

# Temperate southern Australian coastal waters are characterised by surprisingly high rates of nitrogen fixation and diversity of diazotrophs

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Biological dinitrogen ( $N_2$ ) fixation is one mechanism by which specific microorganisms (diazotrophs) can ameliorate nitrogen (N) limitation. Historically, rates of  $N_2$  fixation were believed to be limited outside of the low nutrient tropical and subtropical open ocean, however, emerging evidence suggests that  $N_2$  fixation is also a significant process within temperate coastal waters. Using a combination of amplicon sequencing, targeting the nitrogenase reductase gene (*nifH*), quantitative *nifH* PCR, and  $^{15}N_2$  stable isotope tracer experiments, we investigated spatial patterns of diazotroph assemblage structure and  $N_2$  fixation rates within the temperate coastal waters of southern Australia during Austral autumn and summer. Relative to previous studies in open ocean environments, including tropical northern Australia, and tropical and temperate estuaries, our results indicate that high rates of  $N_2$  fixation ( $10 - 64 \text{ nmol L}^{-1} \text{ d}^{-1}$ ) can occur within the large inverse estuary Spencer Gulf, while comparatively low rates of  $N_2$  fixation ( $2 \text{ nmol L}^{-1} \text{ d}^{-1}$ ) were observed in the adjacent continental shelf waters. Across the dataset, low concentrations of  $NO_3/NO_2$  were significantly correlated with the highest  $N_2$  fixation rates, suggesting that  $N_2$  fixation could be an important source of new N in the region as dissolved inorganic N concentrations are typically limiting. Overall, the underlying diazotrophic community was dominated by *nifH* sequences from Cluster 1 unicellular cyanobacteria of the UCYN-A clade, as well as non-cyanobacterial diazotrophs related to *Pseudomonas stutzeri*, and Cluster 3 sulfate-reducing deltaproteobacteria. Diazotroph community composition was significantly influenced by salinity and  $SiO_4$  concentrations, reflecting the transition from UCYN-A-dominated assemblages in the continental shelf waters, to Cluster 3-dominated

assemblages in the hypersaline waters of the inverse estuary. Diverse, transitional diazotrophic communities, comprised of a mixture of UCYN-A and putative heterotrophic bacteria, were observed at the mouth and southern edge of Spencer Gulf, where the highest  $N_2$  fixation rates were observed. In contrast to observations in other environments, no seasonal patterns in  $N_2$  fixation rates and diazotroph community structure were apparent. Collectively, our findings are consistent with the emerging view that  $N_2$  fixation within temperate coastal waters is a previously overlooked dynamic and potentially important component of the marine N cycle.

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38

**39 Abstract**

40 Biological dinitrogen ( $N_2$ ) fixation is one mechanism by which specific microorganisms  
41 (diazotrophs) can ameliorate nitrogen (N) limitation. Historically, rates of  $N_2$  fixation were  
42 believed to be limited outside of the low nutrient tropical and subtropical open ocean, however,  
43 emerging evidence suggests that  $N_2$  fixation is also a significant process within temperate coastal  
44 waters. Using a combination of amplicon sequencing, targeting the nitrogenase reductase gene  
45 (*nifH*), quantitative *nifH* PCR, and  $^{15}N_2$  stable isotope tracer experiments, we investigated spatial  
46 patterns of diazotroph assemblage structure and  $N_2$  fixation rates within the temperate coastal  
47 waters of southern Australia during Austral autumn and summer. Relative to previous studies in  
48 open ocean environments, including tropical northern Australia, and tropical and temperate  
49 estuaries, our results indicate that high rates of  $N_2$  fixation ( $10 - 64 \text{ nmol L}^{-1} \text{ d}^{-1}$ ) can occur within  
50 the large inverse estuary Spencer Gulf, while comparatively low rates of  $N_2$  fixation ( $2 \text{ nmol L}^{-1}$   
51  $\text{d}^{-1}$ ) were observed in the adjacent continental shelf waters. Across the dataset, low  
52 concentrations of  $NO_3/NO_2$  were significantly correlated with the highest  $N_2$  fixation rates,  
53 suggesting that  $N_2$  fixation could be an important source of new N in the region as dissolved  
54 inorganic N concentrations are typically limiting. Overall, the underlying diazotrophic  
55 community was dominated by *nifH* sequences from Cluster 1 unicellular cyanobacteria of the  
56 UCYN-A clade, as well as non-cyanobacterial diazotrophs related to *Pseudomonas stutzeri*, and  
57 Cluster 3 sulfate-reducing deltaproteobacteria. Diazotroph community composition was  
58 significantly influenced by salinity and  $SiO_4$  concentrations, reflecting the transition from  
59 UCYN-A-dominated assemblages in the continental shelf waters, to Cluster 3-dominated  
60 assemblages in the hypersaline waters of the inverse estuary. Diverse, transitional diazotrophic  
61 communities, comprised of a mixture of UCYN-A and putative heterotrophic bacteria, were  
62 observed at the mouth and southern edge of Spencer Gulf, where the highest  $N_2$  fixation rates  
63 were observed. In contrast to observations in other environments, no seasonal patterns in  $N_2$   
64 fixation rates and diazotroph community structure were apparent. Collectively, our findings are  
65 consistent with the emerging view that  $N_2$  fixation within temperate coastal waters is a  
66 previously overlooked dynamic and potentially important component of the marine N cycle.

67

**68 Introduction**

69 By providing a source of new nitrogen (N), dinitrogen ( $N_2$ ) fixation, the microbially mediated  
70 conversion of  $N_2$  gas to ammonia, represents a fundamental process within oligotrophic marine  
71 environments (Eugster and Gruber, 2012; Karl et al., 2012). Based on global ocean estimates, the  
72 activity of  $N_2$  fixing microorganisms (termed diazotrophs) contributes approximately 160 Tg of  
73 new N to the ocean annually (Wang et al., 2019). However, the majority of empirical  
74 observations contributing towards global  $N_2$  fixation estimates have been derived from tropical  
75 and subtropical oceanic gyres (Luo et al., 2012), which have traditionally been deemed the  
76 principal ecological niche for marine  $N_2$  fixation due to the activity of large filamentous and  
77 heterocystous cyanobacterial diazotrophs (Breitbarth et al., 2007; Goebel et al., 2010; Karl et al.,

78 2002). In contrast, temperate coastal habitats have generally been thought to be enriched in  
79 dissolved inorganic N derived from terrestrial and upwelled sources (Jickells, 1998), thereby  
80 restricting the niche for biological N<sub>2</sub> fixation.

81

82 Temperate coastal waters are some of the most productive regions on Earth, which have  
83 historically been believed to be fuelled by the influx of bioavailable nutrients from rivers,  
84 groundwater, and through the mixing of offshore waters (Jickells, 1998). Often these  
85 hydrodynamic properties result in the development and maintenance of relatively eutrophic  
86 conditions, which in combination with typically cool sea surface temperatures, resulted in the  
87 supposition that diazotrophic growth and activity, particularly for the large filamentous  
88 cyanobacterium *Trichodesmium sp.*, would be limited (Breitbarth et al., 2007; Howarth et al.,  
89 1988; Knapp, 2012). However, largely due to the newly recognised abundance and activity of  
90 non-cyanobacterial diazotrophs and unicellular cyanobacteria (UCYN) outside of the traditional  
91 oceanic niches of N<sub>2</sub> fixation, there has been a recent paradigm shift in the potential importance  
92 of N<sub>2</sub> fixation in temperate coastal regions, where annual N<sub>2</sub> fixation rates have been estimated to  
93 exceed 16 Tg N (Tang et al., 2019b). Therefore, an enhanced understanding of N<sub>2</sub> fixation rates  
94 and patterns in diazotroph diversity and activity within temperate coastal habitats is required to  
95 inform models of marine N availability at regional and global scales.

96

97 The widespread application of molecular surveys targeting the gene encoding a subunit of the  
98 nitrogenase enzyme complex (*nifH*), have greatly expanded the known diversity and global  
99 distribution of numerically important diazotrophs (Cornejo-Castillo et al., 2018; Farnelid et al.,  
100 2011; Langlois et al., 2015; Moisaner et al., 2010; Zehr et al., 1998, 2000, 2003). For example,  
101 *nifH* containing UCYN clades, *Candidatus* Atelocyanobacterium thalassa (UCYN-A), UCYN-B,  
102 and UCYN-C, and putative heterotrophic, non-cyanobacterial diazotrophs, from the gamma-,  
103 delta-, and alphaproteobacteria, have recently been detected throughout the major ocean basins  
104 (Díez et al., 2012; Farnelid et al., 2013; Fernández-Méndez et al., 2016; Gradoville et al., 2017;  
105 Langlois et al., 2015). Investigations into the ecology of these novel diazotrophs have revealed a  
106 range of physiologies and patterns of biogeography. Both free-living (e.g. UCYN-B, and C; Zehr  
107 et al., 2001; Taniuchi et al., 2012; Stenegren et al., 2018) and symbiotic (e.g. UCYN-A) UCYN  
108 groups have been identified, and evidence suggests a diversity of closely related sub-lineages are  
109 representative of distinct ecological niches typically associated with “open ocean” (e.g. UCYN-  
110 A1) and “coastal” (e.g. UCYN-A2) environments (Cornejo-Castillo et al., 2018; Farnelid et al.,  
111 2016; Thompson et al., 2014).

112

113 The isolation of non-cyanobacterial diazotrophs from specific habitats, such as the Peruvian  
114 oxygen minimum zone (Martinez-Perez et al., 2018), and estuarine waters of the Baltic Sea  
115 (Bentzon-Tilia et al., 2014; Farnelid et al., 2014), imply relatively specialised niches for these  
116 organisms. However, genomic analysis of isolates, and metagenome-assembled genomes from  
117 the alphaproteobacteria and Planctomycetes, have revealed the metabolic flexibility of these

118 groups, particularly in regard to their N cycling capabilities (Delmont et al., 2018; Martinez-  
119 Perez et al., 2018). Non-cyanobacterial diazotrophs are distributed throughout tropical and  
120 temperate latitudes and are sometimes the dominant members of diazotrophic communities  
121 (Bombar et al., 2016; Delmont et al., 2018; Langlois et al., 2015; Moisander et al., 2014).  
122 Notably, both non-cyanobacterial diazotrophs and UCYN have recently been identified as  
123 important constituents of temperate and coastal diazotroph communities (Bentzon-Tilia et al.,  
124 2015b; Messer et al., 2015; Mulholland et al., 2012; Needoba et al., 2007; Rees et al., 2009;  
125 Shiozaki et al., 2015a; Short and Zehr, 2007), with their presence often associated with high rates  
126 of N<sub>2</sub> fixation (Bentzon-Tilia et al., 2015b; Tang et al., 2019b).

