

On the enigma of *Palaenigma wrangeli* (Schmidt), a conulariid with a partly non-mineralized skeleton

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Palaenigma wrangeli (Schmidt) is a finger-sized fossil with a tetroradiate conical skeleton; it occurs as a rare component in fossiliferous Upper Ordovician strata of the eastern Baltic Basin and is known exclusively from north Estonia. The systematic affinities and palaeoecology of *P. wrangeli* remained questionable. Here, the available specimens of *P. wrangeli* have been reexamined using environmental scanning electron microscopy and x-ray computed tomography. Additionally, the elemental composition of the skeletal elements has been checked using energy dispersive X-ray spectroscopy. The resulting 2D- and 3D-scans reveal that *P. wrangeli* consists of an alternation of distinct calcium phosphate (apatite) lamellae and originally organic-rich inter-layers. The lamellae form four semicircular marginal pillars, which are connected by irregularly spaced transverse diaphragms. Marginally, the diaphragms and pillar lamellae are not connected to each other and thus do not form a closed conch structure. A non-mineralized or poorly mineralized external conch probably existed originally in *P. wrangeli* but is not preserved in the available material. *P. wrangeli* often co-occurs with conulariids in fossil-rich limestone with mudstone – wackestone lithologies. Based on the new data, *P. wrangeli* can be best interpreted as a poorly mineralized conulariid with a mud-sticking life habit. Here the new conulariid family Palaenigmaidae fam. nov. is proposed as the monotypic taxon for *P. wrangeli*.

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16 17 Abstract

18 *Palaenigma wrangeli* (Schmidt) is a finger-sized fossil with a tetroradiate conical skeleton; it
19 occurs as a rare component in fossiliferous Upper Ordovician strata of the eastern Baltic Basin
20 and is known exclusively from north Estonia. The systematic affinities and palaeoecology of *P.*
21 *wrangeli* remained questionable. Here, the available specimens of *P. wrangeli* have been
22 reexamined using environmental scanning electron microscopy and x-ray computed tomography.
23 Additionally, the elemental composition of the skeletal elements has been checked using energy
24 dispersive X-ray spectroscopy. The resulting 2D-, and 3D-scans reveal that *P. wrangeli* consists
25 of an alternation of distinct calcium phosphate (apatite) lamellae and originally organic-rich
26 inter-layers. The lamellae form four semicircular marginal pillars, which are connected by
27 irregularly spaced transverse diaphragms. Marginally, the diaphragms and pillar lamellae are not
28 connected to each other and thus do not form a closed conch structure. A non-mineralized or
29 poorly mineralized external conch probably existed originally in *P. wrangeli* but is not preserved
30 in the available material. *P. wrangeli* often co-occurs with conulariids in fossil-rich limestone
31 with mudstone – wackestone lithologies. Based on the new data, *P. wrangeli* can be best
32 interpreted as a poorly mineralized conularioid with a mud-sticking life habit. Here the new
33 conularioid family Palaenigmaidae fam. nov. is proposed as the monotypic taxon for *P.*
34 *wrangeli*.

35 36 Introduction

37 The fossil *Palaenigma wrangeli* (Schmidt, 1874) captivates. It is small, less than a small finger
38 in diameter and no more than a couple of centimeters long. *P. wrangeli* has a beautiful
39 tetroradiate symmetry with four, strange horn-like spines or pillars at each corner, and it consists

40 of a shiny, dark-brown calcium phosphate, which cannot be overlooked on a freshly broken
41 Ordovician limestone.

42 The species name refers to Wilhelm F. Baron von Wrangell (1831–1894) (son of the famous
43 seaman Ferdinand von Wrangel), who found this fossil not far from his manor house when he
44 was a young man only to urge twenty years later the Geologist Friedrich K. Schmidt (1832–
45 1908) to solve its mystery. In his original description of the fossil, Schmidt (1874) reported the
46 difficulties in finding more material. It took him two years and hours of focused searching to find
47 another good specimen in a small quarry, where Wrangell guided him and his younger Swedish
48 colleague Jonas. G. O. Linnarson (1841–1881). The quarry exposed the Lyckholmsche Schicht
49 (corresponding to the Nabala and Vormsi Regional stages) and, according to Schmidt (1874),
50 was very rich in conulariids. The dark phosphatic shell of conulariids and the skeleton of *P.*
51 *wrangeli* stand out in the greenish-pale limestone, and if *P. wrangeli* were abundant, it would
52 have been easy to be found by the experienced and dedicated fossil hunters.

53 *P. wrangeli* is generally a rare fossil known exclusively from Estonia, and from Pleistocene
54 erratic blocks from the Åland Islands, Finland and Uppland, Sweden (Holm, 1893). In the
55 palaeontological collections of Estonia only seven specimens have been accumulated until now.
56 The specimen found by Wrangell and the four or so, original specimens collected by Schmidt
57 and Linnarson are unfortunately lost. Two specimens, probably collected by Linnarson, are in the
58 collections of the Naturhistorisk Riksmuseet Stockholm (Sweden). All come from north Estonian
59 light-coloured Upper Ordovician limestone, which is generally poor in skeletal intraclasts (Fig.
60 1).

61 Schmidt (1874) couldn't solve the mystery of *Palaenigma*, for which he created the separate
62 genus *Tetradium*, a name that was already preoccupied by a coral (see Walcott, 1886). He
63 speculated that it could be an operculum of a conulariid. Walcott (1886, p. 224) compared it with
64 the Cambrian calcitic polyplacophoran *Matheva* Walcott. Before, Lindström (1884, p. 41), in his
65 opus magnum on Silurian gastropods of Gotland, excluded any relation with mollusks and
66 curiously suggested that it might be a conulariid infected by a parasitic fungus. Later, Sinclair
67 (1952) placed *Palaenigma* without comment into the Conulariinae, a family of the Conulariida.
68 The conulariid affinities of *P. wrangeli* also appeared unquestionable for Brood (1995), who
69 briefly described the species and interpreted it as a basal part of *Conularia* Sowerby (Brood,
70 1995). The genus, however, does not appear in the Treatise of Invertebrate Palaeontology
71 (Moore & Harrington, 1956) and it was not included in the review and cladistic analysis of the
72 Conulariinae carried out by De Moraes Leme et al. (2008).

73 New finds from a small quarry in central Estonia exposing the Saunja Formation (Nabala
74 Regional Stage) give an opportunity to take the mysterious species under a new scrutiny using
75 modern analytical techniques. Here we describe the new material and review existing specimens
76 available from the Estonian geoscience data platform SARV), and the ~~Naturhistorisk~~
77 Riksmuseet Stockholm (NRM). SARV unites the large palaeontological collections from
78 Estonia, housed at the Department of Geology at Tallinn University of Technology (GIT), the
79 Natural History Museum, University of Tartu, and the Estonian Museum of Natural History.

