

Malformations in Late Devonian brachiopods from the western Junggar, NW China and their potential causes

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Although malformations are found in both extant organisms and the fossil record, they are more rare in the fossil record than in living organisms, and the environmental factors causing the malformations are much more difficult to identify for the fossil record. Two athyrid brachiopod taxa from the Upper Devonian Hongguleleng Formation in western Junggar (Xinjiang, NW China) show distinctive shell malformation. Of 198 *Cleiothyridina* and 405 *Crinisarina* specimens, 18 and 39 individuals were deformed, respectively; a deformity rate of nearly 10%. Considering the preservation status and buried environment of the deformed specimens, and analysis of trace elements and rare earth elements from whole-rock and brachiopod shells, we conclude that the appearance of deformed athyrids may have been related to epi/endoparasites, or a low probability, that slightly high content of heavy metal in the sea.

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Abstract

Although malformations are found in both extant organisms and the fossil record, they are more rare in the fossil record than in living organisms, and the environmental factors causing the malformations are much more difficult to identify for the fossil record. Two athyrid brachiopod taxa from the Upper Devonian Hongguleleng Formation in western Junggar (Xinjiang, NW China) show distinctive shell malformation. Of 198 *Cleiothyridina* and 405 *Crinisarina* specimens, 18 and 39 individuals were deformed, respectively; a deformity rate of nearly 10%. Considering the preservation status and buried environment of the deformed specimens, and analysis of trace elements and rare earth elements from whole-rock and brachiopod shells, we conclude that the appearance of deformed athyrids may have been related to epi/endoparasites, or a low probability, that slightly high content of heavy metal in the sea.

Introduction

Deformities are common in living organisms; the term usually refers to soft body or skeletal tissue malformation of individuals that occur during ontogeny. However, malformations are known from the fossil record too, and have been reported from individuals of different fossil groups, including foraminifera (*Ballent & Carignano, 2008*), trilobites (*Owen, 1985; Babcock, 1993*), brachiopods (*Copper, 1967; He et al., 2017*), bivalves (*Savazzi, 1995*), gastropods (*Lindström & Peel, 2010*), cephalopods (*Mironenko, 2016*), echinoderms (*Thomka, Malgieri & Brett, 2014*), graptolites (*Han & Chen, 1994*), insects (*Vršanský, Liang & Ren, 2012*), conodonts (*Weddige, 1990*), shark teeth (*Itano, 2013*), amphibians (*Witzmann et al., 2013*), reptiles (*Buffetaut et al., 2007*), primate teeth (*Tougaard & Ducrocq, 1999*), and plankton

(*Vandenbroucke et al., 2015; Bralower & Self-Trail, 2016*). In addition to gene mutations or embryonic developmental disorders, deformed fossils may also have resulted from healed injuries and pathology (*Owen, 1985; Babcock, 1993; Kelley Kowalewski & Hansen, 2003; Vinn, 2007, 2008*). Malformed fossils provide important evidence of both organisms-organisms and organisms-environment relationships during geological history. For example, malformed specimens caused by predatory attacks provides us information about the food chain at that time or about the position of prey in the ecological chain (*Kelley Kowalewski & Hansen, 2003*). Moreover, some malformations resulting from diseases or developmental disorders are likely to be related to the habitat of the organism, such as changes in environmental factors, and parasite, viral or bacterial infection.

Although many malformed fossils have been described, deformed fossils (especially macrofossils) are generally rare, with sometimes only one or two specimens known. Therefore, previous studies have generally been limited to description of deformed specimens and simple classification of the cause(s) of malformation, only in a few cases the relationship between deformed specimens and their habitat has been discussed (*Copper, 1967; Vandenbroucke et al., 2015; Bralower & Self-Trail, 2016; He et al., 2017*). Although many deformed fossils are believed to result from developmental disorders or have pathological causes, the environmental factors responsible for the developmental disorders or diseases are generally unknown. The low number of malformed specimens available often limits further study. The Upper Devonian succession in western Junggar, Xinjiang, NW China, contains abundant, well-preserved brachiopods. We collected more than 600 athyrid (*Cleiothyridina* and *Crinisarina*) specimens

from the Upper Devonian Hongguleleng Formation in the Buninuer section, of which nearly ten percent of individuals were deformed. The aims of current study are to explore the biotic or abiotic stressors of these malformed brachiopods, and the effect of change in environmental factors on the brachiopod shells.

