

# The main stage of recovery after the end-Permian mass extinction: Taxonomic rediversification and ecologic reorganization of marine level-bottom communities during the Middle Triassic

Evelyn Friesenbichler<sup>Corresp., 1</sup>, Michael Hautmann<sup>1</sup>, Hugo Bucher<sup>1</sup>

<sup>1</sup> Paläontologisches Institut und Museum, University of Zurich, Zurich, Switzerland

Corresponding Author: Evelyn Friesenbichler  
Email address: evelyn.friesenbichler@pim.uzh.ch

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unimportant for the recovery of marine level-bottom communities, we note that their actual role can only be assessed when tested against predictions of the biotic model.

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Evelyn Friesenbichler<sup>1</sup>, Michael Hautmann<sup>1</sup>, Hugo Bucher<sup>1</sup>

<sup>1</sup> Paläontologisches Institut und Museum, University of Zurich, Zurich, Switzerland

Corresponding Author:

Evelyn Friesenbichler<sup>1</sup>

Karl-Schmid-Strasse 4, Zurich, 8006, Switzerland

Email address: [evelyn.friesenbichler@pim.uzh.ch](mailto:evelyn.friesenbichler@pim.uzh.ch)

# Abstract

The recovery of marine life from the end-Permian mass extinction event provides a test-case for biodiversification models in general, but few studies have addressed this episode in its full length and ecological context. This study analyses the recovery of marine level-bottom communities from the end-Permian mass extinction event over a period of 15 Ma, with a main focus on the previously neglected main phase during the Middle Triassic. Our analyses are based on faunas from 37 lithological units representing different environmental settings, ranging from lagoons to inner, mid- and outer ramps. Our dataset comprises 1562 species, which belong to 13 higher taxa and 12 ecological guilds. The diversification pattern of most taxa and guilds shows an initial Early Triassic lag phase that is followed by a hyperbolic diversity increase during the Bithynian (early middle Anisian) and became damped later in the Middle Triassic. The hyperbolic diversity increase is not predicted by models that suggest environmental causes for the initial lag phase. We therefore advocate a model in which diversification is primarily driven by the intensity of biotic interactions. Accordingly, the Early Triassic lag phase represents the time when the reduced species richness in the wake of the end-Permian mass extinction was insufficient for stimulating major diversifications, whereas the Anisian main diversification event started when self-accelerating processes became effective and stopped when niche-crowding prevented further diversification. Biotic interactions that might drive this pattern include interspecific competition but also habitat construction, ecosystem engineering and new options for trophic relationships. The latter factors are discussed in the context of the resurgence of large carbonate platforms, which occurred simultaneously with the diversification of benthic communities. These did not only provide new hardground habitats for a variety of epifaunal taxa, but also new options for grazing gastropods that supposedly fed from microalgae growing on dasycladaceans and other macroalgae. Whereas we do not claim that changing environmental conditions were generally unimportant for the recovery of marine level-bottom communities, we note that their actual role can only be assessed when tested against predictions of the biotic model.

# Introduction

The end-Permian mass extinction was the most severe extinction event in Earth's history and witnessed the extinction of 81% (Stanley, 2016) to 96% (Raup, 1979) of marine species. The aftermath of this event provides a unique test case to study how life rediversified after such a massive loss of species richness and which evolutionary processes were involved. In this study, these processes were analyzed in the ecological context in which they occurred. Previous studies with a similar scope concentrated on the Early Triassic (e.g. Schubert & Bottjer, 1995; Twitchett & Wignall, 1996; Hofmann et al., 2013; Hofmann, Hautmann & Bucher, 2015; Foster et al., 2015; Hautmann et al., 2015; Foster et al. 2017; Foster, Danise & Twitchett 2017), the time interval of ca. 4 Ma after the end-Permian mass extinction. These studies revealed, among other things, that (1) the recovery of the nekton (e.g. ammonoids and conodonts; Orchard, 2007; Brayard et al., 2009; Brayard & Bucher, 2015) is remarkably different from that of the benthos (Hofmann, Hautmann & Bucher, 2013; Hofmann et al., 2013; Hofmann et al., 2014; Friesenbichler et al. 2019), (2) the main recovery phase of the benthos did not start before the Middle Triassic (Nützel, 2005; Hautmann, 2007; Hofmann et al., 2014; Friesenbichler et al., 2021), (3) benthic ecosystem recovery was not completed before the end of the Middle Triassic (Hausmann & Nützel, 2015) and (4) the resurgence of large carbonate platforms and the recovery of reefs during the Middle Triassic (e.g. Gaetani et al., 1981; Senowbari-Daryan et al., 1993;

*Flügel, 2002*) possibly played an important role in the shift of taxonomic composition of benthic communities (*Friesenbichler et al., 2019*). Despite the fact that the Middle Triassic was obviously a crucial time for the recovery of benthic marine communities, corresponding studies (e.g. *Payne, 2005; Payne et al., 2006; Song et al., 2011; Velledits et al., 2011; Foster & Sebe, 2017; Friesenbichler et al., 2019; Friesenbichler et al., 2021*) are still scarce. Relevant questions in the ecological-evolutionary context include: Was the delay in benthic recovery caused by adverse environmental conditions during the Early Triassic (e.g. *Hallam, 1991; Wignall & Twitchett, 1996; Pruss & Bottjer, 2004; Song et al., 2014*), or is it an intrinsic aspect of the diversification dynamics in a largely vacated ecospace (*Hautmann et al., 2015*)? How did community structures, ecological guilds and ecosystems change during the Middle Triassic, and how did these changes translate into observed diversification patterns? Did the resurgence of large carbonate platforms affect the guild structure, and if yes, which evolutionary processes were involved in the colonization of this new habitat type? What was the relative role of biotic processes such as interspecific competition, habitat alteration, ecosystem engineering and niche construction versus abiotic environmental conditions as controlling factors in diversification and ecosystem evolution? To which diversification model do the empirical patterns conform, and which conclusions does this imply for the diversification of life in general?

## Materials & Methods

This study is based on data of macroinvertebrate occurrences compiled at the species level from palaeontological studies on Early and Middle Triassic lithological units (*Data S1*). We preferred this approach over using data from the Palaeobiology Database because it allowed us to control the correctness of the data. The evaluated studies include primarily monographs with well documented faunal lists, but we also considered studies that provide detailed information about macroinvertebrate occurrences including taxonomic assignments. Prime criteria for our selection of studies are the quality of documentation and the coverage of time intervals and environments. Moreover, we sought to include the most diverse lithological unit of each time interval as representatives for the maximum diversity that was possible at a given stage of recovery (see below).

We did not include any species for which we haven't studied the original description and figures. Taxa reported from thin sections or polished slabs were not considered because taxonomic assignments of these taxa were in most cases difficult to verify. The complete dataset includes 37 lithological units comprising 1562 species belonging to gastropods (695 species), bivalves (587 species), brachiopods (147 species), crinoids (33 species), echinoids (31 species), crustaceans (25 species), ophiuroids (12 species), scaphopods (nine species), serpulids (eight species), microconchids (six species), asteroids (four species), bryozoans (four species) and polyplacophores (one species; *Data S1*). Our approach led to a certain overemphasis of European lithological units (27) in comparison to those from Asia (five) and America (five units from the USA), but the prevailing cosmopolitanism of marine faunas of this time (*Kristan-Tollmann & Tollmann, 1982; Schubert & Bottjer, 1995*) justifies the priority of the quality of the documentation over geographic coverage.

## Data preparation

The investigated lithological units represent different environments (*Data SI*). Investigated environments include lagoons (including carbonate platform interiors), inner ramps (above the fair-weather wave base), mid-ramps (between the fair-weather wave base and the storm wave base) and outer ramps (below the storm wave base). The assignment of faunas to a particular environment is based on the information given by the original authors or inferred from the geological and sedimentological context. Studies on lithological units that represent a range of different environments were only considered if the distribution of species within these environments was clear. Therefore, we did not include the Ladinian part of the Cassian Formation (Italy; *Urlichs, 2017*) in our analyses, although it probably represents the highest benthic diversity of this time interval.

From the collected studies, fossil lists were compiled for each lithological unit. If information about one lithological unit was taken from more than one study, attention was paid on possible synonyms.

Each species was assigned to an ecological guild, usually based on information given by the original authors. In cases where such information was missing, the ecology was inferred from functional morphology, mode of life of Recent species and information from the literature. Species included in this study represent shallow to moderately deep infaunal suspension-feeders, shallow infaunal deposit-feeders, deep infaunal suspension-feeders, endobryssate suspension-feeders, epibryssate suspension-feeders, free-lying suspension-feeders, cemented epifaunal suspension-feeders, pedunculate suspension-feeders, epifaunal herbivores and/or detritus feeders, epifaunal carnivores, erected epifaunal suspension-feeders and epifaunal detritivore-suspension-feeders (Tab. 1). Gastropods, echinoids and ophiuroids were collectively assigned to epifaunal herbivores and/or detritus-feeders and epifaunal detritivore-suspension-feeders, respectively. We are aware that this might be an over-generalization, but the shell morphology of these taxa does not provide evidence for more specialized feeding in extinct species, and the ecology of Recent species justifies this assignment for most species. Species with uncertain ecology were considered for calculating species richness but were not included in the ecological analyses.

Time resolution is provided at the substage level. For lithological units that extend across substages, we assumed constant diversity across substages if the literature suggests that this assumption is justified (e.g. data from the Germanic basin; *Schmidt, 1928; Schmidt, 1938*).

## Data analyses

Conventional diversity studies (e.g. *Sepkoski, 1984; Sepkoski, 1997*) through deep time have relied on higher taxonomic levels (genera, families, orders) as surrogates for the species level; however this approach may cause a strong bias (*Benton, 2001*). Unfortunately, compiling reliable global species level curves is virtually impossible because of the higher incompleteness of the fossil record for species than for higher taxa and the necessity of eliminating synonyms. We suggest here a feasible solution for circumventing these problems. We use the maximum species richness among all lithological units in a given time interval as a surrogate for the potentially highest recovery stage that biosphere could reach at this time point after the end-Permian mass extinction (Figs 1 and 2). In other words, for each substage and ecological guild, the lithological unit that contains the highest number of species was used as a reference point. Using this

approach, synonyms are not a major issue because it is irrelevant for the diversity of a given lithological unit whether the same species is described from another lithological unit under a different name.

Furthermore, we investigated relative changes in the ecological composition throughout the Early and Middle Triassic (Figs 3 and 4). For this purpose, the guild-species diversity (number of species per guild) for each substage was calculated. To do so, we summarized the number of species that occurred in all lithological units per substage. The sum of species was corrected for double counts, meaning that a species that occurred in several lithological units per substage was only counted once.

## Results

Figure 1 shows the huge differences in species richness between Early and Middle Triassic benthic communities that hold across different environments (Fig. 2). The apparent decrease of species richness in the Longobardian (late Ladinian; Figs 1 and 2) is an artefact of the scarcity of data that would disappear e.g. if time-resolved data from the Ladinian part of the Cassian Formation were available (see Methods). Currently, only one lithological unit represents the middle and late Longobardian in Figs 1 and 2. Scarce data may also underlie changes in guild-species diversity seen in the Smithian (early Olenekian) and the Aegean (early Anisian; Fig. 3).

Epifaunal herbivores and/or detritus feeders (i.e. gastropods) diversified most quickly in the Middle Triassic. Shallow to moderately deep and deep infaunal, endobyssate, epibyssate, free-lying, cemented and pedunculate suspension-feeders as well as shallow infaunal deposit-feeders started diversifying at the same time (Fig. 5) but their species richness levelled out at lower plateaus.

Bivalves, gastropods and brachiopods were the most diverse taxa in Early and Middle Triassic benthic communities (*Data S1*). Bivalves had a relatively low and constant species richness throughout the Early Triassic and diversified quickly in the Bithynian (early middle Anisian) followed by a plateau (Figs 6A and 7). Except for the Smithian Sinbad Formation, the species richness of gastropods was low in the Early Triassic and started to increase rapidly in the Bithynian and their species richness started to exceed that of bivalves from the Pelsonian (late middle Anisian) onwards (Figs 6B and 7). Brachiopods also diversified in the Bithynian but not to the same extent as bivalves and gastropods did, and their species richness maintained a Middle Triassic plateau (Figs 6C and 7). Especially gastropods but also brachiopods were relatively diverse in lithological units that are associated with carbonate platforms (red lines in Figs 6B and 6C). Fig. 6D shows a change in taxonomic composition from bivalve-dominated Early Triassic to increasingly more gastropod-dominated Middle Triassic faunas. Almost all lithological units associated with carbonate platforms are dominated by gastropods, which was already noticed by *Friesenbichler et al. (2019)*.

The guild-species diversity of epifaunal herbivores and/or detritus-feeders (i.e. gastropods) changed strongest among all ecological guilds during the studied time interval. During the Early Triassic they were the third most diverse guild but they started to diversify quickly during the Bithynian and became almost instantaneously the dominating guild in Middle Triassic benthic

communities (Fig. 3). The only environment where this trend is not observed are outer ramps, where their species richness remained constant (Figs 2D and 4D). Figures 5A and 5B show that their rapid diversification started in the Bithynian and continued throughout the Anisian.

Epibyssate suspension-feeders (chiefly pteriomorphian bivalves) were the most diverse guild in the Induan and the Spathian (late Olenekian). During the Bithynian their absolute richness increased whereas their relative richness decreased (Figs 1, 3, 5C and 5D) and continued to decrease slightly throughout the Middle Triassic. This decrease is also apparent in lagoons, inner and mid-ramps, whereas in outer ramps their relative diversity increased slightly (Fig. 4).

Shallow to moderately deep infaunal suspension-feeders (bivalves and inarticulate brachiopods) were among the most diverse guilds during the Early Triassic but their relative richness decreased gradually after the Aegean (Figs 1 and 3). The same trend occurred in the guild-species diversity in lagoons, mid- and outer ramps; however, in mid-ramps the relative species richness of shallow to moderately deep infaunal suspension-feeders was higher in the late Illyrian (late Anisian) to late Fassanian (early Ladinian) interval than in the late Pelsonian to early Illyrian (Figs 4A, 4C and 4D). In inner ramp settings, the relative amount of shallow to moderately deep infaunal suspension-feeders decreased already during the Spathian. Their main recovery phase started in the Bithynian but the diversification rate seemed to slow down afterwards (Figs 5C and 5D).

