

Refining the marine reptile turnover at the Early-Middle Jurassic transition (#53076)

1

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Refining the marine reptile turnover at the Early-Middle Jurassic transition

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Even though a handful of long-lived reptilian clades dominated Mesozoic marine ecosystems, several biotic turnovers drastically changed the taxonomic composition of these communities. A seemingly slow paced, within-System turnover took place across the Early-Middle Jurassic transition, seeing the demise of early neoichthyosaurians, rhomaleosaurid plesiosaurians and early plesiosauroids in favour of ophthalmosaurid ichthyosaurians and cryptoclidid and pliosaurid plesiosaurians, clades that will dominate the Late Jurassic and, for two of them, the entire Early Cretaceous as well. The fossil record of this turnover is however extremely poor, and this change of dominance appears to be spread across the entire middle Toarcian-Bathonian interval. We describe a series of ichthyosaurian and plesiosaurian specimens from successive geological formations in Luxembourg and Belgium that detail the evolution of marine reptile assemblages across the Early-Middle Jurassic, within a single sub-basin. These fossils reveal the continuing dominance of large rhomaleosaurid plesiosaurians, microcleidid plesiosaurians, and *Temnodontosaurus*-like ichthyosaurians up to the latest Toarcian, indicating that the structuration of the upper tier of marine ecosystems remained essentially constant up to the very end of the Early Jurassic. These fossils also suddenly record ophthalmosaurid ichthyosaurians and cryptoclidid plesiosaurians by the early Bajocian. As a result, the Early-Middle Jurassic marine reptile turnover appears restricted to the sole Aalenian stage, reducing the uncertainty of its duration, at least for ichthyosaurians and plesiosaurians, to 4 instead of 14 million years.

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2 Middle Jurassic transition

3

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15

16 Abstract

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18 ecosystems, several biotic turnovers drastically changed the taxonomic composition of
19 these communities. A seemingly slow paced, within-System turnover took place across
20 the Early-Middle Jurassic transition, seeing the demise of early neoichthyosaurians,
21 rhomaleosaurid plesiosaurians and early plesiosauroids in favour of ophthalmosaurid
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26

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30 fossils reveal the continuing dominance of large rhomaleosaurid plesiosaurians,
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35 early Bajocian. As a result, the Early-Middle Jurassic marine reptile turnover appears
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37 ichthyosaurians and plesiosaurians, to 4 instead of 14 million years.

38

39

40 Introduction

41 A series of diapsid clades dominated marine ecosystems during the entire Mesozoic (e.g.
42 Motani, 2009; Benson, 2013; Pyenson, Kelley & Parham, 2014; Kelley & Pyenson,
43 2015). This long-term dominance was, however, pulsed by a series of extinctions and
44 turnovers (Massare, 1987; Callaway & Massare, 1989; Bardet, 1994, 1995; Benson et al.,
45 2010; Benson & Druckenmiller, 2014), the most-studied of which being those happening
46 during the latest Triassic (Thorne, Ruta & Benton, 2011; Benson, Evans &
47 Druckenmiller, 2012; Fischer et al., 2014c; Dick & Maxwell, 2015; Moon & Stubbs,
48 2020) and at the Jurassic-Cretaceous boundary (Tennant et al.; Fischer et al., 2012, 2013;
49 Benson & Druckenmiller, 2014; Young et al., 2014a; Tennant, Mannion & Upchurch,
50 2016; Zverkov et al., 2018).

51

52 The dynamics of within-System marine reptile turnovers are less well known, despite
53 evident modifications of assemblages outside System boundaries, such as during the early
54 Late Cretaceous (Bakker, 1993; Bardet et al., 2008; Fischer et al., 2016, 2018). The
55 composition of marine reptile communities clearly changed across the Early-Middle
56 Jurassic transition (Maxwell, Fernández & Schoch, 2012; Vincent et al., 2013). Indeed,
57 early neochthyosaurian ichthyosaurs and microcleidid plesiosaurians seemingly met
58 their demise and rhomaleosaurid plesiosaurians went near extinction (Maxwell,
59 Fernández & Schoch, 2012; Fischer et al., 2013; Benson, Zverkov & Arkhangelsky,
60 2015), while ophthalmosaurid ichthyosaurs, thalattosuchian crocodylomorphs,
61 cryptoclidid plesiosauroids, and pliosaurids radiated (Motani, 1999a; Ketchum & Benson,
62 2011a; Fischer et al., 2013; Benson & Druckenmiller, 2014; Moon, 2017; Foffa, Young
63 & Brusatte, 2018). These newly diversified clades will then dominate the Late Jurassic
64 and beyond.