Materials and methods

The material studied in this paper was collected from the Upper Devonian Hongguleleng Formation in western Junggar, Xinjiang. The Hongguleleng Formation is a widely distributed marine unit near the Devonian–Carboniferous boundary in western Junggar. The formation is divided into three members: the Lower Member is composed of thin bioclastic limestones, muddy limestones and shales; the Middle Member is mainly made up of fine pyroclastic rocks with a few sandy and muddy limestones; and the Upper Member consists of calcareous clastic rocks with a small amount of bioclastic limestones ([Hou et al., 1993](#)). The formation is mostly Famennian in age ([Ma et al., 2017](#); [Zong et al., 2020](#)). In contrast to the general scarcity of fossils after the Late Devonian Upper Kellwasser event in other parts of the world, the Hongguleleng Formation is very rich in many types of the early Famennian fossils, such as [acritarches](#), bivalves, brachiopods, bryozoans, cephalopods, chondrichthyans, conodonts, conulariids, corals, echinoderms, gastropods, ostracods, plants, radiolarians, spores, trace fossils, and trilobites. Therefore, western Junggar is considered to have been a refugium during the Frasnian–Famennian extinction event ([Liao, 2002](#)).

Brachiopods occur in all three members of the Hongguleleng Formation. Brachiopod abundance and diversity is highest in the Lower Member, with the groups present including

Productida, Orthida, Rhynchonellida, Athyridida and Spiriferida ([Zong et al., 2016](#); [Zong & Ma, 2018](#)). Athyrids are most abundant in the Lower Member, with only a few athyrid specimen recovered from the base of the Middle Member and the limestone interlayer of the Upper Member ([Zong et al., 2016](#)). All the athyrids studied in this paper were extracted from the bioclastic limestone in the upper part of the Lower Member of the Hongguleleng Formation in the Buninuer section, 15 km north of Hoxtolgay town. This section is located about 14 km southwest of the Bulongguoer section, which is the type section of the Hongguleleng Formation ([Hou et al., 1993](#)). The lithology and fossil assemblages of the Buninuer section are the same as those of the stratotype section, the upper part of the Lower Member of the Hongguleleng Formation of both sections were deposited in a distal storm lithofacies sedimentary environment ([Fan & Gong, 2016](#)). A total of 603 athyrids in two genera (*Crinisarina* and *Cleiothyridina*) are non-flattened specimens with well-preserved dorsal and ventral valves. Although athyrids occur in other beds of the Hongguleleng Formation, no malformed specimens were found in those levels.

The 603 specimens include a wide range of size and may include individuals representing different growth stages ([Supplemental file 1](#)). We divided the shell length (L) into six size classes: $5\text{ mm} \leq L < 10\text{ mm}$; $10\text{ mm} \leq L < 15\text{ mm}$; $15\text{ mm} \leq L < 20\text{ mm}$; $20\text{ mm} \leq L < 25\text{ mm}$; $25\text{ mm} \leq L < 30\text{ mm}$ and $30\text{ mm} \leq L < 35\text{ mm}$, and counted the number of malformed specimens in each class. The length of all athyrids were measured by a vernier caliper. All photographs were taken using a Nikon D5100 camera with a Micro-Nikkor 55 mm f3.5 lens.

To explore whether athyrid deformities were caused by environmental factors, trace and

rare earth elements of whole-rock samples from specific levels within the Hongguleleng Formation were measured. Samples BL-1 and BL-2 were obtained from the lower part of the Lower Member of the Hongguleleng Formation, which yielded abundant undeformed athyrids. Samples BL-3 and BL-4 came from the upper part of the Lower Member, from where the deformed fossils described in this paper were obtained. Samples BL-5, BL-6 and BL-7 were collected from the Middle Member of the Hongguleleng Formation; only a few athyrids occurred at the bottom of this member, and the group almost disappeared above that level. Sample B9b-1 was from the Upper Member, which yielded a small number of athyrids. Besides, the trace elements and rare earth elements of four athyrid shells were measured. Samples CL69 and CR178 are undeformed shells, while CLJ13 and CRJ29 are malformed shells. All samples were ground into powder and analyzed in the ALS Minerals/ALS Chemex (Guangzhou) Co. Ltd. Rare earth and trace elements were fused with lithium borate, and quantitatively analyzed by ICP-MS with Elan 9000 Perkin Elmer that was made in America. The Ce_{anom} is equal to $\lg[3Ce_n/(2La_n+Nd_n)]$, and Ce_n , La_n and Nd_n were NASC-normalized of Ce, La and Nd, respectively.

