

A large pterosaur limb bone from the Kaiparowits Formation (late Campanian) of Grand Staircase-Escalante National Monument, Utah, USA (#54186)

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A large pterosaur limb bone from the Kaiparowits Formation (late Campanian) of Grand Staircase-Escalante National Monument, Utah, USA

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Pterosaurs were widespread globally during the Late Cretaceous, but their fossils are comparatively rare in terrestrial depositional environments. A large pterosaur bone from the Kaiparowits Formation (late Campanian, ~76–74 Ma) of southern Utah, USA, is tentatively identified as an ulna, although its phylogenetic placement cannot be precisely constrained beyond Pterosauria. The element measures over 36 cm in preserved maximum length, indicating a comparatively large individual with an estimated wingspan between 4.3–5.9 m, the largest pterosaur yet reported from the Kaiparowits Formation. The size estimate places the individual at approximately the same wingspan as the holotype for *Cryodrakon boreas* from the penecontemporaneous Dinosaur Park Formation of Alberta. Thus, relatively large pterosaurs occurred in terrestrial ecosystems in both the northern and southern parts of Laramidia (western North America) during the late Campanian.

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Abstract

Pterosaurs were widespread globally during the Late Cretaceous, but their fossils are comparatively rare in terrestrial depositional environments. A large pterosaur bone from the Kaiparowits Formation (late Campanian, ~76–74 Ma) of southern Utah, USA, is tentatively identified as an ulna, although its phylogenetic placement cannot be precisely constrained beyond Pterosauria. The element measures over 36 cm in preserved maximum length, indicating a comparatively large individual with an estimated wingspan between 4.3–5.9 m, the largest pterosaur yet reported from the Kaiparowits Formation. The size estimate places the individual at approximately the same wingspan as the holotype for *Cryodrakon boreas* from the penecontemporaneous Dinosaur Park Formation of Alberta. Thus, relatively large pterosaurs occurred in terrestrial ecosystems in both the northern and southern parts of Laramidia (western North America) during the late Campanian.

Introduction

Pterosaurs were a widespread component of terrestrial ecosystems during the Late Cretaceous, reconstructed as filling a variety of ecological niches (Barrett et al., 2008; Witton & Naish, 2008). However, the comparative rarity of skeletal material in most formations, due in part to strong taphonomic influences and other geological biases, have limited studies of this clade and clouded interpretations of pterosaur paleobiology and paleoecology (Butler et al., 2012; Butler, Benson & Barrett, 2013; Dean, Mannion & Butler, 2016). Thus, even isolated and incomplete bones can provide important information for establishing the distribution and general morphological attributes of pterosaurs (e.g., Kellner et al., 2019).

The Kaiparowits Formation preserves rocks deposited along the eastern margin of Laramidia during the late Campanian (~76–74 Ma; Roberts, Deino & Chan, 2005; Roberts et al., 2013), with significant exposures within Grand Staircase-Escalante National Monument (GSENM) in southern Utah. A rich fossil record includes exquisitely preserved specimens for numerous tetrapods, including birds, non-avian dinosaurs, crocodylomorphs, turtles, mammals, amphibians, and lepidosaurs (see Titus & Loewen, 2013, and references therein). Pterosaurs are known from only a handful of specimens. An isolated manual phalanx was the first published record of a pterosaur from the Kaiparowits Formation (Farke & Wilridge, 2013), suggesting a fairly small (<3 m wingspan) individual. A potential pteranodontoid metacarpal was later reported (McCormack & Sertich, 2016), as well as the incomplete but associated skeleton of an azhdarchid (Carroll, Farke, Chai, et al., 2017), both of which await formal description.

Here I report on RAM 22574 (Figure 1), an isolated ?ulna from the largest pterosaur (4.3–5.9 m estimated wingspan) yet known from the Kaiparowits Formation. Although it is only a single bone, this element demonstrates the **ecological** and morphological breadth occupied by

this clade in southern Utah during the Late Cretaceous, and the general distribution of large pterosaurs across terrestrial environments during the late Campanian of western North America.