127  
128 In the coastal waters surrounding the Australian continent, bioavailable sources of N are  
129 regularly depleted (Condie and Dunn, 2006). Significant rates of N<sub>2</sub> fixation have been observed  
130 throughout much of the tropical northern seas surrounding Australia (Bonnet et al., 2015; Messer  
131 et al., 2016, 2017; Montoya et al., 2004; Raes et al., 2014) and in the subtropical waters of the  
132 eastern Indian Ocean (Raes et al., 2015). High levels of diversity in *nifH* phylotypes have been  
133 detected throughout these regions, including important contributions by *Trichodesmium*  
134 *erythraeum*, UCYN-A, and the gammaproteobacterial group, Gamma A (Moisander et al., 2014;  
135 Bonnet et al., 2015; Messer et al., 2017). In contrast, our understanding of the importance of N<sub>2</sub>  
136 fixation within the temperate waters along the southern coastline of Australia, which are  
137 dominated by large inverse estuaries, is severely limited.

138  
139 Inverse estuaries represent unique ecosystems within the coastal zone of arid climates (Eyre,  
140 1998), where an excess of evaporation over precipitation results in the formation of strong  
141 positive salinity gradients from marine at the mouth to hypersaline at the head (Nunes Vaz et al.,  
142 1990; Pritchard, 1952). In contrast to classical estuaries, inverse estuaries receive little to no  
143 freshwater input (Eyre, 1998; Smith and Veeh, 1989), and can become seasonally isolated from  
144 the adjacent continental shelf-waters when density fronts restrict oceanic inflow at the mouth  
145 (Petruševics et al., 2011). Consequently, inverse estuaries can experience relatively oligotrophic  
146 conditions, giving rise to nutrient limitation in some systems (Middleton et al., 2013; Smith and  
147 Veeh, 1989).

148  
149 Previously, we detected the presence of UCYN-A sub-lineages, UCYN-A1 and UCYN-A2,  
150 within the inverse estuary Spencer Gulf (Messer et al., 2015), an ecologically and economically  
151 important region of the south Australian marine environment (Deloitte Access Economics,  
152 2017). Despite the fact that Spencer Gulf is typically oligotrophic, and primary production is  
153 reportedly N limited (Middleton et al., 2013), productive aquaculture industries and commercial  
154 fisheries are housed within the region, and the shallow waters act as foraging and nursery  
155 grounds for > 30 species of threatened, protected, and iconic marine macro-organisms (Robbins  
156 et al., 2017). Seagrass-based N<sub>2</sub> fixation has historically been suspected to be an important  
157 source of new N in the shallow upper region (Smith and Veeh, 1989), however, how pelagic

158 productivity is maintained within the low-nutrient waters of Spencer Gulf remains an open  
159 question. To test the hypothesis that  $N_2$  fixation is a significant process within the temperate  
160 coastal waters of southern Australia, we investigated spatial and seasonal dynamics of  $N_2$   
161 fixation activity, and diazotroph diversity, in Spencer Gulf and the surrounding shelf waters.  
162  
163

## 164 **Materials & Methods**

### 165 **Sample collection**

166 Surface seawater samples were collected from Spencer Gulf, a large inverse estuary within the  
167 South Australian Gulf System ( $\sim 22\,000\text{ km}^2$ ), and from adjacent continental shelf waters.  
168 Spencer Gulf is characterised by steep gradients in sea surface temperatures and salinity and  
169 demonstrates marked differences in physicochemical characteristics during autumn/winter and  
170 spring/summer (Nunes and Lennon, 1986; Nunes Vaz et al., 1990; Petrusevics, 1993). Therefore,  
171 samples were collected during two contrasting seasons, in the Austral autumn between 28<sup>th</sup> April  
172 - 8<sup>th</sup> May 2014, and in the Austral summer between 2<sup>nd</sup> - 9<sup>th</sup> December 2014. Although  
173 considered oligotrophic, Spencer Gulf hosts productive aquaculture industries, which have been  
174 implicated in localised nutrient enrichment (Fernandes et al., 2007; Lauer et al., 2009). To  
175 capture local environmental variability, sampling was performed along a latitudinal gradient  
176 within the estuary, from an offshore site situated near Kangaroo Island [35.84S, 136.45E] on the  
177 continental shelf, through to the hypersaline region in the north of Spencer Gulf. Four locations  
178 inside Spencer Gulf were selected, including, Spencer Gulf mouth [35.25S, 136.69E] and three  
179 locations along the edge of the basin, southern Gulf [34.377S, 136.11E], mid-Gulf [33.92S,  
180 136.58E] and northern Gulf [33.04S, 137.59E] (Figure 1).  
181

182 Sampling at the mouth and shelf sites were conducted from on-board the *RV Ngerin* in  
183 conjunction with routine monitoring for the Southern Australian node of the Integrated Marine  
184 Observing System (IMOS). Samples were collected at the Kangaroo Island National Reference  
185 Station (NRSKAI; referred to as “shelf” hereafter), and SAM8SG mooring locations (referred to  
186 as “mouth” hereafter) (Lynch et al., 2014). A shore-based sampling protocol was adopted for the  
187 southern, mid, and northern Gulf sampling sites, whereby surface samples were collected from  
188 jetties (piers), approximately 227, 154 and 440 m from the shore, respectively. In all cases, 60 L  
189 of water was collected from  $\sim 1\text{ m}$  below the surface using a plastic bucket. Buckets and sample  
190 storage carboys were rinsed three times with sample water prior to filling and washed with 10 %  
191 HCl and MilliQ between stations. The temperature and salinity of each sample was immediately  
192 recorded using a multi-parameter portable meter (WTW Proline Multi 3320; Xylem Analytics,  
193 Germany).  
194

### 195 **Dissolved inorganic nutrient analyses**

196 To determine ambient concentrations of dissolved inorganic nutrients, including  $\text{NO}_3^- + \text{NO}_2^-$ ,  
197  $\text{PO}_4^{3-}$  and  $\text{SiO}_4^{4-}$  (hereafter referred to as  $\text{NO}_3/\text{NO}_2$ ,  $\text{PO}_4$ , and  $\text{SiO}_4$  respectively), subsamples (45

198 ml) were collected in triplicate 50 ml Falcon tubes from each sampling site. Samples were  
199 immediately frozen at -20 °C and kept frozen prior to analysis. A Flow Injection Analyser  
200 (Lachat QuikChem 8000) was used to determine concentrations of NO<sub>3</sub>/NO<sub>2</sub>, PO<sub>4</sub>, SiO<sub>4</sub> in the <  
201 0.45 µm filtrate from each thawed sample, with a limit of detection of 0.01 µM.

202

### 203 **Particulate carbon, nitrogen, and δ<sup>15</sup>N analyses**

204 To provide an estimate of the concentrations of particulate carbon (C) and N in the planktonic  
205 material and the natural abundance of the <sup>15</sup>N isotope in particulate matter (δ<sup>15</sup>N, N<sub>2</sub> fixation  
206 incubation T<sub>0</sub>), subsamples of between 2 – 4 L of surface seawater were collected from each site.  
207 Subsamples were filtered onto GF/F grade 0.7 µm filters (Whatman, Kent, UK) which were  
208 previously individually packaged in aluminium foil and pre-combusted at 450 °C for 4 h.  
209 Samples were stored double contained in two snap-lock bags and kept frozen at -20 °C, prior to  
210 being dried for 48 h at 60 °C. As previously described in Messer et al. (2017), filters were  
211 analysed on an elemental analyser (Thermo Finnigan MAT ConFlo IV) coupled to an isotope  
212 ratio mass spectrometer (IRMS; Thermo Finnigan Delta XP; limit of detection = 15 µg N per  
213 filter) at the Research Corporation of the University of Hawaii.

214

### 215 **Biological N<sub>2</sub> fixation incubations**

216 To measure rates of N<sub>2</sub> fixation activity among planktonic diazotrophs, we performed stable  
217 isotope tracer addition experiments at each site with <sup>15</sup>N-labelled N<sub>2</sub> gas. Acid-clean (10 % HCl)  
218 4 L Nalgene incubation bottles were rinsed three times with seawater from the site prior to being  
219 filled to over-flowing via silicone tubing, then capped with rubber septa head-space free. <sup>15</sup>N<sub>2</sub>  
220 gas (3 ml, 98 atom%, Sigma-Aldrich, Australia, lot SZ1670V, 2013 batch) was injected into each  
221 incubation bottle prior to inversion 100 times to disperse the gas bubble.

222

223 Samples for whole community N<sub>2</sub> fixation (bulk seawater) and unicellular N<sub>2</sub> fixation (< 10 µm  
224 size fraction) were incubated in triplicate at *in situ* sea surface temperature using aquaria heaters  
225 and water circulation pumps attached to an outdoor 60 L plastic incubator, which was exposed to  
226 a natural diurnal light cycle at surface seawater intensity. <sup>15</sup>N<sub>2</sub> incubations were terminated via  
227 filtration after 24 h, by directly filtering the entire contents onto a pre-combusted (450 °C for 4 h;  
228 packaged in aluminium foil) GF/F grade 0.7 µm filter (Whatman; whole community), or through  
229 a 10 µm polycarbonate membrane filter (Isopore, Merck Millipore) onto a pre-combusted GF/F  
230 grade 0.7 µm filter (Whatman; unicellular size fraction). Enriched filters were stored double  
231 contained in two snap-lock bags to prevent any possible cross-contamination and kept frozen at -  
232 20 °C prior to analysis.

233

234 Following methods described in Messer et al. (2017), <sup>15</sup>N<sub>2</sub> amended GF/F filters were dried (60  
235 °C for 48 h) separately to natural abundance samples to prevent cross-contamination. Total  
236 particulate N and C and isotopic composition were determined on an elemental analyser (Thermo  
237 Finnegan MAT ConFlo IV) coupled to an IRMS (Thermo Finnegan Delta XP, limit of detection

238 = 15 µg N per filter) at the Research Corporation of the University of Hawaii. Assimilation rates  
239 were calculated following Montoya et al. (1996). An atom% enrichment equivalent to 75 % of  
240 the theoretical for a 24 hour incubation was used as the enrichment factor for volumetric rate  
241 calculations to account for the incomplete dissolution of the  $^{15}\text{N}_2$  gas bubble (Großkopf et al.,  
242 2012; Mohr et al., 2010), following Messer et al. (2017).