80

81 **Methods**

82 The specimens were investigated with a GE phoenix vltomelx s X-ray computed tomography
 83 (CT) device at the Geological Survey of Finland in Espoo, Finland. The samples were imaged
 84 using an accelerating voltage of 80-100 kV and a tube current of 120-220 μA , for a tube power
 85 of 12-22 W. Tube power was kept low enough to avoid spot size – related blurring for the
 86 obtained resolutions of 12-20 μm . 0.1 mm of Cu was used as a beam filter in most scans. 2200-
 87 2500 angle steps were used and at each angle the detector waited for a single exposure time and
 88 then took an average over three exposures, with the single exposure time varying between 500-
 89 1000 ms. This resulted in total scan times of 73-167 minutes. The obtained projections were
 90 reconstructed using GE phoenix datoslx and investigated using ThermoFisher PerGeos 2020.2.
 91 The specimens have also been subjected for microstructural observation using the environmental
 92 scanning electron microscope (ESEM) FEI Quanta 250 housed at the Institute of Earth Sciences
 93 in Sosnowiec, Poland. The specimens have been inspected in uncoated states in low vacuum
 94 conditions using back-scattered (BSE) imaging. Both transverse and longitudinal sections of the
 95 specimens have been investigated. The elemental composition of building structures and layers
 96 have been checked using an energy dispersive X-ray spectroscopy (EDS) detector. ESEM and
 97 CT images have been graphically improved by adjusting whole image Gamma and Contrast
 98 levels using Affinity Photo Version 1.9.2 graphical software. Herein, a few descriptive terms are
 99 used, which are mainly borrowed from the literature about conulariids: Periderm denotes the
 100 exoskeleton of conulariids. Carinae are broad, internal thickenings of the periderm that can be
 101 situated on the sides of the periderm or as keel-like, continuous thickenings at the corners of the
 102 periderm. In many conulariids there are multiple kinds of internal thickenings, collectively
 103 assigned by Bischoff (1978) and Van Iten (1992) to 11 types of internal midline (interradial)
 104 structures and two types of internal corner (perradial) structures. Septa are longitudinal walls,
 105 keels, and deep ridges in the interior of the periderm positioned at the midline. Diaphragms are
 106 horizontal truncations of the periderm well above the apex. At the position of a diaphragm the
 107 periderm tapers to an imperforate, usually adapically convex transverse wall, sometimes also
 108 called the “apical wall”, or “schott” Van Iten (1991).

109 The compilation of conulariid specimens is based on a search in the SARV database (accessed
 110 08.04.2021,

111 [http://geocollections.info/specimen?specimen_number_1=1&specimen_number=&collection_id_1=1&collection_id=&classification_1=2&classification=&taxon_1=2&taxon=conulari&name_geology_1=1&name_geology=&country_1=1&country=&locality_1=1&locality=&stratigraphy_1=11&stratigraphy=Haljala%20Stage&id_1=5&id=&depth_since_1=12&depth_since=&depth_to_1=13&depth_to=&agent_1=1&agent=&reference_1=1&reference=&original_type_1=1&original_type=&part_1=1&part=&date_taken_since_1=12&date_taken_since=&date_taken_to_1=13&date_taken_to=&db%5B%5D=1&db%5B%5D=2&db%5B%5D=3¤tTable=specimen&maxSize=5&page=1&paginateBy=25&sort=locality_locality_en&sortdir=DESC\)](http://geocollections.info/specimen?specimen_number_1=1&specimen_number=&collection_id_1=1&collection_id=&classification_1=2&classification=&taxon_1=2&taxon=conulari&name_geology_1=1&name_geology=&country_1=1&country=&locality_1=1&locality=&stratigraphy_1=11&stratigraphy=Haljala%20Stage&id_1=5&id=&depth_since_1=12&depth_since=&depth_to_1=13&depth_to=&agent_1=1&agent=&reference_1=1&reference=&original_type_1=1&original_type=&part_1=1&part=&date_taken_since_1=12&date_taken_since=&date_taken_to_1=13&date_taken_to=&db%5B%5D=1&db%5B%5D=2&db%5B%5D=3¤tTable=specimen&maxSize=5&page=1&paginateBy=25&sort=locality_locality_en&sortdir=DESC)

119 The electronic version of this article in Portable Document Format (PDF) will represent a
120 published work according to the International Commission on Zoological Nomenclature (ICZN),
121 and hence the new names contained in the electronic version are effectively published under that
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125 through any standard web browser by appending the LSID to the prefix <http://zoobank.org/>. The
126 LSID for this publication is: urn:lsid:zoobank.org:act:3E1F98CB-C7DD-458B-AC45-
127 FE366030DC02. The online version of this work is archived and available from the following
128 digital repositories: PeerJ, PubMed Central and CLOCKSS. The Supplemental Information (S1–
129 S5) for this article is available at www.zenodo.org/doi/10.23729/3eaf1aeb-5e0c-4704-9e18-a925688f810b and the
130 complete set of CT-figures and videos made for this publication is available at
131 <https://www.doi.org/10.23729/3eaf1aeb-5e0c-4704-9e18-a925688f810b>.
132

133 Geological setting

134 All specimens of *P. wrangeli* described herein have been collected from localities in north
135 Estonia, exposing Upper Ordovician strata either in natural outcrops or in drill cores (Fig. 1A).
136 The sediments of north Estonia are tectonically nearly undisturbed and palaeogeographically
137 represent the eastern part of the Baltic Palaeobasin of the Baltica Palaeocontinent (Männil, 1966;
138 Jaanusson, 1979; Nestor and Einasto, 1997). During the Late Ordovician the sedimentary
139 deposition in north Estonia was dominated by limestone and marlstone in temperate to tropical
140 marine settings (Cocks and Torsvik, 2005; Dronov and Rozhnov, 2007). The area comprises the
141 North Estonian Facies Belt or North Estonian Shelf which, toward the south, grades into the
142 Livonian Basin (Jaanusson 1979; Nestor & Einasto, 1997, Fig. 1B). The sediments of the North
143 Estonian Shelf are predominantly neritic to shallow marine and individual sedimentary packages
144 are locally divided by long depositional hiatus and partially by erosional horizons (Raukas and
145 Teedumäe, 1997). A well-established regional chronostratigraphic, lithostratigraphic and
146 biostratigraphic scheme allows for high resolution correlation of the north Estonian Upper
147 Ordovician sediments (e.g., Raukas & Teedumäe, 1997; Nõlvak et al., 2006; Meidla et al., 2014,
148 Fig. 1C).

149 Material

151 The type locality of *P. wrangeli* was provided by Schmidt (1874) as a quarry belonging to Küti
152 (German “Kurküll”), a manor near Viru-Jaagupi in northeastern Estonia. According to (Schmidt,
153 1858) the quarry was located south-west of manor house near Aruküla village (German
154 “Arroküll”). The quarry was abandoned a long time ago and today is untraceable, its former
155 location is indicated by the place name Lubjaahju (Estonian for Lime Kiln) (59°11'43.2"N
156 26°30'00.4"E) It exposed a pale-grey limestone of the Nabala and Vormsi stages (Rõõmusoks,
157 1966). The quarry was repeatedly visited by Schmidt (1858, 1874) because of its fossil richness.
158 The abundance of conulariids was specifically mentioned and listed by Schmidt (1858, 1874)

159 and is also documented in the SARV by an impressive number of more than 60 conularid
160 specimens from the old Küti quarry. According to Rõõmusoks (1966) the richness is mainly
161 limited to the Vormsi Stage; he listed brachiopods (mainly *Sampo hiiuensis* (Öpik), *Ilmarinia*
162 *sinuata* (Pahlen), *Kiaeromena* (*Bekkeromena*) *vormsina* Rõõmusoks), hyoliths, gastropods,
163 heliolitid tabulates, rugose corals, receptaculitids, and trilobites (mainly *Toxochasmops*
164 *vormsiensis* Rõõmusoks). Two specimens of *P. wrangeli* from Küti are available from the
165 collections of the NRM (NRM-Mo 153045, 153046). The lithological information available from
166 matrix of these specimens is consistent with an origin from the Kõrgessaare Formation, Vormsi
167 Stage. The Kõrgessaare Formation consists of an argillaceous, heavily bioturbated, greenish to
168 yellowish pale-coloured mud-wackestone (Oraspõld and Kala, 1980).