65

66 However, the fossil record of marine reptiles across the middle Toarcian – Callovian
67 interval is extremely poor (Benson et al., 2010; Maxwell, Fernández & Schoch, 2012)
68 (Figure 1) and no region preserves fossiliferous successions spanning this transition.
69 These biases blur both the tempo and the severity of this turnover. We tackle this issue by
70 describing and analysing the composition of the late Toarcian-Bajocian marine reptile
71 assemblages of Luxembourg and Belgium. **We show that typical Early Jurassic forms**
72 such as rhomaleosaurid and microcleidid plesiosaurians and **non-baracromian**
73 ichthyosaurians persisted in mid latitudes up to the latest Toarcian and that
74 ophthalmosaurid ichthyosaurians and cryptoclidid plesiosaurians, the dominant Late
75 Jurassic clades, suddenly appear during the early Bajocian, providing a refined timeline
76 in a single sub-basin for what appears as an abrupt, Aalenian-restricted turnover.

77

78

79 Material & Methods

80

81 Upper Toarcian of Luxembourg: Thouarsense, Pseudoradiosa, and Aalensis zones

82 Towards the end of the Early Jurassic, the area of present-day Luxembourg was located
83 in the northeast Paris Basin within the northwestern Peri-Tethys Ocean (e.g. Pieńkowski
84 et al., 2008; Schintgen & Förster, 2013). In a shallow, near-coastal sea between the
85 former landmasses of the Rhenish Massif in the north and the Vosges and Black Forest in
86 the south, a thick siliciclastic succession was deposited during the late Toarcian, with
87 clayey to silty sediments in the lower parts and silty to sandy, iron-grain rich sediments in
88 the upper part following a general coarsening upwards trend (Siehl & Thein, 1989). The
89 material described herein includes a single, semi-articulated post-cranial ichthyosaur
90 skeleton from the so-called “Couches à *Astarte volizi*” (Dittrich, 1993), a clay-rich
91 siltstone corresponding to level 11 of Delsate and Weis (2010), dated to the late Toarcian
92 Thouarsense Ammonite Zone of the Grandcourt Formation.

93

94 All the other late Toarcian marine reptile remains from Luxembourg described herein
95 were found in the overlying Minette ironstone Formation (Lucius, 1945). Deposition of
96 the Minette successions took place under the influence of strong tidal currents (Teyssen,
97 1984) during the upper Toarcian Pseudoradiosa and Aalensis ammonite Zones (Di
98 Cencio & Weis, 2020). The Aalenian Opalinum ammonite zone has not yet been detected
99 in the area and probably forms a local hiatus (Maubeuge, 1972; Guérin-Franiatte & Weis,
100 2010).

101

102 Upper Toarcian of Belgium: Dispansum-Pseudoradiosa-Aalensis zones

103 The Mont-Saint-Martin Formation crops out in southern Belgium and includes iron-rich
104 beds forming a slightly diachronous equivalent of the Minette ironstones in Luxembourg.
105 These ironstones contain all the marine reptiles and were deposited in a shallow marine
106 setting during the late Toarcian (Dispansum to Aalensis ammonite zones; Boulvain et
107 al., 2001). One fossil (IRSNB Vert-06455-0001) was found just outside the Belgian
108 border, in the Meurthe-et-Moselle department, north-eastern France (Figure 2).

109

110 Middle-upper Aalenian of Luxembourg: Murchisonae and Concavum zones

111 The Toarcian part of the Minette ironstones in Luxembourg is locally overlain by a few
112 meters of iron-rich marly sandstones and thin beds with calcareous nodules deposited
113 during the Aalenian Murchisonae to Concavum ammonite zones (Guérin-Franiatte &
114 Weis, 2010; Sadki, Weis & Braun, 2020). These deposits form the onset of a
115 transgressive succession with increasingly finer sedimentation leading to the so-called
116 “marnes micacées”, a succession of poorly lithified silty claystones with thin layers of
117 phosphatic nodules, dated to the lowermost Bajocian Discites ammonite zone (Guérin-
118 Franiatte & Weis, 2010).

119

120 **Lower Bajocian of Luxembourg: Humphriesianum Zone**

121 The youngest marine strata in Luxembourg correspond to a succession of calcarenites,
122 marl-limestone alternations and coral limestones deposited on the northeast margin of the
123 Burgundy carbonate platform at the end of the lower Bajocian. The marine reptile
124 remains described herein were all recovered from the so-called “Marnes sableuses
125 d’Audun-le-Tiche”, dated to the lower Bajocian Humphriesianum ammonite zone
126 (Delsate et al., 2018; Popov, Delsate & Felten, 2019).

127

128 **Fossilised diversity data**

129 We extracted the number of fossil collections of Jurassic (Hettangian-Tithonian) marine
130 reptiles (Ichthyosauria, Plesiosauria, Thalattosuchia, Pleurosauria, Angolachelonia) from
131 the paleobiology database (<https://paleobiodb.org>) on the 26th March 2020. We binned
132 these data per geological stage and generate a plot of the number of collections over time
133 in R v3.6.2 (R Core Team, 2016), sing the geoscale v2.0 package (Bell, 2015).