Abbreviations. FHSM, Sternberg Museum of Natural History, Hays, Kansas, USA; NSM, National Science Museum, Tokyo, Japan; RAM, Raymond M. Alf Museum of Paleontology at The Webb Schools, Claremont, California, USA; SMK, Staatliches Museum für Naturkunde, Karlsruhe, Germany; TMM, Texas Memorial Museum, Austin, Texas, USA; TMP, Royal Tyrrell Museum of Palaeontology, Drumheller, Alberta, Canada.

Geological Setting

RAM 22574 was collected at locality RAM V2005022 (colloquially known as the “Cripe Site”), within the middle unit of the Kaiparowits Formation. This site is a multi-taxon bonebed including multiple associated elements from a tyrannosaurid, at least two hadrosaurids, testudines, and a small (~3 m wingspan) azhdarchid pterosaur (Farke et al., 2016; Carroll, Farke, Chai, et al., 2017).

The bonebed at RAM V2000522 measures over 1 meter in thickness and is interpreted as including at least three main depositional events, with evidence of reworking and slight movement of bones towards the top of the sequence. RAM 22574 was collected at the stratigraphic top of the quarry, ~1.2 meters above the lowest fossil, in a sandy mudstone with extensive clay rip-up clasts and plant debris. The articular surfaces of the bone show some pre-depositional damage, potentially due to fluvial abrasion or decomposition. The fossil was found within 1 m of an azhdarchid associated skeleton (RAM 15445), partially in the same bedding plane. However, RAM 22574 is from a much larger individual, in that its ulna measures 36 cm

long, versus a radius length (which should be roughly equivalent to ulna length) of around 20 cm for RAM 15445.

Materials & Methods

Collection and preparation. RAM 22574 was collected during the summer of 2016, using standard paleontological excavation techniques. Observations in the field showed that it was damaged prior to burial (described in more detail below). Paraloid B-72 was used to stabilize the fossil in the field, and as a consolidant and glue in the preparation lab. The fossil was mechanically prepared using pneumatic hand tools of various sizes, with final preparation completed using dental picks. Fieldwork was completed under US Bureau of Land Management paleontology permits UT06-012E-GS and UT06-001S-GS, and the fossil is reposit at the Raymond M. Alf Museum of Paleontology, Claremont, California, USA.

Comparisons and measurements. Linear measurements of RAM 22574 were collected with a digital calipers, to the nearest millimeter or 0.1 mm (Figure 2; Table 1). Circumference was measured to the nearest millimeter with a cloth measuring tape. For anatomical comparisons, casts of *Montanazhdarcho minor* (MOR 691) and *Quetzalcoatlus northropi* (TMM 41450) were compared directly with RAM 22574.

Wingspan estimation. The wingspan of RAM 22574 was approximated by scaling from relatively complete wings of pterodactyloid pterosaurs. Here, wing length is calculated as the sum of all sequential forelimb bone lengths (humerus, ulna, metacarpal IV, phalanges in digit IV). Wingspan is approximated by doubling wing length. As noted by Hone and Benton (2007), this neglects the width of the torso, but that is offset in part by the flexion of the wings in life. Data (Table 1) were taken from measurements published by Unwin et al. (2000) and Bennett

(2001a). Assuming that RAM 22574 was an ulna, each wing was scaled by ulna size for that specimen. To reduce concerns about allometry, only specimens in the approximate size range of RAM 22574 were used (<25% difference in ulnar length). Because RAM 22574 was slightly “telescoped”, two calculations were run—one with the bone length as preserved, and another adding an additional 15 mm to the bone length to accommodate the effects of the crushing. Finally, the wingspan of *Cryodrakon boreas* (TMP 1992.83.4) was estimated by scaling the holotype humerus (measurements from Godfrey & Currie, 2005) relative to the humerus for *Quetzalcoatlus* sp. (TMM 42422).

Results

Systematic Paleontology

Archosauria (Cope, 1869)
Pterosauria (Kaup, 1834)
Pterosauria indet.

Referred material. RAM 22574, a nearly complete ?ulna (Figure 1).

Locality and horizon. Locality RAM V2005022, the “Cripe Site,” located within Grand Staircase-Escalante National Monument, Garfield County, Utah. The site is situated in the middle unit of the Kaiparowits Formation, which is late Campanian in age (Roberts, Deino & Chan, 2005; Roberts et al., 2013).

Identification. This fossil is identified confidently as a pterosaur limb bone, and tentatively as an ulna. Because the descriptive terminology hinges upon these assumptions, I first address the underlying logic.