169 Two specimens (Department of Geology at Tallinn University of Technology, GIT 812-34, GIT
170 612-35) were collected at the Sutlema quarry, west of Sutlema village, Rapla County, central
171 Estonia (59°10'26.28"N, 24°37'2.62"E). The active quarry exposes the Saunja Formation (Nabala
172 Stage) and the Kõrgessaare Formation (Vormsi Stage). Both specimens came from the Saunja
173 Formation. At Sutlema the Saunja Formation contains a rich fauna and flora, dominated by green
174 algae (*Vermiporella* Stolley, *Coelosphaeridium* Roemer, and an unidentified delicate dendroid
175 form), gastropods (large *Murchisonia*-like forms, *Hormotoma insignis* Eichwald) and sponges.
176 Additionally, bivalves, brachiopods [*Kiaeromena* (*Bekkeromena*) *ilmari* Rõõmusoks],
177 cephalopods, conulariids, receptaculitids, rugose corals, trilobites, stromatoporoids, and dendritic
178 graptolites (*Dictyonema* sp.) occur. The rich fauna of the quarry needs a detailed taxonomic
179 examination. The Saunja Formation is more than 10 m thick at Sutlema and consists of a
180 bioturbated, massively bedded, light-colored mud-wackestone, typical for the Baltic Limestone
181 Facies (Kröger et al., 2019).

182 Three additional specimens come from drillcores, with little information on co-occurring fauna:
183 Specimen GIT 655-1 was collected from Kükita 24 drillcore (58°48'18.9"N 26°56'32.5"E), c. 4
184 km south of Mustvee, Mustvee Parish, west of Lake Peipsi, north-east Estonia, from depth 84.35
185 m, Tudulinna Formation, Vormsi Stage. The faunal content of the Vormsi interval of the
186 drillcore is remarkably rich and comprises a delicate dendroid bryozoa (*Stictopora* sp.), a
187 trilobite (*Isotelus* sp.), a hyolithid (*Dorsolinevitus vomer* Holm), a conulariid, and the putative
188 cnidarian *Sphenothallus* (Vinn and Kirsimäe, 2014).

189 Specimen GIT 655-2 was collected from Ellavere drillcore (59° 0'52.42"N, 26° 1'24.89"E), c. 8
190 km south-east-east from Järva-Jaani, Järva County, north-east Estonia, depth 92.70 m, from the
191 Nabala Stage. At the same horizon occurs a bellerophontid [*Megalomphala crassa* (Koken)], and
192 a brachiopod [*Cyrtonotella kuckersiana* cf. *kuckersiana* (Wysogorski)]. The specimen GIT 655-2
193 occurs in a greenish - grey, bioturbated argillaceous skeletal of uncertain stratigraphy.

194 Specimen GIT 655-3 was collected from Mustvee 2322 drillcore, 3 km west of Mustvee, a
195 village at the shore of the Lake Peipsi, north-east Estonia (58°50'5.41"N, 26°53'19.79"E), depth
196 69.15 m, from an interval within the Pirgu Stage. It occurs in a greenish gray, bioturbated,
197 nodular argillaceous limestone of the Adila Formation.

198 Two specimens have been detected in the collections after completion of the analyses for this
199 review: Specimen GIT 575-43, from Mäemetsa Quarry, Harju County, Saunja Formation,
200 Nabala Stage; and specimen GIT 655-4, from Pala 70 drillcore at 143.90 m, Jõgeva County,
201 Pirgu Stage.

202

203 **Results**

204 **Morphology**

205 The available specimens show some generalities in skeletal morphology. All specimens consist
206 of four equidistant marginal pillars with diameters of up to 3 mm. The distance of the pillars
207 increases at a constant angle of c. 13° toward a maximum preserved periderm diameter of c. 10–
208 11 mm (Specimen GIT 612-34, Fig. 2). The four pillars are apically interconnected by irregularly
209 spaced transverse diaphragms, which are slightly irregularly curved toward the apex of the pillars
210 (Fig. 3). The pillars have a roughly semicircular cross section, which results from a relatively
211 loose and irregular cone-in-cone succession of superimposed tubular shell layers exclusively on
212 the inner side of the pillars (Fig. 3B–C, Fig. 4C, Videos S1–S2). The centers of the outer surface
213 of the pillars are not covered with a continuous shell layer but expose, as a quasi-cross section,
214 the complete succession of layers (Figs 3C–D). This results in a longitudinally carinate
215 appearance of the outer surface of the pillars.

216 The diaphragms are continuations of individual pillar layers or sheets, with the oldest and
217 apicalmost diaphragms representing the most distal, oldest pillar layers. The thickness of the
218 diaphragms is similar to that of the laminae of the pillars, c. 10–80 μm . The shape of the
219 diaphragms can be deeply conically curved, such as in specimen Mo 153045 (Fig. 4), or shallow
220 bowl-shaped, such as in specimen GIT 655-3 (Fig. 5).

221 The transverse shape of the septa is nearly quadratically and the pillars are positioned at or near
222 the four centers of the square margins, which would correspond to the midlines of the four
223 periderm-faces of a conulariid (Fig. 4C, D).

224 The height of the individual cones of the pillars is more than what is preserved in the available
225 specimens and thus exceeds 15 mm. Hence, the skeletal material accreted in form of clearly
226 distinguishable, separate layers or sheets from the outer margins of the conch toward its center.
227 The apical end of the skeleton is open, and the pillars are not in contact with each other at their
228 apical tip. The first septum occurs at a face width of 6 mm in specimen GIT 612-34 and at a face
229 width of 8.5 mm in specimen Mo 153045.

230 In specimen GIT 655-3 the pillars are additionally thickened by massive flange-like skeletal
231 sheets, which merge toward the periderm center with thick diaphragms (Fig. 5). A similarly
232 thickened pillar section is preserved in specimen Mo 153045 (Figs 4A, B).

233 Notably in specimen Mo 153045, GIT 655-1, and GIT 655-3 skeletal fragments of thin cone
234 shaped skeletal sheets or walls with a fragile lattice-like texture are preserved in close proximity
235 of *P. wrangeli* (Figs 6A–C). These sheets in specimen Mo 153045 are longitudinally bent or
236 folded forming sharp angles and flat faces. The shape of the sheets and the lattice-like structure
237 is similar to co-occurring conulariid periderms (Fig. 6C, Video S3). In specimen GIT 655-1,

238 fragments of a finely transversely annulated or ribbed sheet are preserved near the outer margin
239 of a pillar (Figs 7A, D).

240

241 **Microstructure**

242 ESEM observations of *P. wrangeli* reveal that different parts of the skeleton have a similar
243 microstructure, consisting of several distinct thin, phosphatic (fluoroapatite as evidenced from
244 EDS, see Article S4) lamellae (Figs 7). This is similar as in conulariids and *Sphenothallus*
245 (Brood, 1995; Vinn & Kirsimäe, 2015; Ford et al., 2016). The thickness of particular lamellae
246 varies, ranging from c. 10 to 80 μm . As in *Sphenothallus*, the boundaries between particular
247 lamellae may be more or less sharp. In several places, within a single lamella much thinner (c.
248 0.5 to 0.8 μm thick) laminae occur which may mark here a primary lamination. The
249 microstructure of particular lamellae seems to be homogeneous, composed of tiny phosphate
250 crystals. Sometimes, however, within particular solid lamellae, empty spaces may occur (Figs
251 7G–I). Such spaces have a limited extent and are filled by microcrystalline calcium phosphate. In
252 some areas the interspaces between successive laminae contain phosphatic aggregations of thin
253 (c. 1.5–1.8 μm in diameter), branching and diverging filaments. Some of the laminae also
254 possess pores and empty chimney-like structures, with an inner diameter up to 4 μm (Fig. 7H).
255 In the distalmost part of the skeleton of the specimen GIT 655-2 an extremely thin (up to 1 μm)
256 outermost layer occurs, on which tiny (c. 16 μm in diameter), circular bumps (papillae) occur
257 (Figs 7B, E, F). These structures may be isolated or associated in small groups. In some places,
258 additionally smaller wrinkle-like structures (shrinkage features?) also occur (Fig. 7F). The
259 wrinkled and papillate layer is covered from the inside by homogeneous skeletal layers devoid of
260 such structures.