134 **Centra proportions**

135 We gathered a series of measurements on centra, recording their position within the
136 vertebral column (Table 1). We treated and plotted these data in R v3.6.2 (R Core Team,
137 2016) to visualize changes through time, if any.

138

139 **Institutional abbreviations**

140 **IRSNB:** Royal Belgium Institute of Natural Sciences, Brussels, Belgium. **MNHNL:**
141 Muséum national d’histoire naturelle du Luxembourg, Luxembourg-ville, Luxembourg.
142 **ULg-PA:** Collections de paléontologie animale de l’Université de Liège, Liège, Belgium.

143

144

145 **Comparative Descriptions**

146

147 **Thouarsense zone fauna, upper Toarcian, Luxembourg**

148

149 ICHTHYOSAURIA De Blainville, 1835

150 PARVIPELVIA Motani, 1999

151 THUNNOSAURIA Motani, 1999

152 BARACROMIA Fischer et al., 2013

153 Baracromia indet.

154 (Figure 3)

155

156 A single specimen (MNHN TM212) is recorded from the Thouarsense zone: it is a
157 partial articulated ichthyosaur comprising fifty-two centra, several ribs and gastralia, and
158 two partial forefins, from the Dudelange locality. All the preserved pre-apical centra are
159 markedly rounded and anteroposteriorly short, bearing close similarities with Jurassic

160 thunnosaurians such as *Stenopterygius* and *Ophthalmosaurus* (Buchholz, 2001; Massare
161 et al., 2006). Posterior dorsal centra maintain clear bicipital rib articulations and all ribs
162 have a 8-shaped cross-section, which also suggest that this specimen is a thunnosaurian
163 (Sander, 2000; V.F. pers. obs. on NHMUK 2003, R5465, R498). The posterior apical
164 centra are strongly compressed laterally, suggesting the presence of a lunate tailfin;
165 although is the angle of the tailbent is unknown. The humerus is incompletely-preserved,
166 lacking the capitulum. The dorsal trochanter is proximodistally short, as it does not
167 extend to mid-shaft, while the deltopectoral crest is elongated and parallel to the long axis
168 of the humerus. The small dorsal trochanter suggest that this taxon is not an
169 ophthalmosaurid (Motani, 1999a). The humerus forms a prominent anterior process that
170 possesses a conspicuous, anterodorsally facing facet. This facet exhibits a pitted bone
171 texture which suggests the presence of a cartilaginous cap in vivo. While many Early
172 Jurassic ichthyosaurians possess a mesiodistally-thick anterodistal expansion of the
173 humerus (*Temnodontosaurus*, *Leptonectidae*, *Hauffiopteryx*, *Stenopterygius* (Johnson,
174 1979; McGowan & Milner, 1999; McGowan, 2003; McGowan & Motani, 2003; Maisch,
175 2008; Caine & Benton, 2011; Martin et al., 2012), this expansion here forms a flat facet
176 but not for articulation with an anterior accessory epipodial element (the rounded and
177 notched anterior surface of the radius confirms the absence of such an ment). This
178 combination of features rules out earliest neoichthyosaurians and ophthalmosaurids (e.g.
179 Maisch & Matzke, 2000; McGowan & Motani, 2003) and has only been reported in the
180 poorly known, ?late Toarcian-lower Bajocian taxon *Dearcmhara shawcrossi* (Brusatte et
181 al., 2015), although the facet is smaller and deeper in *D. shawcrossi*. Epipodial and
182 proximal elements closely resemble those of *Stenopterygius* spp. and *Chacaicosaurus*
183 *cayi* (Johnson, 1979; Fernández, 1994; Maxwell, Fernández & Schoch, 2012), being
184 polygonal and dorsoventrally thick, with no spatium interosseum. The radius and the
185 radiale are notched as in many early neoichthyosaurians (e.g. Huene, 1922; McGowan,
186 1974; Maxwell, 2012; Lomax & Massare, 2016). The intermedium forms two distal
187 facets unequal in size: a large, distally facing facet for articulation with distal carpal 3 and
188 a smaller, posterodistally facing facet for articulation with distal carpal 4 (the
189 ‘latipinnate’ condition, although the latipinnate-longipinnate dichotomy fails to capture
190 what is essentially a trait with continuous spectrum; see Mazin (1982) and Motani
191 (1999b) for discussions). The specimen MNHN TM212 is regarded as a non-
192 ophthalmosaurid baracromian, resembling *Stenopterygius* spp., *Chacaicosaurus cayi*, and
193 *Dearcmhara shawcrossi*.