Results

Of the 603 athyrid fossils, macroscopic deformities were detected in 57 specimens. The most common teratomorphy is obvious asymmetry on the left and right sides of the shells (Fig. 1B-E, G-J), significantly different from common, undeformed specimens (Fig. 1A, F). Malformation is more obvious on the dorsal valves, and is mainly visible as significantly widening or narrowing on one side of the shell (Fig. 1D1, G1, H1, J1). Near the anterior border

of the dorsal valves, the grooves on both sides of the fold are significantly different from those of undeformed specimens, with some grooves being wider (Fig. 1B1, C1, E1, I1), others being narrower, and some almost disappearing (Fig. 1D1, G1). On the ventral valves, in addition to the unequal size on either side of the shell, the sulcus is slightly curved in some malformed specimens (Fig. 1B2, E2, I2). In frontal view, the asymmetry is more obvious, and is mainly manifested as different depths and widths of the grooves on both sides of the fold (Fig. 1B3–E3, G3–J3); for example, the grooves on one side of some specimens become deeper and wider, up to twice as much as those on the non-deformed side (Fig. 1B3). In addition, the grooves of some specimens become shallower and narrower (Fig. 1H3), even almost disappearing on one side of a few specimens (Fig. 1D3, G3). In the malformed specimens, the commissure on the front of the dorsal and ventral valves forms irregular wavy lines, markedly different from the regular wavy lines in undeformed specimens (Fig. 1A3, F3).

Of the 198 specimens of *Cleiothyridina* and 405 of *Crinisarina*, 18 (9.1%) *Cleiothyridina* and 39 (9.63%) *Crinisarina* were malformed. The overall malformation rate was 9.45%, nearly one-tenth of all specimens (Fig. 2A). In all malformed specimens, the distribution of malformations is asymmetric on the shells in dorsal view, malformations occur in right side of 25 shells of *Crinisarina*, but there are only 14 in left side of shells. For the *Cleiothyridina*, malformation occur in right side of 14 shells, while in left side of 3 shells, and in both sides of one shell (Supplemental file 1). Moreover, malformed individuals occur in almost all size classes (Fig. 2B, C). The malformation percentages of *Crinisarina* are 5.56% ($10\text{ mm} \leq H < 15\text{ mm}$), 8.5% ($15\text{ mm} \leq H < 20\text{ mm}$), 13.7% ($20\text{ mm} \leq H < 25\text{ mm}$) and 11.1% ($25\text{ mm} \leq H < 30\text{ mm}$); those

of *Cleiothyridina* are 9.68% ($10 \text{ mm} \leq H < 15 \text{ mm}$), 9.3% ($20 \text{ mm} \leq H < 25 \text{ mm}$) and 14.8% ($25 \text{ mm} \leq H < 30 \text{ mm}$). Thus, shell-malformation occurs at different athyrid growth stages, and the probability of deformity is higher in larger specimens. That indicates a higher probability of deformity during advanced ontogenetic stages of the studied brachiopod taxa.

Discussion

Western Junggar is part of the Central Asian Orogenic Belt (*Buckman & Aitchison, 2004; Windley et al., 2007*). This region experienced strong tectonic activity during the Paleozoic, resulting in different degrees of metamorphism or deformation of the Paleozoic strata in the study area (*Xu et al., 2009; Gong & Zong, 2015; Wang & Zhang, 2019*). Athyrids exhibiting left-right asymmetry might have resulted from tectonic deformation; however, there is no obvious stratal deformation in the Hongguleleng Formation in the Buninuer section. This section has yielded fossils (e.g., trilobites, crinoids, and corals) that are well-preserved in three dimensions (*Fig. 3A*), which obviously differs from specimens obtained from distorted strata affected by tectonic deformation (*Fig. 3B*). Moreover, asymmetry was not detected in other brachiopods from the same layer, so the deformed specimens were not affected by the tectonic activity. A very small number of athyrid specimens cracked before lithification of the sediment are also significantly different from these asymmetric specimens, and they can be easily distinguished (*Fig. 3D*).

