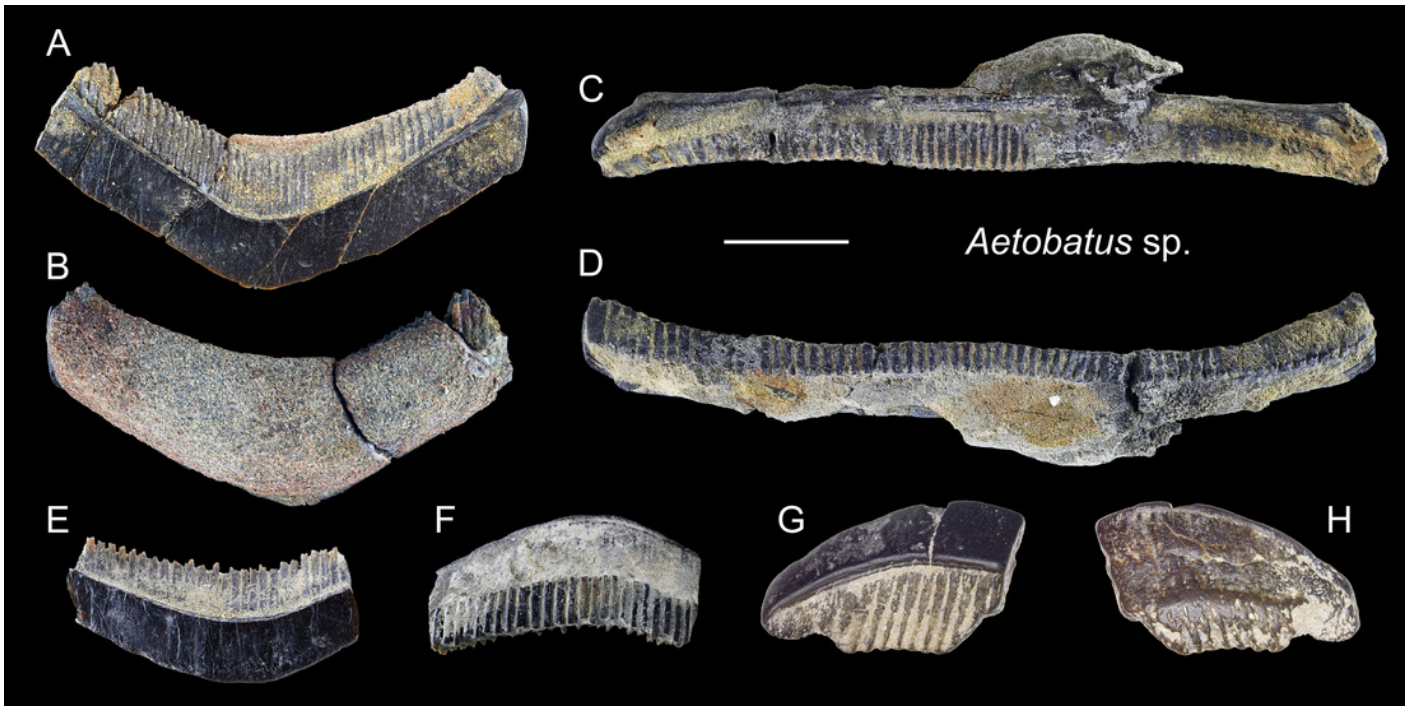

Figure 20

Teeth of *Myliobatis* sp. from the early Pleistocene of Liuchungchi Formation of Niubu, southern Taiwan.

A, B, C, D, ASIZF0100582; E, F, G, H, ASIZF0100587; I, J, K, L, ASIZF0100586; M, N, CMM F0395; O, P, CMM F2855; Q, R, CMM F0393; S, T, CMM F0398. A, E, I, M, O, Q, S = occlusal views; B, F, J, N, P, R, T = basal views; C, G, K = lingual views; D, H, L = labial views. Scale bars = 1 cm.