At the beginning of the Spathian, the relative richness of endobyssate suspension-feeders (bivalves) started to increase, but after the Aegean, it decreased continuously (Fig. 3). The same trend occurred in lagoons and outer ramps. In inner and mid-ramp settings, the relative amount of endobyssate suspension-feeders was more or less constant after the Spathian (Fig. 4). The species richness of endobyssate suspension-feeders increased constantly during the Early Triassic, followed by a rapid rise in species richness in the Bithynian. However, after this increase their species richness remained constant (Figs 5C and 5D).

The relative amount of pedunculate suspension-feeders (articulate brachiopods) fluctuated. They were relatively diverse during the Griesbachian (early Induan) and after the Bithynian but in between they were rather uncommon. After the early Illyrian they became gradually less diverse (Fig. 1). In Early Triassic mid-ramp settings, they are only recorded from the Griesbachian. Their relative importance was high during the Bithynian, but afterwards they became less common. In outer ramps from the Early Triassic, this guild is only reported from the Griesbachian and Dienerian (late Induan); however, in Middle Triassic outer ramps pedunculate suspension-feeders were the most diverse guild (Fig. 4D). They experienced a rapid rise in diversification in the Bithynian (Figs 5C and 5D).

Other guilds were rather uncommon during the Early and Middle Triassic or not represented in all time slices (Figs 1 and 3). Deep infaunal suspension-feeders (siphonate bivalves) are first reported from the Spathian (Fig. 3) and showed a weak increase in species richness in the Bithynian (Figs 5E and 5F). The species richness of shallow infaunal deposit-feeders (i.e. nuculid bivalves and scaphopods) and free-lying epifaunal suspension-feeders (bivalves) increased in the Bithynian but stayed constant during the remaining of the Middle Triassic (Figs 5E and 5F). Cemented epifaunal suspension-feeders (e.g. oysters, serpulids, microconchids) and erected epifaunal suspension-feeders (i.e. crinoids) also started to diversify during the Bithynian,



but their later diversification pattern is poorly documented (Figs 5G and 5H). Epifaunal detritivore-suspension feeders (i.e. ophiuroids) and epifaunal carnivores (i.e. asteroides and crustaceans) started to diversify in the Aegean or Bithynian (Figs 5G and 5H).

A remarkable aspect of the guild analysis is that the explosive Anisian diversity increase was not associated with the evolution of new benthic guilds, confirming previous observations by *Foster & Twitchett (2014)*.

## Discussion

### Possible biases

Our data indicate an explosive increase in benthic diversity at the beginning of the Middle Triassic that followed an extended Early Triassic lag phase (Fig. 1). Can this pattern result from a bias in preservation, rock exposure or research history?

A preservation bias in Early Triassic strata has often been proposed as a possible reason for the generally low species richness at that time (e.g. *Erwin, 1996; Wignall & Benton, 1999; Peters & Foote, 2002; Nützel & Schulbert, 2005; Hautmann et al., 2011*), but well preserved Early Triassic benthic faunas have been reported worldwide, e.g. from the western USA (*Batten & Stokes 1986; Hautmann & Nützel 2005; Nützel & Schulbert 2005; Pruss, Payne & Westacott, 2015; Brayard et al. 2017*), South China (*Kaim et al., 2010; Hautmann et al., 2011; Hautmann et al., 2015; Foster et al. 2019*), Russia (*Shigeta et al., 2009*) and Pakistan (*Wasmer et al. 2012; Kaim et al., 2013*); even silicified faunas are known from the Early Triassic of Oman (*Twitchett et al., 2004; Wheeley & Twitchett, 2005; Oji & Twitchett, 2015*) and Svalbard (*Foster, Danise & Twitchett, 2017*). However, all these well preserved Early Triassic benthic communities are much less diverse than communities from comparable settings in the Middle Triassic, so their low diversity is probably a primary signal.

Except for taxonomic oversplitting and synonyms (see Methods), the history of research and the size of exposure influence our knowledge on species richness. In our data, the most extreme potential bias in this respect is between the stratigraphically adjacent data points from the Bithynian *Tubiphytes*-Limestone Member (Romania) and the Jena Formation (Germany), which corresponds to the strongest increase in diversity within the studied time interval. The described fauna from the *Tubiphytes*-Limestone Member comes from only 1.5 m<sup>3</sup> of rock material (*Grădinaru & Gaetani, 2017; Nützel, Kaim & Grădinaru, 2018; Friesenbichler et al., 2021*), and rarefaction analysis indicates that the bivalve fauna is actually insufficiently sampled (*Friesenbichler et al., 2021*). In contrast, the Germanic Triassic looks back on a long history of research and is represented by a large outcrop area, which should correspond to a significantly higher completeness of sampled biodiversity. However, correcting the data point from the *Tubiphytes*-Limestone Member towards a higher diversity would solely shift the starting point of the main diversity increase from the middle to the early Bithynian, but not alter the overall shape of the diversity trajectory.

### Extrinsic or intrinsic control?

The observation of an Early Triassic lag phase in the recovery from the end-Permian mass extinction is not new. *Schubert & Bottjer (1995)* were among the first to demonstrate this delay

in rediversification based on regional data of post-extinction communities in the western USA. They suggested that the “long pre-radiation period” could be due (1) to the extraordinarily magnitude of the end-Permian mass extinction and/or (2) the persistence of environmental stress. The first explanation implies an evolutionary slow-down of diversification rates that correlates with the extinction magnitude, possibly because there is a positive feedback between species richness and rates of speciation, which becomes ineffective if diversity falls below a critical threshold (Solé *et al.*, 2010; Hautmann *et al.*, 2015). The second (environmental) explanation has found disproportionately more attention in the literature, which offers manifold scenarios of poisonous, acidic, oxygen deficient and lethally hot oceans that prevented life from re-diversifying after the great dying at the end of the Permian (e.g. Wignall & Twitchett, 1996; Payne *et al.* 2007; Joachimski *et al.*, 2012; Sun *et al.*, 2012; Song *et al.*, 2014). Support for the second explanation has been sought in the demonstration of adverse environmental conditions during the Early Triassic from geochemical or palaeontological proxies (e.g. Payne *et al.*, 2010; Sanei, Grasby & Beauchamp, 2012; Schobben *et al.*, 2014; Tian *et al.*, 2014; Rothman *et al.*, 2014; Wei *et al.*, 2015), but the idea of pervasive hostile conditions in the Early Triassic has also been criticized for being at variance with a variety of observations. These include: (1) benthic communities from various environmental settings and palaeogeographically distant regions show little indications for unusual environmental stress (Twitchett *et al.*, 2004; Hautmann *et al.*, 2011; Hautmann *et al.*, 2015; Hofmann, Hautmann & Bucher, 2013; Hofmann *et al.*, 2013; Hofmann *et al.*, 2014); (2) ichnofaunas were diverse and complex soon after the extinction event in different palaeolatitudes (Beatty, Zonneveld & Henderson, 2008; Hofmann *et al.*, 2011); (3) shell sizes of many Early Triassic gastropods reached large sizes (Brayard *et al.*, 2010; Brayard *et al.*, 2011a; Brayard *et al.*, 2015); (4) diversification of ammonoids was extremely rapid (Brayard *et al.*, 2009); (5) biomass productivity of marine benthos was high in spite of low taxonomic diversity (Brosse *et al.*, 2019) and (6) marine apex predators were present throughout the Early Triassic (Scheyer *et al.*, 2014). This criticism does not deny the existence of environmental stress in the Early Triassic, but it suggests that it was locally and temporarily restricted and had differential impacts on the benthos and nekton.

Another class of explanations assumes that a breakdown of primary production in the Early Triassic caused a collapse of the food pyramid. According to these models, the low diversity in the Early Triassic corresponds to a time of reduced primary production (e.g. Grasby *et al.*, 2020) and the recovery of marine life required a successive rebuilding of trophic levels (Chen & Benton, 2012, p. 379). However, Chen & Benton (2012) suggested that this model predicts a logistic rediversification curve, which would be at variance with the extended Early Triassic lag phase. Palaeontological data also demonstrate the presence of apex predators in the Early Triassic (Chen & Benton, 2012; Scheyer *et al.* 2014), which indicates that trophic levels between the base and the top of the food pyramid were at least partly intact. We add here that low primary production per se does not provide an explanation for low diversity. Theory rather predicts that selection for fitness in resource-poor environments favours specialized, efficient populations and therefore results in high diversity (Valentine, 1971). The high diversity of Recent coral reefs and deep-sea communities (Grassle & Maciolek, 1992; Veron *et al.*, 2009) are well-known examples that conform to this prediction.

In contrast to the idea of environmentally driven delay (EDD) models, the biotic interaction (BI) model does not only provide an explanation for the delayed recovery but it also makes a prediction for the diversification pattern that followed the initial lag phase. Mathematically, the

BI model can be written as a differential equation that can be solved numerically. It contains a hyperbolic term that accounts for the positive feedback of species interactions on rates of diversification, which is usually combined with a damping term that limits the increase when the effects of niche pre-emption and ecological crowding become dominant (*Solé et al., 2010*, equation 10; *Hautmann et al., 2015*, equation 1). In contrast to the more familiar logistic model (*Sepkoski, 1978; Sepkoski, 1984*), the shape of the resulting curve is decidedly asymmetrical (compare Figs 8A and 8C). It displays an extended early (left) branch, which represents the prolonged lag phase when feedback processes were ineffective due to low species richness, followed by an explosive increase that stops relatively suddenly when ecospace is filled (Fig. 8C). However, the exact shape of the curve depends upon two parameters: (1) the starting diversity that can be expressed as percentage of the carrying capacity or as the percentage of surviving species, and (2) the proportionality factor, which represents the feedback strength. Depending on these two parameters, which are insufficiently known, the shape of the hyperbolic damped diversification curve can appear relatively similar to a logistic shape under certain conditions, but a qualitative difference between the two models always remains. This difference can be visualized in semilogarithmic plots, where a hyperbolic curve is reflected by an initial increase in the slope of the diversification curve, whereas a logistic curve is represented by a decreasing slope (Figs 8B and 8D).

The majority of the semilogarithmic plots presented in this study (Figs 5 and 7) is conform to the hyperbolic-damped (= BI) model. This applies to the diversification curves of bivalves, gastropods and brachiopods in general, bivalves from mid-ramp settings, as well as shallow infaunal deposit-feeders (i.e., nuculid bivalves and scaphopods) and free-lying epifaunal suspension feeders (some bivalves), with their sudden increase in species richness during the Bithynian that is followed by a plateau, which indicates a hyperbolic-damped increase in species richness. Likewise, the diversification trends of gastropods in mid-ramp settings, epibyssate suspension-feeders (some bivalves), epifaunal herbivores and/or detritus-feeders (i.e., gastropods), erected epifaunal suspension-feeders (i.e., crinoids) and pedunculate suspension-feeders (i.e., articulate brachiopods) show a sudden increase in species richness in the Bithynian, which is as indicative of a hyperbolic trajectory. The data from some other guilds do not provide an unequivocal support for either the hyperbolic or the logistic model. This concerns shallow to moderately deep infaunal, endobysate and cemented epifaunal suspension-feeders (e.g., some bivalves, serpulids) as well as epifaunal detritivore-suspension-feeders (i.e., ophiuroids). The remaining taxa and guilds (i.e., brachiopods of mid-ramp settings, deep infaunal suspension-feeders and epifaunal carnivores like asteroids and crustaceans) do not match with any model, possibly because of their scarcity in our dataset.

A conclusion from these observations is that they suggest a strong effect of biotic interactions on rates of diversification. This is the prime difference to the standard logistic model, which makes a neutral assumption with respect to biotic interactions. In the logistic model, the initially exponential increase is solely an effect of an increasing number of species that each have constant diversification rates. This assumption is clearly incompatible with the explosive diversification revealed in our dataset (Fig. 1). Notably, this is not an isolated observation. *Miller & Sepkoski (1988)* found “hyperexponential bursts” in the Phanerozoic diversity curve of bivalves during the Ordovician diversification and following the end-Permian and end-Cretaceous mass extinctions, which likely represent times of hyperbolic increase. *Markov &*

*Korotayev (2007)* even proposed that the Phanerozoic biodiversity curve from which *Sepkoski (1978, 1984)* derived his model is actually better described by a hyperbolic model.

Implicit in these finds is that we cannot expect an early rediversification if the feedback mechanism between richness and rates of diversification was ineffective after an extreme diversity crash, whether or not the environment was favourable. Does this conclusion refute the EDD model? It partly does, because it means that at least a portion of the lag phase is always attributed to BI. It is therefore incorrect to ascribe the full length of the lag phase to EDD. However, it is currently also uncertain whether the full length of the Early Triassic lag phase can be explained solely by the BI model. The length of the lag phase in this model depends upon two parameters, the starting diversity and the proportionality factor, which are not precisely known. It is therefore possible that adverse environmental conditions delayed or interrupted the early recovery whereas the BI model explains a subsequent extension of the lag phase. Clarifying the exact parameters in the BI model equation for different taxa is required for estimating the length of the intrinsic delay and thus the potential for EDD as an additional explanation for the Early Triassic lag phase.

From an evolutionary viewpoint, the most relevant aspect is the nature of biotic interactions that cause positive feedbacks between species richness and rates of diversification. *Solé et al. (2010)* were vague with respect to possible feedback processes, suggesting that increasing numbers of ecological interactions might provide the context for new opportunities to speciate, but they also noted that this process might operate too slowly to explain explosive diversification. Alternatively, they proposed that the growing number of potential interactions might drive the increase in the number of species. *Hautmann et al. (2015)* reconsidered the problem and identified interspecific competition as a potential main driver of hyperbolic-damped diversification curves. Accordingly, promotion of niche differentiation is the dominant outcome of interspecific competition until a critical level of ecological saturation has been reached, at which further division of niche space requires increasingly elaborated adaptations that are more and more unlikely to evolve. Above such a saturation limit, outcompeting of existing species becomes the prevalent effect of competition, which damps further diversification. A strength of this model is that it correctly predicts different diversification trajectories for taxa with supposedly different intensities of interspecific competition, for which slowly recovering bivalves contrasted by explosively rediversifying ammonoids have been cited as an example (*Hautmann et al., 2015*). The fact that the phase of the hyperbolic diversity increase was not associated with the origin of new guilds further supports the competition-driven scenario, because competition is most intense between species with similar lifestyles. However, *Hautmann et al. (2015)* also noted that many other aspects in addition to competition affect actual patterns of diversification. In the following, we discuss some possible additional factors that were linked to the resurgence of carbonate platforms in the Middle Triassic.