194

195

196 **Pseudoradiosa zone fauna, upper Toarcian, Luxembourg**

197

198

PLESIOSAURIA De Blainville, 1835

199

Rhomaleosauridae Nopcsa, 1928

200 Rhomaleosauridae indet.
201 (Figure 4)
202
203 Two large plesiosaurian propodials have been recovered from Pseudoradiosa zone: one
204 (MNHN DOU307) from the “Couche grise” of the Rodange locality, and one (MNHN
205 KA109) from the “Couche noire” of the Esch-sur-Alzette locality.
206
207 MNHN DOU307 is a complete left propodial (proximal-distal length: 390mm; distal
208 width: 210mm) with a weak posterior and dorsal curvature. The shaft is as wide as the
209 capitulum, and there is no anterior expansion indicating that this propodial is an early
210 plesiosaurian humerus (Storrs, 1997; Bardet, Godefroit & Sciau, 1999; O’Keefe, 2004;
211 Smith & Vincent, 2010; Benson, Evans & Druckenmiller, 2012; Vincent & Storrs, 2019).
212 The dorsal tuberosity is as anteroposteriorly wide as the capitulum and is weakly
213 demarcated. A bulge with muscle scars is present on the ventral surface of the shaft, 30
214 mm distally to the capitulum. A very short preaxial flange is present distally, forming a
215 small, anterodistally facing triangular facet. The radial/tibial facet is flat and marks an
216 angle of ca. 45° with the ulnar/fibular facet. This facet is convex and rugose, indicating
217 the presence of an extensive cartilage layer. A larger postaxial flange is present, forming
218 distally a semioval facet that is hardly discernible from the ulnar/fibular facet. There is no
219 distal ridge, unlike in the cryptoclidid *Colymbosaurus* (Benson & Bowdler, 2014). The
220 absence of a marked curvature (the anterior surface is straight in dorsal and ventral
221 views) and the presence of a conspicuous postaxial flange suggest rhomaleosaurid
222 affinities (Smith & Vincent, 2010; Smith & Benson, 2014; Smith, 2015) (with the
223 exception of *Lindwurmia thiuda* Vincent & Storrs, 2019), which are reinforced by the
224 large size of the propodial.
225
226 MNHN KA109 is the distal end of a plesiosaurian propodial; the anterior-posterior
227 asymmetry suggests it is also a humerus. It resembles MNHN DOU307 in lacking a
228 preaxial flange and the postaxial flange is smaller than in MNHN DOU307. The distal
229 surface is markedly rounded in dorsoventral view. A prominent, strongly pitted
230 anteroposterior ridge textures of the distal surface, obliterating radial/tibial and
231 ulnar/fibular facets. The postaxial flange forms a triangular and concave facet posterior to
232 the distal ridge. A distal ridge is regarded as an autapomorphic feature of *Colymbosaurus*
233 (Benson & Bowdler, 2014; Arkhangelsky et al., 2019) but the lack of an extensive
234 postaxial flange precludes such a referral; the size of the humerus and its similarities with
235 MNHN DOU307 suggest rhomaleosaurid affinities.
236
237 **Dispansum-Pseudoradiosa-Aalensis zones, upper Toarcian, Belgium**
238
239 ICHTHYOSAURIA De Blainville, 1835

279 IRSNB Vert-00000-00800 is a partial left ichthyosaurian surangular, probably associated
280 with the angular IRSNB Vert-06462-0002. It consists of posterior third, where the
281 angular in a mesiolaterally-flattened bone with a slightly convex lateral surface and a
282 slightly concave medial surface for accommodation of the Meckelian canal. The dorsal
283 surface of the surangular is saddle-shaped, forming a long and shallow concavity anterior
284 to the coronoid process. The preserved morphology of IRSNB Vert-00000-00800 is not
285 diagnostic, but we assign it to *Parvipelvia* indet. because of its probable association with
286 the angular IRSNB Vert-06462-0002 (see below).

287

288 IRSNB Vert-06462-0002 is a partial left ichthyosaurian angular, bearing the typical
289 double-grooved dorsal surface. The lateral wall of the angular is dorsoventrally short
290 anteriorly, and only increases in size posterior to the level of the coronoid process. This
291 narrow lateral exposure of the angular suggests that this specimen does not belong to
292 Ophthalmosauridae, as this clade is notably characterized by large exposure of the
293 angular in lateral view (Motani, 1999a).

