

# A novel spatio-temporal two-stage standardized weighting scheme for regional drought analysis

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Drought is a complex phenomenon that occurs due to insufficient precipitation. It does not have immediate effects, but sustained drought can affect the hydrological, agriculture, economic sectors of the country. Therefore, there is a need for efficient methods and techniques that properly determine drought events. Considering the significance and importance of drought monitoring methodologies, the current study proposed a Novel weighting scheme, known as Spatio-Temporal Two-Stage Standardized Weighting Scheme (STTSSWS) for regional drought analysis. The potential of the weighting scheme is based on the steady-state probabilities. Further, in the first stage of the proposed scheme, the steady-state probabilities are calculated for various stations at a one-month time scale to assign weights for varying drought classes. However, in the second stage (stage two) these weights are further divided based on spatiotemporal characteristics to obtain new weights for the various drought classes in the selected region. The proposed weighting scheme is validated to the six selected stations of the northern area, Pakistan. The outcomes of the proposed weighting scheme may accurately assign spatiotemporal weights for selected states and provides efficient information for the selected region.



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## 24 1. Introduction

25 Drought is a creeping phenomenon, influencing more individuals than other natural  
26 hazards (Botai et al., 2027; Mousavikhah et al., 2020; Bhunia et al., 2020; Elhoussaoui et al., 2021).  
27 It is a slowly evolving and multifaceted disaster, often poorly understood in the perspective of  
28 regional climatic, hydrological, and human environment (Mukherjee et al., 2018; Vicente-Serrano  
29 et al., 2020). Drought is a recurring natural hazard appearing in all climatic zones worldwide and  
30 significantly influences social and economic well-being, ecological environment, and agricultural  
31 sectors (Gunalp et al., 2015; Hagenlocher et al., 2019; Shah and Mishra, 2020). In its explicit  
32 form, drought can be described as the water discrepancy that appears in several types, including  
33 agricultural, meteorological, hydrological, and socio-economic drought (Lai et al., 2019; Jiang &  
34 Wang, 2019). The meteorological drought occurs due to insufficient precipitation, and  
35 insufficiency in soil water supply triggers the agricultural drought. If drought distribution further  
36 continues via the hydrological cycle, a deficiency in surface or groundwater evolves, causing  
37 hydrological drought.

38 Further, the monitoring, modeling, and prediction of meteorological droughts are of utmost  
39 significance. Because the meteorological drought becomes the root for other drought types due to  
40 insufficient precipitation (Wu et al., 2017; Lai et al., 2019; West et al., 2019). Moreover, accurate  
41 valuation of meteorological drought brings useful information for decision-makers worldwide  
42 working in several fields associated with agriculture, hydrology, industrial, and water-budget  
43 managers to formulate precautionary measures and develop future planning (Alizadeh and Nikoo,  
44 2018). Further, various studies have described the importance of monitoring regional drought  
45 (Wilhite et al., 2000; Zhai & Feng, 2009; Zhang et al., 2012; Van Lanen et al., 2016; Santos et al.,

46 2019). The regional monitoring of drought has significantly influenced the country's economy and  
47 other human activities (Zhai & Feng, 2009; Mousavikhah et al., 2020). The regional drought  
48 monitoring highlights those issues if they could improve at the regional level before the events  
49 occur, then potential adverse effects of drought can be minimized in the future (Wilhite et al.,  
50 2000; Santos et al., 2019; Pontes Filho et al., 2019; Pontes Filho et al., 2020). More comprehensive  
51 and accurate drought monitoring can be possible by applying suitable tools and techniques  
52 according to climatic conditions. Based on the various climatic conditions, several drought indices  
53 have been used for drought monitoring. These indices require proper and effective recording  
54 related to drought occurrences. For instance, more accurate estimation of drought indices requires  
55 appropriate gauge stations with suitable records for regional drought. The indices contain essential  
56 information that can be used for the analysis that considerably improves forecasting and early  
57 warning policy about future drought occurrences (Tsakiris, et al., 2007; Niemeyer, 2008; Hayes et  
58 al., 2011; Mukherjee et al., 2018).

59         Moreover, various studies from the literature have discussed the drought indices. Several  
60 studies have developed some new drought indices (McKee et al., 1993; Tsakiris, et al., 2007;  
61 Vicente- Serrano et al., 2010). The development in drought mentoring is leading to enhancing the  
62 capabilities of drought monitoring more precisely and accurately. The standardized indices are  
63 used for the categorization. The estimation of the drought indices is based on the various  
64 parameters (precipitation, temperature, etc.). However, the preference for calculating the indices  
65 is based on climatic conditions of the available data (Niemeyer, 2008; Hayes et al., 2011). For  
66 instance, an index based on the precipitation, which is known as Standardized Precipitation Index  
67 (SPI) proposed by McKee et al., 1993, the Reconnaissance Drought Index (RDI) developed by  
68 Tsakiris, et al., 2007, the Vicente- Serrano et al. (2010) has proposed an index which is called as

69 Standardized Precipitation Evapotranspiration Index (SPEI). The spatiotemporal characteristics  
70 in the stations situated in a homogenous climatic area arise several issues in data analysis  
71 preliminaries. These issues are discussed in various studies available in the literature (Maybank et  
72 al., 1995; Umran, 1999; SIRDAŞ & Sen, 2003; Wang et al., 2020; Zhou et al., 2020).

73 Furthermore, drought is a complicated phenomenon; its occurrences are complex.  
74 Therefore, there is a need to develop some efficient techniques in drought monitoring policies to  
75 handle the spatiotemporal behavior of the drought (Maybank et al., 1995; SIRDAŞ & Sen, 2003;  
76 Wang et al., 2020). Since the estimation of spatiotemporal characteristics is of utmost importance  
77 in drought monitoring. The information obtained from the spatiotemporal characteristics can be  
78 used for significant modeling and drought prediction. Therefore, an intense spatiotemporal  
79 analysis is required to assimilate the spatiotemporal information of the selected homogenous  
80 region (Dabanlı et al., 2017; Caloiero, et al., 2018; Zhou et al., 2020). In this regard, we aimed to  
81 develop a new weighting scheme for the homogenous region. The weighting scheme, known as  
82 Spatio-Temporal Two-stage Standardized Weighting Scheme (STTSSWS), is validated on six  
83 selected stations of the northern area of Pakistan. The proposed weighting scheme provides more  
84 efficient and accurate information about the selected stations.

## 85 **2. Methods**

### 86 2.1. Standardized Drought Index

87 The various Standardized Drought Indices (SDI) have been used to monitor drought (Alley,  
88 1985; Narasimhan & Srinivasan, 2005; Stagge et al., 2015; Eslamian et al., 2017; Niaz et al., 2020).  
89 The two familiar drought indices SPI and SPEI are used in the present analysis. The SPI index was  
90 proposed by (McKee et al., 1993). It has been commonly used for drought monitoring. Several

91 studies have used SPI for drought monitoring (Stagge et al., 2015; Eslamian et al., 2017; Pathak  
92 and Dodamani, 2019; Hagenlocher et al., 2019; Niaz et al., 2020). The calculation of SPI is very  
93 simple; only the precipitation data is used for its calculation. The SPI can be calculated at various  
94 time scales. The calculation and the availability of the data of the precipitation are relatively easy;  
95 therefore, the SPI is commonly used worldwide. Further, the SPEI is a multi-scalar drought index  
96 that attracted significant attraction in drought estimation. The SPEI was developed by Vicente-  
97 Serrano et al. (2010) that obtains the simplicity in temporal characterization and considers as an  
98 extension of SPI. SPEI evaluates the effects of evaporative demand on drought and is computed  
99 by considering both precipitation and potential evapotranspiration. The SPEI can be calculated by  
100 fitting various probability distributions based on the climatic conditions of the selected stations.  
101 More detailed information concerning the SPEI computation can be acquired in Vicente Serrano  
102 et al. (2010) and Beguería et al. (2014).