RAM 22574 is clearly hollow, which for the Late Cretaceous restricts possible identifications to either Theropoda or Pterosauria. The extremely thin cortical bone (between 0.7 and 1.7 mm) is distinct to pterosaurs (with the exception of dsungaripterids; Unwin, 2003), particularly for elements of this size. Thus, the identification to Pterosauria is quite confident.

Because key parts of RAM 22574 were damaged prior to fossilization (see below), identification of this bone within the skeleton is less certain. It is clearly a limb bone (rather than vertebra, ribs, or cranial material), but does not match well with morphology expected for any of the hind limb elements. There is nothing that resembles either the head or distal end of the femur, and the overall robustness of the bone (proportions of length vs. width) differs sharply from that seen in the tibia and fibula for typical Late Cretaceous pterosaurs of this size range (e.g., Bennett, 2001b; Averianov, 2010). A humerus can be excluded on the basis of a lack of a deltopectoral crest or the bulbous distal articular surface processes. Neither articular end shapes or element proportions fall within what would be expected for metacarpals or phalanges, regardless of clade. Thus, a radius or ulna seems to be the most likely identification.

The more heavily pre-depositionally damaged end of RAM 22574 shows topographic complexity that differs from what is seen in typical pterosaur radii (Wellnhofer, 1991; Bennett, 2001b; Veldmeijer, 2003). A prominent protrusion matches with a similar feature seen on *Montanazhdarcho minor* (MOR 691, personal observation on cast) and also material referred to *Cryodrakon boreas* (TMP 1965.14.398; Godfrey & Currie, 2005; Hone, Habib & Therrien, 2019). The relatively restricted nature of this protrusion differs from the more elongate anterior tuberosity seen on a radius referred to *Azhdarcho lancicollis* (Fig. 26 in Averianov, 2010). Finally, large pterosaur radii (especially in Late Cretaceous clades; e.g., *Quetzalcoatlus northropi*, TMM 41450) tend to show substantial tapering on the shaft relative to the proximal

and distal ends. RAM 22574, at least relative to the more complete end, does not show substantial tapering, which is more consistent with an ulna than a radius.

The overall proportions (length vs. width) of RAM 22574 also are more similar to that of an ulna than a radius. For instance, in *Quetzalcoatlus northropi* (TMM 41450) shows that radii tend to be far more slender (proportionately) than ulnae. Similar proportions are seen in *Montanazhdarcho minor*, *Azhdarcho lancicollis*, and also in *Pteranodon* spp. (Bennett, 2001b; McGowen et al., 2002; Averianov, 2010), as a few examples.

Assuming that RAM 22574 is an ulna, it can be further identified as a right element, based on the position of a roughened area on the proximal end that may represent the bicipital tuberosity (Figure 1B, C). This area should be on the ventral edge of the element's anterior surface.

Description. RAM 22574 measures 366 mm in maximum preserved length. As outlined above, this description assumes that RAM 22574 is a right ulna, so that directional and anatomical terminology follow accordingly. Both proximal and distal ends were slightly abraded prior to fossilization, with the proximal end most damaged. Part of this end was broken open prior to burial, as evidence by some rip-up clay clasts inside the bone as well as field observations of the incomplete element.

The proximal end of the element is mostly incomplete, preserving only a portion of what is interpreted as the ventral cotylus. The cotylus projects ventrally, with much of its surface abraded away. A roughened patch of bone around 20 mm distal to the peak of the cotylus may represent the bicipital tuberosity, for attachment of *m. biceps brachii* (Figure 1B, C). There is no evidence of pneumatic foramina on the proximal end of RAM 22574, but the area where such foramina would be expected is broken.

The shaft broadens gradually from the proximal to the distal end of the element (Figure 1C, F, G), although this appears accentuated in part by crushing. At mid-shaft, the cross-section is oval and elongated in the dorso-ventral direction (width/height ratio of 0.51; see Table 1). At the distal third of the bone, cortical thickness ranges from 1.0 to 1.6 mm; at the proximal third, it is 1.7 mm thick. At least part of the shaft is mildly telescoped through proximo-distal distortion, as shown by displacement around the shaft. This would add another 15 mm or so to the total length of the bone. The distal third of the shaft in RAM 22574 is flattened on its ventral surface. This is somewhat accentuated by crushing, but appears to be at least partly original morphology.