243

#### 244 **Collection, preservation, and extraction of microbial nucleic acids**

245 In order to concentrate microbial cells for nucleic acid extraction, amplicon sequencing, and  
246 quantitative PCR, triplicate 2 L samples were filtered onto 0.2 µm membrane filters (Durapore,  
247 EMD Merck Millipore, Billerica, MA, USA). Filters were stored at -20 °C during the field  
248 sampling (~ 2 weeks) and transported to the laboratory on dry ice before being stored at -80 °C  
249 until extraction. The MoBio PowerWater DNA isolation kit (MoBio Laboratories, Carlsbad, CA,  
250 USA; now Qiagen) was used to extract microbial community DNA, following the  
251 manufacturer's guidelines including an additional incubation with solution PW1 (10 min at 60  
252 °C) prior to 10 min of bead beating, to ensure complete cell lysis.

253

#### 254 ***nifH* amplicon sequencing and analyses**

255 To determine the diversity of diazotrophic bacterioplankton, a fragment of the *nifH* gene was  
256 amplified using a nested protocol and the degenerate primers *nifH3* and *nifH4*, and *nifH1* and  
257 *nifH2* (Zani et al., 2000; Zehr and Turner, 2001), largely following methods previously described  
258 (Messer et al., 2015, 2016, 2017). The following PCR reaction conditions were used to amplify  
259 the *nifH* gene: 95 °C (2 min) followed by 30 cycles of 95 °C (1 min), 48 °C (1 min) and 72 °C (1  
260 min) followed by 72 °C (10 min). Amplification was confirmed using gel electrophoresis,  
261 replicates were pooled, and the resultant fragment was sequenced using the 454 FLX Titanium  
262 pyrosequencing platform (Roche, Nutley, NJ, USA) at Molecular Research LP (Shallowater, TX,  
263 USA).

264

265 The Quantitative Insights Into Microbial Ecology (QIIME) (Caporaso *et al.*, 2010) open source  
266 software was used to process *nifH* pyrosequencing reads. Briefly, sequences were de-multiplexed  
267 and the low-quality sequences were removed ( $q < 25$  and  $< 200$  bp in length) using default  
268 parameters. Chimeric sequences were removed using USEARCH61 with default parameters  
269 against an unaligned version of a curated *nifH* reference database exported from Arb  
270 (downloaded from: [http://www.zehr.pmc.ucsc.edu/nifH\\_Database\\_Public/](http://www.zehr.pmc.ucsc.edu/nifH_Database_Public/); Heller et al., 2014;  
271 Zehr et al., 2003). The remaining high-quality reads were clustered at 99 % sequence identity  
272 using UCLUST, whereby sequences within 1 % identity of the most abundant read were  
273 classified as operational taxonomic units (OTUs; Edgar, 2010). A representative sequence set  
274 was generated based on the most abundant sequence comprising an OTU. The PyNAST aligner  
275 tool (Caporaso et al., 2010a) was used with default parameters to BLAST and pairwise align  
276 representative *nifH* OTU sequences to those from the aligned version of the same curated *nifH*  
277 database used for chimera removal, providing putative taxonomy and “best hits” to primarily

278 uncultured environmental sequences (Heller et al., 2014; Zehr et al., 2003). Any potential stop  
279 codons and frameshifts in the *nifH* sequences were identified using the FrameBot tool from the  
280 FunGene pipeline using default parameters (Fish et al., 2013). As part of this pipeline, taxonomy  
281 was assigned to the closest representatives within the Ribosomal Database Project's *nifH*  
282 database based on amino acid identity (AAI) and sequence alignment (Fish et al., 2013). Finally,  
283 an OTU by sample matrix was generated, in which each sample was rarefied to the lowest  
284 number of sequences per sample (3068) and singletons were removed prior to downstream  
285 analyses.

286

### 287 **Quantification of UCYN-A *nifH* genes**

288 Based on our previous observations (Messer et al., 2015), we hypothesised that UCYN-A would  
289 be the most important diazotrophic group within Spencer Gulf and the adjacent continental shelf  
290 waters. In order to determine UCYN-A abundance, previously designed TaqMan qPCR probes  
291 (*Table S1*) were utilised to quantify the UCYN-A1 (Langlois et al., 2008) and UCYN-A2  
292 (Thompson et al., 2014) clades. qPCR standards were either cloned into the P-Gem T Easy  
293 Vector (Promega, Sydney, NSW, Australia) following the manufacturer's guidelines (UCYN-  
294 A2) as previously described (Messer et al., 2017), or synthesised into the PUC-57 Amp  
295 (Genewiz) vector (UCYN-A1). The *nifH* gene inserts were then amplified from the plasmid  
296 DNA using plasmid specific PCR primers targeting the M13 binding site of the vector. A band of  
297 the correct size was purified from an electrophoresis gel using the Isolate II Gel/PCR Purification  
298 kit (Bioline, Eveleigh, NSW, Australia). DNA was then quantified using a Qubit Fluorometer  
299 and serially diluted to generate a standard curve incorporating  $10^7$  to  $10^1$  *nifH* copies.

300

301 qPCR reactions were performed as previously described in Messer et al. (2017). Specifically,  
302 template DNA was diluted 1:5 using nucleic-acid-free H<sub>2</sub>O to prevent inhibition and 5 µl of the  
303 template dilution was subsequently used in the qPCR assay. Each qPCR reaction included 200  
304 nM of forward and reverse primer, 100 nM of TaqMan probe, 2x TaqMan Master Mix II, and 3  
305 µl of nucleic-acid-free H<sub>2</sub>O. Samples were analysed in triplicate, with additional triplicate  
306 technical replicates and triplicate no template negative controls (5 µl nucleic-acid-free H<sub>2</sub>O),  
307 alongside the relevant standards (also analysed in triplicate). Reaction conditions were optimised  
308 for each primer and probe set using a combination of temperature, annealing time, and extension  
309 time gradients on a StepOnePlus™ Real-Time PCR machine (software v2.3; Applied  
310 Biosystems, Thermo Fisher Scientific, Scoresby, Victoria, Australia). The final optimal reaction  
311 conditions were identified to be: 50 °C (5 min), 95 °C (10 min) and 40 cycles of 95 °C (15 sec)  
312 and 64 °C (60 sec) for UCYN-A2; and 95 °C (10 min) followed by 40 cycles of 95 °C (15 sec),  
313 55 °C (15 sec) and 72 °C (60 sec) for UCYN-A1. Linear regression analyses of quantification  
314 cycle (C<sub>q</sub>) versus log<sub>10</sub> *nifH* gene copies demonstrated that the UCYN-A2 assay had a mean R<sup>2</sup>  
315 of 0.999 and an efficiency between 99.2 - 99.9 % and the UCYN-A1 assay had a mean R<sup>2</sup> of  
316 0.993 and an efficiency between 92.0 - 98.7 %. The C<sub>q</sub> limit of detection and quantification for  
317 each assay was identified to be equivalent to ~1 - 10 *nifH* copies per reaction.

318

**319 Statistical analyses**

320 Prior to testing for significant differences between “season” and “site”, environmental data, N<sub>2</sub>  
321 fixation rates, and qPCR data were checked for normality and homogeneity of variance using the  
322 Shapiro-Wilk and Brown-Forsythe tests respectively (SPSS, IMB Statistics 24). Data meeting  
323 these criteria were tested for significance using a one-way ANOVA, while a Kruskal Wallis  
324 ANOVA on ranks was used for data that failed to meet the stipulations of normality (SPSS, IMB  
325 Statistics 24). Pearson correlation coefficients and significance values were calculated (SPSS,  
326 IMB Statistics 24) between biological and environmental variable pairs across the entire dataset,  
327 and independently for samples collected in Austral autumn or summer.

328

329 Statistical analyses of diazotroph community dissimilarity were performed using the PRIMER 7  
330 + PERMANOVA software. The final OTU by sample matrix was square-root transformed and a  
331 Bray Curtis resemblance matrix was generated. Significant differences between *nifH* amplicon  
332 sequencing profiles were explored using the non-parametric Analysis of Similarity (ANOSIM)  
333 test, using either “season” or “site” as a factor, while the contribution of each OTU to the  
334 observed dissimilarity between sampling sites was determined using Similarity Percentage  
335 analysis (SIMPER). In addition, a distance-based linear model (DistLM) was generated from the  
336 Bray-Curtis resemblance matrix, using the corresponding site-specific environmental metadata as  
337 predictor variables. Relationships between the environmental predictor variables and diazotroph  
338 community composition were also investigated using the BEST, biota and environment  
339 (BIOENV) test, using Spearman rank correlation.

340

341 The multivariate relationships between individual diazotroph OTUs, environmental metadata,  
342 and N<sub>2</sub> fixation rates were explored using a negative binomial many-generalised linear model  
343 (Wang et al., 2012). The model was performed using the mvabund (v.4.1.3) package (Wang et  
344 al., 2012) in R (v4.0.2) and R studio (v1.3.959) (R Core Team, 2013). The *nifH* OTU by sample  
345 matrix was input as count data and converted to an mvabund object prior to model creation using  
346 the ‘manyglm’ function. The analysis of deviance table was generated using the ‘anova’ function  
347 with ‘p.uni = adjusted’ selected to correct for the effect of multiple testing.

348

349

**350 Results****351 Environmental characteristics of Spencer Gulf and shelf waters**

352 Consistent with the inverse estuarine nature and seasonal variability of Spencer Gulf, patterns in  
353 sea surface temperature (SST) and salinity exhibited a clear transition from cooler oceanic  
354 conditions in southern shelf waters, to warmer and hypersaline conditions in the northern region  
355 of the Gulf (Figure 1; Table 1). Across this gradient, SST ranged from 18 °C to ~23 °C, while  
356 salinity increased from 36 at the mouth to ≥ 40 at the northern site (Table 1). During the Austral  
357 autumn, SST was typically lower than SST observed during the summer (Table 1), with mean

358 temperature ( $\pm$  standard deviation) across the five sites,  $18.9 \pm 0.8$  °C relative to  $21.0 \pm 1.8$  °C,  
359 respectively. In contrast, the salinity profile of Spencer Gulf was highly similar during both the  
360 Austral autumn and summer across the five sampling sites, with means for each season ( $\pm$   
361 standard deviation) of  $37.3 \pm 1.65$  and  $37.1 \pm 1.81$ , respectively.