103 2.2. The proposed weighting Scheme: The Spatio-Temporal Two-Stage Standardized Weighting  
104 Scheme (STTSSWS).

105 Drought causes severe damages worldwide. However, drought monitoring policies need a  
106 deep knowledge regarding the spatial and temporal distribution of drought risk at the local or  
107 regional level. Therefore, in this perspective, we propose STTSSWS, the innovative methodology  
108 giving a better evaluation and management of drought monitoring, especially for spatial and  
109 temporal characteristics of the region. The STTSSWS is based on steady-state probabilities. The  
110 steady-state probabilities can be defined as the average probability that the system remains in a  
111 certain state after many transitions.

112 Moreover, in a Markov process, it can be more explicitly defined as the probabilities are  
113 approached the steady-state probabilities after some periods have been passed. Moreover, detailed

114 mathematical explanations related to the steady-state probabilities of the Markov chain are  
 115 presented in Stewart (2009). Further, in STTSSWS, Steady-state probabilities are used as weights  
 116 in two stages. In the first stage of the calculation of STTSSWS, the weights for all drought classes  
 117 are calculated using steady-state probabilities as a weighting scheme (Niaz et al., 2020). However,  
 118 in the second stage, the weights obtained from steady-state probabilities for varying classes are  
 119 used to calculate new weights from the spatiotemporal distributions of the data. The second stage  
 120 has two phases; in the first phase, the weights are computed using temporal information for each  
 121 station separately. The calculation of the first phase is given in Equation 1.

$$122 \quad T_{December}(P_{(mi)(Skardu)}) = \frac{S_{(mi)Skardu}}{\sum_{i=1}^n S_{(mi)Skardu}}, \quad i = 1,2,3,\dots,47 \quad \text{and} \quad m = 1,2,\dots,6 \quad (1)$$

123 Where  $T_{December}(P_{(mi)(Skardu)})$  is indicating the probabilities (weights) for varying  
 124 drought classes at the time December in Skardu station. The  $i$  is indicating the specific month  
 125 (say, December of 1971, December of 1972 and so forth) varying over the designated period (from  
 126 January 1971 to December 2017). And  $m$  denoting the drought classes which are studied in this  
 127 study (say,  $m = ("1$  (Extremely Wet (EW)), 2 (Severely Dry (SW)), 3 (Median dry (MW)), 4  
 128 (Normal Dry (ND)), 5 (Median Dry (MD)), 6 (Severely Dry (SD)), and 7 (extremely Dry (ED)))").  
 129 The severity of the drought classes are illustrated according to (Niaz et al., 2020). The steady-  
 130 state probabilities for the period from January 1971 to December 2017 contains all Decembers at  
 131 the station (Skardu) are signified by  $S_{(mi)(Skardu)}$ . And  $\sum_{i=1}^n S_{(mi)(Skardu)}$  signifying that the  
 132 steady-state weights are added for December over the designated period with monthly data the  
 133 Skardu station with several drought classes. Moreover,  $n$  indicating the total months of December  
 134 (i.e 47) in Skardu for the designated period. For instance, the  $S_{(mi)(Skardu)}$  are computed with  
 135 several drought classes for Skardu station for the various months of December, then the

136 denominator contemplates the sum of the weights achieved from steady state probabilities for the  
 137 varying drought classes of Decmeber in the selected study period at Skardu station. More plainly,  
 138 it can be quantified as we have 47 December in the selected data. Hence, the weights of 47  
 139 December are being included (added) to the denominator for Skardu for seven drought classes.  
 140 Now, the monthly weights for other months (January, up to November) with these selected classes  
 141 are evaluated on the same rationale. To have sidestep from the complication of the mathematical  
 142 equalities, we only presented Skardu station for December in the first phase of the STTSSWS.  
 143 However, this can be extended for the other stations over varying months. Furthermore, the  
 144 methodology is presented in the second phase to get spatiotemporal qualities of the selected  
 145 drought classes. Thus, the spatiotemporal weights for these selected drought classes can be  
 146 obtained as follows,

$$147 \quad ST_{December}(P_{(mi)(Skardu)}) = \frac{T_{December}(P_{(mi)(Skardu)})}{\sum_{j=1}^M Q_{mij}}, \quad i = 1,2,3,\dots,47 \quad \text{and} \quad j = 1,2,\dots,6 \quad (2)$$

148 Where in Equation (2) takes monthly spatiotemporal weights for various drought classes at Skardu  
 149 station.  $ST_{December}(P_{(mi)(Astor)})$  is representing the probabilities computed from spatiotemporal  
 150 information (spatiotemporal weights) for varying drought classes in December at Skardu station.  
 151 Further, the weights  $T_{December}(P_{(mi)(Skardu)})$  which were calculated from Equation (1), are being  
 152 further divided by the  $\sum_{j=1}^M Q_{mij}$ . And where the quantity (observation)  $Q_{mij}$  precieved from the  
 153 varying selected drought classes ( $m$ ) at various selected stations ( $j$ ), and the total number of  
 154 selected station are denoted by  $M$  (i.e  $M = 6$ ). The STTSSWS uses spatiotemporal characteristics  
 155 of the selected stations and provides imperative knowledge about drought occurrences in a  
 156 homogenous region. The obtained information from the STTSSWS can be used to build substantial  
 157 drought monitoring procedures, techniques and methodologies.

### 158 3. Application

159           The six meteorological stations of the Northern Areas of Pakistan (Figure 1) are selected  
160 in STTSSWS for the regional drought analysis. Northern Area is a geographic area that has a  
161 group of three mountain ranges, the Himalayas, Karakoram, and the Hindu Kush, which cover  
162 most of the region (Rasul, G., 2011). Many of the world's tallest peaks are found in this region,  
163 including K-2, Nanga Parbat, and Rakaposhi. The average altitude of Karakorum is (6,100 M),  
164 Hindukush (7,690 M) and Himalaya (8,848 M) (Latif, Yet al., 2020). These high altitudes of  
165 mountains frequently deliver a significant portion of precipitation (Rasul et al., 2011; Bocchiola  
166 & Diolaiuti, 2013; Adnan et al., 2017). Further, the precipitation and the temperature of this region  
167 have substantial effects on the country's other regions (Bocchiola & Diolaiuti, 2013; Adnan, M et  
168 al., 2017). Therefore, the precipitation and temperature of the selected region are used in  
169 STTSSWS to substantiate drought occurrences. Drought is a typical phenomenon that directly or  
170 indirectly affects agriculture, forestry, cattle, fisheries, banking, energy, transportation, and growth  
171 rate (Anjum et al., 2010; Mazhar et al., 2015). Therefore, proper monitoring and appropriate  
172 drought mitigation techniques are required to minimize drought effects. The STTSSWS provides  
173 spatiotemporal information at the regional level. The information obtained from STTSSWS can  
174 be used to mitigate adverse drought effects. The STTSSWS provides weights for the severity of  
175 drought classes. The class which receives greater weight, its chances of occurrence are high in the  
176 selected region. The weights obtained from STTSSWS can be used for further drought monitoring  
177 strategies and mitigation policies.