Epibionts can cause malformation of their hosts; such phenomena have also been reported for shelly fossils (e.g., *Zatoń & Borszcz, 2013; Mironenko, 2016*). Some athyrids from the Upper Devonian of western Junggar bear epibionts, such as corals and bryozoans. However, epibionts

are not found on any malformed specimens; on the contrary, specimens bearing epibionts are all undeformed shells (Fig. 3C). Therefore, it is unlikely that these athyrid teratomorphies were caused by epibionts. However, endoparasitic organism cannot be ruled out as a teratogenic factor (e.g., Savazzi, 1995; Vinn, Wilson & Toom, 2014), as well as epiparasitic shell-less organisms and microorganisms, they are likely to cause deformities in these athyrid shells from western Junggar. For the specific identity of endoparasites, shell-less organisms and microorganisms, it is difficult to confirm because their poor preservation in the fossil record. Some malformed brachiopod were caused by predators, which presented the fractures, indentations, and scars on the shells, and often accompanied by repaired signs (e.g., Alexander, 1986; Kowalewski, Flessa & Marcot, 1997; Happer, 2005; Vinn, 2017). In these malformed athyrid shells from western Junggar, except for a pair of indentations on the opposite valves of specimen BGEG-CLJ01 (Fig. 3H), no wounds or scars were found on other specimens. Most specimens only showed left-right asymmetry of the shells, reflecting that predation is not the main cause of the malformation, but predators may preyed on the malformed athyrids.

Malformations of organisms may also be related to their living environment. Changes in certain environmental factors, such as oxygen deficiency or excessive organic matter, heavy metals, and toxic elements, often lead to the malformation or even death of organisms. The high number of malformed athyrids from the Upper Devonian strata might be related to the marine environment in western Junggar at that time. Excessive organic matter is a common factor, and eutrophication has been identified in the Late Devonian sea (Murphy, Sageman & Hollander, 2000). Suttner et al. (2014) found that there were no significant changes in the total organic

carbon (TOC) content through the Lower Member of the Hongguleleng Formation at its type locality, i.e., the TOC content of the beds with deformed specimens was basically the same as that of sediments with only undeformed specimens. [Copper \(1967\)](#) studied deformities of the Devonian brachiopod *Kerpina* in the Eifel region, Germany, and concluded that the variations in the shell-morphology resulted from the influence of bottom currents on the immobile *Kerpina*, which had a thick, short pedicle. However, this mechanism cannot be used to explain the deformity of these athyrids in western Junggar, because large numbers of other benthic organisms (i.e., brachiopods, corals, bryozoans, and stromatoporoids), which are all undeformed in the same layer. In addition, the fossils preserved in the upper part of the Lower Member of the Hongguleleng Formation are relatively complete, and there is no evidence of strong bottom currents, so the influence of bottom currents can be excluded as a teratogenic factor. [Hoel \(2011\)](#) found the shells of brachiopod *Pentlandina loveni*, from the Sliurian Höglint Formation in Gotland (Sweden), are commonly markedly asymmetric, and some groups of shells occur in tight clusters, each apparently attached to other shells of the same species. He interpreted these asymmetrical shells resulted from the limited space for growth, i.e., overcrowded conditions. However, all specimens from western Junggar are isolated, instead of tight clusters or attached to other shells. Furthermore, if they are living in overcrowded space, the distribution of malformation should be random or almost uniform on both sides of the shells, but the malformations commonly occur on the right side of athyrids shells (dorsal view) from western Junggar ([Supplemental file 1](#)), so the overcrowded conditions also can be excluded.

Marine hypoxia ~~would~~ lead to brachiopod deformities. [He et al. \(2017\)](#) for example,

proposed that hypoxia was a major factor in the miniaturization of brachiopods during the end-Permian in southern China. U/Th and Ce_{anom} are often used as indicators of marine hypoxia (Jones & Manning, 1994; Carmichael et al., 2014, 2016). For the sediments of the Hongguleleng Formation at the Buninuer section, the U/Th ratio of sample BL-4 fell into the oxic range, whereas that of sample BL-3 fell into the dysoxic range; both Ce_{anom} values were near the oxic–anoxic boundary (Fig. 4, Supplemental file 2). In addition, shallow-marine benthic fossils, such as corals, trilobites, brachiopods and echinoderms, occur in abundance in the same horizon together with malformed athyrids, and the beds lack sedimentary indicators of anoxia (such as black shale), so hypoxia presumably did not occur during deposition of the upper part of the Lower Member of the Hongguleleng Formation.