362

363 Concentrations of dissolved inorganic nutrients were relatively stable between the southern shelf  
364 and northern Spencer Gulf waters. Indeed,  $\text{NO}_3/\text{NO}_2$  concentrations were always  $< 0.05$   $\mu\text{M}$ , and  
365  $\text{PO}_4$  concentrations were generally low, ranging from 0.01 (i.e., limit of detection) to 0.08  $\mu\text{M}$   
366 across the five sampling locations (Table 1). Mean ( $\pm$  standard deviation)  $\text{NO}_3/\text{NO}_2$  and  $\text{PO}_4$   
367 concentrations were similar between the two sampling periods, at  $0.02 \pm 0.01$  and  $0.03 \pm 0.02$   
368  $\mu\text{M}$  during Austral autumn, and  $0.03 \pm 0.01$  and  $0.04 \pm 0.03$   $\mu\text{M}$  during Austral summer,  
369 respectively. Conversely, concentrations of  $\text{SiO}_4$  showed a sharp increase from the southern shelf  
370 to northern Gulf waters, ranging from 0.22 up to 1.10  $\mu\text{M}$  (Table 1). While mean  $\text{SiO}_4$   
371 concentrations were typically elevated during Austral autumn compared to summer, at  $0.49 \pm$   
372  $0.36$  and  $0.38 \pm 0.14$   $\mu\text{M}$ , respectively.

373

#### 374 **Biological $\text{N}_2$ fixation rates in temperate southern Australia**

375 Measurable rates of  $\text{N}_2$  fixation occurred at all sites during both the Austral autumn and summer,  
376 but rates were highly heterogeneous ranging from 2  $\text{nmol L}^{-1} \text{d}^{-1}$  to 64  $\text{nmol L}^{-1} \text{d}^{-1}$  (Figure 2).  
377 Across the entire dataset, no significant differences were observed between whole community  
378 (WC) and unicellular size fraction (USF)  $\text{N}_2$  fixation rates (Kruskal-Wallis test,  $H = 0.32$ , d.f. =  
379 1,  $n = 30$ ,  $P = 0.574$ ), indicating that the unicellular size fraction contributed the majority of the  
380 observed  $\text{N}_2$  fixation activity. Overall, no significant differences in  $\text{N}_2$  fixation rates were  
381 observed between incubations conducted during Austral autumn compared to summer (Kruskal-  
382 Wallis test,  $H = 1.397$  and  $1.931$ , d.f. = 1,  $P = 0.237$  and  $0.165$ , for WC and USF respectively;  $n$   
383 = 15 per season). During both Austral autumn and summer, WC and USF  $\text{N}_2$  fixation rates were  
384 highly correlated, with Pearson correlation coefficients ( $r$ ) of 0.85 and 0.76 respectively, further  
385 supporting the proposition that the unicellular size fraction contributed the majority of the  
386 observed  $\text{N}_2$  fixation activity

387

388 When grouped by “site” as opposed to “season”,  $\text{N}_2$  fixation rates exhibited significant spatial  
389 heterogeneity (One-way ANOVA,  $P \leq 0.001$ ,  $F = 37.38$ , d.f. = 4,  $n = 6$  per site). The lowest rates  
390 of  $\text{N}_2$  fixation during both the Austral autumn and summer occurred in the southern shelf waters,  
391 with maximum rates at this site reaching only  $8 \pm 2$   $\text{nmol L}^{-1} \text{d}^{-1}$  (mean  $\pm$  standard deviation;  
392 Figure 2). In contrast,  $\text{N}_2$  fixation rates peaked in the waters at the mouth of Spencer Gulf, where  
393 they reached  $64 \pm 3$  and  $40 \pm 19$   $\text{nmol L}^{-1} \text{d}^{-1}$ , in Austral autumn and summer respectively  
394 (Figure 2). Relative to rates observed at the mouth of Spencer Gulf,  $\text{N}_2$  fixation rates decreased at  
395 the southern and mid-western sites of the gulf during both autumn and summer (Figure 2).  $\text{N}_2$   
396 fixation rates then showed a notable increase at the northern-gulf site, reaching  $29 \pm 4$   $\text{nmol L}^{-1}$   
397  $\text{d}^{-1}$  during Austral autumn, and  $14 \pm 10$   $\text{nmol L}^{-1} \text{d}^{-1}$  during Austral summer (Figure 2).

398

399 N<sub>2</sub> fixation rates were significantly correlated with low concentrations of NO<sub>3</sub>/NO<sub>2</sub> (Pearson's r:  
400 -0.53; P = 0.002, n = 30, USF). This relationship was maintained when considering only Austral  
401 summer samples (r: -0.64; P = 0.01; n = 15, USF), but not when only considering those collected  
402 during Austral autumn. In contrast, during the Austral autumn N<sub>2</sub> fixation rates were positively  
403 correlated to PO<sub>4</sub> concentrations (r: 0.73; P = 0.002; n = 15, WC). No significant relationships  
404 were observed between N<sub>2</sub> fixation and SST or salinity, despite clear spatial gradients in these  
405 environmental parameters (Table 1).

406

#### 407 **Diversity and composition of *nifH* containing bacterioplankton**

408 After rarefaction to 3068 sequences per sample and the removal of singletons, between 159 and  
409 332 *nifH* OTUs were detected at each sampling site. The diversity of *nifH* containing  
410 bacterioplankton increased along the latitudinal gradient of Spencer Gulf, whereby Shannon's  
411 Diversity (H') was lowest in the southern shelf waters, where H' = 1.95 and 2.86 and peaked at  
412 the mid-western edge of Spencer Gulf, where H' = 4.97 and 4.37, during Austral autumn and  
413 summer respectively (Table S2). Despite the site-specific differences in diazotroph diversity,  
414 mean H' across the Gulf was approximately equal for both sampling seasons, whereby H' = 3.67  
415 during austral autumn, and H' = 3.58 during austral summer.

416

417 Phylogenetic analyses of *nifH* sequences demonstrated that the most abundant OTUs (n = 25),  
418 equivalent to ~53 % of total sequences and between 15 and 82 % of sequences for any given  
419 sample, comprised a mixture of Cluster 1 and Cluster 3 diazotrophs at ≥ 83 % amino acid  
420 identity (AAI; Table S3). A Bray-Curtis resemblance matrix of rarefied *nifH* sequence data was  
421 used to compare diazotroph community composition within and between the southern shelf  
422 waters and Spencer Gulf sampling locations, revealing significant spatial variability in  
423 diazotroph assemblage structure (ANOSIM, R: 0.59, P = 0.005). SIMPER analysis revealed 99.7  
424 % and 100 % community dissimilarity between northern Gulf diazotroph assemblages and those  
425 in the shelf waters and at the mouth of the Gulf, respectively. Diazotroph assemblages in the  
426 shelf waters and mouth were dominated by five OTUs identified to be the UCYN-A1 open ocean  
427 ecotype (OTU51120, OTU3535, OTU45147, OTU7980, and OTU1115; Fig. S1), which  
428 collectively represented 75 % and 56 % of sequences at the shelf during Austral autumn and  
429 summer respectively (Figure 3). Similarly, these OTUs comprised 54 % and 58 % of sequences  
430 at the mouth during Austral autumn and summer (Figure 3). Correspondingly, diazotroph  
431 communities in the southern shelf waters and at the mouth of the gulf shared the greatest  
432 similarity in composition, with SIMPER analysis revealing only 63.4 % dissimilarity between  
433 these populations. The dissimilarity between the shelf waters and the mouth was largely driven  
434 by the coastal and open ocean ecotypes UCYN-A2 and UCYN-A4 (OTU9097 and OTU67260,  
435 respectively), which were collectively present at higher relative abundances at the mouth (Figure  
436 3).

437

438 Spencer Gulf communities showed a decline in the abundance of UCYN-A OTUs, and a greater  
439 proportion of sequences associated with non-cyanobacterial diazotrophs, along with a small  
440 proportion of OTUs closely related to filamentous cyanobacteria such as *Trichodesmium*  
441 *erythraeum* (Figure 3; Table S3). The average relative abundance of two UCYN-A1 open ocean  
442 group OTUs, (OTU51120 and OTU3535), were identified by SIMPER analysis as the main  
443 drivers of community dissimilarity between the shelf waters, Spencer Gulf mouth, and northern  
444 gulf diazotroph assemblages. At the southern gulf site, a transitional community was observed,  
445 which comprised UCYN-A1 and UCYN-A2 (12 - 44% of sequences), *Pseudomonas stutzeri* (7  
446 %), *Desulfovibrio aespoeensis* (8 - 12 %), *Coralimargarita akajimensis* (7 %), and  
447 *Desulfonatrosopira thiodismutans* (7 %). In contrast, at the northern site the community was  
448 primarily comprised of OTUs related to *Desulfovibrio aespoeensis* (10 - 28 %), *Pseudomonas*  
449 *stutzeri* (11 %), and Verrucomicrobiae (11 %; Figure 3). SIMPER analysis identified the  
450 *Desulfovibrio aespoeensis* OTU (OTU41624; 96 % AAI similarity) as also being responsible for  
451 the between-site discrimination of the diazotroph community, with this OTU absent from  
452 assemblages detected in the southern waters. Interestingly, only a small proportion of the most  
453 abundant 25 OTUs were represented at the mid-Spencer Gulf site (15 - 38 %) and northern-  
454 Spencer Gulf site (32 - 36 %) especially during the Austral autumn. Instead, overall low  
455 abundance OTUs, which were typically unique to these sites (i.e., OTUs representing < 0.5 % of  
456 total sequences), were responsible for the high alpha diversity associated with these sites.

457  
458 Across the dataset, several variables were identified as having a significant effect on the relative  
459 abundance and composition of diazotrophic bacterioplankton within Spencer Gulf and the  
460 adjacent shelf waters. These included NO<sub>3</sub>/NO<sub>2</sub> (P = 0.002), N<sub>2</sub> fixation by the unicellular size  
461 fraction (P = 0.007) and the whole community (P = 0.016), salinity (P = 0.017), temperature (P =  
462 0.037), particulate nitrogen (PN; P = 0.039), and PO<sub>4</sub> (P = 0.048; Many GLM, Table S4). Only  
463 three of these predictors displayed significant relationships (adjusted P-value < 0.1) with  
464 individual OTUs, including PN (1 OTU), N<sub>2</sub> fixation by the unicellular size fraction (14 OTUs),  
465 and salinity (4 OTUs; Table S4).