### **Habitat construction and biotic interactions on resurging large carbonate platforms**

The recovery of large carbonate platforms after the end-Permian mass extinction started in the Anisian (early Middle Triassic; *Gaetani et al., 1981; Senowbari-Daryan et al. 1993; Berra, Rettori & Bassi, 2005*), and coincided with the main diversification of most benthic marine invertebrate taxa (Figs 5 and 7). One of the oldest Triassic carbonate platforms is represented by the *Tubiphytes*-Limestone Member of the Caerace Formation in Romania. This Bithynian

biostrome is dominated by the micro-encruster *Tubiphytes* and large volumes of synsedimentary cements (Popa, Panaiotu & Grădinaru, 2014) and provided the basis for a rich hardground community that is much more diverse than contemporaneous level-bottom faunas (Grădinaru & Gaetani, 2017; Forel & Grădinaru, 2018; Nützel, Kaim & Grădinaru, 2018; Friesenbichler et al., 2021). Its high amount of newly described species is potentially related to the new habitat type. This applies to the dominant taxa, i.e., epibyssate bivalves, gastropods and brachiopods (Data S1) that were well adapted to the hardgrounds provided by the *Tubiphytes*-microbial buildup (Friesenbichler et al., 2021). Further Middle Triassic *Tubiphytes*-dominated buildups are known from the Iranian carbonate mounds of Nakhlak (late Bithynian; Berra et al., 2012), the Italian Camorelli platform (Bithynian-Pelsonian; Gaetani & Gorza, 1989) and Dont Formation (Pelsonian-Illyrian; Blendinger, 1983; Fois & Gaetani, 1984), the Hungarian Aggtelek reef (Pelsonian-Ladinian; Velledits et al., 2011; Velledits, Hips & Péro, 2012) and Chinese reefs (Anisian; Enos, Wei & Yan, 1997; Lehrmann 1999; Enos et al., 2006; Payne et al., 2006a; Payne et al. 2006b; Lehrmann et al. 2007). The latter flourished on the Great Bank of Guizhou, an isolated carbonate platform in the Nanpanjiang Basin, and represent the oldest Triassic platform margin reefs (Lehrmann, Wei & Enos, 1998; Payne et al. 2006a; Payne et al. 2006b).

Reefs existed throughout the Triassic, but their composition changed stepwise (see Martindale, Foster & Velledits, 2019). The oldest Triassic reefs are represented by microbial-metazoan reefs that formed immediately after the end-Permian mass extinction (e.g. Lehrmann et al., 2001; Flügel, 2002; Pruss et al., 2006; Wu et al., 2007; Kiessling & Simpson, 2011; Kershaw et al., 2011; Ezaki, Liu & Adachi, 2012; Yang et al., 2015, Friesenbichler et al. 2018). The first metazoan reefs are reported from the Olenekian (late Early Triassic) and represent small sponge biostromes and bivalve build-ups (Pruss, Payne & Bottjer, 2007; Brayard et al., 2011b; Marenco et al., 2012). Reefs developed quickly in the middle Anisian (Flügel, 2002) and in addition to *Tubiphytes* other organisms such as sphinctozoans (segmented calcareous sponges), dasycladacean algae, corals, and in some cases also bivalves became important reef-builders (Flügel, 2002; Fürsich & Hautmann, 2005; Senowbari-Daryan & Link, 2011). These organisms acted as ecological engineers and created new niches that provided opportunities for additional species. Well-documented examples are bivalve-crinoid reef mounds from the Trochitenkalk Formation (middle Illyrian, Germany; e.g. Hagdorn, 1978; Sellwood & Fürsich, 1981; Hagdorn & Mundlos, 1982; Flügel, 2002; Hagdorn, 2004; Hagdorn & Nitsch, 2009), where shell beds were incrustated by cementing bivalves that provided hardgrounds for the colonization by crinoids. These dm-scaled bioherms provided new habitats for brachiopods, serpulids, other bivalves and boring worm-like organisms.

The sudden increase in species richness of gastropods is another feature of the Middle Triassic main stage of marine recovery (Figs. 1, 2, 3, 4, 5A, 5B, 6B and 7) that was at least partly linked to the resurgence of large carbonate platforms. As shown by Friesenbichler et al (2019, fig. 14), the ratio of gastropod versus bivalve species is much higher in carbonate platforms than in other settings. However, the data of Friesenbichler et al. (2019b) might even underestimate the true signal. Roden et al. (2020) compared the diversity of the Cassian Formation, an exceptional Triassic lagerstätte, with the diversity of the Wetterstein Formation, which is environmentally comparable and approximately of the same age. They found that all mollusks were underrepresented in the Wetterstein Formation by factor 7 in comparison to the Cassian Formation, but gastropods by the astonishing factor 87. The cause for the affinity of gastropods to carbonate platform facies remains speculative, not at least because the life habit and feeding

mode of gastropods cannot be deduced from the shell morphology. In spite of this limitation, we put forward the hypothesis that the quick diversification of gastropods in carbonate platform environments might have been related to the proliferation of dasycladacean algae, which were dominant carbonate producers in many settings where gastropods became diverse.

Palaeoecological studies have previously suggested that at least some Triassic gastropods lived in association with macroalgae (Fürsich & Wendt, 1977; Sellwood & Fürsich, 1981; Stiller, 2001; Hagdorn, 2004; Nützel & Schulbert, 2005; Diedrich, 2010; Urlichs, 2014). This conclusion was derived from the ecology of Recent macroalgae (e.g. Davies, 1970; Taylor, 1971; Thomassin, 1971; Brasier, 1975; Poulicek, 1985; Sánchez-Moyano et al., 2000; Chemello & Milazzo, 2002; Antoniadou & Chintiroglou, 2005; Pitacco et al., 2014; Duarte et al., 2015; Chiarore et al., 2017), on which gastropods grazed (e.g. Underwood, 1980; Johnson & Mann, 1986; Williams, 1993) and thereby enhance the photosynthetic capacity of the host algae (Amsler et al., 2015). The Triassic fossil record of non-calcifying algae is virtually non-existent, but the diversification and abundance of dasycladaceans at the beginning of the Middle Triassic is well established. In our dataset, only 5 out of 37 lithological units contain information about the number of Dasycladacea species associated with the investigated faunas (e.g., Salomon, 1895; Ogilvie Gordon, 1927; Schmidt, 1928; Schmidt, 1938; Granier & Grgasović, 2000; Russo et al., 2000; Emmerich et al., 2005; Piros & Preto, 2008). Given the small number of datapoints, the positive correlation between the richness of dasycladaceans and gastropods is statistically not robust ( $r^2 = 0.25$ ;  $p = 0.32$ ), but it is strengthened by an indirect line of evidence. Middle Triassic gastropod diversity decreased towards deeper water settings (Figs 2 and 4), along with declining light. Because light limits the distribution of algae, it is certain that this trend was matched by decreasing macroalgae abundance.

## Conclusions

Data on species richness from 37 Early and Middle Triassic lithological units containing level-bottom communities indicate that most of the analyzed taxa and guilds followed a hyperbolic-damped diversity trajectory, with an extended Early Triassic lag phase followed by an explosive increase in diversity at the beginning of the Middle Triassic that levelled out in the Ladinian. The Early Triassic delay in rediversification of benthic organisms has conventionally been attributed to ongoing environmental stress at that time, but the explosive (hyperbolic rather than exponential) diversity increase at the beginning of the Middle Triassic cannot be explained by relaxation of environmental stress. Our data therefore support a model in which the intensity of biotic interactions determines the rate of diversification. Accordingly, the Early Triassic delay in rediversification represents the time in which the intensity of biotic interactions was too low for driving a major diversification because of the dramatic loss of species during the end-Permian mass extinction. Conversely, the hyperbolic Middle Triassic diversity increase occurred when self-accelerating processes became effective, and it stopped when niche-crowding prevented further diversification. This pattern is in agreement with a competition-driven diversification model that predicts a reversal in the effects of interspecific competition on diversification rates from accelerating to damping. Apart from interspecific competition, other ecological feedbacks might have stimulated the hyperbolic increase in diversity during the Anisian. Carbonate platforms, which reappeared simultaneously with the main stage of benthic rediversification, provided the ecological context for rich hardground communities. The coincidence of the resurgence of carbonate platforms and the evolution of reef-builders with the

main diversification of benthic faunas suggests a positive loop effect, which is confirmed by several case studies. When reefs diversified, other reef-builders became abundant and acted as ecological engineers and niche constructors, providing new habitats for additional species. The quick diversification of gastropods in Middle Triassic carbonate platforms might be related to the proliferation of macroalgae, which provided much improved opportunities for grazing. A strength of the biotic interaction model is that it correctly predicts both an extended lag phase and a hyperbolic increase in diversity thereafter, whereas environmental scenarios fail to predict the second aspect.

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## References

- ABERHAN, M. 1994. Guild-Structure and Evolution of Mesozoic Benthic Shelf Communities. *Palaios*, **9**, 516–545.
- AMSLER, M. O., HUANG, Y. M., ENGL, W., MCCLINTOCK, J. B. and AMSLER, C. D. 2015. Abundance and diversity of gastropods associated with dominant subtidal macroalgae from the western Antarctic Peninsula. *Polar Biology*, **38**, 1171–1181.
- ANTONIADOU, C. and CHINTIROGLOU, C. 2005. Biodiversity of zoobenthic hard-substrate sublittoral communities in the Eastern Mediterranean (North Aegean Sea). *Estuarine, Coastal and Shelf Science*, **62**, 637–653.
- BATTEN, R. L. and STOKES, W. M. L. 1986. Early Triassic gastropods from the Sinbad member of the Moenkopi Formation, San Rafael Swell, Utah. *AmMusNov*, **2864**, 1–33.
- BEATTY, T. W., ZONNEVELD, J. P., HENDERSON, C. M. 2008. Anomalously diverse Early Triassic ichnofossil assemblages in Northwest Pangea: a case for a shallow-marine habitable zone. *Geology* **36**, 771–774.
- BENTON, M. J. 2001. Biodiversity on land and in the sea. *Geological Journal*, **36**, 211–230.
- BERRA, F., BALINI, M., LEVERA, M., NICORA, A. and SALAMATI, R. 2012. Anatomy of carbonate mounds from the Middle Anisian of Nakhlak (Central Iran): architecture and age of a subtidal microbial-bioclastic carbonate factory. *Facies*, **58**, 685–705.
- BERRA, F., RETTORI, R. and BASSI, D. 2005. Recovery of carbonate platform production in the Lombardy Basin during the Anisian: paleoecological significance and constrain on paleogeographic evolution. *Facies*, **50**, 615–627.
- BLENDINGER, W. 1983. Anisian sedimentation and tectonics of the M. Pore - M. Cernera area

- 519 (Dolomites). *Rivista Italiana di Paleontologia e Stratigrafia*, **89**, 175–208.
- 520 BRASIER, M. D. 1975. Ecology of Recent sediment-dwelling and phytal foraminifera from the  
521 lagoons of Barbuda, West Indies. *Journal of Foraminiferal Research*, **5**, 42–62.
- 522 BRAYARD, A. and BUCHER, H. 2015. Permian-Triassic Extinctions and Rediversifications. In  
523 KLUG, C., KORN, D., DE BEATS, K., KRUTA, I., MAPES, R. H. and LANDMAN, N. H.  
524 (eds.) *Ammonoid Paleobiology: From Anatomy to Ecology*, Springer-Verlag, 465–473 pp.
- 525 BRAYARD, A., ESCARGUEL, G., BUCHER, H., MONNET, C., BRÜHWILER, T.,  
526 GOUDEMAND, N., GALFETTI, T. and GUÉX, J. 2009. Good genes and good luck:  
527 Ammonoid diversity and the end-permian mass extinction. *Science*, **325**, 1118–1121.
- 528 BRAYARD, A., KRUMENACKER, L. J., BOTTING, J. P., JENKS, J. F., BYLUND, K. G.,  
529 FARA, E., VENNIN, E., OLIVIER, N., GOUDEMAND, N., SAUCÈDE, T.,  
530 CHARBONNIER, S., ROMANO, C., DOGUZHAIEVA, L., THUY, B., HAUTMANN, M.,  
531 STEPHEN, D. A., THOMAZO, C. and ESCARGUEL, G. 2017. Unexpected Early Triassic  
532 marine ecosystem and the rise of the Modern evolutionary fauna. *Science Advances*, **3**,  
533 e1602159.
- 534 BRAYARD, A., MEIER, M., Escarguel, G., FARA, E., NÜTZEL, A., OLIVIER, N.,  
535 BYLUND, K. G., JENKS, J. F., STEPHEN, D. A., HAUTMANN, M., VENNIN, E. and  
536 BUCHER, H. 2015. Early Triassic Gulliver gastropods: Spatiotemporal distribution and  
537 significance for biotic recovery after the end-Permian mass extinction. *Earth-Science*  
538 *Reviews*, **146**, 31–64.
- 539 BRAYARD, A., NÜTZEL, A., STEPHEN, D. A., BYLUND, K. G., JENKS, J. and BUCHER,  
540 H. 2010. Gastropod evidence against the Early Triassic Lilliput effect. *Geology*, **38**, 147–  
541 150.
- 542 BRAYARD, A., NÜTZEL, A., KAIM, A., ESCARGUEL, G., HAUTMANN, M., STEPHEN,  
543 D. A., BYLUND, K. G., JENKS, J. and BUCHER, H. 2011a. Gastropod evidence against  
544 the Early Triassic Lilliput effect: reply. *Geology*, **39**, e233.
- 545 BRAYARD, A., VENNIN, E., OLIVIER, N., BYLUND, K. G., JENKS, J., STEPHEN, D. A.,  
546 BUCHER, H., HOFMANN, R., GOUDEMAND, N. and ESCARGUEL, G. 2011b.  
547 Transient metazoan reefs in the aftermath of the end-Permian mass extinction. *Nature*  
548 *Geoscience*, **4**, 693–697.
- 549 BROSSE, M., BUCHER H., BAUD, A., FRISK, A., GOUDEMAND, N., HAGDORN, H.,  
550 NÜTZEL, A., WARE, D. and HAUTMANN, M. 2019. New data from Oman indicate  
551 benthic high biomass productivity coupled with low taxonomic diversity in the aftermath of  
552 the Permian-Triassic Boundary mass extinction. *Lethaia*, **52**, 165–187.
- 553 CHEMELLO, R. and MILAZZO, M. 2002. Effect of algal architecture on associated fauna:  
554 some evidence from phytal molluscs. *Marine Biology*, **140**, 981–990.
- 555 CHEN, Z.-Q. and BENTON, M. J. 2012. The timing and pattern of biotic recovery following the