High levels of heavy metals or toxic elements can also lead to malformation of organisms soft and hard tissue, as has been proven for a large number of living organisms (Wang, Yang & Wang, 2009; Ma et al., 2011; Zhao et al., 2017; Lasota et al., 2018; Riani, Cordova & Arifin, 2018). For example, when Cu and Zn were added to the water for feeding the foraminiferan *Ammonia beccarii*, the organisms developed deformities (Sharifi, Croudace & Austin, 1991), and when scallops were placed in wastewater from a gold mine with concentrations of 14% and 50% for 6 h, the deformity rate increased by 6% and 21%, respectively (Ma et al., 2011). Sediments have been demonstrated to be an important source of heavy metals for benthic animals (Wang, Stupakoff & Fisher, 1999). The levels of heavy metals and toxic elements through the Hongguleleng Formation are presented in Fig. 4 and Supplemental file 2. The levels of some heavy metals (e.g., cadmium, lead, barium, and zinc) in the layer with the malformed athyrid

shells are relatively high compared to the levels in the lower part of the Lower Member, particularly cadmium and lead. The levels of cadmium are 0.02 and 0.03 ppm in the lower part of the Lower Member, and are 0.03 and 0.04 ppm in the horizon that yielded the deformed fossils. The lead contents are 1.3 and 2.2 ppm in the lower part, but 5.1 and 5.2 ppm in the upper part. In the Middle Member of the Hongguleleng Formation, where athyrids almost disappeared, the heavy metal levels are even higher. In the Upper Member, where athyrids reappear, the heavy-metal content decreases again (Fig. 4). Thus, the abundance of athyrids is negatively correlated with the levels of heavy metals, but the number of malformed specimens is positively correlated with that, especially that of lead (Fig. 4). Furthermore, the levels of most heavy metals and toxic elements are slightly higher in malformed shells than in undeformed shells, particularly lead, silver, cobalt and arsenic (Supplemental file 3). In the studied specimens of teratomorphic *Cleiothyridina* and *Crinisarina*, the lead, silver, cobalt, and arsenic levels were all higher than those of undeformed shells. In *Cleiothyridina* the levels of lead, silver, cobalt, and arsenic were 2.3, 0.01, 2.0, and 2.9 ppm, respectively, in undeformed shells, but 2.6, 0.03, 2.5, and 3.6 ppm, respectively, in malformed shells. In *Crinisarina*, the levels of lead, silver, cobalt, and arsenic in undeformed shells were 1.9, <0.01, 2.6, and 2.2 ppm, respectively, whereas in the deformed specimens, the levels were 7.2, 0.01, 3.2, and 2.5 ppm, respectively (Supplemental file 3). However, the contents of cobalt and other heavy metals (i.e., copper, chromium, and vanadium) did not change significantly in the sediments from the undeformed athyrids-bearing layers to the layers containing malformed specimens, and the contents of arsenic and silver even decreased in the layers containing malformed athyrids (Fig. 4). Therefore, these heavy metals or toxic

elements may not be related to the deformity of athyrids. Only the lead content in sediments is positively correlated with the athyrid shell deformation, but negatively correlated with the abundance, while the lead content of malformed shells is higher than that of undeformed shells.

Lead is a common type of marine heavy-metal, and excessive lead content in the sea often leads to malformation or even death of shellfish, or at least affects their growth (Li, Sun & Li, 2011). In Upper Devonian strata of western Junggar, the deformity rate of athyrids is nearly one in ten. In addition, the heavy metal (especially lead) content in sediments is higher than that of sediments containing only undeformed shells, and the levels of some heavy metals are higher in deformed shells than in undeformed shells. However, the difference of lead content in the shells and surrounding rocks of undeformed and malformed brachiopods is still within the same order of magnitude, and no deformities have been found in other brachiopods taxa except these two genera. Therefore, it is not entirely certain that slightly higher lead content could cause deformities in these athyrid shells. In the future, geochemical analysis of more athyrid shells and a comprehensive comparison of heavy metal content from adjacent horizon and other brachiopods may be able to provide a more cogent teratogenic factor.

Conclusions

Some specimens of two athyrid genera, *Cleiothyridina* and *Crinisarina*, from the Upper Devonian Hongguleleng Formation in western Junggar are obviously deformed, mainly in the form of asymmetry of the left and right sides of the shells. The deformity rate is nearly 10% of specimens. Malformation is apparent in individuals of different sizes, with larger individuals being more likely to exhibit malformation. Based on the study of the burial state and preserved

environment of the fossils, and geochemical analysis of the sediments and athyrid shells and comparison with rock material from horizons that did not contain teratomorphic specimens, we hypothesize that the deformities were possibly caused by unidentified epi/endoparasites, or a low probability, that slightly high heavy-metal (specifically lead) in the sea, rather than eutrophication, bottom current activity, overcrowded conditions, hypoxia or other factors.