466  
467 Approximately 33 % of the spatial variation in diazotroph community dissimilarity could be  
468 explained by ambient salinity and SiO<sub>4</sub> concentrations (DistLM R<sup>2</sup>: 0.33; salinity F = 2.24, P =  
469 0.001; SiO<sub>4</sub> F = 1.56, P = 0.028; n = 10). The importance of salinity and SiO<sub>4</sub> in structuring the  
470 diazotroph community was further confirmed by BEST/BIOENV analyses, resulting in a  
471 significant (P = 0.01, n = 10) coefficient, Rho = 0.67, using Spearman's Rank correlation. The  
472 sequential addition of the environmental parameters, PN, NO<sub>3</sub>/NO<sub>2</sub>, and PO<sub>4</sub>, reduced the  
473 strength of the correlation to 0.56, 0.52, and 0.49, respectively. In contrast to the observed spatial  
474 heterogeneity in diazotroph assemblage structure, no significant differences in diazotroph  
475 community dissimilarity were observed between the Austral autumn and summer sampling times  
476 (ANOSIM, R: -0.12, P = 0.80).

477

## 478 **Abundance of UCYN-A1 and UCYN-A2 *nifH* genes**

479 qPCR derived abundances of UCYN-A1 and UCYN-A2 *nifH* genes demonstrated higher  
480 abundances of these organisms in shelf waters and at the more southern sites of Spencer Gulf  
481 (Figure 4). Specifically, the maximum mean abundance of UCYN-A1 occurred in the southern  
482 shelf waters during Austral summer, whereby  $5.4 \pm 4.7 \times 10^4$  *nifH* copies L<sup>-1</sup> were detected  
483 (Figure 4). Similarly, UCYN-A2 also reached maximum abundance in the shelf waters during  
484 Austral summer, with mean *nifH* copies  $1.9 \pm 1.4 \times 10^4$  L<sup>-1</sup> (Figure 4).

485  
486 Across all sampling locations, UCYN-A1 was significantly more abundant during the Austral  
487 summer compared to autumn (Mann Whitney test, U: 69, P < 0.05). While UCYN-A2  
488 abundances did not differ significantly between the Austral autumn and summer sampling.  
489 Across the entire dataset, UCYN-A1 abundance was positively correlated with concentrations of  
490 PO<sub>4</sub> (r: 0.39; P = 0.03, n = 30). In contrast, overall UCYN-A2 abundance was not significantly  
491 correlated with any of the measured environmental parameters. However, when analysed by  
492 “season”, UCYN-A2 abundance was negatively correlated to SST during both Austral autumn  
493 and summer (n = 15 per season; r: -0.55 and r: -0.53, P = 0.03 and 0.04, respectively). In  
494 addition, during Austral autumn UCYN-A2 abundance was negatively correlated to PO<sub>4</sub>  
495 concentrations (r: -0.53; P = 0.04; n = 15). Despite the potential importance of salinity in  
496 structuring the overall diazotroph community, no significant relationships were observed  
497 between UCYN-A qPCR derived abundances and salinity.

498  
499

## 500 **Discussion**

501 Increasing evidence suggests that temperate coastal waters may be overlooked hotspots of N<sub>2</sub>  
502 fixation activity (Mulholland et al., 2012, 2019; Tang et al., 2019b). Determining the distribution  
503 and activity of diazotrophs, and the environmental processes that influence them within coastal  
504 zones is therefore important to further our understanding of N availability across diverse marine  
505 environments. Compared to previous studies in temperate and tropical estuarine environments,  
506 where maximum N<sub>2</sub> fixation rates of 30 - 85 nmol L<sup>-1</sup> d<sup>-1</sup> have been observed (Ahmed et al.,  
507 2019; Bentzon-Tilia et al., 2015b; Bhavya et al., 2016), here we report relatively high rates of N<sub>2</sub>  
508 fixation in temperate coastal waters of southern Australia within the inverse estuary Spencer  
509 Gulf. We show that N<sub>2</sub> fixation rates, diazotroph diversity, and community structure, can vary  
510 considerably across relatively small spatial scales, however the dynamics of N<sub>2</sub> fixation were  
511 relatively stable across two contrasting seasons. Our findings suggest that N<sub>2</sub> fixation, possibly  
512 mediated by UCYN-A and non-cyanobacterial diazotrophs, may provide an important source of  
513 fixed N to support primary production within the oligotrophic, temperate coastal waters of  
514 southern Australia.

515

## 516 **N<sub>2</sub> fixation in temperate coastal environments**

517 Recent efforts to determine the importance of  $N_2$  fixation as a source of new N within temperate  
518 coastal waters have revealed  $N_2$  fixation activity in these regions is similar to, and at times higher  
519 than, rates reported for tropical and subtropical open ocean environments (Mulholland et al.,  
520 2019; Tang et al., 2019b). For example, maximum  $N_2$  fixation rates of 65, 130, and 100  $nmol L^{-1}$   
521  $d^{-1}$  have recently been observed in coastal waters of the north-eastern, mid-, and western Atlantic  
522 Ocean respectively (Fonseca-Batista et al., 2019; Mulholland et al., 2019; Tang et al., 2019b). In  
523 environments representing the traditional niche of  $N_2$  fixation, such as the North Pacific  
524 Subtropical Gyre (NPSG) and the Eastern South Pacific (ESP), maximum  $N_2$  fixation rates have  
525 been reported to be considerably lower at  $\leq 20 nmol L^{-1} d^{-1}$  (Böttjer et al., 2017; Gradoville et al.,  
526 2017; Shiozaki et al., 2017).

527

528 In the temperate coastal waters of southern Australia, we observed relatively high rates of  $N_2$   
529 fixation, with a maximum  $N_2$  fixation rate of 64  $nmol L^{-1} d^{-1}$ . This observation is similar in  
530 magnitude to the high  $N_2$  fixation rates reported for the tropical oligotrophic seas of northern  
531 Australia (Bonnet et al., 2015; Messer et al., 2016), and is almost double maximum  $N_2$  fixation  
532 rates previously reported for tropical estuarine systems (31 - 34  $nmol L^{-1} d^{-1}$ ; Ahmed et al., 2019;  
533 Bhavya et al., 2016). The lowest rates of  $N_2$  fixation (2  $nmol L^{-1} d^{-1}$ ) occurred in the continental  
534 shelf waters. This finding is comparable to observations from other continental shelf ecosystems  
535 where  $N_2$  fixation rates are typically lower than those observed in sites closer to the coast  
536 (Mulholland et al., 2012; Shiozaki et al., 2015a; Singh et al., 2019). Intermediate rates of  $N_2$   
537 fixation (10 - 45  $nmol L^{-1} d^{-1}$ ) were measured within Spencer Gulf, and these rates are placed  
538 within the upper end of those previously reported for other temperate coastal, and tropical  
539 estuarine waters (Ahmed et al., 2019; Bentzon-Tilia et al., 2015b; Bhavya et al., 2016;  
540 Mulholland et al., 2012; Rees et al., 2009; Shiozaki et al., 2015a). Importantly, our findings  
541 demonstrate that  $N_2$  fixation in the temperate waters of southern Australia are similar to, and can  
542 exceed, those observed in the NPSG and ESP (Gradoville et al., 2017).

543

544 While  $N_2$  fixation rates demonstrated clear spatial patterns in their magnitude between southern  
545 shelf and Spencer Gulf waters, we observed relatively consistent  $N_2$  fixation rates across  
546 opposing seasons. This is in contrast to previous seasonal observations of  $N_2$  fixation from  
547 distinct marine environments, where  $N_2$  fixation rates are typically higher during spring/summer  
548 than autumn/winter and are accompanied by shifts in the abundance of different diazotrophic  
549 taxa (Bentzon-Tilia et al., 2015b; Böttjer et al., 2017; Fernandez et al., 2015; Mulholland et al.,  
550 2019). We hypothesised that seasonal differences in  $N_2$  fixation rates would occur within  
551 Spencer Gulf and shelf waters due to the known seasonality in physico-chemical characteristics,  
552 such as temperature, salinity, and dissolved nutrients, which ultimately influence the distribution  
553 and activity of marine diazotrophic microorganisms (Monteiro et al., 2011; Moore et al., 2013;  
554 Ward et al., 2013). However, while limited in replication, we observed relatively stable site-  
555 specific physico-chemical conditions between the two contrasting seasons, and no significant  
556 differences in the composition of the underlying diazotrophic community. While limited in scope

557 to two time-points, our observations suggest that relatively high N<sub>2</sub> fixation rates can be  
558 maintained within Spencer Gulf while favourable conditions prevail. In future, increased  
559 sampling resolution is required to define the seasonal dynamics of N<sub>2</sub> fixation within the  
560 temperate coastal waters of southern Australia.

561

### 562 **Regional significance of biological N<sub>2</sub> fixation**

563 Our previous research indicated that the pelagic microbial community of Spencer Gulf includes a  
564 diverse array of diazotrophic clades (Messer et al., 2015). However, the presence of diazotrophic  
565 groups cannot solely be used as evidence for the importance of pelagic N<sub>2</sub> fixation, as the  
566 physiological process is tightly regulated (Paerl et al., 1987). To the best of our knowledge, our  
567 observations of N<sub>2</sub> fixation within the pelagic realm of Spencer Gulf represent the first N<sub>2</sub>  
568 fixation measurements from a temperate inverse estuary. Our N<sub>2</sub> fixation rate measurements  
569 support our hypothesis that pelagic N<sub>2</sub> fixation may provide a supply of fixed N within Spencer  
570 Gulf and the southern shelf waters, at considerably high rates relative to tropical and subtropical  
571 open ocean environments. In an earlier study, Middleton et al. (2013) estimated the influx of  
572 bioavailable N (in the form of NO<sub>3</sub> and NH<sub>4</sub>) within Spencer Gulf to be 16.9 kilotonnes yr<sup>-1</sup>,  
573 including anthropogenic N sources and mixing of upwelled nutrients from continental shelf  
574 waters. This estimate did not include biological N<sub>2</sub> fixation as a source of N, using their estimate  
575 of the volume of Spencer Gulf (4.58 x 10<sup>14</sup> L), the N<sub>2</sub> fixation rates measured herein could  
576 theoretically contribute an additional 23-149 kilotonnes N yr<sup>-1</sup>, albeit assuming consistent daily  
577 N<sub>2</sub> fixation rates for a given site. Indeed, an accurate N budget would require extensive  
578 additional N<sub>2</sub> fixation rates, with the appropriate modifications to the bubble method used to  
579 measure N<sub>2</sub> fixation (White et al., 2020). Nevertheless, based on our estimates, we propose that  
580 the process of biological N<sub>2</sub> fixation could be one mechanism by which productivity is  
581 maintained throughout the region.