- 556 end-Permian mass extinction. *Nature Geoscience*, **5**, 375–383.
- 557 CHIARORE, A., FIORETTI, S., MECCARIELLA, A., SACCONI, G. and PATTI, P. 2017.  
558 Molluscs community associated with the brown algae of the genus *Cystoseira* in the Gulf of  
559 Naples (South Tyrrhenian Sea). *BioRxiv*, **nicht peer**.
- 560 COHEN, K. M., FINNEY, S. C., GIBBARD, P. L. and FAN, J.-X. 2013. The ICS international  
561 chronostratigraphic chart. *Episodes*, **36**, 199–204.
- 562 DAVIES, G. R. 1970. Carbonate bank sedimentation, eastern Shark Bay, Western Australia.  
563 *Memoirs of the American Association of Petroleum Geologists*, **13**, 85–168.
- 564 DIEDRICH, C. G. 2010. Palaeoecology of *Placodus gigas* (Reptilia) and other placodontids -  
565 Middle Triassic macroalgae feeders in the Germanic Basin of central Europe - and evidence  
566 for convergent evolution with *Sirenia*. *Palaeogeography, Palaeoclimatology,*  
567 *Palaeoecology*, **286**, 287–306.
- 568 DUARTE, R. C. S., MOTA, E. L. S., ALMEIDA, I. C. S., PESSANHA, A. L. M.,  
569 CHRISTOFFERSEN, M. L. and DIAS, T. L. P. 2015. Gastropods associated to three reef  
570 macroalgae with different architectures. *Strombus*, **22**, 5–18.
- 571 EMMERICH, A., ZAMPARELLI, V., BECHSTÄDT, T. and ZÜHLKE, R. 2005. The reefal  
572 margin and slope of a Middle Triassic carbonate platform: the Latemar (Dolomites, Italy).  
573 *Facies*, **50**, 573–614.
- 574 ENOS, P., LEHRMANN, D. J., WEI, J., YU, Y., JIAFEI, X., CHAICHIN, D. H., MINZONI,  
575 M., BERRI, A. C. and MONTGOMERY, P. 2006. Triassic Evolution of the Yangtze  
576 Platform in Guizhou Province, People's Republic of China. *Geological Society of America*  
577 *Special Paper*, **417**, 1–105.
- 578 ENOS, P., WEI, J. I. and YAN, Y. . 1997. Facies distribution and retreat of Middle Triassic  
579 platform margin, Guizhou Province, South China. *Sedimentology*, **44**, 563–584.
- 580 ERWIN, D. H. 1996. Understanding biotic recoveries: extinction, survival and preservation  
581 during the end-Permian mass extinction. In JABLONSKI, D., ERWIN, D. H. and LIPPS, J.  
582 H. (eds.) *Evolutionary Paleobiology*, Chicago Press, Chicago, 398–418 pp.
- 583 EZAKI, Y., LIU, J. B. and ADACHI, N. 2012. Lower Triassic stromatolites in Luodian County,  
584 Guizhou Province, South China: evidence for the protracted devastation of the marine  
585 environments. *Geobiology*, **10**, 48–59.
- 586 FLÜGEL, E. 2002. Triassic reef patterns. *Special Publication - Society for Sedimentary Geology*,  
587 **72**, 391–463.
- 588 FOIS, E. and GAETANI, M. 1984. The recovery of reef-building communities and the role of  
589 cnidarians in carbonate sequences of the Middle Triassic (Anisian) in the Italian Dolomites.  
590 *Paleontographica Americana*, **54**, 191–200.

- 591 FOREL, M.-B. and GRĂDINARU, E. 2018. First report of ostracods (Crustacea) associated with  
592 Bithynian (Anisian, Middle Triassic) *Tubiphytes*-microbial reef in the North Dobrogean  
593 Orogen (Romania). *Papers in Palaeontology*, **4**, 211–244.
- 594 FOSTER, W. J. and SEBE, K. 2017. Recovery and diversification of marine communities  
595 following the late Permian mass extinction event in the western Palaeotethys. *Global and*  
596 *Planetary Change*, **155**, 165–177.
- 597 FOSTER, W. J., DANISE, S. and TWITCHETT, R. J. 2017. A silicified Early Triassic marine  
598 assemblage from Svalbard. *Journal of Systematic Palaeontology*, **15**, 851–877.
- 599 FOSTER, W. J., DANISE, S., PRICE, G. D. and TWITCHETT, R. J. 2017. Subsequent biotic  
600 crises delayed marine recovery following the late Permian mass extinction event in northern  
601 Italy. *PLoS ONE*, **12**, e0172321.
- 602 FOSTER, W. J., DANISE, S., SEDLACEK, A., PRICE, G. D., HIPS, K. and TWITCHETT, R.  
603 J. 2015. Environmental controls on the post-Permian recovery of benthic, tropical marine  
604 ecosystems in western Palaeotethys (Aggtelek Karst, Hungary). *Palaeogeography,*  
605 *Palaeoclimatology, Palaeoecology*, **440**, 374–394.
- 606 FOSTER, W. J., LEHRMANN, D. J., HIRTZ, J. A., WHITE, M., YU, M., LI, J. and  
607 MARTINDALE, R. C: 2019. Early Triassic benthic invertebrates from the Great Bank of  
608 Guizhou, South China: Systematic palaeontology and palaeobiology. *Papers in*  
609 *Palaeontology*, **5**, 613–656.
- 610 FOSTER, W. J. and TWITCHETT, R. J. 2014. Functional diversity of marine ecosystems after  
611 the Late Permian mass extinction event. *Nature Geoscience*, **7**, 233–238.
- 612 FRIESENBIHLER, E., HAUTMANN, M., GRĂDINARU, E. and BUCHER, H. 2021. A  
613 highly diverse bivalve fauna from a Bithynian (Anisian, Middle Triassic) *Tubiphytes*-  
614 microbial buildup in North Dobrogea (Romania). *Papers in Palaeontology*, **7**, 447–495.
- 615 FRIESENBIHLER, E., HAUTMANN, M., NÜTZEL, A., URLICHS, M. and BUCHER, H.  
616 2019. Palaeoecology of Late Ladinian (Middle Triassic) benthic faunas from the  
617 Schlern/Sciliar and Seiser Alm/Alpe di Siusi area (South Tyrol, Italy). *Paläontologische*  
618 *Zeitschrift*, **93**, 1–29.
- 619 FRIESENBIHLER, E., RICHOZ, S., BAUD, A., KRYSTYN, L., SAHAKYAN, L.,  
620 VARDANYAN, S., PECKMANN, J., REITNER, J. and HEINDEL, K. 2018. Sponge-  
621 microbial build-ups from the lowermost Triassic Chanakhchi section in southern Armenia:  
622 Microfacies and stable carbon isotopes. *Palaeogeography, Palaeoclimatology,*  
623 *Palaeoecology*, **490**, 653–672.
- 624 FÜRSICH, F. T. and HAUTMANN, M. 2005. Bivalve reefs from the Upper Triassic of Iran.  
625 *Annali dell'Università degli Studi di Ferrara, Museologia Scientifica e Naturalistica,*  
626 **Special volume in honour of Carmela Loriga**, 13–23.
- 627 FÜRSICH, F. T. and WENDT, J. 1977. Biostratinomy and palaeoecology of the cassian

- 628 formation (Triassic) of the Southern Alps. *Palaeogeography, Palaeoclimatology,*  
629 *Palaeoecology*, **22**, 257–323.
- 630 GAETANI, M. and GORZA, M. 1989. The Anisian (Middle Triassic) carbonate bank of  
631 Camorelli (Lombardy, southern Alps). *Facies*, **21**, 41–56.
- 632 GAETANI, M., FOIS, E., JADOUL, F. and NICORA, A. 1981. Nature and evolution of Middle  
633 Triassic carbonate buildups in the Dolomites (Italy). *Marine Geology*, **44**, 25–57.
- 634 GÖTZ, A., SZULC, J. and FEIST-BURKHARDT, S. 2005. Distribution of sedimentary organic  
635 matter in Anisian carbonate series of S Poland: evidence of third-order sea-level  
636 fluctuations. *International Journal of Earth Sciences (Geologische Rundschau)*, **94**, 267–  
637 274.
- 638 GRĂDINARU, E. and GAETANI, M. 2017. Spathian to Bithynian (Early to Middle Triassic)  
639 brachiopods from North Dobrogea (Romania) Spathian to Bithynian (Early to Middle  
640 Triassic) brachiopods from North Dobrogea (Romania). *11th Romanian Symposium on*  
641 *Palaeontology*, 49.
- 642 GRANIER, B. R. C. and GRGASOČIĆ, T. 2000. Les Algues Dasycladales du Permien et du  
643 Trias - Nouvelle tentative d'inventaire bibliographique, géographique et stratigraphique.  
644 *Geologia Croatica*, **53**, 1–197.
- 645 GRASBY, S. E., KNIES, J., BEAUCHAMP, B., BOND, D. P., WIGNALL, P. and SUN, Y.  
646 2020. Global warming leads to Early Triassic nutrient stress across northern Pangea. *GSA*  
647 *Bulletin*, **132**, 943–954.
- 648 GRASSLE, J F. and MACIOLEK, N. J. 1992. Deep-sea species richness: regional and local  
649 diversity estimates from quatitative bottom samples. *The American Naturalist*, **139**, 313–  
650 341.
- 651 HAGDORN, H. 1978. Muschel/Krinoiden-Bioherme im Oberen Muschelkalk (mo1, Anis) von  
652 Crailsheim und Schwäbisch Hall (Südwestdeutschland). *Neues Jahrbuch für Geologie und*  
653 *Paläontologie Abhandlungen*, **156**, 31–86.
- 654 HAGDORN, H. 2004. *Muschelkalkmuseum Ingelfingen*. Edition Lattner, Lattner Photographic  
655 Art GmbH, Heilbronn.
- 656 HAGDORN, H. and MUNDLOS, R. 1982. Autochthonschille im Oberen Muschelkalk  
657 (Mitteltrias) Südwestdeutschlands. *Neues Jahrbuch für Geologie und Paläontologie,*  
658 *Abhandlungen*, **162**, 332–351.
- 659 HAGDORN, H. and NITSCH, E. 2009. The Germanic Triassic in its Southwest German type  
660 area. *6th International Triassic Field Workshop (Pan-European Correlation of the Triassic)*  
661 *Triassic of Southwest Germany 175th Anniversary of the Foundation of the Triassic System*  
662 *by Friedrich von Alberti September 7-11, 2009, Tübingen and Ingelfingen*, 72.
- 663 HALLAM, A. 1991. Why was there a delayed radiation after the end-Palaeozoic extinctions?

- 664 *Historical Biology*, **5**, 257–262.
- 665 HAUSMANN, I. M. and NÜTZEL, A. 2015. Diversity and palaeoecology of a highly diverse  
666 Late Triassic marine biota from the Cassian Formation of north Italy. *Lethaia*, **48**, 235–255.
- 667 HAUTMANN, M. 2007. Steuerungsfaktoren globaler Biodiversifizierungsmuster am Beispiel  
668 frühmesozoischer Muscheln. *Beringeria*, **37**, 61–74.
- 669 HAUTMANN, M. and NÜTZEL, A. 2005. First record of a heterodont bivalve (Mollusca) from  
670 the Early Triassic: palaeoecological significance and implications for the ‘Lazarus Problem’.  
671 *Palaeontology*, **48**, 1131–1138.
- 672 HAUTMANN, M., BAGHERPOUR, B., BROSE, M., FRISK, Å. M., HOFMANN, R., BAUD,  
673 A., NÜTZEL, A., GOUEMAND, N. and BUCHER, H. 2015. Competition in slow  
674 motion: The unusual case of benthic marine communities in the wake of the end-Permian  
675 mass extinction. *Palaeontology*, **58**, 871–901.
- 676 HAUTMANN, M., BUCHER, H., BRÜHWILER, T., GOUEMAND, N., KAIM, A. and  
677 NÜTZEL, A. 2011. An unusually diverse mollusc fauna from the earliest Triassic of South  
678 China and its implications for benthic recovery after the end-Permian biotic crisis. *Geobios*,  
679 **44**, 71–85.
- 680 HOFMANN, R., GOUEMAND, N., WASMER, M., BUCHER, H. and HAUTMANN, M.  
681 2011. New trace fossil evidence for an early recovery signal in the aftermath of the end-  
682 Permian mass extinction. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **310**, 216–  
683 226.
- 684 HOFMANN, R., HAUTMANN, M. and BUCHER, H. 2015. Recovery dynamics of benthic  
685 marine communities from the Lower Triassic Werfen Formation, northern Italy. *Lethaia*,  
686 **48**, 474–496.
- 687 HOFMANN, R., HAUTMANN, M., and BUCHER, H. 2013. A new paleoecological look at the  
688 Dinwoody Formation (Lower Triassic, Western USA): Intrinsic versus extrinsic controls on  
689 ecosystem recovery after the end-Permian mass extinction. *Journal of Paleontology*, **87**,  
690 854–880.
- 691 HOFMANN, R., HAUTMANN, M., BRAYARD, A., NÜTZEL, A., BYLUND, K. G., JENKS,  
692 J. F., VENNIN, E., OLIVIER, N. and BUCHER, H. 2014. Recovery of benthic marine  
693 communities from the end-Permian mass extinction at the low latitudes of eastern  
694 Panthalassa. *Palaeontology*, **57**, 547–589.
- 695 HOFMANN, R., HAUTMANN, M., WASMER, M. and BUCHER, H. 2013. Palaeoecology of  
696 the Spathian Virgin Formation (Utah, USA) and its implications for the Early Triassic  
697 recovery. *Acta Palaeontologica Polonica*, **58**, 149–173.
- 698 JATTIOT, R., BRAYARD, A., BUCHER, H., VENNIN, E., CARAVACA, G., JENKS, J. F. and  
699 BYLUND, K. G. 2018. Palaeobiogeographical distribution of Smithian (Early Triassic)  
700 ammonoid faunas within the Western USA Basin and its controlling parameters.