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179

### 180 3.1. Results

181 The climatological characteristics of the selected region are given in Table 1. The observed  
182 data from these characteristics are used for the classification. The classification for the drought is  
183 completed using SPI and SPEI. The drought classification shows the various levels of drought  
184 severity (Li et al., 2015). For instance, the SDI (SPI and SPEI) value less than or equal to -2  
185 represents the extremely dry and greater than 2 classify as extreme wet condition and so forth  
186 (Table 2). Further, two standardize drought indices are included in the study based on the climatic  
187 conditions of the selected stations, and the standardization of these drought indices is done by  
188 using varying probability distributions. The distribution which is suitable according to climatic  
189 conditions is selected for the standardizations. For instance, at a one-month time scale, the 3p  
190 Weibull distribution shows suitable candidacy for the Astor station. The BIC of 3p Weibull  
191 distribution is -1036.5 which is minimum among other distributions. Therefore, the distribution is  
192 used for the standardization in this station. The 3p Weibull distribution, at a one-month time scale,  
193 shows suitable candidacy for Bunji station with BIC (-1031.0), Gilgit, with BIC (-1097), and for  
194 Skardu with BIC (-735.1). The 4p Beta distribution shows better candidacy at a one-month of SPI  
195 for two stations, including Gupis, and Chilas with BIC -788.7, -805.6, respectively. Further, for  
196 SPEI at a one-month time scale, the Trapezoidal distribution is fitting suitably for station Astor  
197 and Skardu with BIC -710.1 and -664.6, respectively. In Bunji, Gupis, Chilas, and Gilgit, the  
198 Johnson SB distribution is a suitable candidate concerning their minimum BIC values accordingly.  
199 Further, in spatiotemporal analysis based on STSSWS for a homogenous region, the weights of  
200 SPI at a one-month time scale are presented in Table 4. The weights for various drought classes  
201 obtained from STSSWS are defined accordingly. The STSSWS contains the temporal and  
202 spatial information of the whole region and provides more comprehensive and precise results for

203 varying drought classes. For example, in January, for the Astore station at the one-month time  
204 scale of SPI, the SD takes a value of 0.0268. The value shows that the SD has very less likely to  
205 occur in January. However, in January, the ND is more likely to occur in Gilgit station among  
206 other stations of the region with the weight (0.2672). For other stations and months, the weights  
207 for varying drought classes can be observed. Furthermore, STTSSWS weights for a homogenous  
208 region using SPEI at a one-month time scale are presented in Table 5. In the Astore station at the  
209 one-month time scale of SPEI, the MW takes a value of 0.0836. The value shows that the MW has  
210 very less likely to occur in January at Gilgit. However, in January, the ND is more likely to occur  
211 in Skardu station among other stations of the region with the weight (0.1844). Moreover, to avoid  
212 the complexity in the presentation of results, we just presented results for the particular year, which  
213 is 2017. However, the results of selected years can be observed from the proposed scheme.  
214 Furthermore, Figure 2 shows theoretical and empirical distributions for SPI at a one-month time  
215 scale (SPI-1), and theoretical and empirical distributions for SPEI at a one-month time scale (SPEI-  
216 1) for various stations are presented in Figure 3. Figure 4 shows the temporal behavior of SPI-1 at  
217 selected stations. Further, the temporal behavior of SPEI-1 at selected stations can be observed in  
218 Figure 5. The counts plots for the varying drought classes which are computed by SPI and SPEI  
219 are presented in Figure 6. It can be observed from the visualizations that the most prevalent class  
220 among other selected classes is ND. In both SPI and SPEI computes ND from the selected stations.  
221 It can be perceived that the ND class should be considered as an important class for further,  
222 analysis. To obtain the weights for the selected classes on varying stations for the selected indices  
223 the current study provides STTSSWS. The weights obtained from STTSSWS using SPI for seven  
224 drought classes are presented for first six months of the year 2017 in Figure 7. Moreover, the  
225 weights computed from STTSSWS based on SPEI for selected drought classes are presented for

226 first six months of the year 2017 in Figure 8. The weights are presented only for the classes that  
227 available in the corresponding month. For example, the ND occurs in January 2017, the weight  
228 can only be presented for ND and so forth. When the climatic conditions of the stations change,  
229 the proposed STTSSWS will work accordingly to the observed situation. The information obtained  
230 from STTSSWS can be applied to monitor drought more accurately. The obtained weights from  
231 STTSSWS can be used for more precise drought monitoring approaches.

### 232 3.2. Discussion

233 The two drought indices (SPI & SPEI) are considered in the current analysis. The  
234 probability distributions, which are appropriate according to time scales and stations, were selected  
235 for the standardization. The BIC criteria are used to select these probability distributions. Further,  
236 the steady-state probabilities are used for the computation of STTSSWS. In the first stage, the  
237 steady-state probabilities are calculated for each station separately regardless of the spatial  
238 accountability in the calculation. However, in stage two, these steady-state probabilities are used  
239 to calculate spatiotemporal weights for the varying drought classes. Hence, the proposed scheme  
240 used spatial and temporal characteristics of regional drought to calculate weights for the various  
241 drought classes. The class which receives maximum weights from STTSSWS is reflected for  
242 consideration in the region. Moreover, the drought is a typical phenomenon that affects directly or  
243 indirectly the agriculture, forestry, cattle, fisheries, banking, energy, transportation, and growth  
244 rate (Mazhar et al., 2015; Guneralp et al., 2015; Hagenlocher et al., 2019; Lai et al., 2019; Jiang &  
245 Wang, 2019). Therefore, proper monitoring and appropriate drought mitigation techniques are  
246 required to minimize drought effects. The STTSSWS provides the more appropriate  
247 spatiotemporal information at the regional level. The obtained information from STTSSWS can  
248 be used to monitor drought more precisely and accurately. The weights obtained from STTSSWS

249 can be used for further drought monitoring strategies and mitigation policies. The obtained results  
250 from the STTSSWS provide the basis to improve the drought monitoring and forecasting methods  
251 at the regional level.

#### 252 **4. Conclusion**

253 Drought is a multifaceted phenomenon that occurs due to insufficient precipitation. It does  
254 not have instant consequences, but persistent drought for a prolonged time can greatly influence  
255 the agriculture, hydrological, and economic sectors of the country. Since there are required  
256 effective procedures that appropriately identify the drought occurrences and help policymakers to  
257 prepare their plans explicitly at the regional level. Therefore, the present study proposed a Novel  
258 weighting scheme, known as STTSSWS for regional drought analysis. The STTSSWS provides  
259 more accurate and precise spatiotemporal information about drought occurrences. The proposed  
260 weighting scheme is based on the steady-state probabilities. The steady-state probabilities at  
261 various stations are used to obtain new spatiotemporal weights for various drought classes. The  
262 proposed weighting scheme is validated on the six selected stations of the northern area, Pakistan.  
263 The outcomes of the proposed weighting scheme may accurately assign spatiotemporal weights  
264 for selected states and provides efficient information for the selected region.