The distal end of RAM 22574 is slightly abraded, but overall more intact than the proximal end. Its distal margin in anterior view is relatively straight, probably accentuated by abrasion. When viewed end on, the ventral condyle region is more expanded than the dorsal condyle, although it appears that the dorsal condyle is abraded (Figure 1D). A broad depression separates the condyles on the posterior surface of the element, although the extent of this depression appears accentuated by crushing. After accounting for the space between the crushed bone pieces, the width of the distal end is exaggerated by around 10 percent.

Discussion

Unfortunately, the incomplete and crushed nature of RAM 22574 limits substantial interpretation of the element and the animal. In general, neither the radius nor ulna exhibit major diagnostic features in pterosaurs, so the element cannot currently be identified beyond Pterosauria. Nevertheless, RAM 22574 does represent the largest pterosaur bone yet known from the Kaiparowits Formation and only the second formally described element, and is thus useful for establishing the maximum size of pterosaurs in the Kaiparowits ecosystem.

The total wingspan for RAM 22574 is estimated at 4.3 to 5.9 m (Table 2). This places it within the same size range as *Quetzalcoatlus* sp. (TMM 42422; Maastrichtian, Javelina Formation, Texas) or *Cryodrakon boreas* (TMP 92.83.4; late Campanian, Dinosaur Park Formation, Alberta); the ulna of RAM 22574 is roughly the same length as that of TMM 42422. McGowen et al. (2002) estimated a 2.5 m wingspan for *Montanazhdarcho minor* (late Campanian, Two Medicine Formation, Montana), and Sullivan and Fowler (2011) estimated 3.5 m for *Navajodactylus* (late Campanian, Kirtland Formation, New Mexico). Thus, RAM 22574 stands alongside *C. boreas* as one of the largest pterosaurs known from late Campanian-aged terrestrial deposits of North America. Furthermore, it demonstrates that pterosaurs in this size range occurred in ecosystems both in the northern and southern parts of Laramidia (western North America) during the late Campanian. Future discoveries will undoubtedly help clarify phylogenetic relationships between pterosaurs living in terrestrial environments at this time, to see if they were relatively geographically restricted, or if individual species had continent-level ranges. Additionally, additional work is required to determine if large pterosaurs played similar ecological roles across their various environments.

Acknowledgments

This specimen would never have been revealed without the sharp eyes of Jeff Cripe, who discovered the locality where this bone was collected. The 2016 field crew of volunteers and museum staff from the Alf Museum assisted in collection of this fossil, and Jared Heuck skillfully completed preparation. Discussions with Dave Hone, Nate Carroll, and others were helpful in establishing and confirming the identity of this fossil. We thank Valeria Pellicer for her renderings of the bone. Fieldwork in GSENM was completed under US Bureau of Land

Management paleontology permits UT06-012E-GS and UT06-001S-GS, with facilitation by Alan Titus and Greg MacDonald.

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Figure 1

RAM 22574, ulna of Pterosauria indet.

A) ?dorsal; B) ?proximal; C) ?anterior; D) ?distal; E) ?ventral; and F) ?posterior views; with G) showing interpretive drawing of ?posterior view, including missing parts; and H) showing restored view of bone in ?posterior view. Scale bars equal 10 cm. Abbreviations: ?bt, ?bicipital tuberosity; ?dist, ?distal end; ?prox, ?proximal end.