- 701 *Palaeontology*, **61**, 881–904.
- 702 JOACHIMSKI, M. M., LAI, X., SHEN, S., JIANG, H., LUO, G., CHEN, B., CHEN, J. and  
703 SUN, Y. 2012. Climate warming in the latest Permian and the Permian-Triassic mass  
704 extinction. *Geology*, **40**, 195–198.
- 705 JOHNSON, C. R. and MANN, K. H. 1986. The importance of plant defence abilities to the  
706 structure of subtidal seaweed communities: the kelp *Laminaria longicuris* de la Pylaie  
707 survives grazing by the snail *Lacuna vineta* (Montagu) at high population densities. *Journal*  
708 *of Experimental Marine Biology and Ecology*, **97**, 231–267.
- 709 KAIM, A., NÜTZEL, A., BUCHER, H., BRÜHWILER, T. and GOUEMAND, N. 2010. Early  
710 Triassic (Late Griesbachian) gastropods from South China (Shanggan, Guangxi). *Swiss*  
711 *Journal of Geosciences*, **103**, 121–128.
- 712 KAIM, A., NÜTZEL, A., HAUTMANN, M. and BUCHER, H. 2013. Early Triassic gastropods  
713 from Salt Range, Pakistan. *Bulletin of Geosciences*, **88**, 505–516.
- 714 KERSHAW, S., CRASQUIN, S., FOREL, M.-B., RANDON, C., COLLIN, P.-Y., KOSUN, E.,  
715 RICHOSZ, S. and BAUD, A. 2011. Earliest Triassic microbialites in Çürük Dag, southern  
716 Turkey: composition, sequences and controls on formation. *Sedimentology*, **58**, 739–755.
- 717 KIESSLING, W. and SIMPSON, C., 2011. On the potential for ocean acidification to be a  
718 general cause of ancient reef crises. *Global Change Biology*, **17**, 56–67.
- 719 KRISTAN-TOLLMANN, E. and TOLLMANN, A. 1982. Die Entwicklung der Tethystrias und  
720 Herkunft ihrer Fauna (The development of the Triassic within the Tethys realm and the  
721 origin of its fauna). *Geologische Rundschau*, **71**, 987–1019.
- 722 LEHRMANN, D. J. 1999. Early Triassic calcimicrobial mounds and biostromes of the  
723 Nanpanjiang basin, China. *Geology*, **27**, 359–362.
- 724 LEHRMANN, D. J., PAYNE, J. L., PEI, D., ENOS, P., DRUKE, D., STEFFEN, K., ZHANG,  
725 J., WEI, J., ORCHARD, M. and ELLWOOD, B. 2007. Record of the end-Permian  
726 extinction and Triassic biotic recovery in the Chongzuo-Pingguo platform, southern  
727 Nanpanjiang basin, Guangxi, south China. *Palaeogeography, Palaeoclimatology,*  
728 *Palaeoecology*, **252**, 200–217.
- 729 LEHRMANN, D. J., WAN, Y., WEI, J., YU, Y. and XIAO, J. 2001. Lower Triassic peritidal  
730 cyclic limestone: an example of anachronistic carbonate facies from the Great Bank of  
731 Guizhou, Nanpanjiang Basin, Guizhou province, South China. *Palaeogeography,*  
732 *Palaeoclimatology, Palaeoecology*, **173**, 103–123.
- 733 LEHRMANN, D. J., WEI, J. and ENOS, P. 1998. Controls on facies architecture of a large  
734 Triassic carbonate platform; the Great Bank of Guizhou, Nanpanjiang Basin, South China.  
735 *Journal of Sedimentology*, **68**, 311–326.
- 736 LI, M., HUANG, C., HINNOV, L., CHEN, W., OGG, J. and TIAN, W. 2018. Astrochronology

- 737 of the Anisian stage (Middle Triassic) at the Guandao reference section, South China. *Earth*  
738 *and Planetary Science Letters*, **482**, 591–606.
- 739 MARENCO, P. J., GRIFFIN, J. M., FRAISER, M. L., CLAPHAM, M. E., 2012. Paleoecology  
740 and geochemistry of Early Triassic (Spathian) microbial mounds and implications for  
741 anoxia following the end-Permian mass extinction. *Geology*, **40**, 715–718.
- 742 MARKOV, A. V. and KOROTAYEV, A. V. 2007. Phanerozoic marine biodiversity follows a  
743 hyperbolic trend. *Paleoworld*, **16**, 311–318.
- 744 MARTINDALE, E. C., FOSTER, W. J. and VELLEDEITS, F. 2019. The survival, recovery, and  
745 diversification of metazoan reefs following the end-Permian mass extinction event.  
746 *Palaeogeography, Palaeoclimatology, Palaeoecology*, **513**, 100–115.
- 747 MILLER, A. I. and SEPKOSKI, J. 1988. Modeling bivalve diversification: the effect of  
748 interaction on a macroevolutionary system. *Paleobiology*, **14**, 364–369.
- 749 NAWROCKI, J. and SZULC, J. 2000. The Middle Triassic magnetostratigraphy from the Peri-  
750 Tethys basin in Poland. *Earth and Planetary Science Letters*, **182**, 77–92.
- 751 NÜTZEL, A. 2005. Recovery of gastropods in the Early Triassic. *C. R. Palevol*, **4**, 501–515.
- 752 NÜTZEL, A. and SCHULBERT, C. 2005. Facies of two important Early Triassic gastropod  
753 lagerstätten: Implications for diversity patterns in the aftermath of the end-Permian mass  
754 extinction. *Facies*, **51**, 480–500.
- 755 NÜTZEL, A., KAIM, A. and GRĂDINARU, E. 2018. Middle Triassic (Anisian, Bithynian)  
756 gastropods from North Dobrogea (Romania) and their significance for gastropod recovery  
757 from the end-Permian mass extinction event. *Papers in Palaeontology*, 1–36.
- 758 OGILVIE GORDON, M. M. (1927). Das Grödener-, Fassa- und Enneberggebiet in den Südtiroler  
759 Dolomiten. Geologische Bundesanstalt Wien.
- 760 OJI, T. and TWITCHETT, R. J. 2015. The oldest post-Palaeozoic crinoid and Permian-Triassic  
761 origins of the articulata (echinodermata). *Zoological science*, **32**, 211–215.
- 762 ORCHARD, M. J. 2007. Conodont diversity and evolution through the latest Permian and Early  
763 Triassic upheavals. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **252**, 93–117.
- 764 PAYNE, J. L. 2005. Evolutionary dynamics of gastropod size across the end-Permian extinction  
765 and through the Triassic recovery interval. *Paleobiology*, **31**, 269–290.
- 766 PAYNE, J. L., LEHRMANN, D. J., CHRISTENSEN, S., WEI, J. and KNOLL, A. H. 2006a.  
767 Environmental and biological controls on the initiation and growth of a Middle Triassic  
768 (Anisian) reef complex on the Great Bank of Guizhou, Guizhou Province, China. *Palaaios*,  
769 **21**, 325–343.
- 770 PAYNE, J. L., LEHRMANN, D. J., FOLLETT, D., SEIBEL, M., KUMP, L. R., RICCARDI, A.,

- 771 ALTINER, D., SANO, H. and WEI, J. 2007. Erosional truncation of uppermost Permian  
772 shallow-marine carbonates and implications for Permian-Triassic boundary events.  
773 *Geological Society of America Bulletin*, **119**, 771–784.
- 774 PAYNE, J. L., LEHRMANN, D. J., WEI, J. and KNOLL, A. H. 2006b. The pattern and timing  
775 of biotic recovery from the end-Permian extinction on the Great Bank of Guizhou, Guizhou  
776 Province, China. *Palaios*, **21**, 63–85.
- 777 PAYNE, J. L., TURCHYN, A. V., PAYTAN, A., DEPAOLO, D. J., LEHRMANN, D. J., YU,  
778 M. and WEI, J. 2010. Calcium isotope constraints on the end-Permian mass extinction.  
779 *Proceedings of the National Academy of Science*, **107**, 8543–8548.
- 780 PETERS, S. E. and FOOTE, M. 2002. Determinants of extinction in the fossil record. *Nature*,  
781 **416**, 420–424.
- 782 PIROS, O. and PRETO, N. 2008. Dasycladalean algae distribution in ammonoid-bearing Middle  
783 Triassic platforms (Dolomites, Italy). *Facies*, **54**, 581–595.
- 784 PITACCO, V., ORLANDO-BONACA, M., MAVRIČ, B., POPOVIČ, A. and LIPEJ, L. 2014.  
785 Mollusc fauna associated with the Cystoseira algal associations in the Gulf of Trieste  
786 (Northern Adriatic Sea). *Mediterranean Marine Science*, **15**.
- 787 POPA, L., PANAIOTU, C. E. and GRĂDINARU, E. 2014. An early Middle Anisian (Middle  
788 Triassic) *Tubiphytes* and cement crusts-dominated reef from North Dobrogea (Romania):  
789 Facies, depositional environment and diagenesis. *Acta Geologica Polonica*, **64**, 189–206.
- 790 POULICEK, M. 1985. Les mollusques des biocénoses à algues photophiles en Méditerranée: II.  
791 *Analyse du peuplement Cahiers de biologie marine*, **26**, 127–136.
- 792 PRUSS, S. B. and BOTTJER, D. J. 2004. Early Triassic trace fossils of the western United States  
793 and their implications for prolonged environmental stress from the end-Permian mass  
794 extinction. *Palaios*, **19**, 551–564.
- 795 PRUSS, S. B., BOTTJER, D. J., CORSETTI, F. A. and BAUD, A. 2006. A global marine  
796 sedimentary response to the end-Permian mass extinction: examples from southern Turkey  
797 and the western United States. *Earth-Science Reviews*, **78**, 193–206.
- 798 PRUSS, S. B., PAYNE, J. L. and WESTACOTT, S. 2015. Taphonomic bias of selective  
799 silicification revealed by paired petrographic and insoluble analysis. *Palaios*, **30**, 620–626.
- 800 PRUSS, S. B., PAYNE, J. L., BOTTJER, D. J., 2007. *Placunopsis* bioherms; the first metazoan  
801 buildups following the end-Permian mass extinction. *Palaios*, **22**, 17–23.
- 802 RAUP, D. M. 1979. Size of the permo-triassic bottleneck and its evolutionary implications.  
803 *Science*, **206**, 217–218.
- 804 RODEN, V. J., HAUSMANN, I., NÜTZEL, A., SEUSS, B., REICH, M., URLICHS, M.,  
805 HAGDORN, H. and KIESSLING, W. 2020. Fossil liberation: A model to explain high

- 806 biodiversity in the Triassic Cassian Formation. *Palaeontology*, **63**, 85–102.
- 807 ROMANO, C., GOUEMAND, N., VENNEMANN, T., WARE, D., SCHNEEBELI-  
808 HERMANN, E., HOCHULI, P. A., BRÜHWILER, T., BRINKMANN, W. and BUCHER,  
809 H. 2013. Climate and biotic upheavals following the end-Permian mass extinction. *Nature*  
810 *Geoscience*, **6**, 57–60.
- 811 ROTHMAN, D. H., FOURNIER, G. P., FRECH, K. L., ALM, E. J., BOYLE, E. A., CAO, C.,  
812 SUMMONS, R.E. 2014. Methanogenic burst in the end-Permian carbon cycle. *Proceedings*  
813 *of the National Academy of Sciences*, **111**, 5462-5467.
- 814 RUSSO, F., MASTANDREA, A., STEFANI, M. and NERI, C. 2000. Carbonate facies  
815 dominated by syndepositional cements: a key component of Middle Triassic platforms. The  
816 Marmolada case history (Dolomites, Italy). *Facies*, **42**, 211–226.
- 817 SÁNCHEZ-MOYANO, J. E., ESTACIO, F. J., GARCÍA-ADIEGO, E. M. and GARCÍA-  
818 GÓMEZ, J. C. 2000. The molluscan epifauna of the alga *Halopteris scoparia* in southern  
819 Spain as a bioindicator of coastal environmental conditions. *Journal of Molluscan Studies*,  
820 **66**, 431–448.
- 821 SANEI, H., GRASBY, S. E., BEAUCHAMP, B. 2012. Latest Permian mercury anomalies.  
822 *Geology*, **40**, 63–66.
- 823 SCHEYER, T. M., ROMANO, C., JENKS, J. and BUCHER, H. 2014. Early Triassic Marine  
824 Biotic Recovery: The Predators’ Perspective. *PLoS ONE*, **9**, e88987,  
825 doi:10.1371/journal.pone.00889
- 826 SCHMIDT, M. 1928. *Die Lebewelt unserer Trias*. Hohenlohesche Buchhandlung F. Rau,  
827 Öhringen.
- 828 SCHMIDT, M. 1938. *Die Lebewelt unserer Trias - Nachtrag*. Hohenlohesche Buchhandlung F.  
829 Rau, Öhringen.
- 830 SCHOBEN, M., JOACHIMSKI, M. M., KORN, D., LEDA, L., KORTE, C. 2014. Palaeotethys  
831 seawater temperature rise and an intensified hydrological cycle following the end-Permian  
832 mass extinction. *Gondwana Research*, **26**, 675–682.
- 833 SCHUBERT, J. K. and BOTTJER, D. J. 1995. Aftermath of the Permian-Triassic mass  
834 extinction event: Paleoecology of Lower Triassic carbonates in the western USA.  
835 *Palaeogeography, Palaeoclimatology, Palaeoecology*, **116**, 1–39.
- 836 SELLWOOD, B. W. and FÜRSICH, F. T. 1981. Die Trias. In *Palökologie*, Franckh’sche  
837 Verlagshandlung, W. Keller & Co., Stuttgart, 248 pp.
- 838 SENOWBARI-DARYAN, B. and LINK, M. 2011. Hypercalcified segmented sponges  
839 (‘sphinctozoans’) from the upper Triassic (Norian) reef boulders of Taurus Mountains  
840 (southern Turkey). *Facies*, **57**, 663–693.