261

262 **Discussion**

263 **Interpretation of the shell microstructure**

264 The distinct lamellae of the skeleton of *P. wrangeli* partly contain fine irregular vertical
265 perforations, chimney-like structures (Fig. 7H), and additionally in some places the lamellae-
266 interspaces are filled with a layer of fine filamentous phosphatic aggregations (Fig. 7H–I). The
267 chimney-like structures may be interpreted as original pore-like anatomical structures, because
268 these structures appear to be limited to the papillate layer and there the shell lamina are often
269 deflected toward the perforations. If so, it can be hypothesized that the papillate thin layer in
270 specimen GIT 655-2 represents the remnants of the inner side of an external covering ‘periderm’.
271 However, the filled perforations in other areas of the shell are less regular, and the presence of
272 filamentous micro-apatitic aggregations (Fig. 7I, Article S4) in some of the lamellae-interlayers
273 can be best interpreted as a product of microbial and fungal degradation of originally organic-
274 rich laminae. Hence these structures may indicate the presence of originally organic layers in
275 between the phosphatic lamellae, which were post-mortem infested by boring microbial-fungal
276 consortia. Similar, alternating phosphatic-organic layers also occur in skeletons of conulariids
277 (Ford et al. 2016) and *Sphenothallus* (Vinn & Kirsimäe, 2015; Vinn and Mironenko, 2020).

278

279 Systematic affinities

280 The name giving enigma of *P. wrangeli* has two aspects: the first refers to the anatomical
281 interpretation of the skeletal structures, and the second one, which relates to the first one, refers
282 to its systematic affinity. Both mysteries can be partly solved with the new evidence from the
283 examinations performed herein.

284 The preserved skeletal parts of *P. wrangeli* are invariably composed of calcium phosphate
285 (presumably of apatite, such as in conulariids and *Sphenothallus*, Vinn & Kirsimäe, 2015). The
286 3D-reconstruction and ESEM examination of several well-preserved specimens reveals a
287 consistent tetradial symmetry of the *P. wrangeli* skeleton with four semicircular marginal
288 pillars, which are connected by irregularly spaced transverse diaphragms and which form a cone-
289 like skeleton with an angle of c. 13°. The pillars and diaphragms are formed by a cone-in-cone
290 structure of distinct sheets, which are accreted from the outer margin of the entire structure
291 toward the center and from the apex toward the opening of the cone. Marginally, the diaphragms
292 are not connected to each other except at the position of the pillars and thus do not form a closed
293 structure. Similarly, the pillar-layers are open toward its margin and end abruptly at the outer
294 surface of the pillars, resulting in a semicircular pillar cross-section and in a peculiar
295 longitudinally lirated relief of the external pillar surface. The abrupt ending of the skeletal sheets
296 at the margins of the diaphragms and at the external surfaces of the pillars suggests the presence
297 of an organic outer cover or periderm, which is not fossilized, and which served as an attachment
298 structure and matrix for the formation of diaphragms and the external pillar surface.

299 In summary, the skeleton of *P. wrangeli* shares a number of crucial characters, known in its
300 combination only in the Conulariina: 1) the skeleton composed of calcium phosphate, 2)
301 tetradial, slender cone with thickened longitudinal septa at midline position and transverse
302 diaphragms, 3) skeletal sheets forming irregularly and loosely spaced cone-in-cone structures.
303 The poorly preserved, lightly mineralized phosphatic, transversely ornamented walls could be
304 interpreted remains of a periderm. Therefore, *P. wrangeli* can be best interpreted as a conulariid
305 with poorly mineralized marginal conch walls (periderm), phosphatic apical pillars and
306 diaphragms. The pillars with their flat external surfaces can be best interpreted as homologue to
307 the mineralized longitudinal septa at midline position in the Conulariina (see e.g., Ford et al.,
308 2016; de Morales Leme et al., 2008).

309 Taking the general similarities and distinct constructional differences into account, it is evident
310 that *P. wrangeli* should be placed to a separate conulariid family. Here we suggest the new
311 family Palaenigmaidae fam. nov. for Conulariina with steeply pyramidal skeletons with a thin
312 chitinophosphatic periderm that consist of four equidistant marginal pillars, without or with
313 poorly biomineralized outer shell; the apical end of the skeleton is open, and the pillars are not in
314 contact with each other at their apical tip. *P. wrangeli* is the only species of the Palaenigmaidae
315 fam. nov.

316

317 Palaeoecology

318 In his original description, Friedrich Schmidt noticed the extraordinary abundance of co-
319 occurring conulariids with specimens of *P. wrangeli* in the type locality of Kūti, north-east
320 Estonia (Schmidt, 1874). A co-occurrence of *P. wrangeli* with conulariids was described from
321 Baltic Limestone boulders from Sweden (Holm, 1893). And conulariids are also relatively
322 common in the Sutlema quarry, where two specimens of *P. wrangeli* have been found, as well
323 (see above).

324 The compilation of conulariids in the SARV database allows for an investigation of the question
325 whether this co-occurrence of *P. wrangeli* with conulariids represents a general pattern.

326 Conulariids inhabited the eastern part of Baltica basin from the Darriwilian (Kunda Stage)
327 onwards throughout the Silurian. They reached their Ordovician abundance climax within the
328 Haljala Stage with 127 specimens in the collections from 18 different localities. A second
329 abundance peak was reached during the Vormsi Stage, from which 51 specimens from seven
330 different localities are known (Fig. 8). Most of the known specimens of *P. wrangeli*, including
331 the type specimens, are also from the Vormsi Stage. This seems to support the idea that
332 conulariids and *P. wrangeli* shared general habitat preferences and /or preservation pattern.
333 Based on the specimens available for this study, *P. wrangeli* and the Late Ordovician conulariids
334 of the eastern Baltica basin occur preferentially in depositional settings within an originally
335 extraordinarily faunal-rich, calcareous soft substrate habitat.

336 Neither the extreme apices of *P. wrangeli*, nor that of co-occurring conulariids are known.
337 Firmly skeletonized apical holdfast structures occur in Late Ordovician conulariids (Kozłowski,
338 1968; Brood, 995). These holdfasts are discoidal or rootlet-like, indicating differentiated
339 conulariid attachment on hard substrate (discoids) and soft substrate (rootlets). Rootlet-like
340 skeletal appendages are often interpreted as functioning for stabilization and attachment within
341 soft substrate (e.g., Kozłowski, 1968; Seilacher & Macclintock, 2005). Additionally, elongated,
342 stick-like conch forms often occur in mud-sticking bivalves, such as *Pinna* Linnaeus, which
343 shares even more similarities with conulariids in having a subquadratic conch cross section (see
344 e.g., Seilacher, 1984). A mud-sticking original life habit of *P. wrangeli* is therefore highly
345 probable.