294

295 IRSNB Vert-00000-00803 and IRSNB Vert-00000-00804 are large ichthyosaurian
296 centra. IRSNB Vert-00000-00803 is 118 mm high and 48 mm long, giving a height-
297 length ratio of 2.45; IRSNB Vert-00000-00804 is damaged on the edges and will not be
298 measured but is of comparable size. The lateral surfaces of these centra are rounded in
299 anteroposterior view, as in parvipelvians (Merriam, 1908; Maisch & Matzke, 2000), but
300 their size departs from that of Early Jurassic thunnosaurians, even particularly large
301 individuals (e.g. Maxwell, 2012; Lomax & Sachs, 2017). These centra rather match the
302 size and shape of the dorsal-caudal centra of some large non-thunnosaurian
303 neioichthyosaurians such as *Temnodontosaurus* spp. (Godefroit, 1993a; Martin et al.,
304 2012) and *Excalibosaurus costini* (McGowan, 2003). IRSNB Vert-00000-00803
305 possesses a single apophysis (the parapophysis) placed laterally, suggesting it is middle
306 preflexural caudal (Huene, 1922; McGowan & Motani, 2003). IRSNB Vert-00000-00804
307 possesses a diapophysis and a parapophysis placed ventrolaterally and is regarded as a
308 posterior dorsal centrum. The size and shape of IRSNB Vert-00000-00803 and IRSNB
309 Vert-00000-00804 match those of large early neioichthyosaurians and but are regarded as
310 *Parvipelvia* indet. in the absence of unambiguous synapomorphies.

311

312 IRSNB Vert-00000-00808 is an ichthyosaurian anterior caudal centrum. Its height-length
313 ratio is 65 mm / 22 mm = 2.94; this high shape ratio close to 3 might suggest
314 thunnosaurian affinities (Buchholtz, 2001). The lateral surface of the centrum is rounded
315 as in parvipelvians (Merriam, 1908; Maisch & Matzke, 2000) and slightly oval, as the
316 ventral surface is flattened. The parapophyses are located lateroventrally.

317

318 IRSNB Vert-06462-0005 is an ichthyosaurian cervical centrum. Its height-length ratio is
319 63 mm / 24 mm = 2.62. The lateral surface of the centrum is rounded as in parvipelvians
320 (Merriam, 1908; Maisch & Matzke, 2000). A faint ventral keel is present, giving the
321 centrum a pentagonal shape. Such a shape is often seen in the cervical/anterior dorsals of
322 parvipelvians (e.g. Huene, 1922; Fischer et al., 2014b; Lomax, Porro & Larkin, 2019) and
323 is not of diagnostic value at a lower taxonomic level at the current state of our
324 knowledge.

325

326

PLESIOSAURIA De Blainville, 1835
Plesiosauria indet.

330 IRSNB Vert-11312-00007 is a plesiosaurian dorsal centrum, as evidenced by the absence
331 of transverse processes. Its height-length ratio is $67 \text{ mm} / 44 \text{ mm} = 1.52$. The centrum is
332 poorly preserved and we assign it to Plesiosauria indet.

333

334

335

Aalensis zone fauna, upper Toarcian, Luxembourg

337

338 The material from the Aalensis zone of the Minette ironstones is more abundant. We
339 identified 12 ichthyosaurians and 8 plesiosaurians in the MNHL collections. Most of the
340 material consist of isolated centra, complemented by three plesiosaurian propodials, one
341 ichthyosaur coracoid, and one plesiosaurian tooth crown.

342

ICHTHYOSAURIA De Blainville, 1835

PARVIPELVIA Motani, 1999

345

Parvipelvia indet.

346

347

348 M NHNL DOU944 is an ichthyosaurian anterior dorsal centrum from an unknown
349 locality. A central keel is not present but the centrum forms a ventral expansion, giving
350 the centrum a slight pentagonal shape. The diapophysis is small and rounded, as in
351 parvipelvians and unlike in more primitive forms (Merriam, 1908; Huene, 1922) and is
352 strongly protruding and has a rounded distal surface. Shallow dorsoventral ridges connect
353 to the base of the diapophysis, anterior and posteriorly. The parapophysis is flatter and
354 less prominent and its anterior margin merges with the anterior edge of the centrum. Such
355 a pentagonal shape is often seen in parvipelvians (e.g. Huene, 1922; Fischer et al., 2014b;
356 Lomax, Porro & Larkin, 2019) and is not of diagnostic value at a lower taxonomic level
357 at the current state of our knowledge.

358
359 M NHNL DOU998 is an anterior caudal centrum from an unknown locality. The centrum
360 has a rounded outline, as in usually parvipelvian ichthyosaurians (Merriam, 1908; e.g.
361 Maisch & Matzke, 2000; McGowan & Motani, 2003). Its height/length ratio is
362 90/31=2.9. This high ratio close to 3 might suggest thunnosaurian affinities (Buchholz,
363 2001), but in the absence of additional data, this specimen is regarded as Parvipelvia
364 indet.

365
366 M NHNL DOU369 is a series of disarticulated ichthyosaur centra. One originates from
367 the mid-thoracic region and two others from the anterior caudal region, their respective
368 height/length ratios are as follows: 68/32=2.1, 61/35=1.74, and 63/32=1.96. The mid-
369 thoracic resembles M NHNL DOU944. The rather elongated anterior caudal suggest these
370 centra do not belong to a thunnosaurian, but rather to an early parvipelvian (Buchholz,
371 2001).