- 841 SENOWBARI-DARYAN, B., ZÜHLKE, R.; BECHSTÄDT, T., and FLÜGEL, E. 1993. Anisian  
842 (Middle Triassic) Buildups of the Northern Dolomites (Italy): The Recovery of Reef  
843 Communities after the Permian/Triassic Crisis. *Facies*, **28**, 181–256.
- 844 SEPKOSKI, J. J. 1978. A kinetic model of Phanerozoic taxonomic diversity. I. Analysis of  
845 marine orders. *Paleobiology*, **4**, 223–251.
- 846 SEPKOSKI, J. J. 1984. A kinetic model of Phanerozoic taxonomic diversity. III. Post-Paleozoic  
847 families and mass extinctions. *Paleobiology*, **10**, 246–267.
- 848 SEPKOSKI, J. J. 1997. Biodiversity: Past, present, and future. *Journal of Paleontology*, **71**, 533–  
849 539.
- 850 SHIGETA, Y., ZAKHAROV, Y. D., MAEDA, H. and POPOV, A. M. 2009. *The Lower Triassic*  
851 *System in the Abrek Bay area, South Primorye, Russia*. Vol. 38. National Museum of  
852 Nature and Science Monographs.
- 853 SALOMON, W. 1895. Geologische und palaeontologische Studien über die Marmolata.  
854 *Palaeontographica*, **42**, 1–210.
- 855 SOLÉ, R. V., SALDAÑA, J., MONTOYA, J. M. and ERWIN, D. H. 2010. Simple model of  
856 recovery dynamics after mass extinction. *Journal of Theoretical Biology*, **267**, 193–200.
- 857 SONG, H., WIGNALL, P. B., CHEN, Z.-Q., TONG, J., BOND, D. P. G., LAI, X., ZHAO, X.,  
858 JIANG, H., YAN, C., NIU, Z., CHEN, J., YANG, H. and WANG, Y. 2011. Recovery  
859 tempo and pattern of marine ecosystems after the end-Permian mass extinction. *Geology*,  
860 **39**, 739–742.
- 861 SONG, H., WIGNALL, P.B., CHU, D., TONG, J., SUN, Y., SONG, H., HE, W., & TIAN, L.  
862 2014. Anoxia/high temperature double whammy during the Permian-Triassic marine crisis  
863 and its aftermath. *Scientific reports*, **4**, 4132.
- 864 STANLEY, S. M. 2016. Estimates of the magnitudes of major marine mass extinctions in earth  
865 history. *Proceedings of the National Academy of Sciences*, **113**, E6325–E6334.
- 866 STILLER, F. 2001. Fossilvergesellschaftungen, Paläökologie und paläosynökologische  
867 Entwicklung im Oberen Anis (Mittlere Trias) von Qingyan, insbesondere Bangtoupou,  
868 Provinz Guizhou, Südwestchina. *Münstersche Forschung zur Geologie und Paläontologie*,  
869 **92**, 1–523.
- 870 SUN, J., JOACHIMSKI, M. M., WIGNALL, P. B., YAN, C., CHEN, Y., JIANG, H., WANG, L.  
871 and LAI, X. 2012. Lethally Hot Temperatures During the Early Triassic Greenhouse.  
872 *Science*, **338**, 366–370.
- 873 TAYLOR, J. D. 1971. Reef associated molluscan assemblages in the Western Indian Ocean. In  
874 STODDART, D. R. and YONGE, M. (eds.) *Regional Variation in Indian Ocean Coral*  
875 *Reefs*, Symposia of the Zoological Society of London, 501–534 pp.

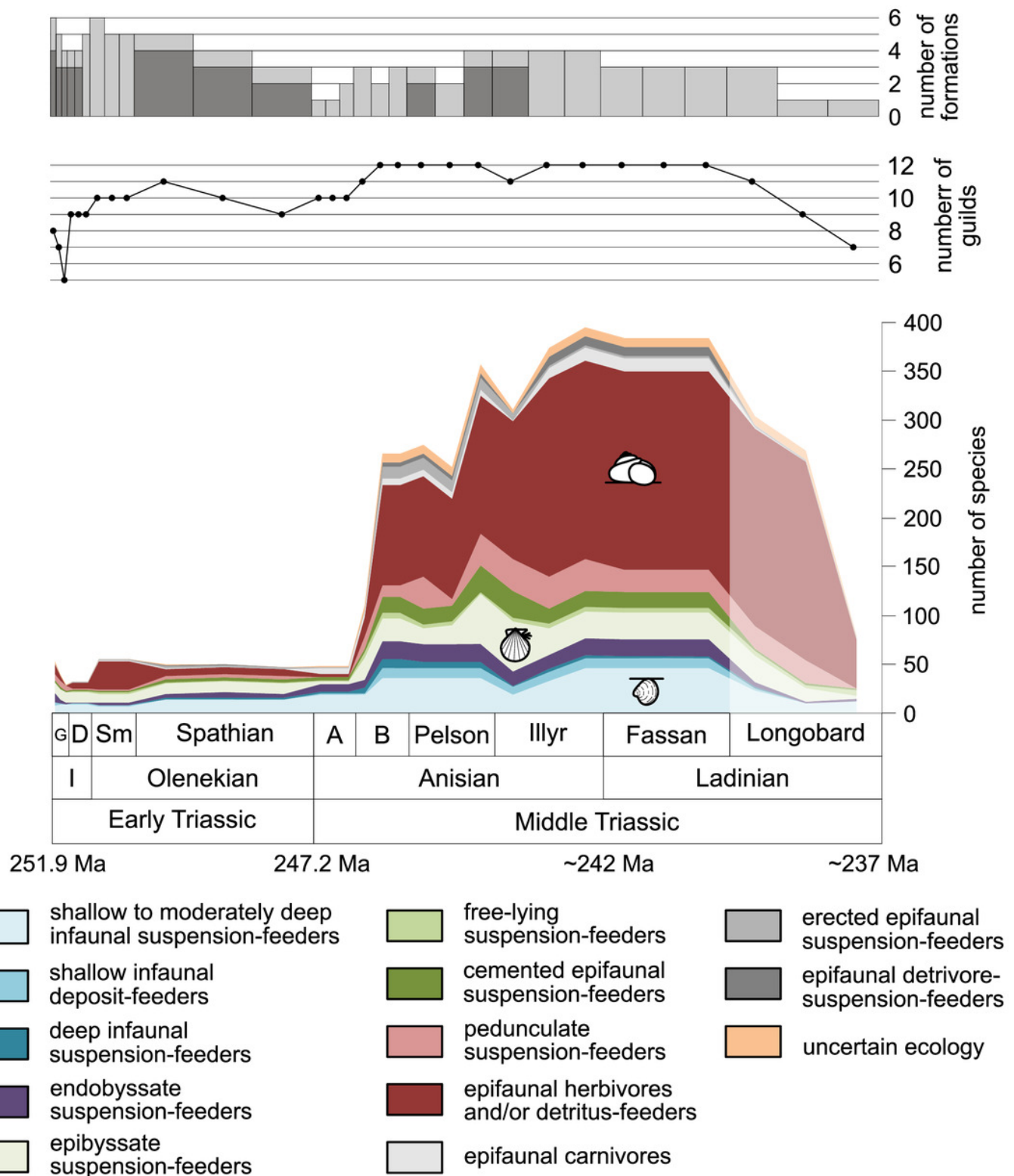
- 876 THOMASSIN, B. A. 1971. Les faciès d'épifauna et d'épiflore des biotopes sédimentaires des  
877 formations coralliennes dans la région de Tuléar (sud-ouest de Madagascar). In  
878 STODDARD, D. R. and YONGE, M. (eds.) *Regional Variation in Indian Ocean Coral*  
879 *Reefs*, London, 501–534 pp.
- 880 TIAN, L., TONG, J., ALGEO, T. J., SONG, H., SONG, H., CHU, D., SHI, L. and BOTTJER, D.  
881 J. 2014. Reconstruction of Early Triassic ocean redox conditions based on framboidal pyrite  
882 from the Nanpanjiang Basin, South China. *Palaeogeography, Palaeoclimatology,*  
883 *Palaeoecology*, **412**, 68–79.
- 884 TWITCHETT, R. J. and WIGNALL, P. B. 1996. Trace fossils and the aftermath of the Permo-  
885 Triassic mass extinction: Evidence from northern Italy. *Palaeogeography,*  
886 *Palaeoclimatology, Palaeoecology*, **124**, 137–151.
- 887 TWITCHETT, R. J., KRYSTYN, L., BAUD, A., WHEELLEY, J. R. and RICHOZ, S. 2004.  
888 Rapid marine recovery after the end-Permian mass-extinction event in the absence of  
889 marine anoxia. *Geology*, **32**, 805–808.
- 890 UNDERWOOD, A. J. 1980. The effects of grazing by gastropods and physical factors on the  
891 upper limits of distribution of intertidal macroalgae. *Oecologia*, **46**, 201–213.
- 892 URLICH, M. 2014. A phylogenetic lineage of *Myophoria kefersteini* subspecies (Bivalvia)  
893 from the Late Ladinian to the Early Carnian (Triassic) in the Dolomites and Julian Alps  
894 (Italy). *Neues Jahrbuch für Geologie und Paläontologie, Abhandlungen*, **274**, 187–207.
- 895 URLICH, M. 2017. Revision of some stratigraphically relevant ammonoids from the Cassian  
896 Formation (latest Ladinian-Early Carnian, Triassic) of St. Cassian (Dolomites, Italy). *Neues*  
897 *Jahrbuch für Geologie und Paläontologie, Abhandlungen*, **283**, 173–204.
- 898 VALENTINE, J. W. 1971. Resource supply and species diversity patterns. *Lethaia*, **4**, 51–61.
- 899 VELLEITIS, F., HIPS, K. and PÉRO, C. 2012. Lower and Middle Triassic succession in  
900 Aggtelek Karst. *IGCP 572 field trip POST 2. June 5-7*, 1–33.
- 901 VELLEITIS, F., PÉRO, C., BLAU, J., SENOWBARI-DARYAN, B., KOVÁCS, S., PIROS, O.,  
902 POCSAI, T., SZÜGYI-SIMON, H., DUMITRICĂ, P. and PÁLFY, J. 2011. The oldest  
903 Triassic platform margin reef from the Alpine - Carpathian region (Aggtelek, NE Hungary):  
904 Platform evolution, reefal biota and biostratigraphic framework. *Rivista Italiana di*  
905 *Paleontologia e Stratigrafia*, **117**, 221–268.
- 906 VERON, J. E., DEVANTIER, L. M., TURAK, E., GREEN, A. L., KININMONTH, S.,  
907 STAFFORD-SMITH, M. and PETERSON, N. 2009. Delineating the coral triangle.  
908 *Galaxea, Journal of Coral Reef Studies*, **11**, 91–100.
- 909 WASMER, M., HAUTMANN, M., HERMANN, E., WARE, D., ROOHI, G., UR-REHMAN,  
910 K., YASEEN, A. and BUCHER, H. 2012. Olenekian (Early Triassic) bivalves from the Salt  
911 Range and Surghar Range, Pakistan. *Palaeontology*, **55**, 1043–1073.

- 912 WEI, H., SHEN, J., SCHOEPFER, S.D., KRYSTYN, L., RICHOS, S., & ALGEO, T.J. 2015.  
913 Environmental controls on marine ecosystem recovery following mass extinctions, with an  
914 example from the Early Triassic. *Earth-Science Reviews*, **149**, 108–135.
- 915 WHEELLEY, J. R. and TWITCHETT, R. J. 2005. Palaeoecological significance of a nes  
916 Griesbachian (Early Triassic) gastropod assemblage from Oman. *Lethaia*, **38**, 37–45.
- 917 WIGNALL, P. B. and BENTON, M. J. 1999. Lazarus taxa and fossil abundances at times of  
918 biotic crises. *Journal of the Geological Society*, **156**, 453–456.
- 919 WIGNALL, P. B. and TWITCHETT, R. J.,1996. Oceanic anoxia and the end-Permian mass  
920 extinction. *Science* **272**, 1155–1158.
- 921 WILLIAMS, G. A. 1993. Seasonal variation in algal species richness and abundance in the  
922 presence of molluscan herbivores on a tropical rock shore. *Journal of Experimental Marine*  
923 *Biology and Ecology*, **167**, 261–275.
- 924 WU, Y., JIANG, H., YANG, W. and FAN, J., 2007. Microbialite of anoxic condition from  
925 Permian-Triassic transition in Guizhou, China. *Science in China Series D: Earth Sciences*,  
926 **50**, 1040–1051.
- 927 YANG, H., CHEN, Z.-Q., WANG, Y., OU, W., LIAO, W. and MEI , X., 2015. Palaeoecology of  
928 microconchids from microbialites near the Permian–Triassic boundary in South China.  
929 *Lethaia*, **48**, 497–508.

# Figure 1

Maximum species richness throughout the Early and Middle Triassic.

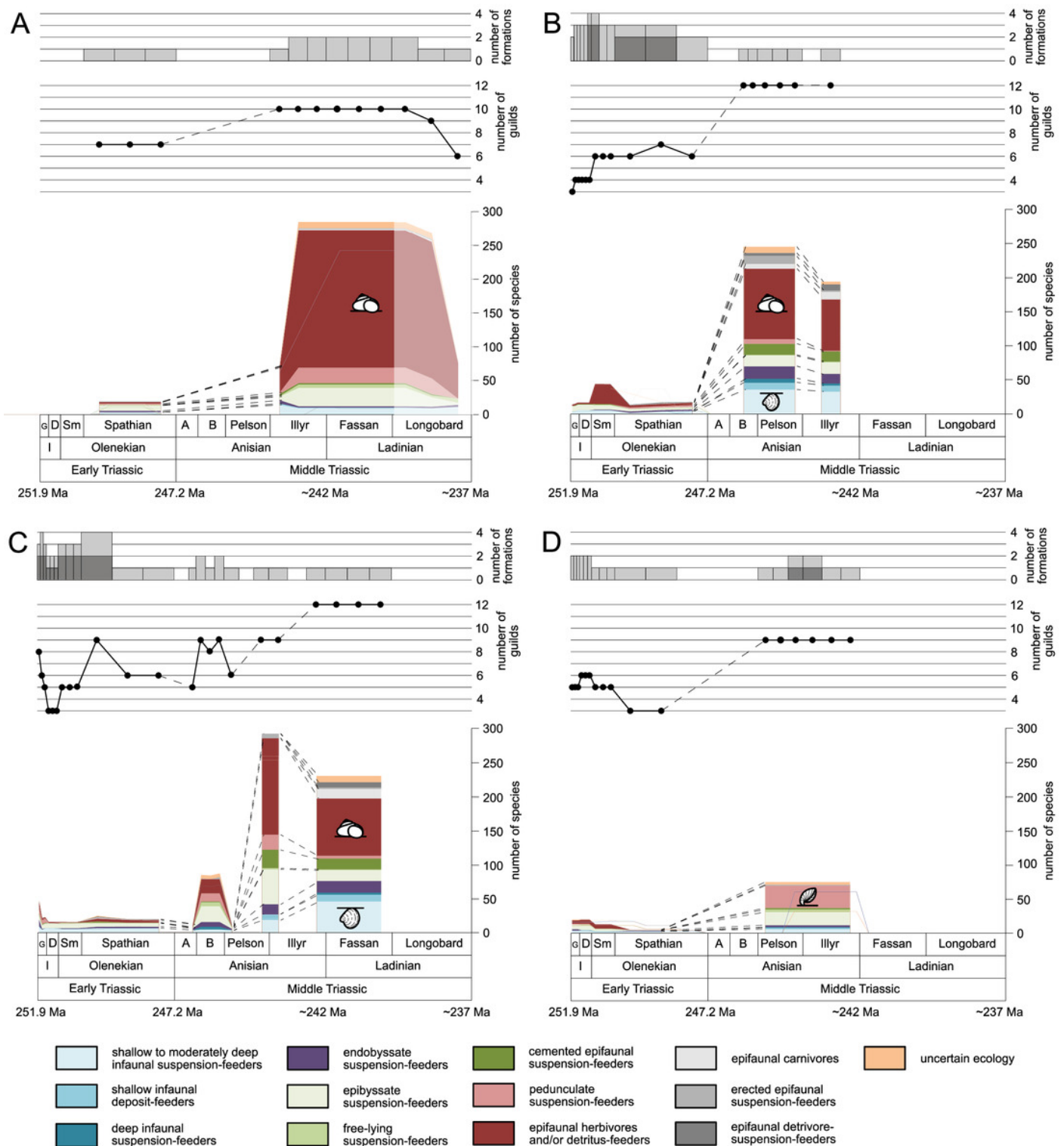
The Longobardian is palish to indicate that the decrease of species richness is not a primary signal. Light grey bars on top indicate the number of lithological units representing each time interval. In cases where information does not come from all lithological units representing the corresponding time interval, the actual number of lithological units from which information comes from is indicated by dark grey bars. Abbreviations: A = Aegean, B = Bithynian, D = Dienerian, Fassan = Fassanian, G = Griesbachian, I = Induan, Illyr = Illyrian, Longobard = Longobardian, Pelson = Pelsonian, Sm = Smithian. Absolute ages according to (Cohen *et al.*, 2013, *updated*). Information about the relative duration of substages come from Nawrocki & Szulc (2000), Götz, Szulc & Feist-Burkhardt (2005), Jattiot *et al.* (2018) and Li *et al.* (2018). Symbols for ecological guilds redrawn after Aberhan (1994). See text for further explanation.