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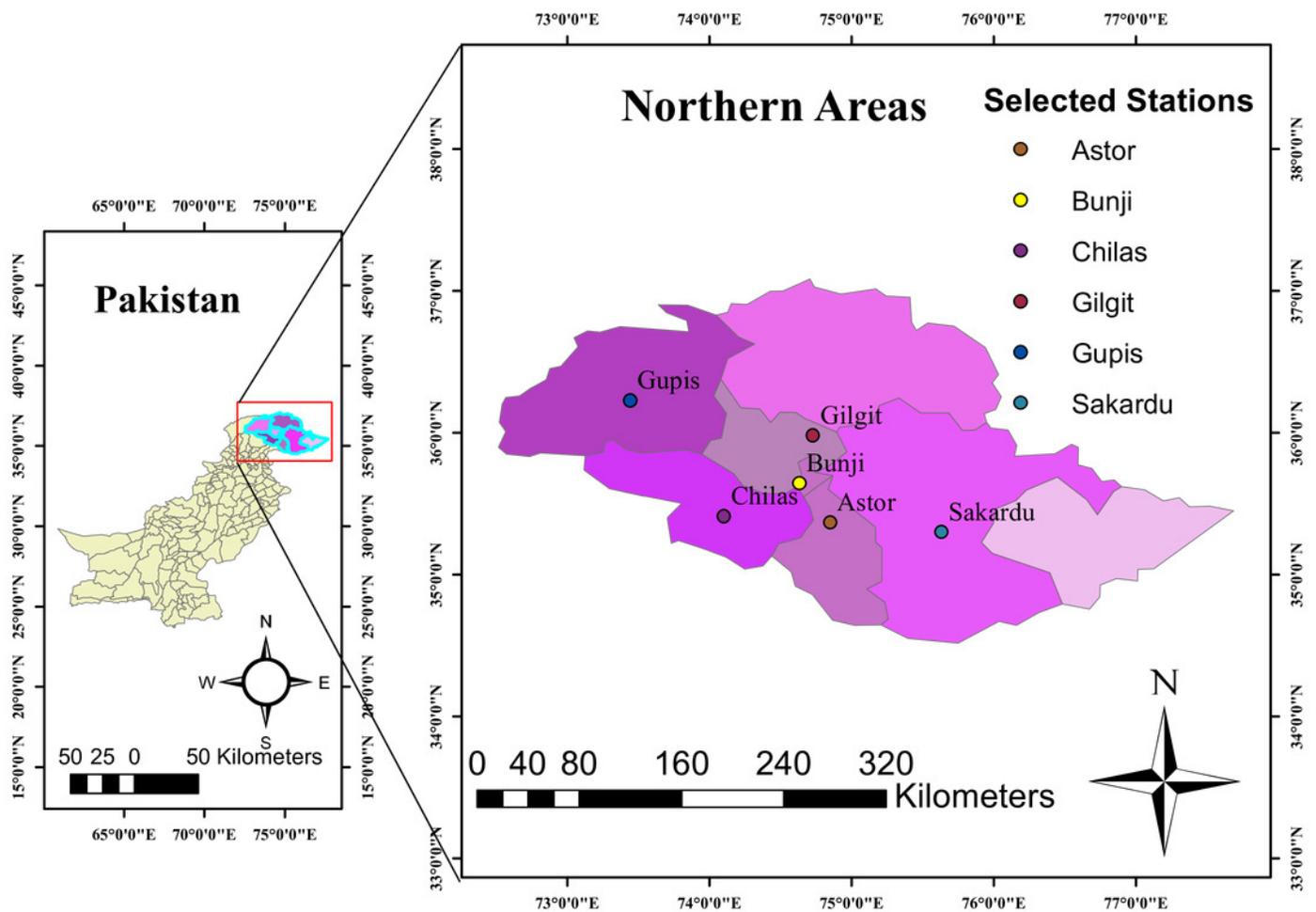
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385

# Figure 1

Geographical locations of the selected stations

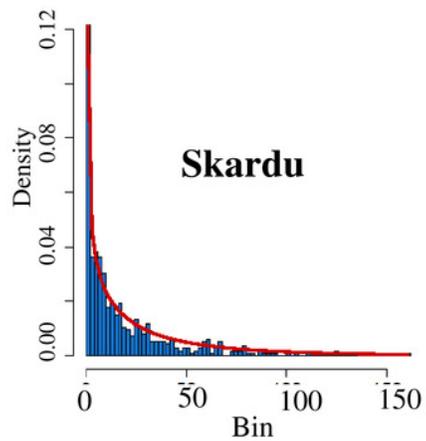
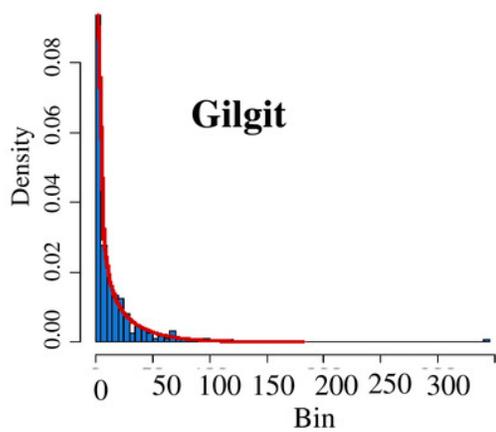
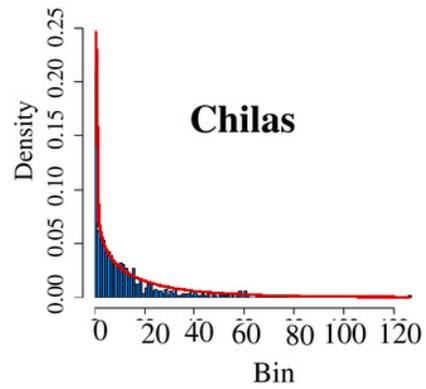
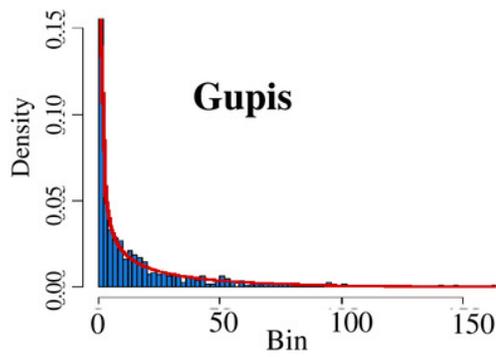
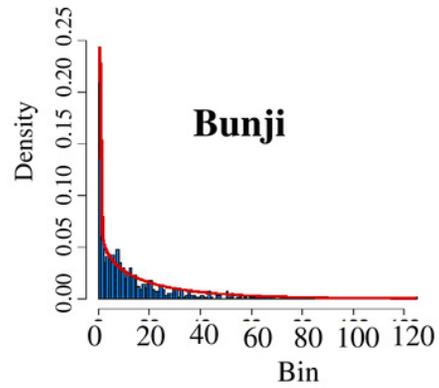
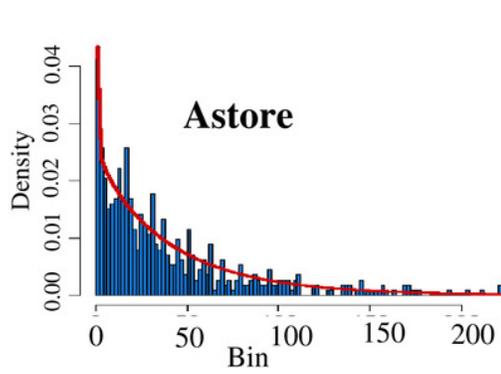
The map of Northern area of Pakistan and Geographical locations of the selected stations.