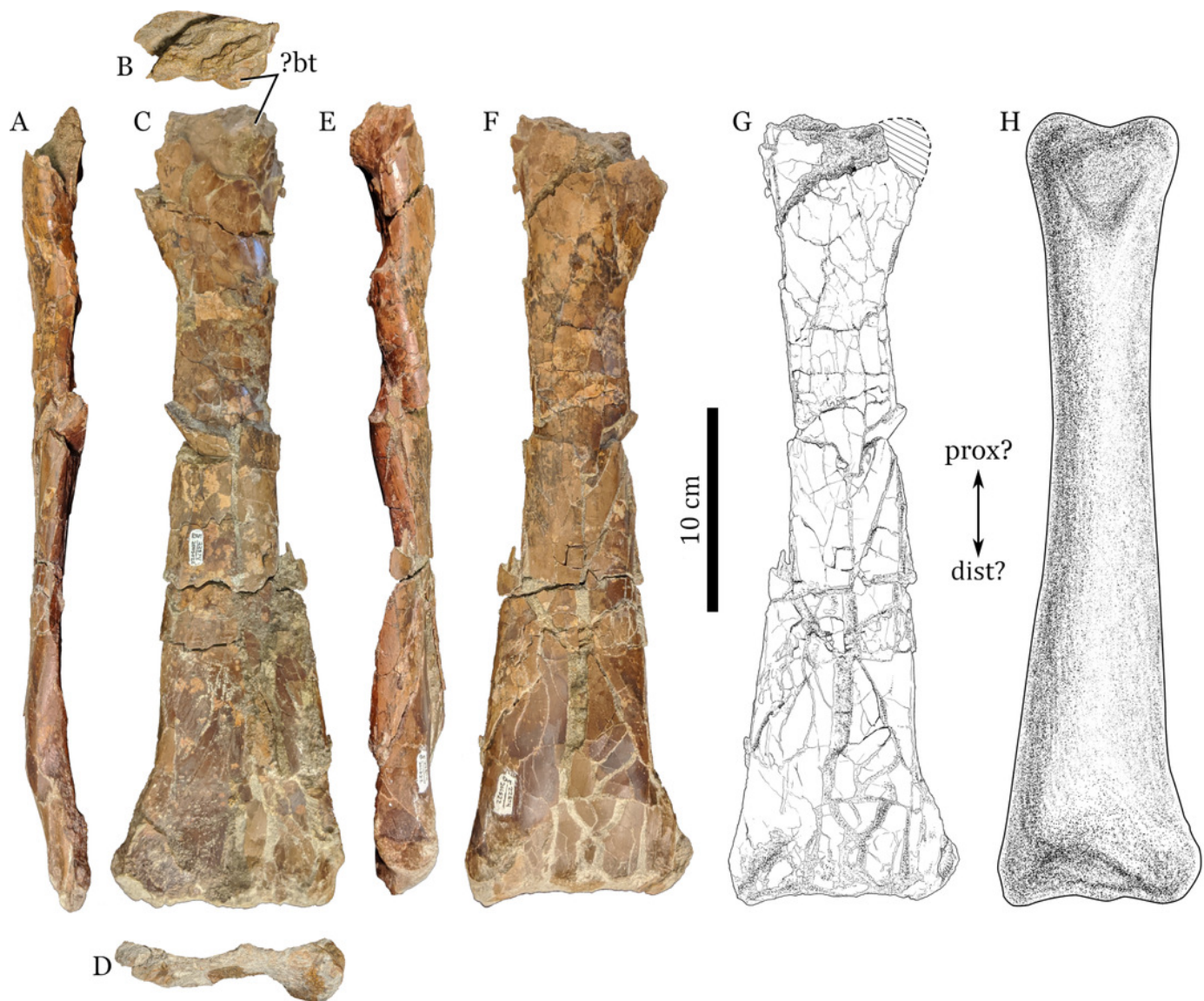


Figure 2

Interpretive drawing of RAM 22574, showing measurements taken here.

Measurements include: 1a, maximum proximo-distal length; 1b, maximum proximo-distal length (adjusting for telescoping); 2, maximum dorso-ventral width of proximal end; 3, minimum antero-posterior width of proximal end; 4, antero-posterior width of proximal end at ventral cotyle; 5, dorso-ventral width at narrowest point of shaft; 6, dorso-ventral width at mid-shaft; 7, antero-posterior width at mid-shaft; 8, circumference at mid-shaft; 9, maximum dorso-ventral width of distal end; 10, maximum antero-posterior width of distal end; 11, minimum antero-posterior width of distal end. Data are provided in Table 1.