372
373 M NHNL DOU353 is a fragmentary left coracoid from an unknown locality. It preserves
374 the lateral part and a fragment of the intercoracoid facet. The ventral surface is
375 moderately saddle-shaped (intercoracoid facet is only 56mm thick while the mesiolateral
376 length is 180mm). The glenoid facet is large (anteroposterior length=75mm) and oval,
377 while the scapular facet forms a small triangle (anteroposterior length= 39mm) and forms
378 a ca. 60° angle with the glenoid facet. The posterolateral emargination is wide and deep,
379 the coracoid surface posterior to the glenoid faces posteriorly and not posterolaterally,
380 similar to the condition seen in *Ichthyosaurus communis*, *Ichthyosaurus anningae*,
381 *Protoichthyosaurus postaxialis*, and many but not all specimens of *Temnodontosaurus*
382 *platyodon* (Home, 1819; Godefroit, 1993a; Lomax & Massare, 2015; Lomax, Porro &
383 Larkin, 2019). The Triassic shastasaurid *Shatasaurus neoscapularis* also possesses a wide
384 posterior notch, but the coracoid of that taxon is different, having a very wide anterior
385 notch (McGowan, 1994). We regard this coracoid as belonging to a non-baracromian
386 parvipelvian.

387
388 M NHNL DOU378 is a very large centrum from an unknown locality. It originates from
389 the posterior dorsal – sacral – anterior caudal region (length: 65mm and diameter:
390 175mm, estimated height: 180-185mm, H/L ratio: 2.79-2.84). Its size departs from the
391 Early Jurassic thunnosaurians (*Ichthyosaurus* and *Stenopterygius*, which have much
392 smaller centra even particularly large individuals (e.g. Maxwell, 2012; Lomax & Sachs,
393 2017)). It rather matches the size and shape of the anteriormost caudal centra of some
394 large non-thunnosaurian neoichthyosaurians such as *Temnodontosaurus* spp. (Godefroit,
395 1993a; Martin et al., 2012) and *Excalibosaurus costini* (McGowan, 2003) (the anterior
396 caudals are much longer in *Eurhinosaurus longirostris* [V.F., pers. obs. on MNHN 1946-
397 20]).

398

399

400

401 PLESIOSAURIA De Blainville, 1835

402 Plesiosauria indet.

403 (Figures 7, 8)

404

405 MNHNL DOU906 is a plesiosaurian tooth crown from the Esch-sur-Alzette locality. It is
406 slender (length/basal diameter= 23/8= 2.875), labiolingually flattened and recurved
407 indicating it does not belong to derived rhomaleosaurids and thalassophoneans (Owen,
408 1865; Smith & Vincent, 2010; Ketchum & Benson, 2011b; Smith & Benson, 2014). The
409 crown bears fine striations along the entire surface.

410

411 MNHNL DOU724 is a plesiosaurian sacral vertebra probably originating from the Esch-
412 sur-Alzette locality. The articular surface is strongly dorsoventral depressed and the
413 centrum is as wide as long. As in MNHNL DOU722, MNHNL DOU723, and MNHNL
414 DOU954, the subcentral foramina are widely spaced and are separated by a well-
415 rounded/convex surface. Similar to the other centra, the dorsal edge of the articular
416 surface is notched by the neural canal. The transverse process is elongated, stout, and
417 points posterolaterally. The rib facet is 8-shaped (waisted), this time with a larger, ear-
418 shaped dorsal part and a smaller, semioval ventral part. The neural canal is also 8-shaped,
419 resembling a ‘plesiosauroid’ pectoral vertebra from the Aalenian-Bajocian of Australia
420 (Kear, 2012). Despite having similar size to the vertebrae we referred to as
421 Rhomaleosauridae indet., none of the feature we highlight above are recorded in
422 *Rhomaleosaurus* (Smith, 2013; Smith & Benson, 2014) and their representation in other
423 derived rhomaleosaurids is unclear. Accordingly, we regard this specimen as Plesiosauria
424 indet.

425

426 Rhomaleosauridae Nopcsa, 1928

427 Rhomaleosauridae indet.

428 (Figures 7, 8)

429

430 MNHNL DOU722 and MNHNL DOU723 are two plesiosaurian pectoral vertebrae from
431 the Esch-sur-Alzette locality. The articular surface of the centrum is oval, with a dorsal
432 notch corresponding to the neural canal. The subcentral foramina are displaced laterally,
433 being located on the ventrolateral surface. There is no ventral keel, but the ventral surface
434 in between the subcentral foramina is clearly transversely convex. The rib facets are oval,
435 connect to the neural arch, and their long axis is vertical; they extend ventrally below the
436 level of the notochordal pit. The semi-circular prezygophyses face dorsomedially and
437 are clearly separated from one another; their ventral margin is set below the centre of the