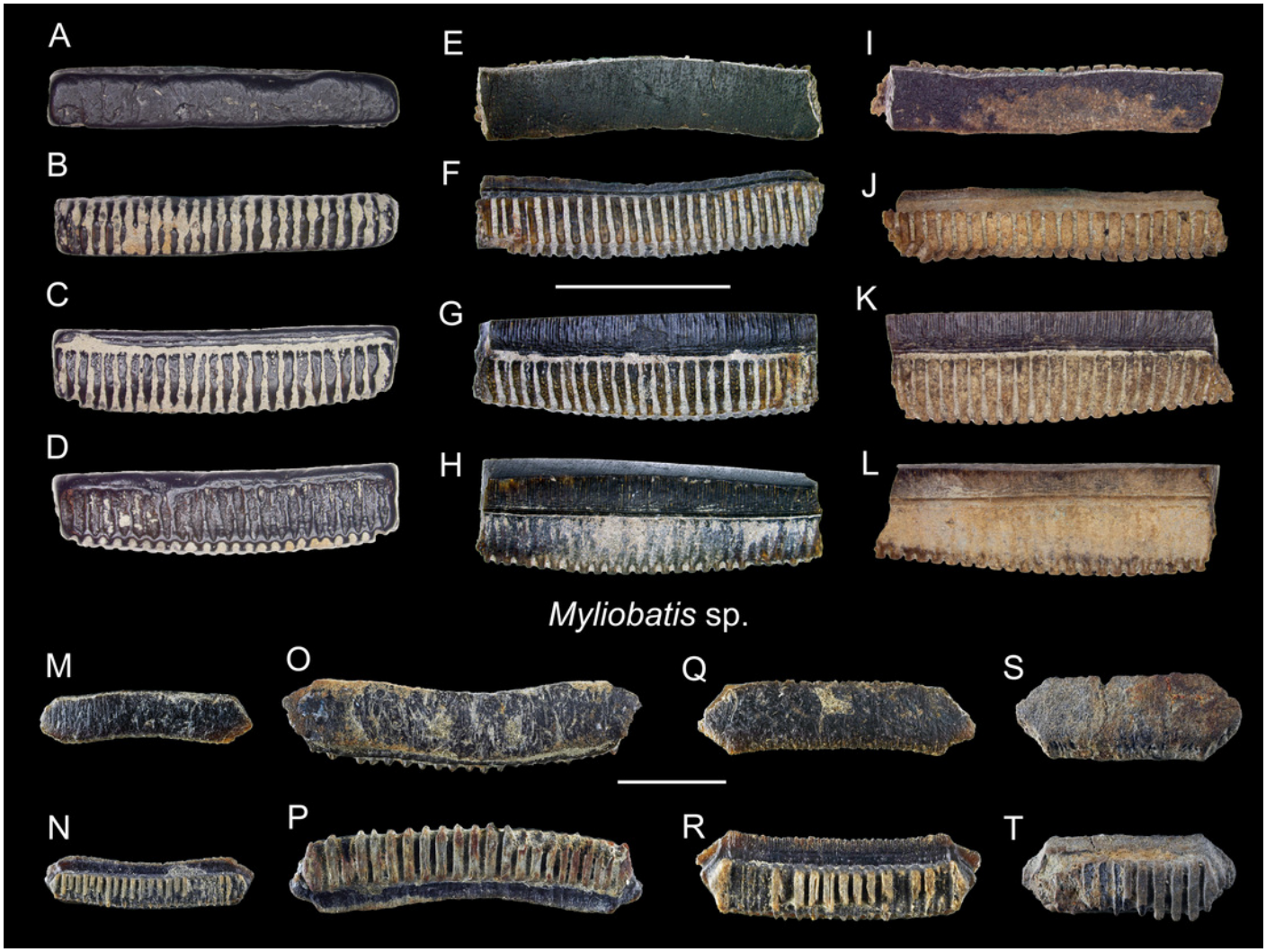

Table 1(on next page)

Elasmobranchs from the early Pleistocene Liuchungchi Formation of Niubu, southern Taiwan.

1 Table 1. Elasmobranchs from the early Pleistocene Liuchungchi Formation of Niubu, southern Taiwan.
2

Order	Family	Taxa	ASIZF	CMM	NTM	Total
Lamniformes	Carchariidae	*Carcharias taurus*	1	1		2
Lamniformes	Lamnidae	*Carcharodon carcharias*	28	25	2	55
Lamniformes	Lamnidae	*Isurus oxyrinchus*	4	1	1	6
Carcharhiniformes	Hemigaleidae	†*Hemipristis serra*	3	3	1	7
Carcharhiniformes	Carcharhinidae	*Carcharhinus altimus*	5	10	2	17
Carcharhiniformes	Carcharhinidae	*Carcharhinus amboinensis*	3	2		5
Carcharhiniformes	Carcharhinidae	*Carcharhinus leucas*	17	53	1	71
Carcharhiniformes	Carcharhinidae	*Carcharhinus limbatus*	16	21	3	40
Carcharhiniformes	Carcharhinidae	*Carcharhinus longimanus*	18	16	2	36
Carcharhiniformes	Carcharhinidae	*Carcharhinus obscurus*	9	15	1	25
Carcharhiniformes	Carcharhinidae	*Carcharhinus plumbeus*	8	42	1	51
Carcharhiniformes	Carcharhinidae	*Carcharhinus sorrah*	1	10		11
Carcharhiniformes	Carcharhinidae	*Carcharhinus tjutjot*	5	14		19
Carcharhiniformes	Carcharhinidae	*Carcharhinus* spp.	88	110	10	208
Carcharhiniformes	Carcharhinidae	*Rhizoprionodon acutus*	2	6		8
Carcharhiniformes	Galeocerdonidae	*Galeocerdo cuvier*	1	5	1	7
Carcharhiniformes	Sphyrnidae	*Sphyrna lewini*		2		2
Myliobatiformes	Dasyatidae	Dasyatidae indet.	2			2
Myliobatiformes	Aetobatidae	*Aetobatus* sp.	32	22	4	58
Myliobatiformes	Myliobatidae	*Myliobatis* sp.	9	20	1	30
Indet.	Indet.	Indet.	25	12		37
Total			277	390	30	697

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Table 2(on next page)

Various diversity indices from the Pleistocene West Pacific elasmobranch assemblages showing high diversity of the present material.

See Supplemental Table S1 for details of the data.

Table 2. Various diversity indices from the Pleistocene West Pacific elasmobranch assemblages showing high diversity of the present material. See Supplemental Table S1 for details of the data.

Location	Age	Species richness	Shannon	Simpson	Fisher's alpha	Reference
Taiwan	early Pleistocene	20	2.4	0.9	3.9	Our study
Sulawesi	Pleistocene	6	1.6	0.8	2.0	Hooijer, 1954
Java	Pleistocene	4	0.9	0.5	1.4	Koumans, 1949
Java	Plio-Pleistocene	11	1.7	0.7	2.9	Yudha et al., 2018
Central Japan	early Pleistocene	2	0.5	0.3	1.2	Karasawa, 1989
Central Japan	middle Pleistocene	14	2.2	0.9	6.0	Kawase & Nishimatsu, 2016
Eastern Japan	Pleistocene	14	2.3	0.9	4.6	Tanaka & Taru, 2022