582

583 It must be noted that the N<sub>2</sub> fixation rates presented herein have been corrected to allow for the  
584 incomplete dissolution of the <sup>15</sup>N<sub>2</sub> gas bubble at 75 % of the theoretical for a 24 hour incubation  
585 (Großkopf et al., 2012; Mohr et al., 2010). However, recent methodological comparisons suggest  
586 no “global factor” exists for rate corrections to the bubble method (Wannicke et al., 2018; White  
587 et al., 2020). Despite the known caveats of the bubble method, this approach was used in the  
588 present study due to the predicted highly dissimilar environmental conditions at each site, which  
589 would be very difficult to replicate with pre-prepared <sup>15</sup>N<sub>2</sub> saturated artificial seawater (Wilson et  
590 al., 2012). In particular, the observed gradient in ambient temperature and salinity, which  
591 determines gas solubility and is accounted for in the rate calculations based on our observations  
592 at each site, would be difficult to anticipate ahead of sample collection. We also note that  
593 contamination of Sigma-Aldrich commercial <sup>15</sup>N<sub>2</sub> gas stocks was reported after our initial study  
594 (Dabundo et al., 2014). Although we cannot explicitly rule out contamination in the batch of <sup>15</sup>N<sub>2</sub>  
595 that we used, assuming the mean values for Sigma Aldrich lot SZ1670V reported in Table 1 of  
596 Dabundo et al. (2014) are consistent across batches, we estimate that potential contamination

597 from  $^{15}\text{NO}_3$ ,  $^{15}\text{NH}_4$ , and  $^{15}\text{N}_2\text{O}$ , would represent an extremely small proportion of additional  $^{15}\text{N}$   
598 in our incubations, equivalent to a total of  $3.2 \times 10^{-7}$  moles. In our experiments, the relative  
599 concentration of  $^{15}\text{N}$  gas added was  $2.7 \times 10^{-4}$  moles. Including this estimate of additional  $^{15}\text{N}$  in  
600 our trace additions, any potential contamination would inflate our  $\text{N}_2$  fixation rates by between  
601  $0.001 - 0.079 \text{ nmol L}^{-1} \text{ d}^{-1}$ , which is within the lower end of the inferred  $\text{N}_2$  fixation rates  
602 resulting from  $^{15}\text{NH}_4$  contamination for 4.5 L incubations, presented in Table 2 of Dabundo et al.  
603 (2014). Moreover, this estimate is within our calculated standard error of mean  $\text{N}_2$  fixation rates  
604 across triplicate samples (equivalent to  $0.05 - 10.77 \text{ nmol L}^{-1} \text{ d}^{-1}$ ). Therefore, any potential  
605 contamination would have a negligible effect on the ultimate  $\text{N}_2$  fixation rates reported herein. In  
606 future work, the modified bubble method should be employed, including additional  
607 determination of  $^{15}\text{N}_2$  atom% enrichment of individual incubation bottles and  $^{15}\text{N}_2$  gas purity, as  
608 recently suggested by the scientific community (Jayakumar et al., 2017; Klawonn et al., 2015;  
609 White et al., 2020).

610

### 611 **Identifying the key players in coastal $\text{N}_2$ fixation**

612 Understanding the abundance and composition of the diazotrophic community underlying  $\text{N}_2$   
613 fixation activity is important for deciphering the potential impact of newly fixed N to a given  
614 region (Mulholland, 2007; Zehr and Kudela, 2011). For instance, throughout tropical and  
615 subtropical open ocean environments,  $\text{N}_2$  fixation by autotrophic diazotrophs such as  
616 *Trichodesmium* sp. will contribute directly to local primary production and may also release  
617 recently fixed  $\text{N}_2$  into the water column to support the growth of non-diazotrophic organisms  
618 (Berthelot et al., 2015; Caffin et al., 2018; Garcia et al., 2007; Glibert and Bronk, 1994;  
619 Mulholland et al., 2004). On the other hand, symbiotic diazotrophs such as the heterocystous  
620 cyanobacterium *Richelia*, which is typically associated with “tropical” phytoplankton species,  
621 transfer fixed  $\text{N}_2$  directly to their eukaryotic phytoplankton host (Foster et al., 2011), and  
622 therefore contribute to new production and carbon sequestration in regions where they are  
623 abundant, such as the NPSG (Karl et al., 2012). While the contribution of newly fixed N by non-  
624 cyanobacterial diazotrophs is not yet clear (Turk-Kubo et al., 2014), their combined high  
625 abundances and widespread transcriptional activity in areas of high  $\text{N}_2$  fixation rates (Bird and  
626 Wyman, 2013; Chen et al., 2018; Langlois et al., 2015; Moisander et al., 2014), indicate that they  
627 could make an important contribution to support primary production in both open ocean and  
628 coastal environments.

629

630 The diversity of diazotrophic organisms detected in the present study indicates that  $\text{N}_2$  fixation  
631 activity may directly and indirectly support primary production within Spencer Gulf. Within the  
632 diverse diazotrophic communities detected, Cluster 1B UCYN-A, and Cluster 1G and Cluster 3  
633 Proteobacteria, dominated diazotroph community profiles. Specifically, we observed high  
634 relative abundances of sequences closely related ( $\geq 96\%$  AAI) to the symbiotic UCYN-A, in  
635 addition to the presumed free-living *Pseudomonas stutzeri* and *Desulfovibrio aespoeensis*, as  
636 well as lower relative abundances of the large filamentous tropical cyanobacterium

637 *Trichodesmium erythraeum*. To date, UCYN-A and gammaproteobacterial diazotrophs (related  
638 to *Pseudomonas stutzeri*), have consistently been observed within temperate coastal diazotroph  
639 communities (Bentzon-Tilia et al., 2015b; Mulholland et al., 2019; Needoba et al., 2007;  
640 Shiozaki et al., 2015b), but they are also key components of subtropical and tropical assemblages  
641 (Bonnet et al., 2015; Langlois et al., 2015; Moisander et al., 2014). Our observations provide  
642 further support for the global significance of these groups, although it must be noted that the  
643 *Pseudomonas stutzeri* OTUs did not cluster with known sequences from the globally distributed  
644 Gamma A clade.

645

646 While the presence of *Trichodesmium erythraeum* was somewhat unexpected due to its tropical  
647 and subtropical distribution (Capone et al., 2005), sequences related to *Trichodesmium* sp. have  
648 previously been observed at temperate latitudes of the Atlantic and Pacific Oceans (Mulholland  
649 et al., 2019; Rivero-Calle et al., 2016; Shiozaki et al., 2015a), and their presence has also been  
650 reported in south Australian waters based on microscopic observations (Paxinos, 2007). We did  
651 not determine the specific activity of *Trichodesmium* sp. within our samples, however, it  
652 comprised up to 17 % of the diazotroph community at the mid-Gulf site during Austral summer.  
653 Owing to its presence and potential importance for both local primary production and N  
654 availability, further investigation into the significance of *Trichodesmium* sp. within temperate  
655 coastal waters is required.

656

657 Consistent with our previous observations of UCYN-A diversity and distribution within Spencer  
658 Gulf (Messer et al., 2015), we observed differences in the abundances of the open-ocean UCYN-  
659 A1 and the coastal UCYN-A2 within and between the southern shelf waters. The emerging sub-  
660 lineage UCYN-A4 (Farnelid et al., 2016) was also detected in our amplicon sequencing profiles  
661 during Austral summer. Due to the similarity between this OTU and the UCYN-A2 qPCR assay  
662 of Thompson et al. (2014), we cannot rule out that our qPCR derived abundances do not contain  
663 a mixture of the UCYN-A2 and UCYN-A4 sub-lineages (Farnelid et al., 2016). Since UCYN-  
664 A1, and to a lesser extent UCYN-A2 (possibly A2/A3/A4 sub-lineages), have recently been  
665 shown to be highly abundant ( $\leq 10^6$  *nifH* copies L<sup>-1</sup>) and reasonably active, fixing N<sub>2</sub> at rates of 6  
666 nmol L<sup>-1</sup> d<sup>-1</sup> in the cold surface waters of the Western Arctic Ocean (Harding et al., 2018),  
667 UCYN-A are highly likely to be important mediators of N<sub>2</sub> fixation within Spencer Gulf and  
668 more broadly across temperate and coastal marine environments.

669

### 670 **What environmental factors influence N<sub>2</sub> fixation in temperate southern Australian** 671 **waters?**

672 Across the global ocean, SST and subsurface minimum dissolved oxygen concentrations have  
673 been identified as the major environmental variables influencing pelagic N<sub>2</sub> fixation rates (Luo et  
674 al., 2014; Tang et al., 2019a). In addition, the availability of dissolved iron, phosphorus, other N  
675 sources (Landolfi et al., 2015; Ward et al., 2013), and grazing by zooplankton (Wang et al.,  
676 2019), have all been identified as factors shaping the distribution and magnitude of marine N<sub>2</sub>

677 fixation. Although limited in scope and replication, in the temperate southern Australian waters  
678 examined here, high  $N_2$  fixation rates during the Austral autumn were significantly correlated  
679 with increased  $PO_4$  concentrations, as was the overall abundance of UCYN-A1 (derived by  
680 qPCR). This is consistent with patterns observed in other temperate coastal waters, where  $N_2$   
681 fixation has previously been shown to be significantly correlated with phosphorus availability  
682 (Tang et al., 2019b). This pattern is also in-line with patterns observed within more oceanic  
683 waters, where  $PO_4$  availability has been shown to influence *nifH* expression and  $N_2$  fixation rates  
684 in experimentally manipulated and natural diazotroph assemblages (Rees et al., 2006; Sañudo-  
685 Wilhelmy et al., 2001; Turk-Kubo et al., 2012; Watkins-Brandt et al., 2011). As phosphorus is an  
686 important constituent of cellular and molecular machinery, there is likely a direct causal  
687 relationship between  $PO_4$  and  $N_2$  fixation, whereby diazotroph abundances and  $N_2$  fixation rates  
688 are increased under P-replete conditions, as has previously been observed for the UCYN-A1-  
689 haptophyte symbiosis (Krupke et al., 2015). Within Spencer Gulf, phytoplankton growth is  
690 estimated to be limited by  $PO_4$  availability year-round (Middleton et al., 2013), indicating that  
691 while higher  $N_2$  fixation rates may provide a source of bioavailable N to the dissolved pool, the  
692 increased diazotrophic activity may deplete  $PO_4$  concentrations for non-diazotrophic  
693 microorganisms.