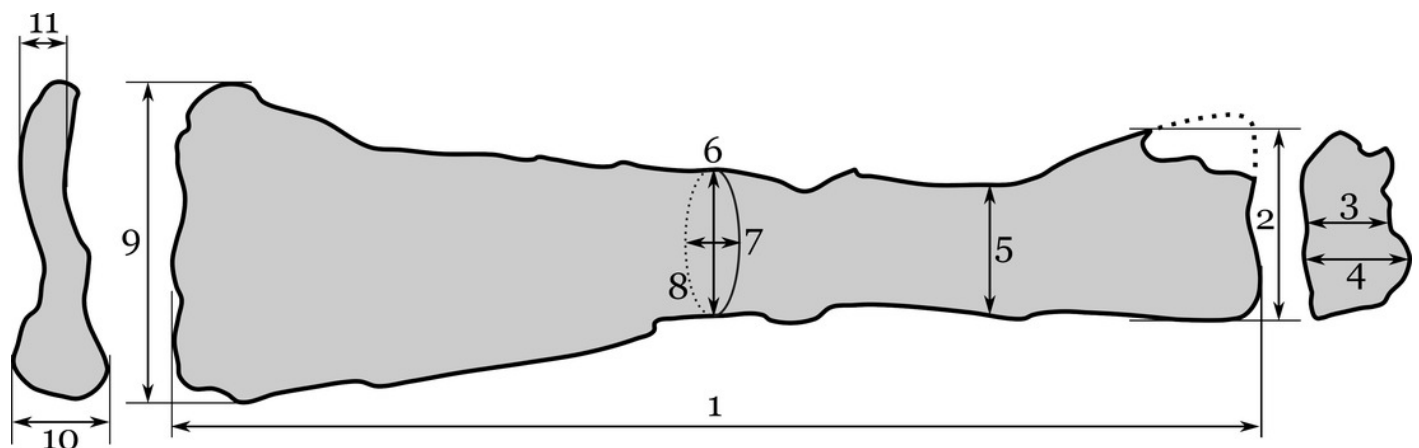


Table 1(on next page)

Measurements of pterosaur ?ulna, RAM 22574, in millimeters.

All measurements were taken with sliding digital calipers, except for 6, which was measured with a cloth measuring tape. See Figure 2 for explanation of measurements.

	Standard	Measurement
1a	Maximum proximo-distal length (as preserved)	366
1b	Maximum proximo-distal length (adjusting for telescoping)	381
2	Maximum dorso-ventral width of proximal end (as preserved)	71
3	Minimum antero-posterior width of proximal end (as preserved)	32
4	Antero-posterior width of proximal end at ventral cotyle	48
5	Dorso-ventral width at narrowest point of shaft	49
6	Dorso-ventral width at mid-shaft (as preserved)	52
7	Antero-posterior width at mid-shaft (as preserved)	27
8	Circumference at mid-shaft (as preserved)	133
9	Maximum dorso-ventral width of distal end (as preserved)	106
10	Maximum antero-posterior width of distal end (as preserved)	31
11	Minimum antero-posterior width of distal end (as preserved)	17

Table 2 (on next page)

Comparative measurements of selected pterodactyloid pterosaurs, with wingspan for RAM 22574 scaled from those measurements.

Two ulna lengths are provided for RAM 22574, representing the element as preserved (the smaller number) and a second estimate accounting for mild telescoping that reduced the preserved length of the element. Measurements for *Pteranodon* are taken from Bennett (2001a), and those for *Cryodrakon* are from Godfrey and Currie (2005); all others are from Unwin et al. (2000). Measurements are in millimeters, except for wingspan estimates, which are in meters. Abbreviations: H, humerus length; MC-IV, metacarpal IV; IV-1,-2,-3,-4, fourth digit manual phalanges 1 through 4; RAM WS, range of wingspans estimated for RAM 22574 based on direct scaling from each specimen; WS, wingspan (calculated by summing forelimb bone lengths and multiplying by 2).

Taxon	Specimen	H	U	MC-IV	IV-1	IV-2	IV-3	IV-4	WS	RAM WS
<i>Arthurdactylus conandoylei</i>	SMK 1132 PAL	230	312	227	445	402	312	275	4.41 m	5.2–5.4 m
<i>Anhanguera santanae</i>	NSM PV 19892	257	384	257	462	387	270	225	4.48 m	4.3–4.5 m
<i>Cryodrakon boreas</i>	TMP 1992.83.4	245							4.57 m (est.)	
<i>Pteranodon</i> sp.	FHSM 184	269	393	583	653	539	390	194	6.04 m	5.6–5.9 m
<i>Quetzalcoatlus</i> sp.	TMM 42422	250	358	620	602	305	156	39	4.66 m	4.8–5.0 m
Pterosauria indet.	RAM 22574		366/ 381							4.3–5.9 m