346

347 **Conclusions**

348 *Paleaenigma wrangeli* (Schmidt, 1874) is a rare fossil known from few specimens collected
349 from Upper Ordovician limestone outcrops across northern and central Estonia and from erratic
350 boulders in Finland and east central Sweden. The systematic affinities of the monotypic
351 *Paleaenigma* were disputed. A thorough analysis of well-preserved specimens with X-ray
352 computed tomography, scanning electron microscopy, and energy dispersive X-ray spectroscopy
353 reveal that the skeleton of *P. wrangeli* is composed of distinct calcium phosphate (apatite)
354 lamellae. The lamellae are partly porous and ornamented with distinct papillae and contain
355 poorly mineralized interlayers. The skeleton consists of four pillars, which are connected by
356 irregularly spaced diaphragms and which are marginally open. The diaphragms are quadratic in
357 transverse view and the pillars are situated at the four sides of the diaphragm squares. In few

358 specimens remains of thin, poorly preserved transversally ornamented apatitic tube-forming
359 walls are preserved near the distal margins of the pillars. Therefore, *P. wrangeli* can be best
360 interpreted as a conulariid with poorly mineralized marginal conch walls (periderm), phosphatic
361 apical pillars at midline position, and diaphragms. The new monospecific family Palaenigmaidae
362 fam. nov. is proposed for *P. wrangeli*. Conulariids often co-occur with *P. wrangeli*. A
363 comparison of other conulariid occurrences in Estonia with *P. wrangeli* occurrences indicates
364 that these fossils are most abundant in depositional settings within an originally extraordinarily
365 faunal-rich, calcareous soft substrate habitat. Based on its general morphology *P. wrangeli* can
366 be interpreted as a poorly mineralized conulariid with a mud-sticking original life habit.

367

368 **Acknowledgements**

369 We are grateful to Mare Isakar (Tartu, Estonia) and Christian Skovsted (Stockholm, Sweden) for
370 help with finding material and early suggestions on how to proceed with the review of the
371 material. Anna Madison (Moscow, Russia) and Aleksey Sokolov, Vadim Glinskiy, and Galina
372 Gataulina (St. Petersburg, Russia) helped to gather information on the (missing) type material.
373 Gennady Baranov (Tallinn, Estonia) and Duncan Matthews (Helsinki, Finland) were supportive
374 and important companions during field trips.

375

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Figure 1

Occurrences of *Palaenigma wrangeli* (Schmidt, 1874) in north Estonia

(A) Map of Baltoscandia with national boundaries and capitals (black dots). (B) Map of Estonia with *P. wrangeli* occurrences discussed herein (yellow dots), and with outline of Late Ordovician facies belts (from Harris et al., 2004). (C) Middle - Late Ordovician Regional stages of Baltoscandia (stars mark occurrences of *P. wrangeli*. Hirn., Hirnantian. Map data: R Package „maps” Version 3.3.0 (<https://cran.r-project.org/web/packages/maps/maps.pdf>).

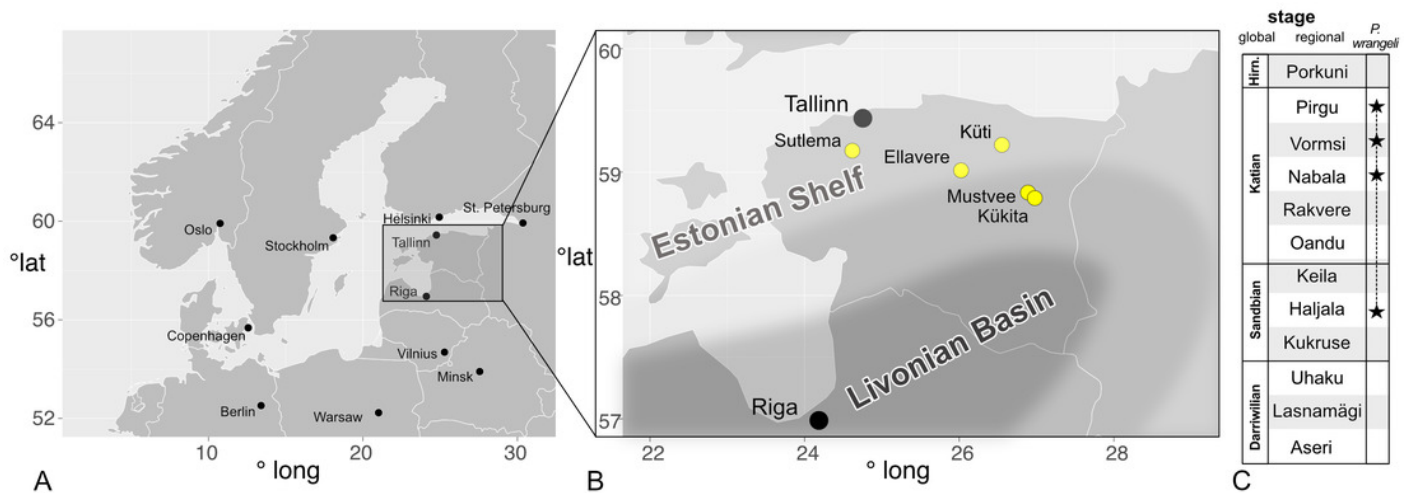


Figure 2

X-ray computed tomography image of *Palaenigma wrangeli* (Schmidt, 1874), specimen GIT 612-34, from Sutlema quarry, west of Sutlema village, Rapla County, Estonia, Nabala Stage.

(A) Lateral view. B. Lateral view 90° rotated along the growth axis relative to A. (C) Lateral view parallel to two pillars. Scale applies to all figures.