# Figure 2

Maximum species richness in different environments.

(A) lagoons, (B) inner ramps, (C) mid-ramps and (D) outer ramps. The Longobardian is palish to indicate that the decrease of species richness is not a primary signal. Light grey bars on top indicate the number of lithological units representing each time interval. In cases where information does not come from all lithological units representing the corresponding time interval, the actual number of lithological units from which information comes from is indicated by dark grey bars. Due to our data selection, the following guilds are not shown in this figure although reported in polished slabs: shallow infaunal deposit-feeders in late Griesbachian-Smithian and middle Spathian inner ramps (one species, respectively) and Smithian mid-ramps (one species), cemented epifaunal suspension-feeders in early-middle Spathian inner ramps (one species) and Dienerian-Smithian outer ramps (two species), epifaunal herbivores and/or detritus-feeders in early-middle Spathian (two species) and late Pelsonian-early Illyrian (five species) outer ramps. Abbreviations, symbols for ecological guilds and information about the relative duration of substages as in Figure 1. See text for further explanation.

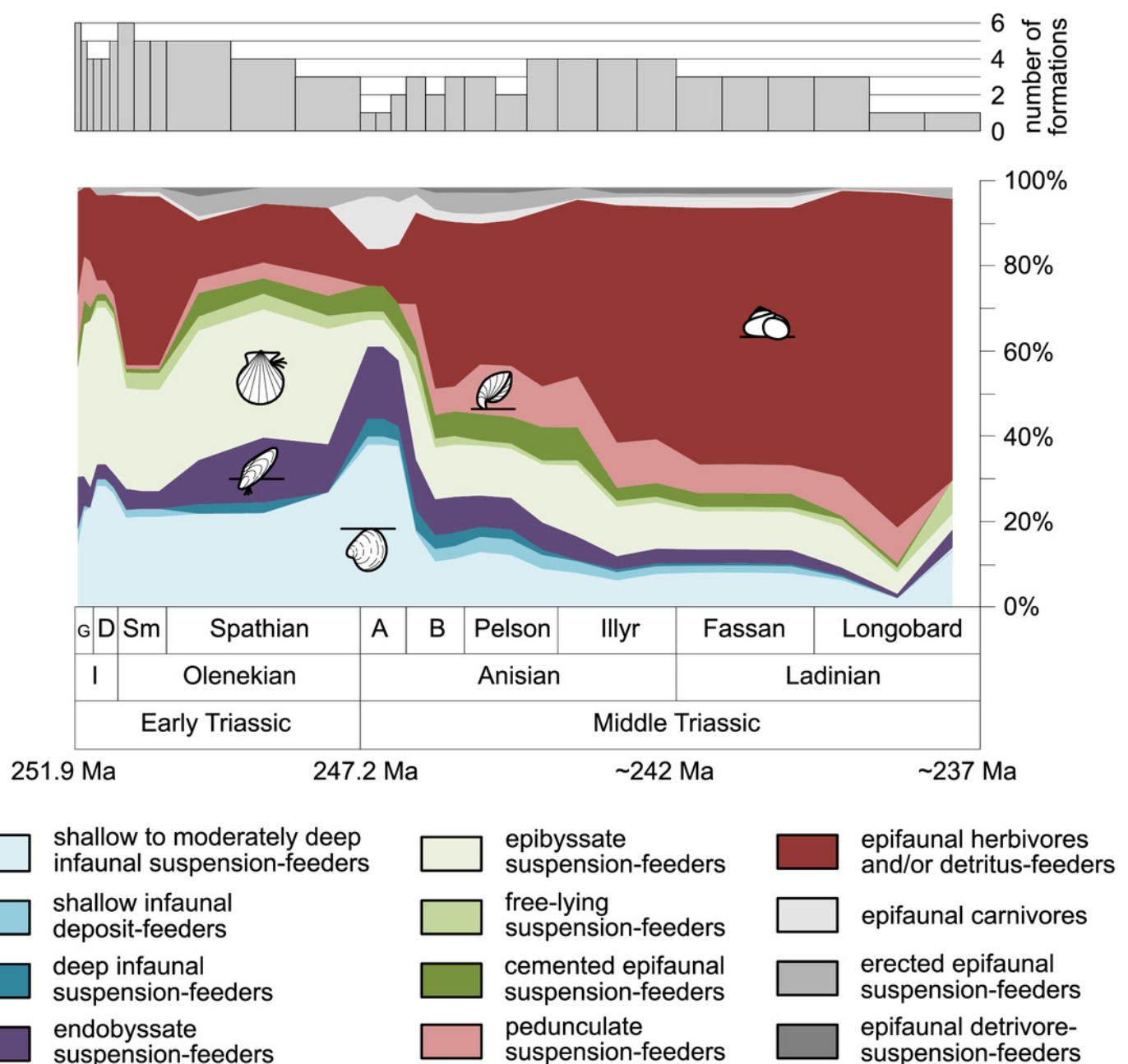


# Figure 3

Guild-species diversity throughout the Early and Middle Triassic.

Bars on top indicate the number of lithological units representing each time interval.

Abbreviations, symbols for ecological guilds and information about the relative duration of substages as in Figure 1. See text for further explanation.

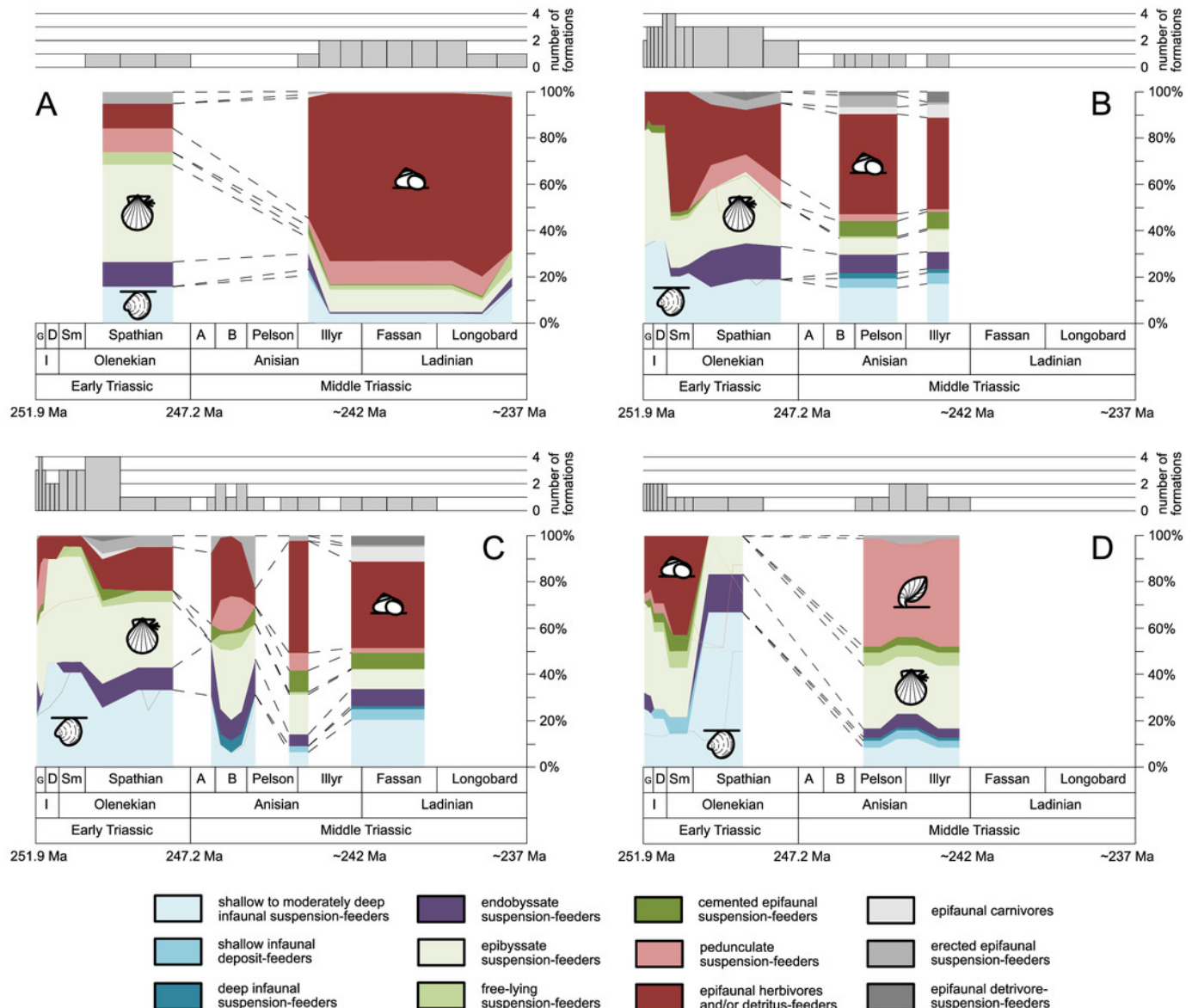




# Figure 4

Guild-species diversity of different environments.

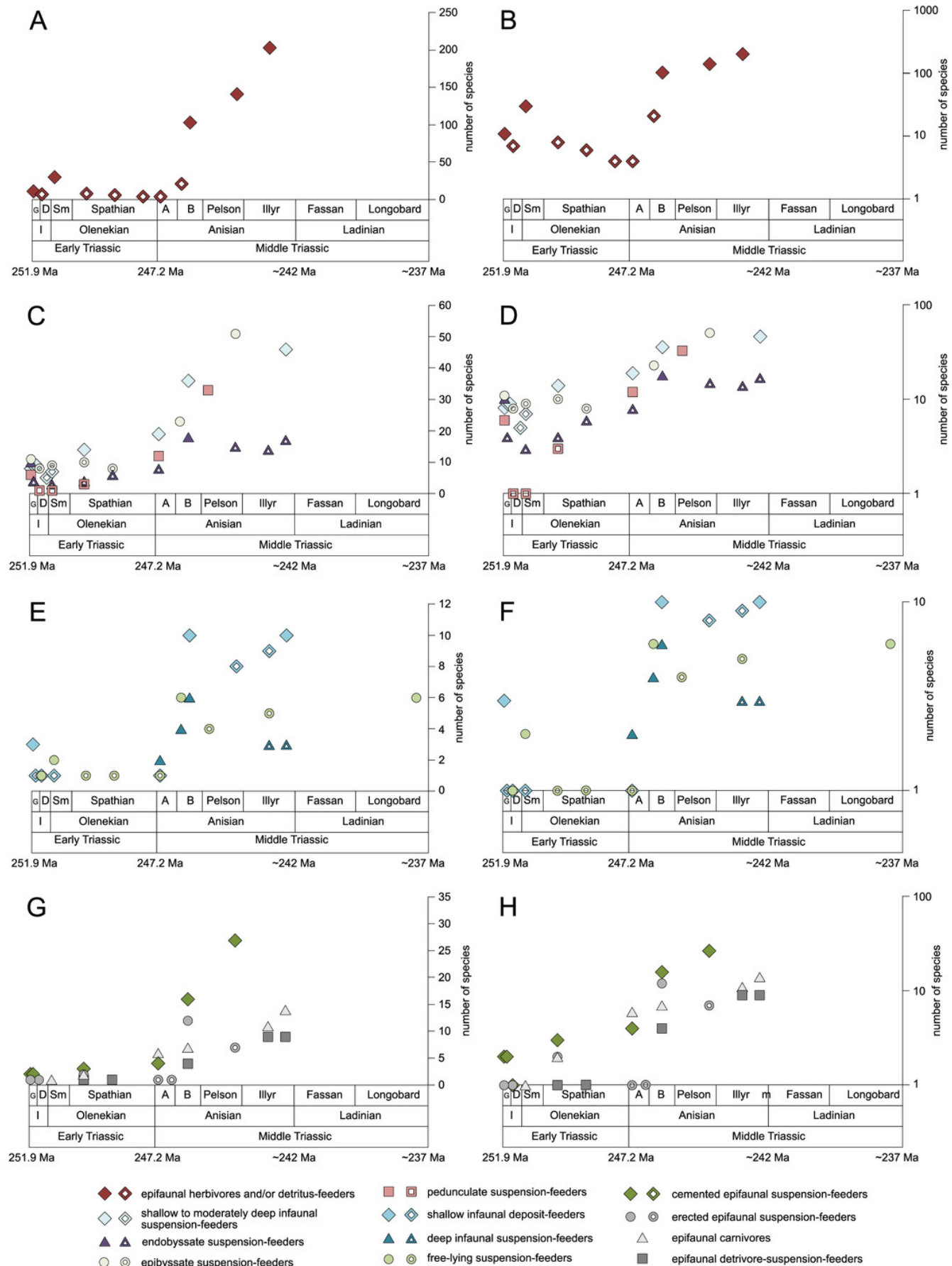
(A) lagoons, (B) inner ramps, (C) mid-ramps and (D) outer ramps. Bars on top indicate the number of lithological units representing each time interval. Due to our data selection, the following guilds are not shown in this figure although reported in polished slabs: shallow infaunal deposit-feeders in late Griesbachian-Smithian and middle Spathian inner ramps (one species, respectively) and Smithian mid-ramps (one species), cemented epifaunal suspension-feeders in early-middle Spathian inner ramps (one species) and Dienerian-Smithian outer ramps (two species), epifaunal herbivores and/or detritus-feeders in early-middle Spathian (two species) and late Pelsonian-early Illyrian (five species) outer ramps. Abbreviations, symbols for ecological guilds and information about the relative duration of substages as in Figure 1. See text for further explanation.