694  
695 In the present study, overall  $N_2$  fixation rates were also negatively correlated with concentrations  
696 of  $NO_3/NO_2$ , which are typically depleted in Gulf waters during Austral summer yet may remain  
697 relatively high on the continental shelf due to a permanent deep nutrient pool (Doubell et al.,  
698 2018). Spencer Gulf and the adjacent continental shelf waters are characterised by a unique  
699 combination of oceanographic and regional circulation processes that create seasonal and  
700 localised east-west gradients in ambient concentrations of key macro- and micro-nutrients,  
701 underpinning variability in microbial productivity (Doubell et al., 2018; Middleton et al., 2013;  
702 van Ruth et al., 2018). During Austral autumn, the density front at the entrance to Spencer Gulf  
703 begins to break down and an influx of continental shelf water, relatively rich in macronutrients,  
704 enters the Gulf along the western edge, while the oligotrophic Gulf water exits from the eastern  
705 side of the mouth (Middleton and Bye, 2007). These north-south and east-west gradients in  
706  $NO_3/NO_2$  and  $PO_4$  concentrations (low to relatively high, respectively) (Middleton et al., 2013),  
707 may explain the observed correlations between  $N_2$  fixation rates and these nutrients. This  
708 suggests that increased  $N_2$  fixation activity may occur due to the low concentrations of  
709 bioavailable N, further indicating that N derived from  $N_2$  fixation could sustain productivity  
710 within the N limited Spencer Gulf region. Recently,  $N_2$  fixation by UCYN-A was shown to occur  
711 even when dissolved inorganic nitrogen sources are replete, and may even be stimulated by  
712 increased  $NO_3$  concentrations (Mills et al., 2020), highlighting the complexity of factors  
713 governing  $N_2$  fixation activity in the environment. Collectively, our observations in fact represent  
714 the classic nutrient regime within which diazotrophs gain a competitive advantage over non-  
715 diazotrophic microorganisms (Ward et al., 2013), utilising excess  $PO_4$  and fixing  $N_2$  to support  
716 growth.

717

718 While  $\text{PO}_4$  and  $\text{NO}_3/\text{NO}_2$  were correlated with rates of  $\text{N}_2$  fixation at the sites examined in this  
719 study, they were not significant predictors of diazotroph assemblage structure. Rather, the  
720 structure of the underlying diazotroph community was significantly influenced by the prevailing  
721 salinity and  $\text{SiO}_4$  concentrations. Regional variability in  $\text{SiO}_4$  concentrations may reflect abiotic  
722 indicators of different water masses, and may drive distinct differences in the composition of  
723 microbial assemblages (Foster et al., 2007; Weber et al., 2017). The observed transition towards  
724 increased non-cyanobacterial diazotrophs in the upper shallow waters of the Gulf could be  
725 indicative of their redistribution from the sediment or seagrass microbiome (Brown et al., 2003;  
726 Lehnen et al., 2016), and warrants further exploration of their specific activity, source and  
727 contribution to N cycling in Spencer Gulf.

728

729 Salinity is a major structuring factor for estuarine microbial communities, driving the transition  
730 from freshwater- to marine-adapted lineages (Bouvier and del Giorgio, 2002; Jeffries et al.,  
731 2016; Kirchman et al., 2005), and influencing rates of biogeochemical nutrient cycling (Bernhard  
732 et al., 2007; Bhavya et al., 2016). Unlike classical estuaries, inverse estuaries such as Spencer  
733 Gulf experience hypersaline conditions at the head of the estuary and marine salinities at the  
734 mouth, which has previously been shown to influence the overall composition of specific  
735 cyanobacterial ecotypes (Messer et al., 2015). In the present study, hypersaline regions of  
736 Spencer Gulf were associated with an increase in the relative abundance of non-cyanobacterial  
737 diazotrophs and a decrease in the abundance of UCYN-A at sites with salinities  $> \sim 37$  PSU,  
738 which may reflect an inhibitory effect of high salinity on UCYN-A and its eukaryotic host. In  
739 contrast, members of the deltaproteobacteria, related to the Cluster 3 diazotrophs observed in the  
740 present study, have previously been shown to be moderately halophilic (Gam et al., 2009;  
741 Warthmann et al., 2005), and their increased relative abundances at the northern most stations of  
742 Spencer Gulf suggests they are likely to be halotolerant.

743

744

## 745 **Conclusions**

746 This study provides further evidence that marine  $\text{N}_2$  fixation is not limited to tropical and  
747 subtropical open ocean environments, yet is widespread throughout diverse, temperate  
748 ecosystems, which have previously been overlooked as hotspots of  $\text{N}_2$  fixation activity. Our  
749 results indicate that  $\text{N}_2$  fixation is influenced by an interplay of physical and chemical  
750 environmental variables, which may have direct and indirect effects on the distribution and  
751 activity of diazotrophs in coastal waters. Our data revealed notable stability in  $\text{N}_2$  fixation across  
752 contrasting seasons, suggesting that the oligotrophic conditions of southern Australian coastal  
753 waters promote diazotrophy within the region. Notably, our findings suggest that pelagic  $\text{N}_2$   
754 fixation, mediated by UCYN and non-cyanobacterial diazotrophs, could provide a greater source  
755 of fixed N than upwelled and anthropogenic bioavailable N within the coastal waters of southern  
756 Australia.

757

758

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765

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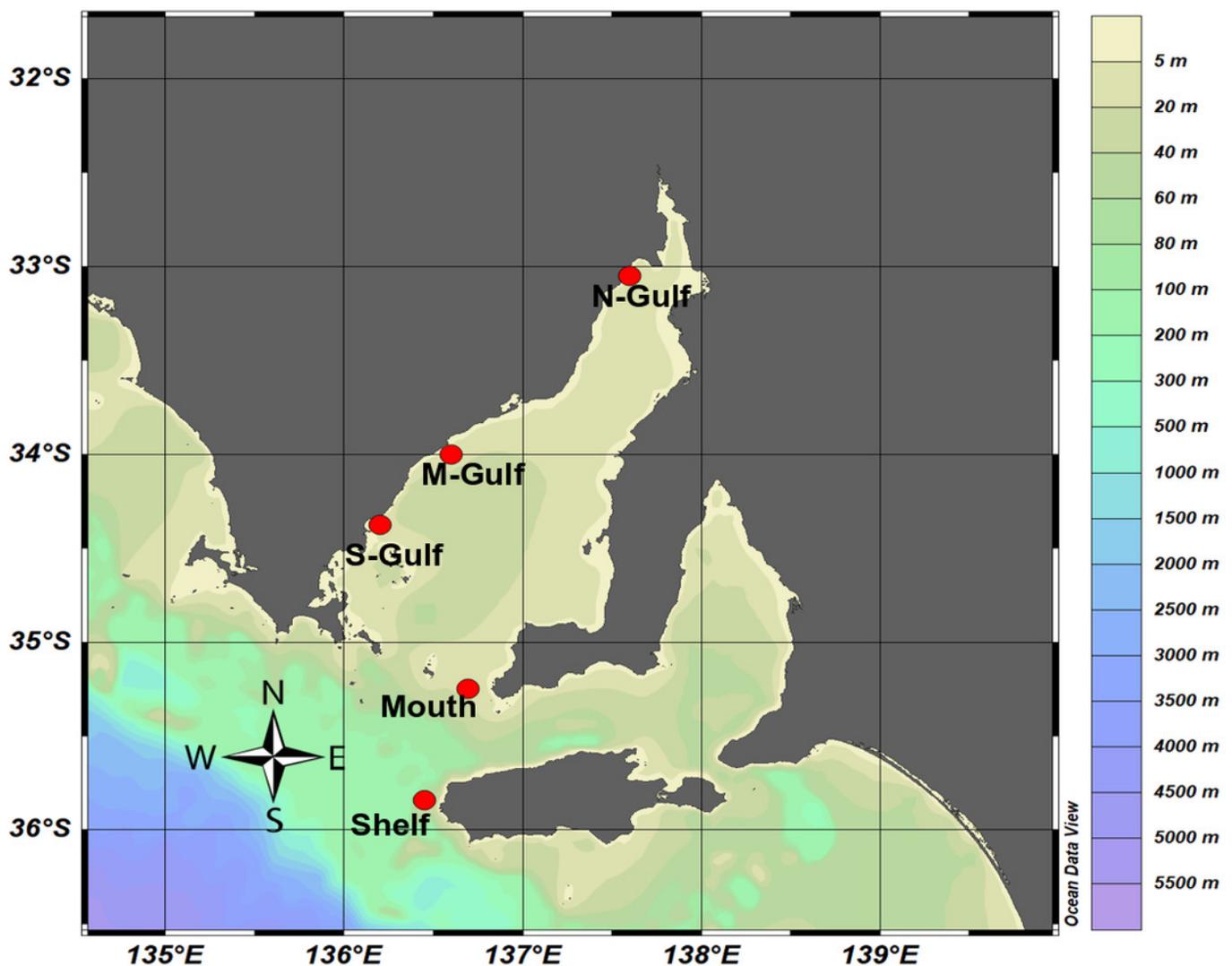
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1199

## Figure 1

Sampling locations within Spencer Gulf and the adjacent continental shelf waters.