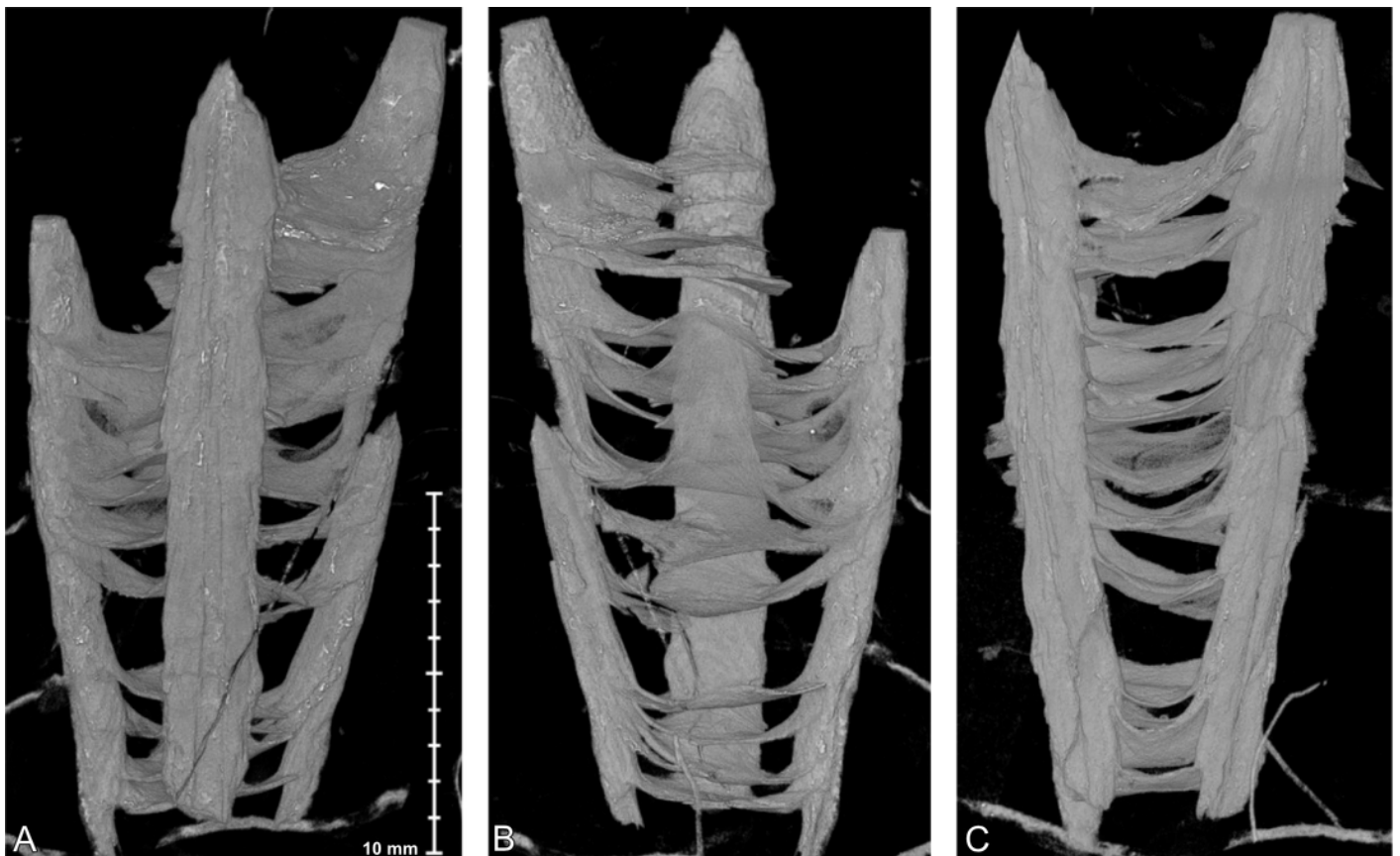


Figure 3

X-ray computed tomography cut of *Palaenigma wrangeli* (Schmidt, 1874)

(A), (B) Specimen GIT 612-34, from Sutlema quarry, west of Sutlema village, Rapla County, Estonia, Nabala Stage. A. Sagittal cut. B. Transverse cut. (C), (D) Specimen NRM-Mo 153045, from Küti quarry, near Viru-Jaagupi in northeastern Estonia, Vormsi Stage. C. Transverse cut. D. Sagittal cut. Note the irregular spacing of the diaphragms and the continuation of the diaphragm - pillar layers in A and D, and the open half-circle cross section shape of the pillars in C and D. Scale applies to all figures.

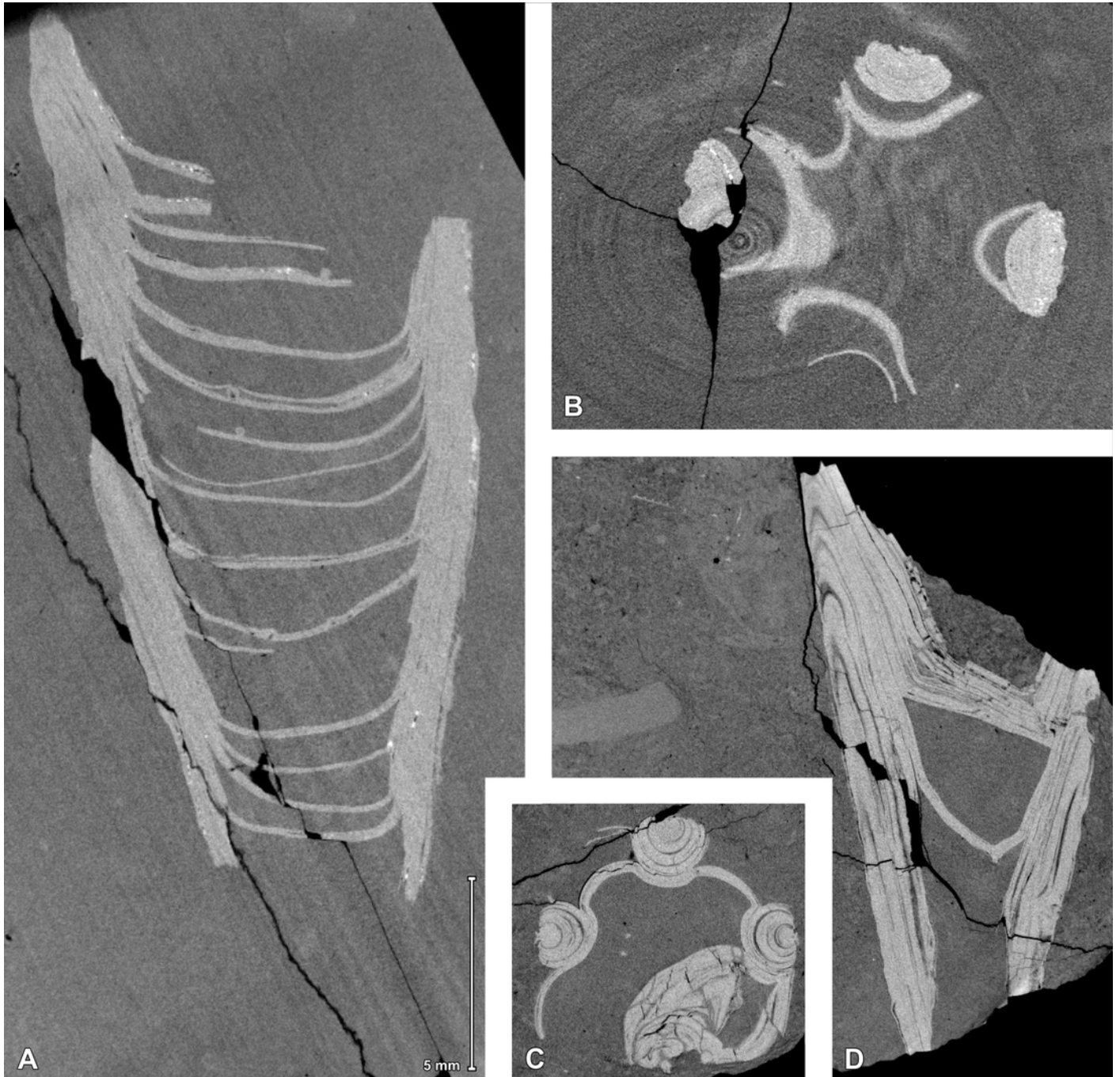


Figure 4

X-ray computed tomography image of *Palaenigma wrangeli* (Schmidt, 1874), specimen NRM-Mo 153045, from Küti quarry, near Viru-Jaagupi in northeastern Estonia, Vormsi Stage

(A) Lateral view, scale applies to A, (B) Lateral view 180° rotated along the growth axis relative to A. C. Adapical view, scale applies to C, D. (D) Adoral view.