## Figure 2

Histograms of selected distributions for various stations

Theoretical vs. empirical histograms of selected distributions for selected stations

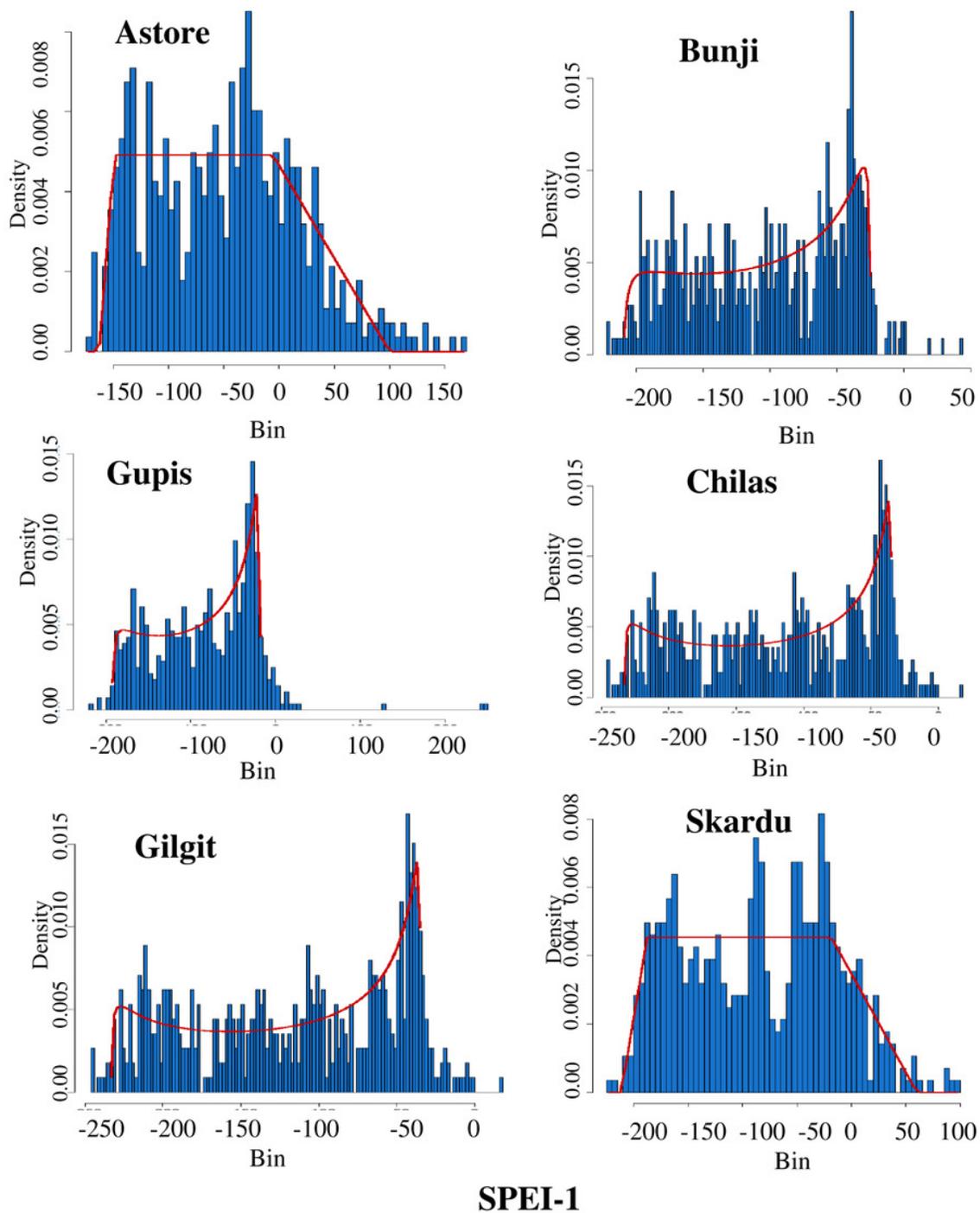


**SPI-1**

## Figure 3

Various histograms of selected distributions for varying stations

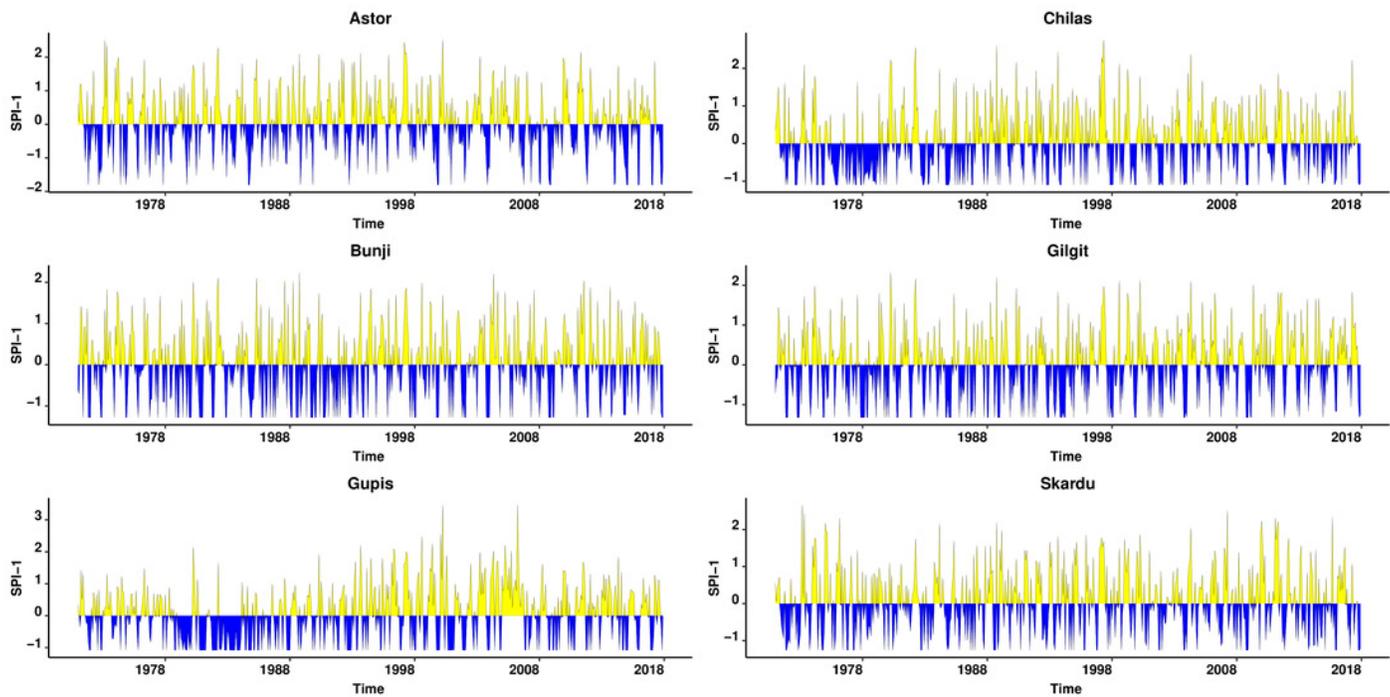
Theoretical vs. empirical histograms of selected distributions for selected stations