438 neural canal, even though the neural canal is large. Anteriorly, the neural spine forms a
439 basal triangular cavity containing a paired ridge. The height of this trough is about 1/3 of
440 the total height of the neural spine. On the posterior surface, the basal region of the neural
441 spines forms a deep tear-shaped concavity that extends for about one half of the total
442 height of the neural spine, and appears similar to the condition of Microcleididae (Bardet,
443 Godefroit & Sciau, 1999; Schwermann & Sander, 2011) and *Rhomaleosaurus* (Smith &
444 Benson, 2014). However, the neural spines are much shorter than in microclidids (e.g.
445 Owen, 1865; Bardet, Godefroit & Sciau, 1999). The neural spine possesses a posterior
446 ridge while its anterior surface is concave, as in *Rhomaleosaurus thortoni* (Smith &
447 Benson, 2014). The neural spine thickens dorsally and its dorsal half is posteriorly
448 inclined, unlike in *Westphaliasaurus simonensis* (Schwermann & Sander, 2011). We
449 assign these vertebrae to Rhomaleosauridae indet.

450

451 MNHNL DOU954 is a posterior pectoral/anterior dorsal vertebra from an unknown
452 locality. It is essentially similar to MNHNL DOU722 and MNHNL DOU723. The main
453 difference is the existence of ventral expansion of the centrum, even if a clear ventral keel
454 remains absent. This ventral expansion gives the articular surface of the centrum a heart
455 shape. The neural spine is concave anteriorly and bears a median ridges similar to the
456 condition of *Rhomaleosaurus thortoni* (Smith & Benson, 2014). We assign this vertebra
457 to Rhomaleosauridae indet.

458

459 MNHNL DOU558 from the Belvaux locality and MNHNL DOU324A and MNHNL
460 DOU324B (broken, juvenile) from the Esch-sur-Alzette locality are three plesiosaurian
461 humeri. These strongly resemble MNHNL DOU307: the shaft is long and slightly curved
462 posteriorly and dorsally, the dorsal tuberosity is wide, a boss with muscle scars is present
463 ventrally to the capitulum, the preaxial lamella is very short and forms a small triangular
464 distal facet, the postaxial lamella is longer and forms a semioval posterodistal facet,
465 radial/tibial and ulnar/fibular facets form a 45–55° angle and no distal ridge is present.
466 The only difference is the position of the dorsal tuberosity, which is offset ventrally,
467 creating a flattened, dorsomedially-facing surface between the dorsal part of the
468 capitulum and the ventral part of the dorsal tuberosity. The larger propodial (MNHNL
469 DOU324A) measures 331mm in proximal distal length and 168mm is maximal distal
470 width. The smaller ‘adult’ propodial (MNHNL DOU558) measures 309mm in
471 proximodistal length and 150mm in distal width.

472

473 Microcleididae Benson et al. 2012
474 *Microcleidus* Watson, 1909
475 cf. *Microcleidus*
476 (Figure 7)

477

478 M NHNL DOU978 is a plesiosaurian cervical centrum from an unknown locality. The
479 centrum is slightly wider than high and is markedly elongated (height=42mm,
480 width=53mm, length=55mm) unlike in derived rhomaleosaurids and thalassophoneans,
481 which have anteroposteriorly short centra (Owen, 1865; Smith & Vincent, 2010;
482 Ketchum & Benson, 2011b; Smith & Benson, 2014). The ventral surface is flattened as in
483 *Microcleidus* (Owen, 1865; Vincent et al., 2017) and bears two subcentral foramina.
484 There is no ventral keel. The lateral surface bears a (very) faint anteroposterior ridge; a
485 more conspicuous lateral ridge is present in *Microcleidus* spp. (Owen, 1865; Bardet,
486 Godefroit & Sciau, 1999; Vincent et al., 2017). The neural arch is fused with the centrum.
487 The orientation of bone fibers on the right side suggests that the suture is V-shaped,
488 which is mainly seen in rhomaleosaurids (Benson, Evans & Druckenmiller, 2012) and in
489 the early pliosaurid *Hauffiosaurus* (Benson et al., 2011). The neural canal is very small in
490 diameter, resulting in zygapophyses that are located close to the centrum, as in
491 *Microcleidus* (Owen, 1865; Bardet, Godefroit & Sciau, 1999) and unlike in *Plesiosaurus*
492 *dolichodeirus* (Owen, 1865), *Westphaliasaurus simonensis* (Schwermann & Sander,
493 2011), *Hauffiosaurus* spp. (White, 1940; Benson et al., 2011). The prezygapophyses are
494 clearly separated, medially-inclined and their ventral surface is located more ventrally
495 than the centre of the neural canal. The rib facets is 8-shaped, being formed of two semi-
496 ovals separated by an anterior and a posterior furrows, as in *Plesiosaurus dolichodeirus*
497 (Owen, 1865; Storrs, 1997) and unlike in many microcleidids (Owen, 1865; Schwermann
498 & Sander, 2011), *Hauffiosaurus* (White, 1940), and cryptoclidids (Andrews, 1910a;
499 Brown, 1981; Knutsen, Druckenmiller & Hurum, 2012a,b). The rib facet in posterior
500 cervical centra is separated by a groove *Microcleidus melusinae* (Vincent et al., 2017).
501 This centrum thus bears strong similarities with *Microcleidus* and is regarded as cf.
502 *Microcleidus*.
503
504