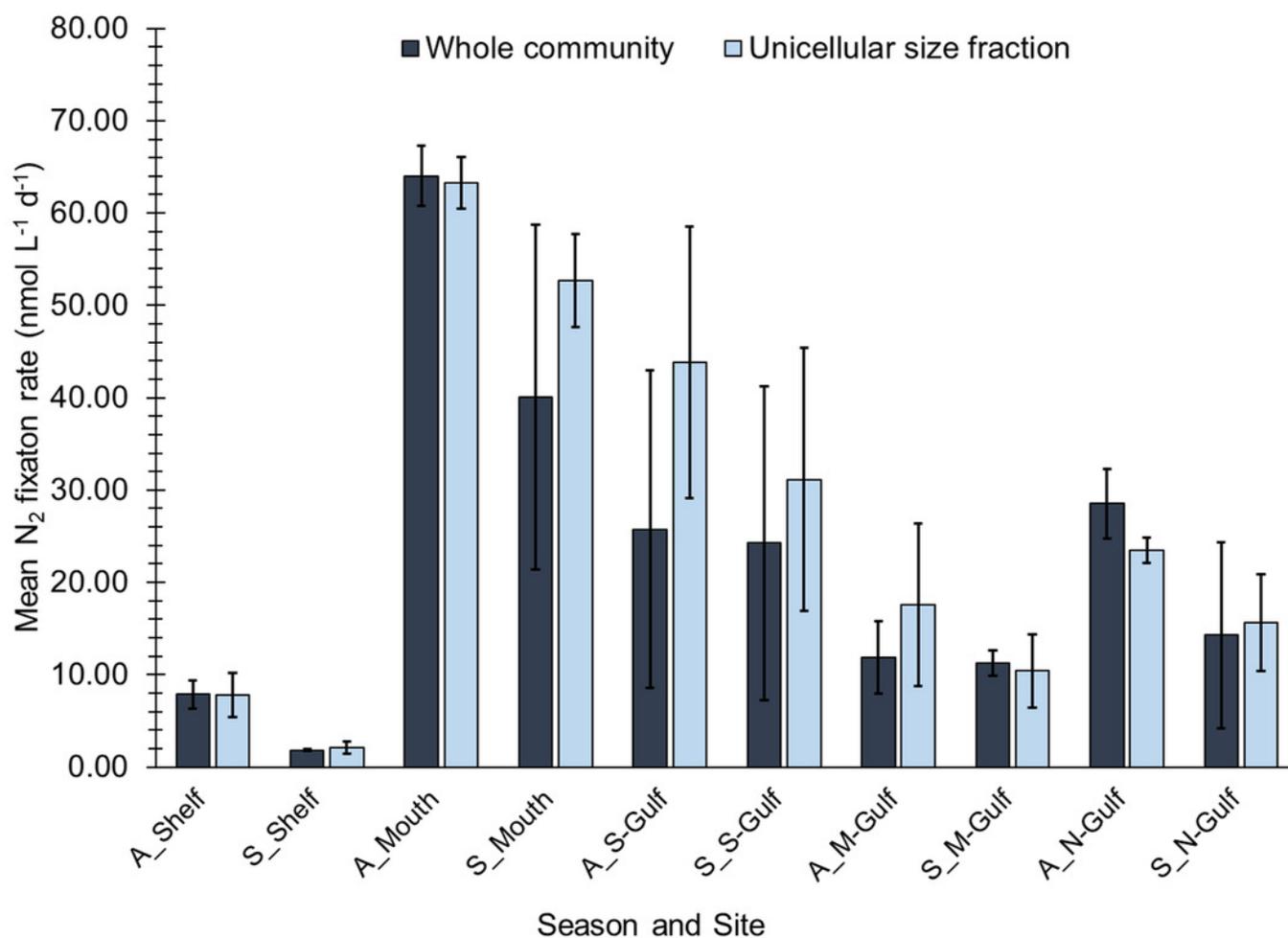
Samples were collected from the Kangaroo Island National Reference Station (Shelf), Spencer Gulf mouth (Mouth), south western edge (S-Gulf), mid western edge (M-Gulf), and northern Spencer Gulf (N-Gulf), with ocean bathymetry shown as a colour chart generated using Ocean Data View (Schlitzer, 2018).



## Figure 2

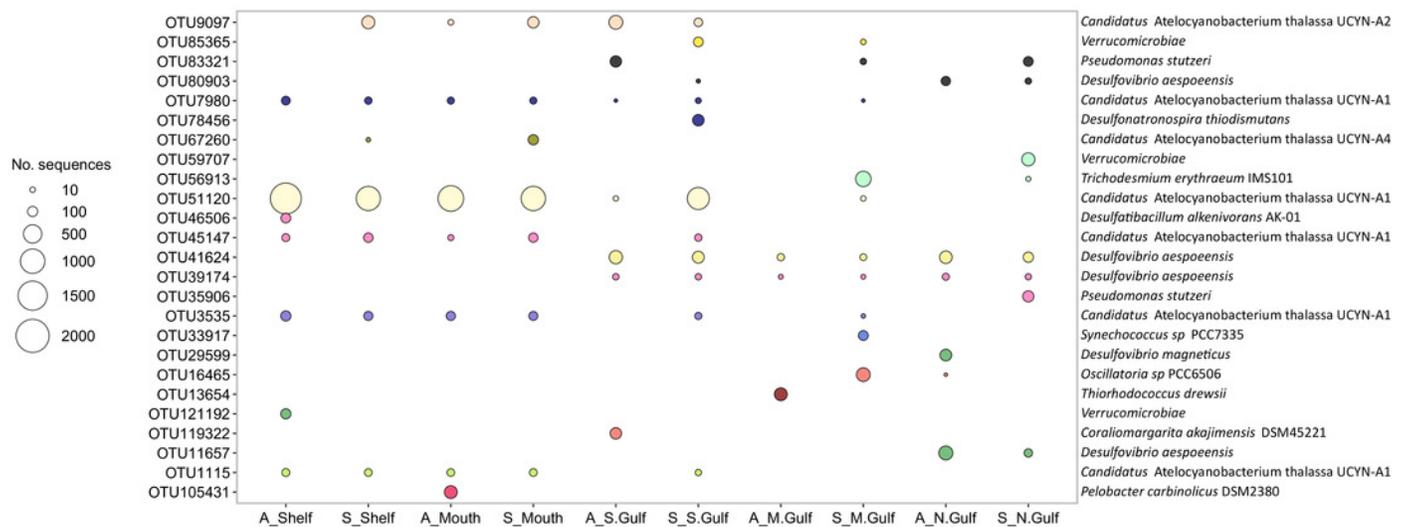
Biological N<sub>2</sub> fixation rates measured during Austral autumn and summer in south Australian coastal waters.

Rates have been corrected to account for the incomplete dissolution of the <sup>15</sup>N<sub>2</sub> gas bubble (see Methods). Error bars represent the standard deviation about the mean (n = 3).



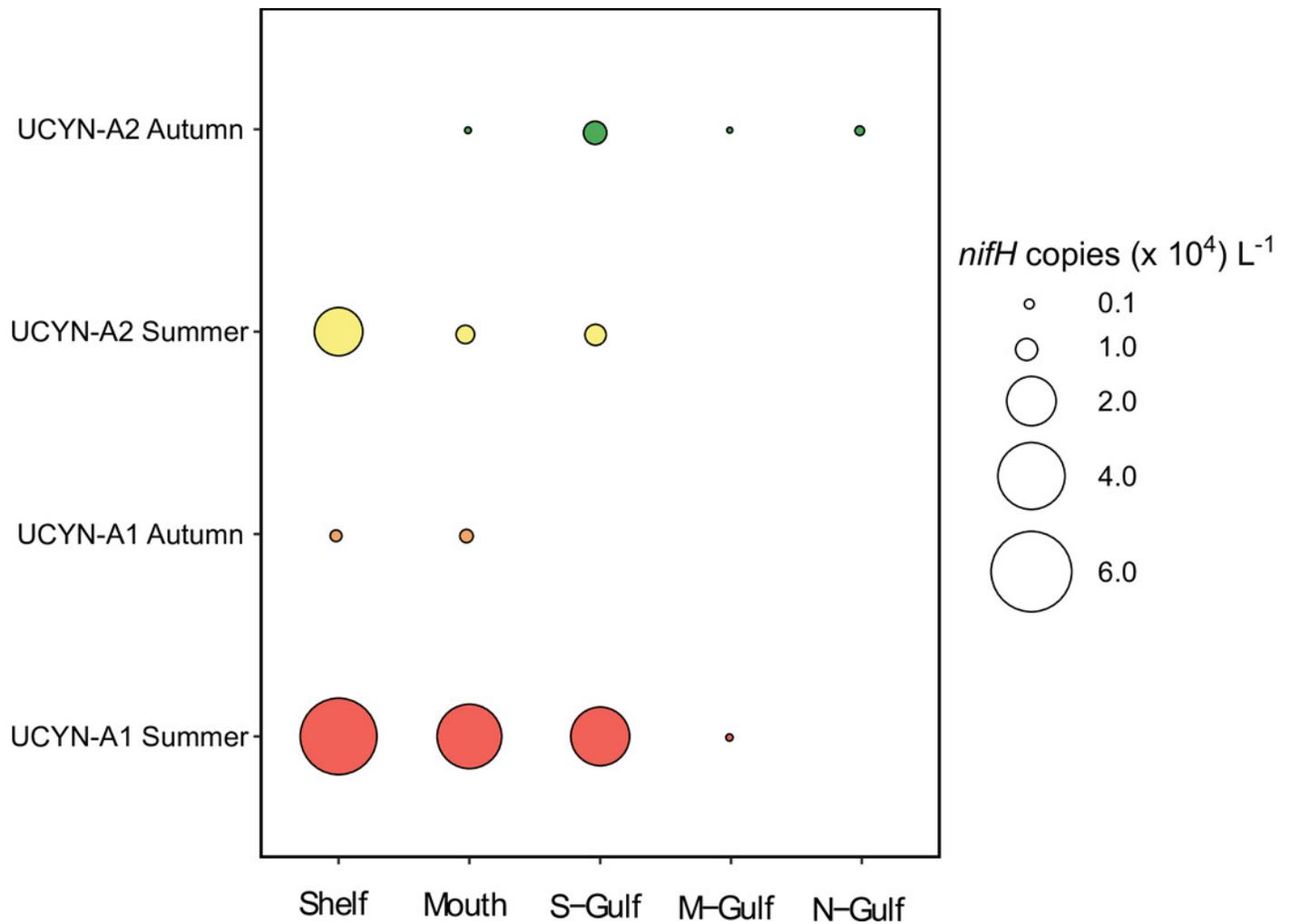
## Figure 3

Relative abundance of the top 25 *nifH* OTUs and their taxonomic assignment (closest representative) detected within south Australian coastal waters during Austral autumn (A\_) and summer (S\_).



## Figure 4

Mean qPCR derived abundances of UCYN-A1 and UCYN-A2 in south Australian coastal waters (n = 3).



**Table 1** (on next page)

Physico-chemical metadata associated with each sampling site.

Abbreviations: A = Autumn, S = Summer; Temp. = sea surface temperature; PC = particulate carbon; PN = particulate nitrogen. Sampling Time refers to the local time at the point of sample collection (Australian Eastern Standard Time).

Sample	Sampling Time	Temp. (°C)	Salinity	NO <sub>3</sub> /NO <sub>2</sub> (μM)	PO <sub>4</sub> (μM)	SiO <sub>4</sub> (μM)	PC (μg)	PN (μg)
A_Shelf	14:30	18.9	36.0	0.04	0.03	0.22	373	50.3
A_Mouth	8:00	18.7	36.0	0.02	0.06	0.36	377.8	53.2
A_S-Gulf	15:00	18	37.0	0.01	0.01	0.25	419.6	54.6
A_M-Gulf	7:30	18.8	37.7	0.01	0.02	0.52	341.7	37.8
A_N-Gulf	16:00	20.1	40.0	0.04	0.03	1.1	389.9	47.9
S_Shelf	6:30	18.7	36.0	0.02	0.08	0.24	1184.2	36.7
S_Mouth	9:00	19.6	36.0	0.01	0.04	0.24	762.7	44.6
S_S-Gulf	15:50	22.3	36.5	0.03	0.02	0.52	376.8	45.5
S_M-Gulf	7:55	21.1	36.9	0.03	0.05	0.53	755.5	62.8
S_N-Gulf	16:00	23.1	40.3	0.04	0.02	0.39	342.3	42.8

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