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Figure 1

Athyrids from the upper part of the Lower Member of the Hongguleleng Formation (Upper Devonian) in western Junggar

(A-E) *Crinisarina*, figure (A) (specimen number BGEG-CR324) is a undeformed specimen; figures (B-E) (specimen numbers BGEG-CRJ05, BGEG-CRJ10, BGEG-CRJ18 and BGEG-CRJ17) are malformed specimens, with malformations indicated by white arrows. (F-J) *Cleiothyridina*, figure (F) (specimen number BGEG-CL98) is a undeformed specimen; figures (G-J) (specimen numbers BGEG-CLJ06, BGEG-CLJ01, BGEG-CLJ15 and BGEG-CLJ13) are malformed specimens, with malformations indicated by white arrows. All scales are 10 mm.

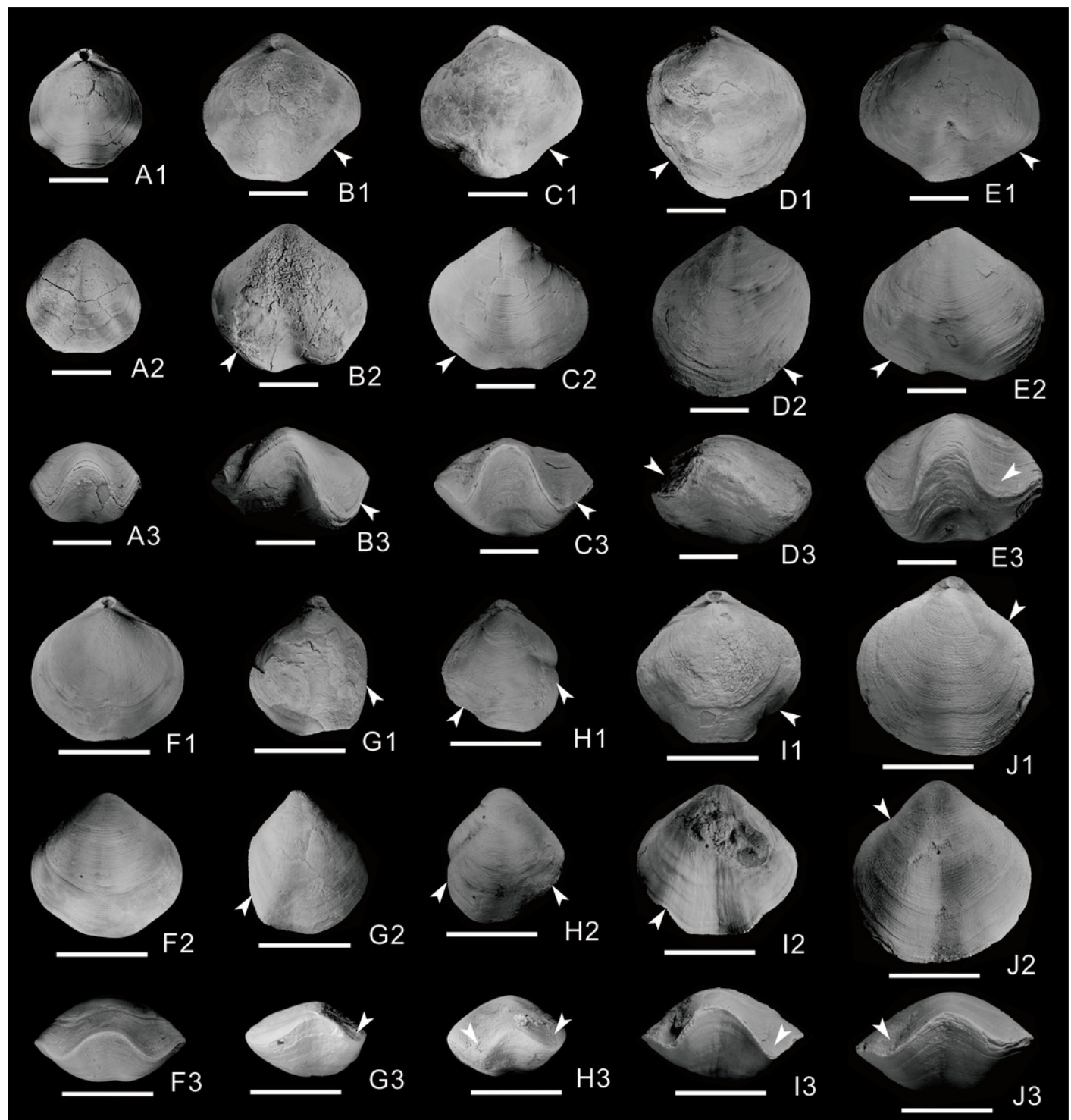


Figure 2