## Figure 4

Temporal plots using SPI at one-month time scale

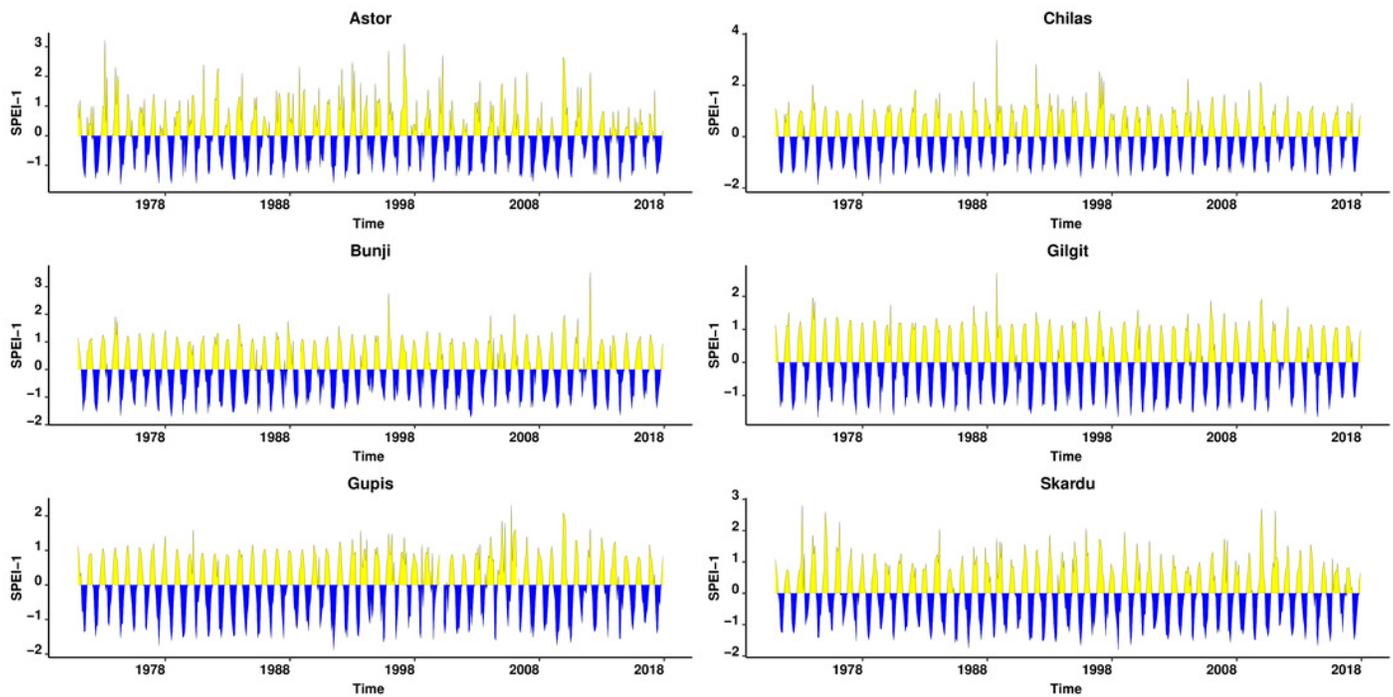
The temporal plots using SPI at one-month time scale for the selected stations



## Figure 5

The temporal plots using SPEI at one-month time scale

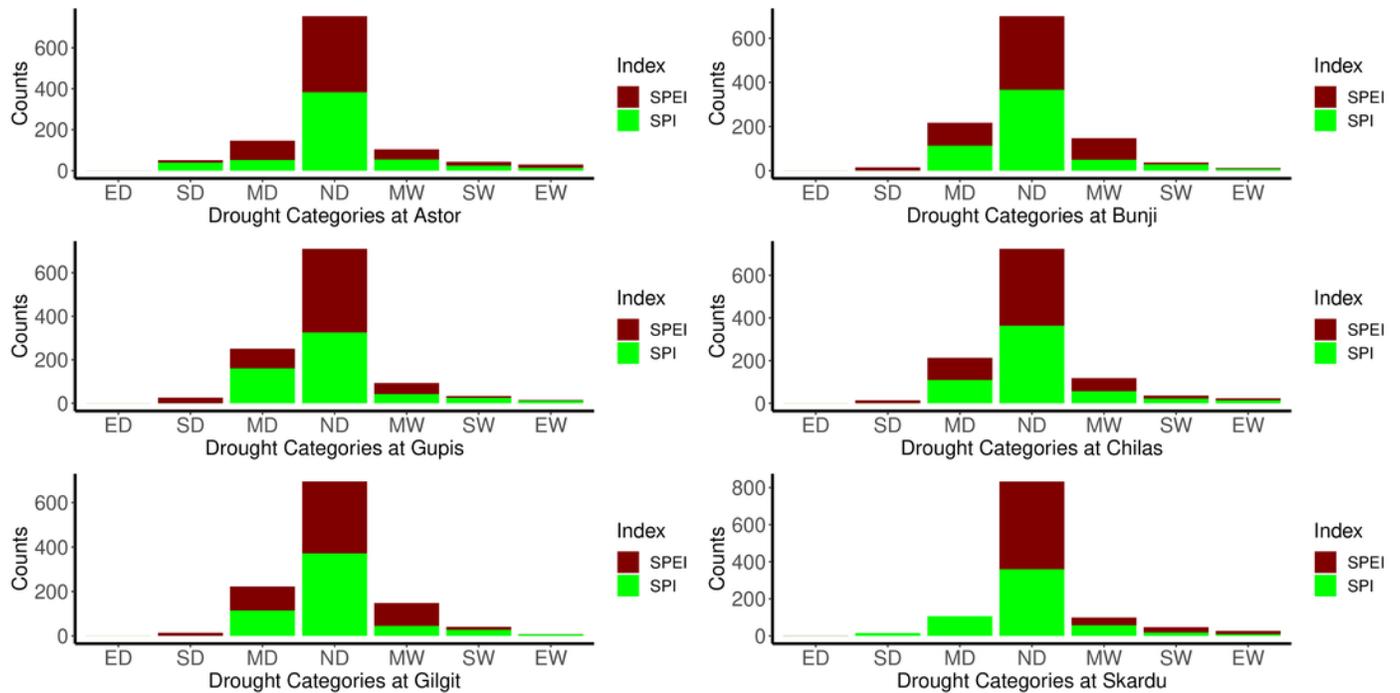
The temporal plots using SPEI at one-month time scale the selected stations