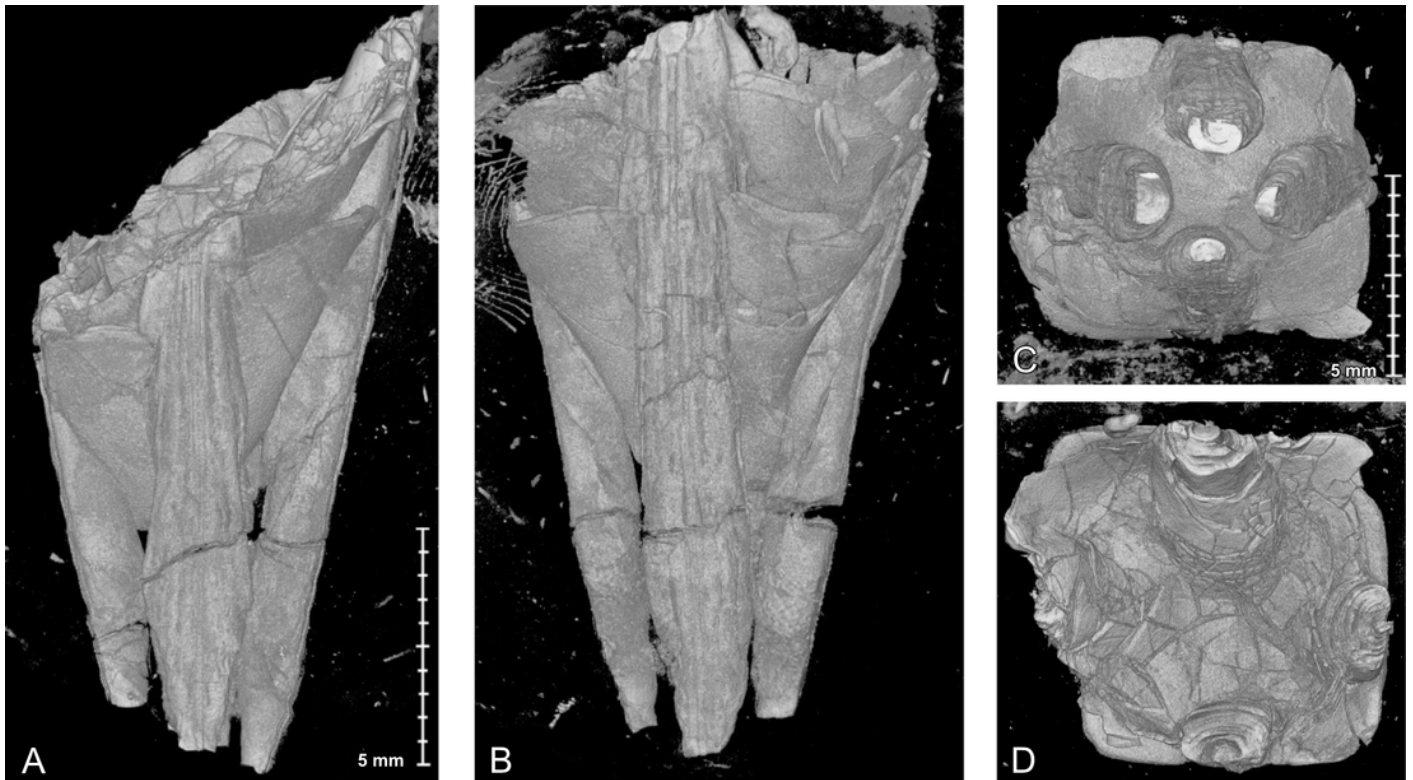


Figure 5

X-ray computed tomography image of *Palaenigma wrangeli* (Schmidt, 1874), specimen GIT 655-3, from Mustvee 2322 drillcore, west of Mustvee, north-east Estonia, Pirgu Stage

(A) Lateral view. (B) Lateral view 90° rotated along the growth axis relative to A. (C) Lateral view 180° rotated along the growth axis relative to A. Scale applies to all figures.

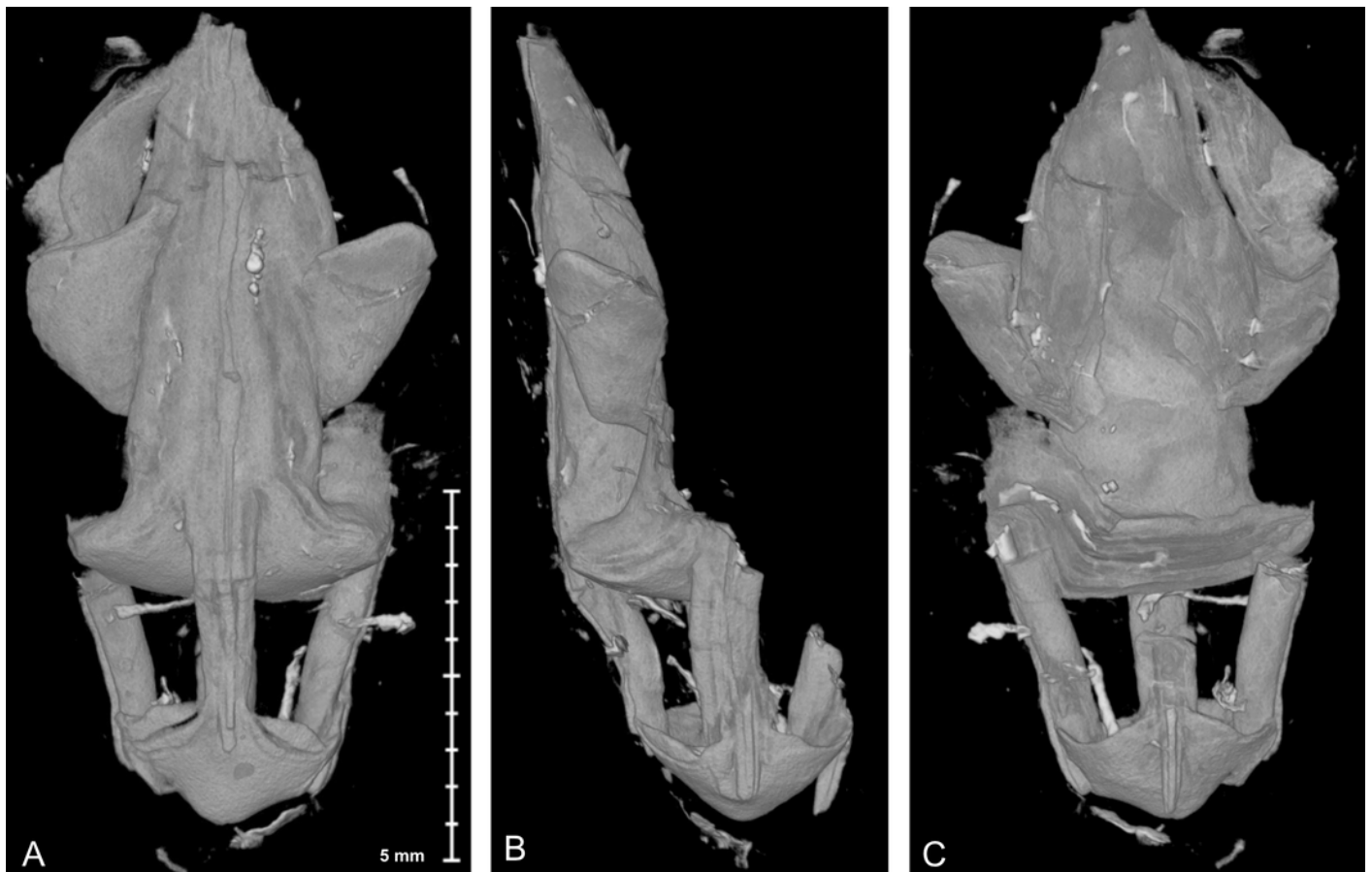


Figure 6

X-ray computed tomography image of *Palaenigma wrangeli* (Schmidt, 1874) and conulariid fragments

(A) Lateral view of specimen GIT 655-3 with unidentified skeletal debris in surrounding sediment matrix and fragment of conulariid (upper right). (B) Lateral view of specimen NRM-Mo 153045 with unidentified skeletal debris in surrounding sediment matrix and fragment of conulariid (upper right). (C) Lateral view of conulariid GIT 812-35-1, from Sutlema quarry, Nabala stage, note also the disc-like shadow of a crinoid ossicle in the sediment matrix.