# Figure 5

Maximum species richness throughout the Early and Middle Triassic of ecological guilds.

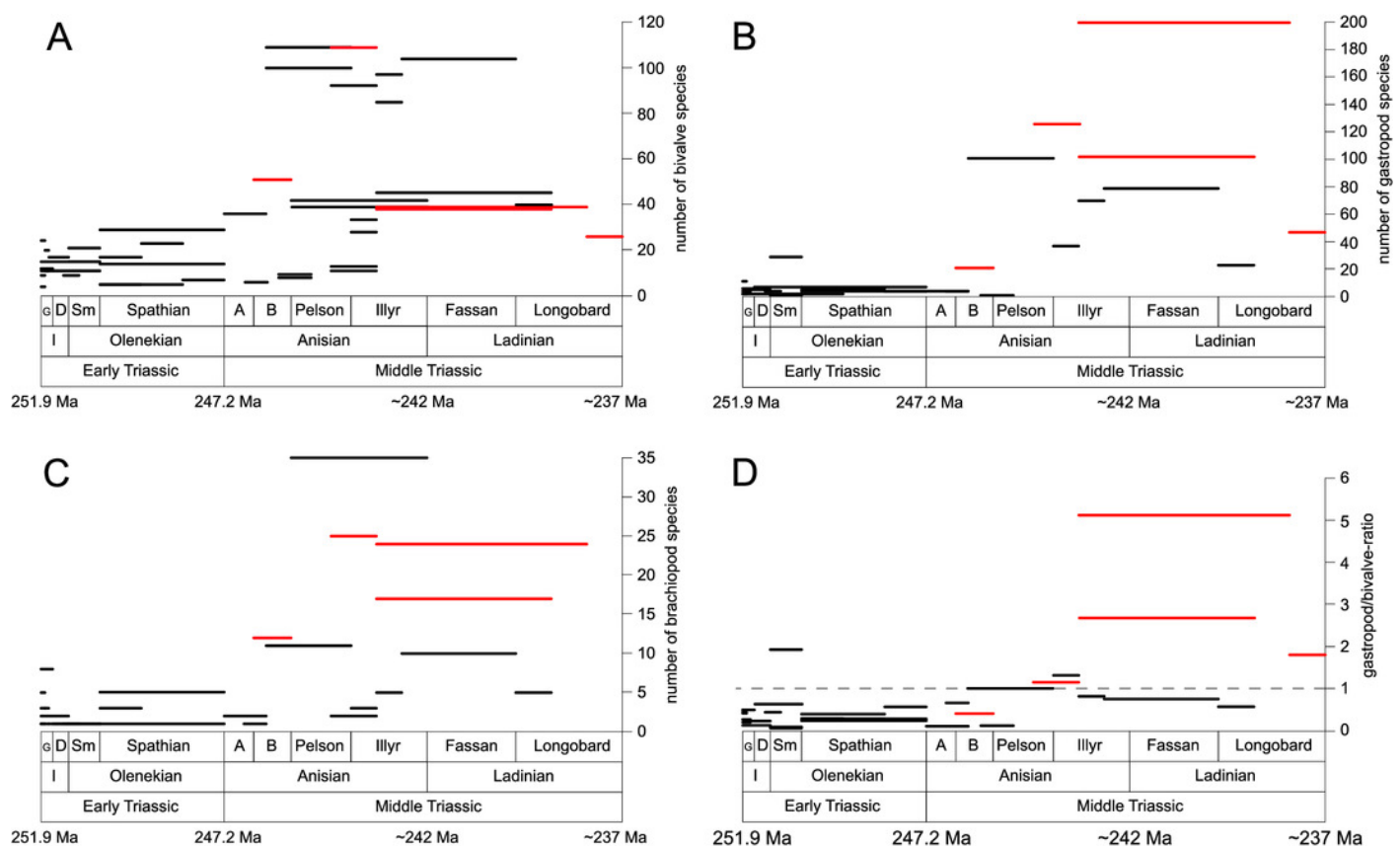
(A-B) epifaunal herbivores and/or detritus-feeders, (C-D) shallow to moderately deep infaunal suspension-feeders, endobyssate suspension-feeders, epibyssate suspension-feeders and pedunculate suspension-feeders, (E-F) shallow infaunal deposit-feeders, deep infaunal suspension-feeders, free-lying suspension-feeders and cemented microcarnivores, (G-H) cemented epifaunal suspension-feeders, erected epifaunal suspension-feeders, epifaunal carnivores and epifaunal-detritivore-suspension-feeders. Number of species illustrated with linear scale (A, C, E and G) and logarithmic scale (B, D, F and H). Filled symbols represent faunas that are more or as diverse as older ones, whereas empty symbols represent faunas that are less diverse than older ones. Abbreviations and information about the relative duration of substages as in Figure 1. See text for further explanation.



# Figure 6

Species richness of the most diverse taxa and the gastropod/bivalve-ratio throughout the Early and Middle Triassic.

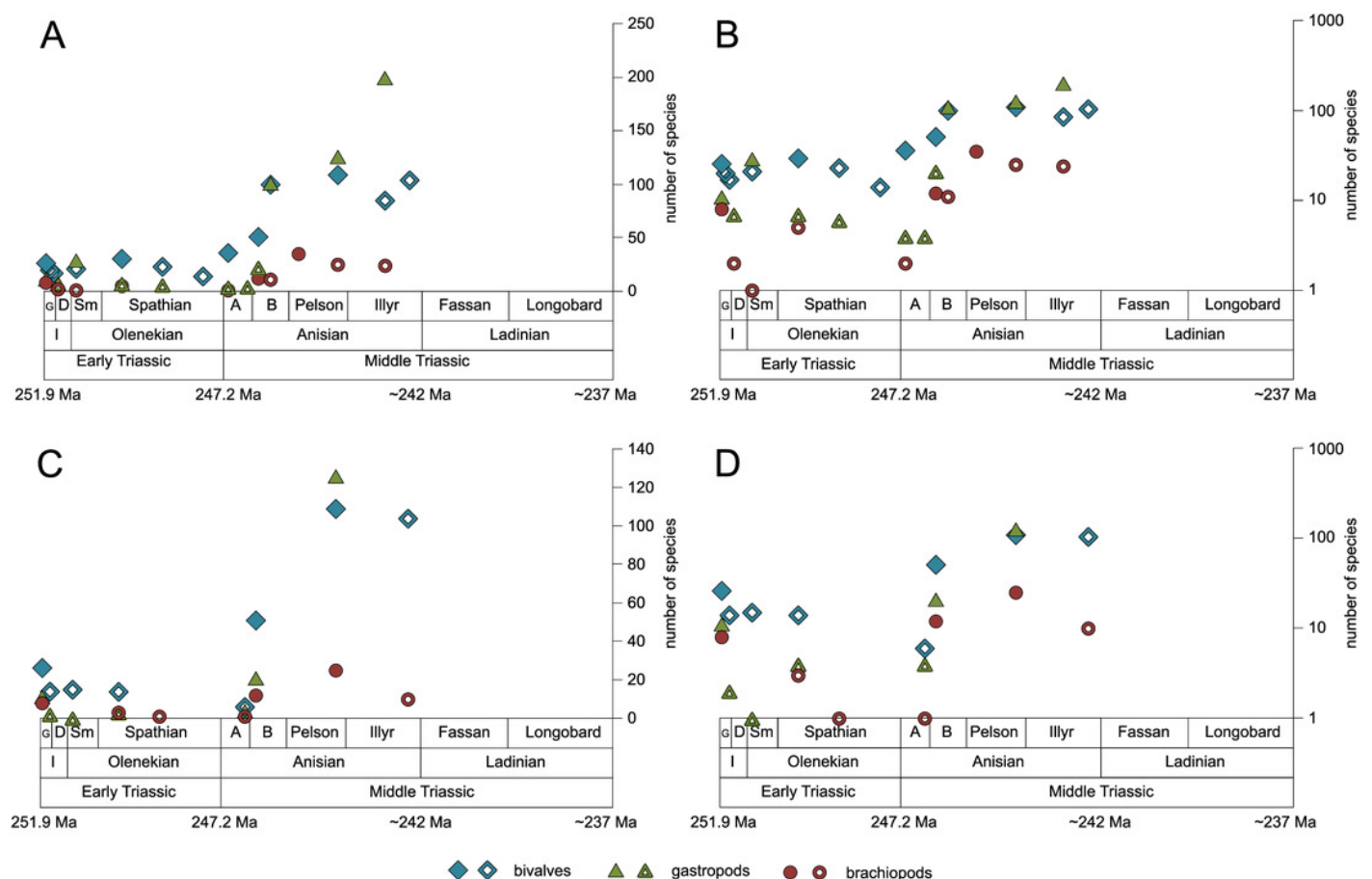
(A) bivalves, (B) gastropods, (C) brachiopods and (D) the gastropod/bivalve-ratio. Each line represents one lithological unit. Red lines mark lithological units and ratios associated with carbonate platforms. Abbreviations and information about the relative duration of substages as in Figure 1. (A) modified after *Friesenbichler et al. (2021)* and (D) modified after *Friesenbichler et al. (2019b)*. See text for further explanation.



# Figure 7

Maximum species richness throughout the Early and Middle Triassic of bivalves, gastropods and brachiopods.

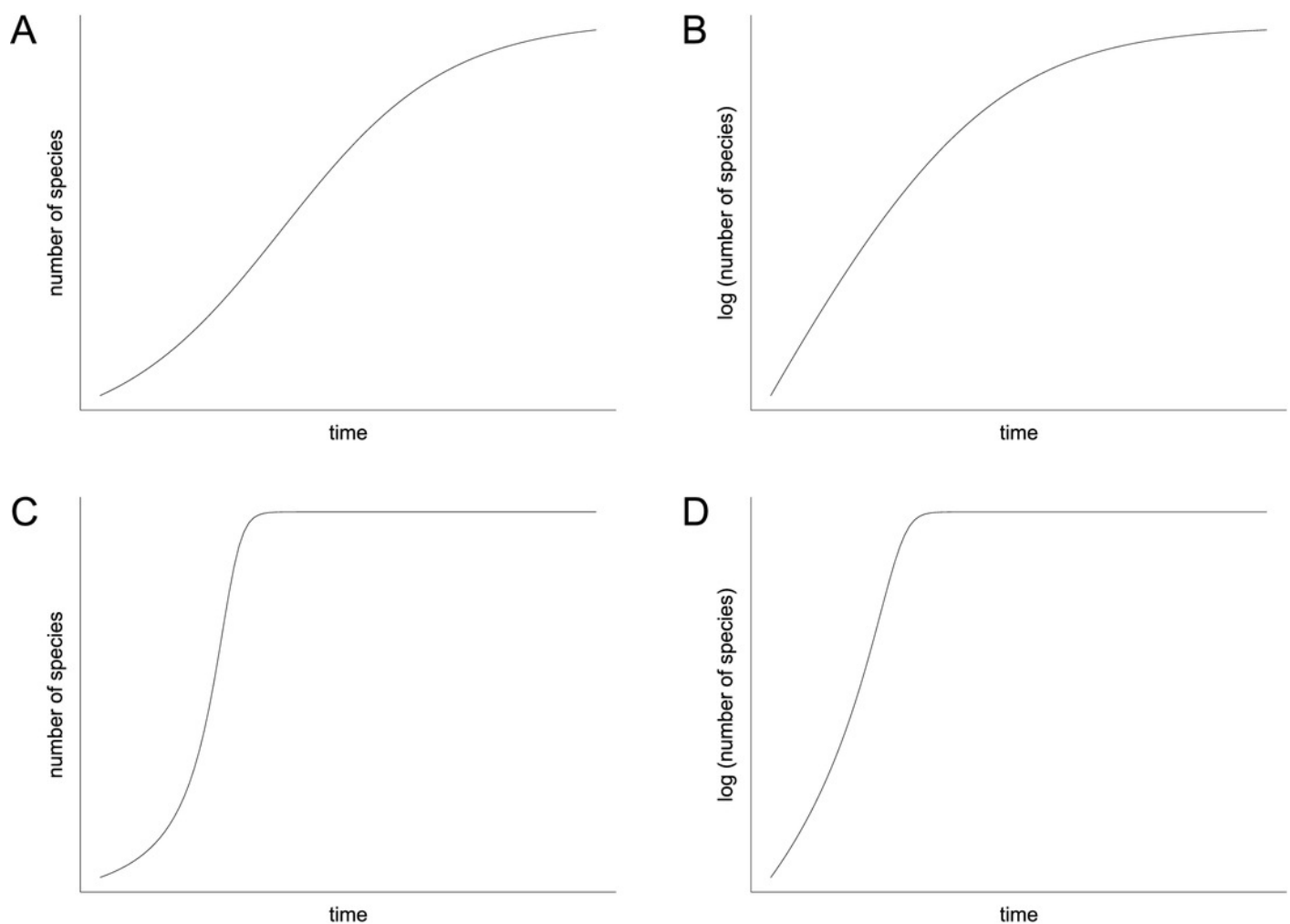
(A-B) in general and (C-D) in mid-ramps. Number of species illustrated with linear scale (A and C) and logarithmic scale (B and D). Filled symbols represent faunas that are more or as diverse as older ones, whereas empty symbols represent faunas that are less diverse than older ones. Abbreviations and information about the relative duration of substages as in Figure 1. See text for further explanation.



# Figure 8

Diversification curves.

(A-B) Logistic with number of species illustrated with linear scale (A) and logarithmic scale (B). (C-D) Hyperbolic-damped with number of species illustrated with linear scale (C) and logarithmic scale (D).



**Table 1** (on next page)

List of ecological guilds and representing taxa.



<b>Ecological guild</b>	<b>Taxa</b>
Shallow to moderately deep infaunal suspension-feeders	Bivalves, inarticulate brachiopods
Shallow infaunal deposit-feeders	Bivalves, scaphopods
Deep infaunal suspension-feeders	Bivalves
Endobyssate suspension-feeders	Bivalves
Epibyssate suspension-feeders	Bivalves
Free-lying epifaunal suspension-feeders	Bivalves
Cemented epifaunal suspension-feeders	Bivalves, serpulids, inarticulate brachiopods, bryozoans, microconchids
Pedunculate suspension-feeders	Articulate brachiopods
Epifaunal herbivores and/or detritus-feeders	Gastropods, echinoids, polyplacophors
Epifaunal carnivores	Asteroids, crustaceans
Erected epifaunal suspension-feeders	Crinoids
Epifaunal detritivore-suspension-feeders	Ophiuroids