## Figure 6

Counts plots for selected stations

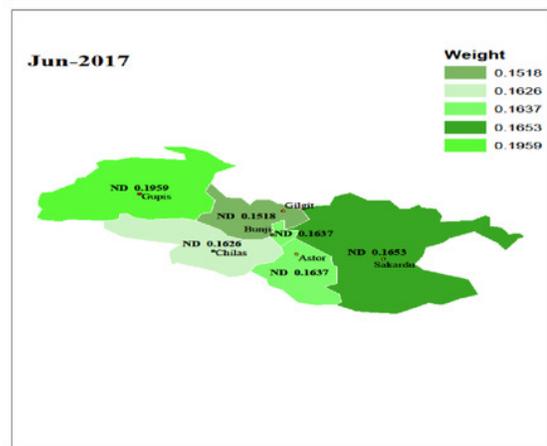
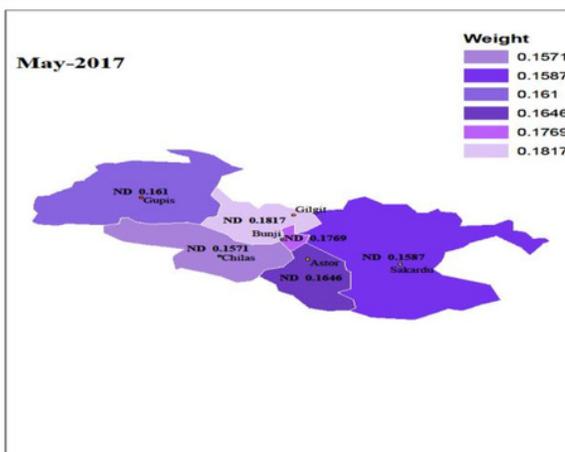
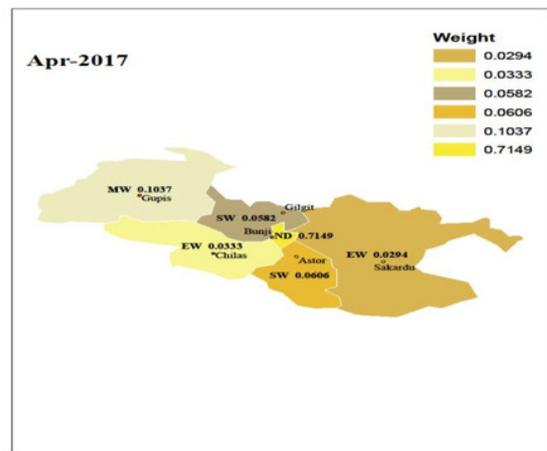
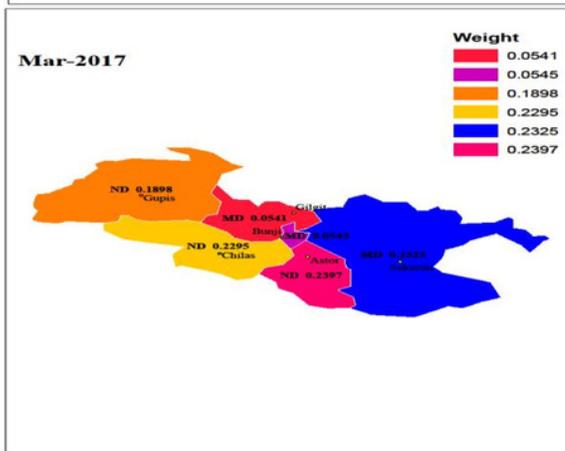
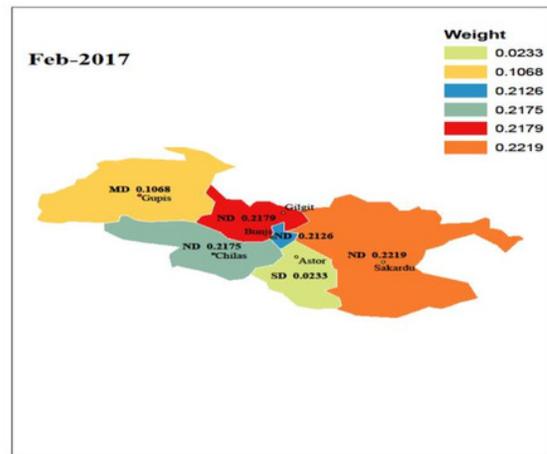
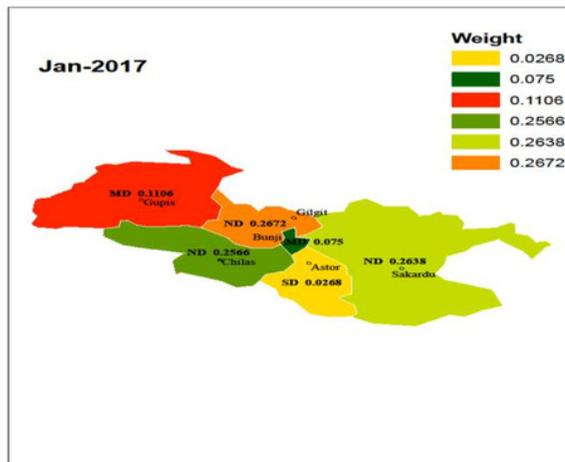
The counts plots at one-month time scale using SPI and SPEI for selected stations



## Figure 7

Weights obtained from STTSSWS using SPI for seven drought classes

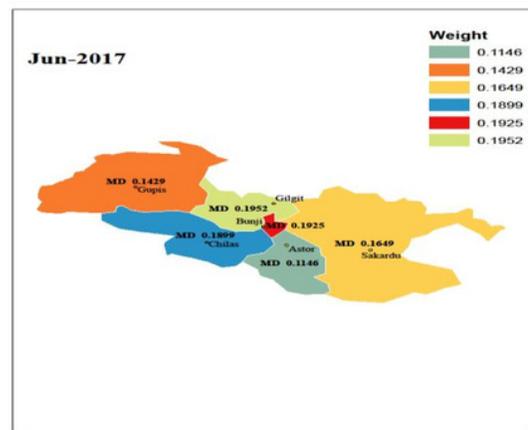
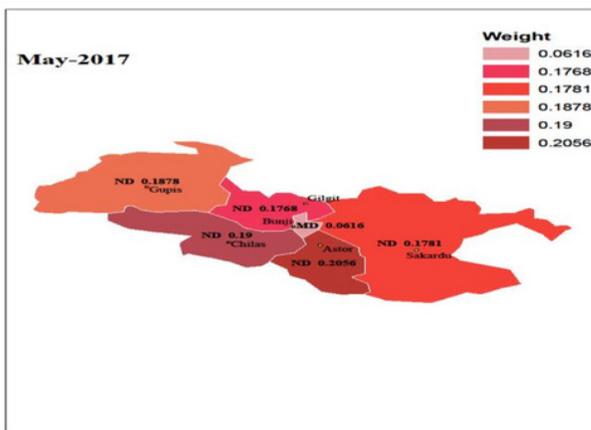
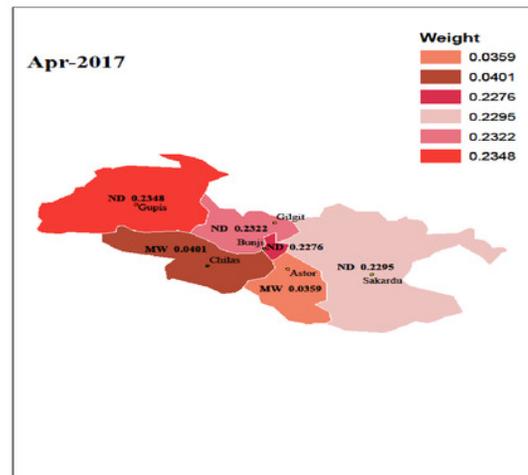
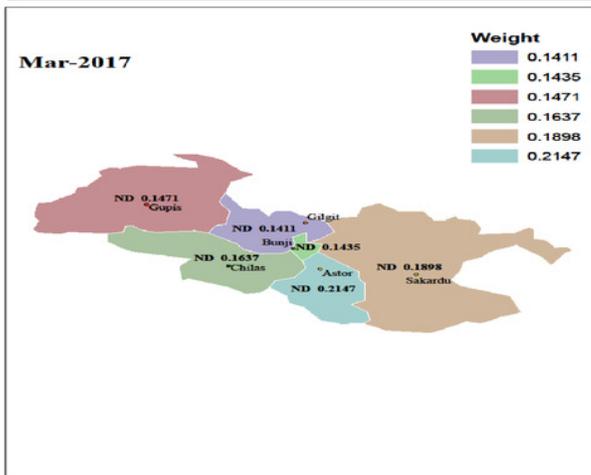
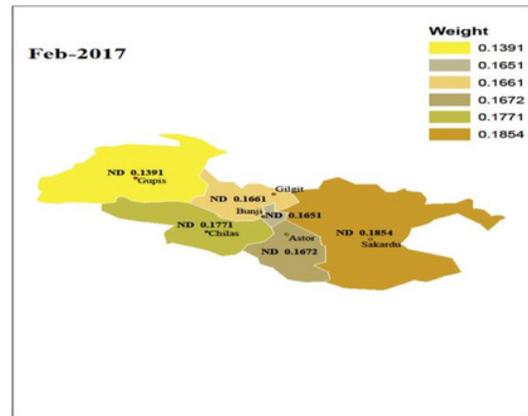
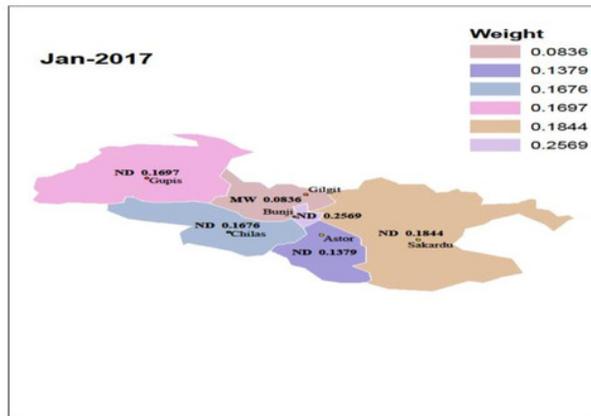
The weights obtained from STTSSWS using SPI for seven drought classes are presented for first six months of the year 2017