Histogram showing the number of malformed athyrids (A) and the distribution of undeformed and malformed specimens in different size classes (B-C) from the Upper Devonian Hongguleleng Formation in western Junggar

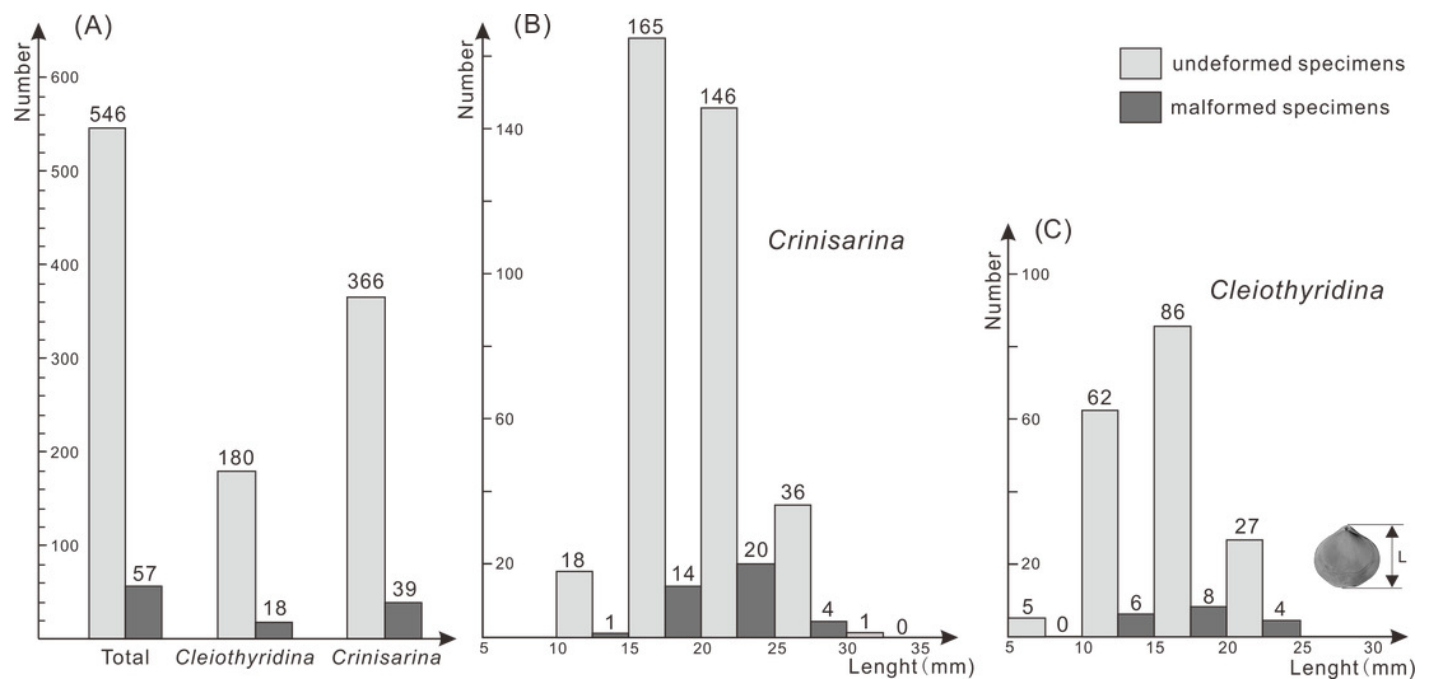


Figure 3

Some crinoid stems and athyrids from western Junggar

(A) Three-dimensional crinoid stem preserved in muddy limestone from the upper part of the Lower Member of the Hongguleleng Formation in the Buninuer section; (B) flattened crinoid stem preserved in the calcareous siltstone distorted by tectonic activity, Carboniferous Hala'ate Formation, western Junggar; (C) coral parasitizing a undeformed shell of *Crinisarina* (specimen number BGEG-CR44) from the Lower Member of the Hongguleleng Formation; (D) *Cleiothyridina* (specimen number BGEG-CL56) cracked before lithification of the sediment, obviously different from the malformed specimens. All scales are 10 mm.