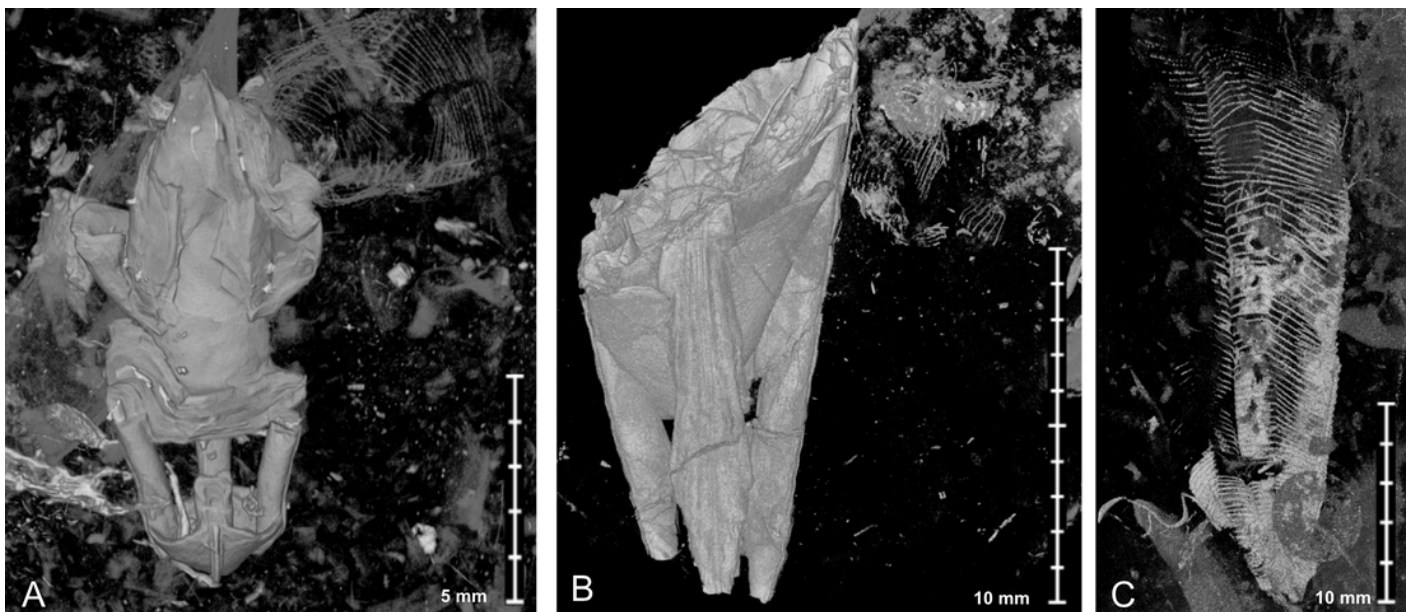


Figure 7

Scanning electron images of *Palaenigma wrangeli* (Schmidt, 1874)

(A), (D) Specimen GIT 655-1, from Kükita 24 drillcore, Mustvee Parish, north-east Estonia, Vormsi Stage, showing fragment of a transversally ornamented periderm? near external surface of a pillar. (B-C), (E-I) Specimen GIT 655-2, from Ellavere drillcore, Järva County, north-east Estonia, Nabala Stage. B, E, F. Area with distinctive papillate surface and with wrinkles (arrow in F). G-I. Details showing lamellate conch cross section with empty or filamentous interspaces (arrow in G). Note also the chimney like structures (arrow in H).

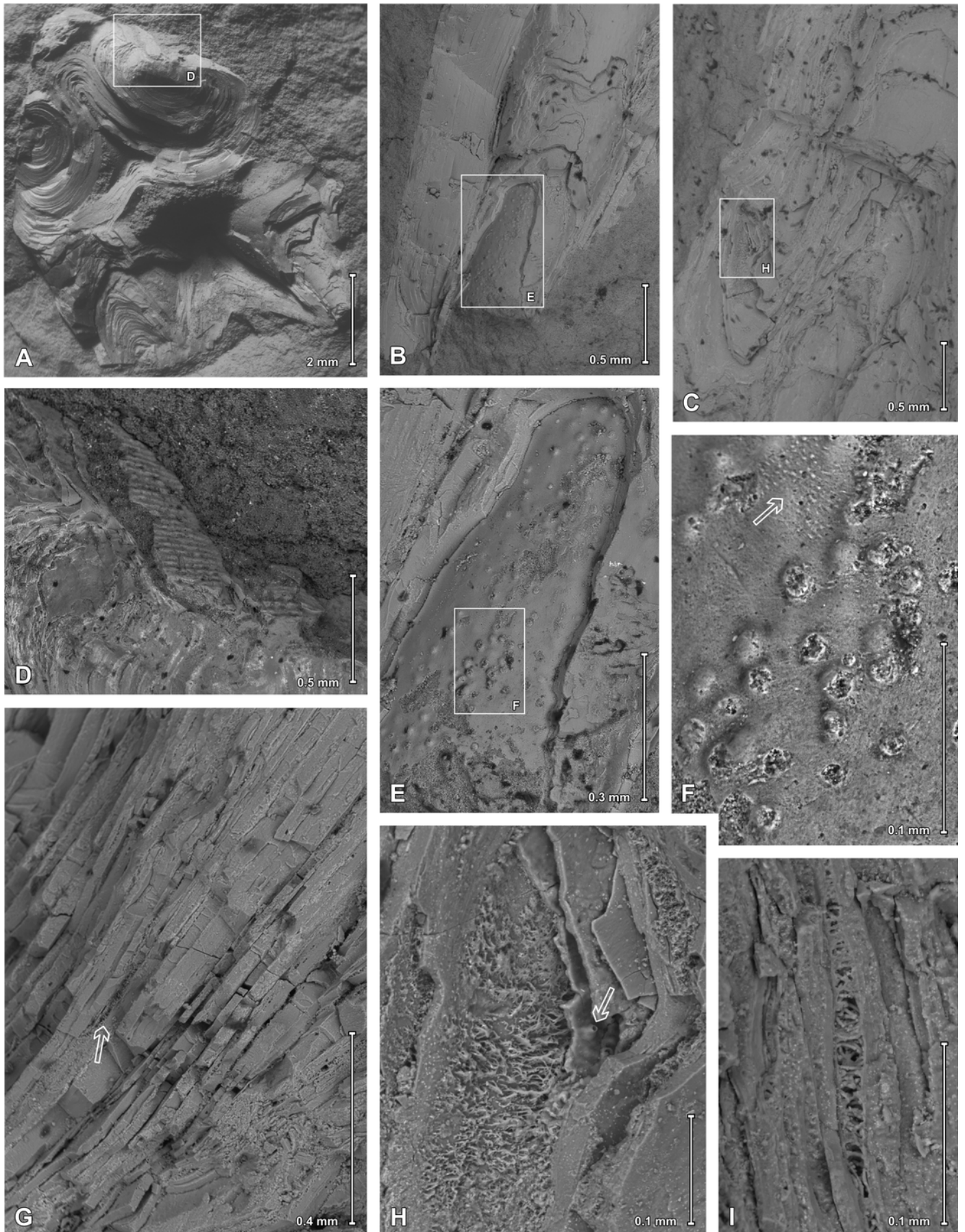


Figure 8

Abundance (= frequency of occurrences) of conulariids in Estonian Ordovician strata and stratigraphic occurrence of *Palaenigma wrangeli* (Schmidt, 1874)

Hirn., Hirnantian; n, number. Data: downloaded from SARV at 08.04.2021(see also Methods section and Data S5)