## Figure 8

The weights obtained from STTSSWS using SPEI for seven drought classes

The weights obtained from STTSSWS using SPEI for seven drought classes are presented for first six months of the year 2017



**Table 1** (on next page)

The climatology characteristics of selected period

The climatology characteristics during the period 1971–2017 of six selected stations

1

Variable	Station	Mean	1st Quartile	Median	3rd Quartile	Kurtosis	St.Dev
Precipitation	Astor	39.34	10.80	25.70	52.62	3.01	41.93
	Bunji	13.56	1.30	7.10	17.10	7.55	18.90
	Gupis	15.94	0.00	5.70	19.38	51.38	30.21
	Chilas	15.85	0.95	7.00	19.32	8.88	23.53
	Gilgit	11.75	1.10	6.05	14.72	9.93	16.57
	Skardu	19.51	2.30	9.10	26.75	5.60	25.90
Maximum Temperature	Astor	15.76	7.38	16.70	23.86	-1.36	8.65
	Bunji	23.82	15.78	24.95	32.02	-1.32	8.98
	Gupis	18.86	10.32	19.70	27.40	-1.31	9.46
	Chilas	26.41	17.68	27.35	35.62	-1.38	9.66
	Gilgit	24.07	15.60	25.15	32.80	-1.33	9.21
	Skardu	18.83	9.95	20.05	27.90	-1.25	9.82
Minimum Temperature	Astor	4.10	-2.43	4.30	10.70	-1.23	7.48
	Bunji	11.22	3.78	11.50	17.70	-1.24	7.80
	Gupis	6.37	-1.10	6.90	13.32	-1.27	8.06
	Chilas	14.32	5.68	14.30	23.20	-1.41	9.08
	Gilgit	7.63	0.60	7.75	13.53	-1.24	7.30
	Skardu	4.79	-2.73	5.55	11.80	-1.18	8.36

2

**Table 2** (on next page)

Drought categorization based on the values of the SDI

Categorization of drought based on the values of the SDI

1

SDI	Major drought classes
SDI $\geq 2$	Extremely Wet (EW)
SDI $> 1.5$ & SDI $\leq 2$	Severely (SW)
SDI $> 1$ & SDI $\leq 1.5$	Median Wet (MW)
SDI $> -1$ & SDI $\leq 1$	Normal Dry (ND)
SDI $> -1.5$ & SDI $\leq -1$	Median Dry (MD)
SDI $> -2$ & SDI $\leq -1.5$	Severely Dry (SD)
SDI $\geq -2$	Extremely Dry (ED)

**Table 3** (on next page)

BIC values for selected distributions

The selected probability distribution of SPI-1 and SPEI-1 and their BIC values for selected stations

1

	Astore		Bunji		Gupis	
index	Distribution	BIC	Distribution	BIC	Distribution	BIC
SPI	3p Weibull	-1036.5	3p Weibull	-1031.0	4p Beta	-788.7
SPEI	Trapezoidal	-710.1	Johnson SB	-1248.4	Johnson SB	-977.6
	Chilas		Gilgit		Skardu	
index	Distribution	BIC	Distribution	BIC	Distribution	BIC
SPI	4P Beta	-805.6	3P Weibull	-1097.4	3P Weibull	-735.1
SPEI	Johnson SB	-594.7	Johnson SB	-1213.2	Trapezoidal	-664.6

**Table 4**(on next page)

The weights calculated from STTSSWS using SPI-1 for varying drought classes at six stations

The weights calculated from STTSSWS using SPI-1 for varying drought classes at six stations are given. The STTSSWS used the monthly data which has range from 1971 to 2017.

However, here, the obtained weights from STTSSWS for the specific year 2017 at first six months (Jan, Feb, March, Apr, May, and Jun) are presented only. The bold value in SPI-1 shows that the the class has more likely to occur in various months of 2017. The values indicate that the ND is prevailing in this year among other drought classes for the selected stations

1

## SPI-1

Station Month	Astor		Bunji		Gupis		Chilas		Gilgit		Skardu		Sum
	Category	Weight	Category	Weight	Category	Weight	Category	Weight	Category	Weight	Category	Weight	
Jan	SD	0.0268	MD	0.0750	MD	0.1106	ND	0.2566	ND	<b>0.2672</b>	ND	0.2638	1
Feb	SD	0.0233	ND	0.2126	MD	0.1068	ND	0.2175	ND	0.2179	ND	<b>0.2219</b>	1
Mar	ND	<b>0.2397</b>	MD	0.0545	ND	0.1898	ND	0.2295	MD	0.0541	ND	0.2325	1
Apr	SW	0.0606	ND	<b>0.7149</b>	MW	0.1037	EW	0.0333	SW	0.0582	EW	0.0294	1
May	ND	0.1646	ND	0.1769	ND	0.1610	ND	0.1571	ND	<b>0.1817</b>	ND	0.1587	1
Jun	ND	0.1609	ND	0.1637	ND	0.1959	ND	0.1626	ND	0.1518	ND	<b>0.1653</b>	1

**Table 5** (on next page)

The weights calculated from STTSSWS using SPEI-1 for various drought classes at six stations

The weights calculated from STTSSWS using SPEI-1 for various drought classes at six stations are presented. The STTSSWS utilized the monthly data which has range from 1971 to 2017. However, here, the weights obtained from STTSSWS for the particular year 2017 in first six months (Jan, Feb, March, Apr, May, and Jun) are given only. The bold value in SPI-1 representing that the the classes have more likely to occur in first six months of 2017. The values suggest that the ND is more prevalent in this year among other drought classes for the selected

## SPEI-1

Station Month	Astor		Bunji		Gupis		Chilas		Gilgit		Skardu		Sum
	Category	Weight	Category	Weight	Category	Weight	Category	Weight	Category	Weight	Category	Weight	
Jan	ND	0.1379	ND	0.2569	ND	0.1697	ND	0.1676	MW	0.0836	ND	<b>0.1844</b>	1
Feb	ND	0.1672	ND	0.1651	ND	0.1391	ND	0.1771	ND	0.1661	ND	<b>0.1854</b>	1
Mar	ND	0.2147	ND	0.1435	ND	0.1471	ND	0.1637	ND	0.1411	ND	<b>0.1898</b>	1
Apr	MW	0.0359	ND	0.2276	ND	0.2348	MW	0.0401	ND	<b>0.2322</b>	ND	0.2295	1
May	ND	<b>0.2056</b>	MD	0.0616	ND	0.1878	ND	0.1900	ND	0.1768	ND	0.1781	1
Jun	MD	0.1146	MD	0.1925	MD	0.1429	MD	0.1899	MD	<b>0.1952</b>	MD	0.1649	1

1