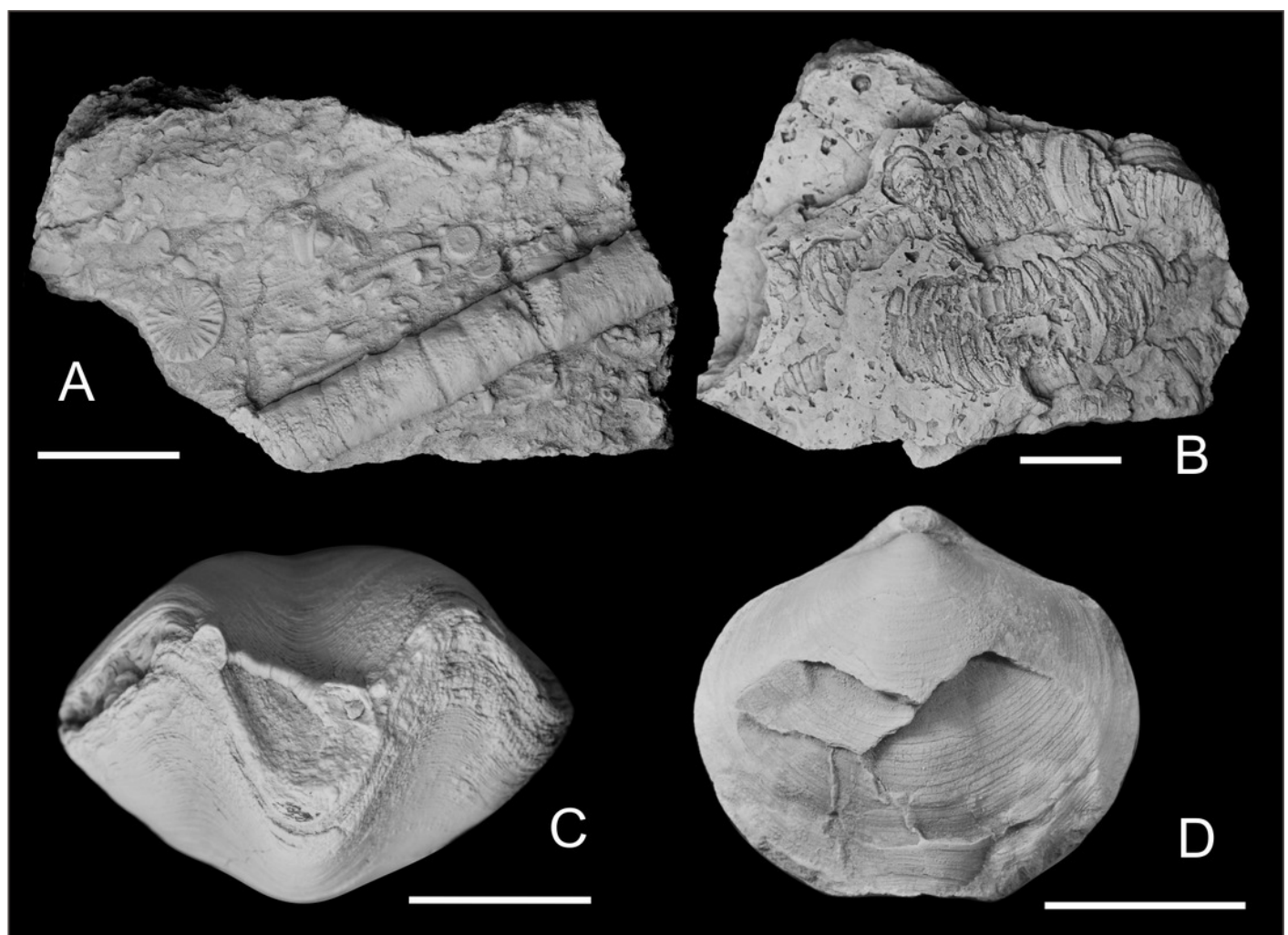


Figure 4

U/Th ratio, Ce_{anom} and distribution of heavy metals and toxic elements in whole rocks in the Upper Devonian Hongguleleng Formation in western Junggar