505 Upper Aalenian fauna

506

507 ICHTHYOSAURIA De Blainville, 1835

508 Ichthyosaura indet.

509 (Figure 9)

510
511 M NHNL HU242 (Concavum zone) is a small partial ichthyosaur rostrum from te
512 Rumelange locality. It contains the anterior part of the left dentary (or right premaxilla),
513 showing the labial and lingual walls. Two teeth are preserved in situ. The root and the
514 acellular cementum ring have a circular cross-section and are perfectly smooth externally,
515 unlike in *Ichthyosaurus*, *Protoichthyosaurus*, and *Temnodontosaurus* (McGowan, 1973;
516 Godefroit, 1993a; Vincent et al., 2014; Brusatte et al., 2015; Lomax, Porro & Larkin,
517 2019) and there is no evidence for plicidentine (unlike other Jurassic ichthyosaurians for

518 which this feature is known (Maxwell, Caldwell & Lamoureux, 2012)), but this latter
519 feature cannot be used for a taxonomic purpose at the present state of our knowledge. The
520 teeth are of a maximal diameter of 8mm and estimated 20mm apicobasal length; the size
521 and shape of the tooth fits within the range of early baracromians (Godefroit, 1993b,
522 1994a; Fernández, 1994; Maisch, 2008; Maxwell, Fernández & Schoch, 2012) but in the
523 lack of additional evidence, this material is here referred to Ichthyosauria indet.

524

525 PLESIOSAURIA De Blainville, 1835

526

Plesiosauria indet.

527

(Figure 9)

528

529 MNHNL HU384 (Murchisonae zone) is a small caudal centrum of a plesiosaurian of a
530 juvenile individual, originating from the Tétange locality. Its height/width ratio is
531 32/36mm = 0.88. Two small triangular chevron facets are present ventrally, merging with
532 the edge of the articular surface. The ventral surface is flat. A rounded rib facet is present
533 on the lateral surface. The neural arc is disconnected from the centrum and absent; the
534 facets appear diamond shaped. Sacral vertebrae usually have less diagnostic features than
535 other vertebrae in plesiosaurians, as could be inferred from their respective number of
536 phylogenetic characters (Benson & Druckenmiller, 2014); this, coupled to the immaturity
537 of the specimen would make an assignment ambiguous and we refer this material to
538 Plesiosauria indet.

539

540 PLESIOSAUROIDEA Gray, 1825

541

Plesiosauroidea indet.

542

(Figure 9)

543

544 MNHNL BU157 (Concavum or Murchisonae zone) is a moderately elongated
545 plesiosaurian cervical vertebra from the Rumelange locality. The shape of the centrum
546 suggests it does not belong to derived rhomaleosaurids and thalassophoneans, which
547 have anteroposteriorly short centra (Owen, 1865; Smith & Vincent, 2010; Ketchum &
548 Benson, 2011b; Smith & Benson, 2014). The ventral surface is flattened and bears two
549 subcentral foramina. No ventral keel is present. The edge of the centrum is rugose, as
550 sometimes seen in various plesiosaurian clades (Owen, 1865; Seeley, 1874a; Fischer et
551 al., 2020). An anteroposteriorly elongated bulge (rather than a ridge) is present on the
552 dorsolateral surface. This bulge is separated from the rib facet by a median concave area.
553 The rib facet is roughly oval but is very strongly waisted by an anterior and a posterior
554 notch, giving the facet a marked 8 shape; this structure recalls – but is more marked than
555 in – the late Toarcian centra MNHNL DOU978. The neural arch is fully fused to the
556 centrum and no suture is discernible. A pair of supracentral foramina is present on the
557 floor of foramen magnum. While MNHNL DOU978 could clearly be attributed to

558 *Microcleidus*, this vertebra – although superficially similar – is less elongated, possesses
559 a lateral bulge rather than a ridge, and does not preserve the zygapophyses. As a result, a
560 referral to Microclididae is too ambiguous. Nevertheless, the presence of an 8-shaped rib
561 facet precludes a referral to cryptoclidids (Andrews, 1910a; Brown, 1981; Knutsen,
562 Druckenmiller & Hurum, 2012a,b). Accordingly, this specimen is referred to as a non-
563 cryptoclidid plesiosauroid.

564

565 **Lower Bajocian fauna**

566

567 ICHTHYOSAURIA De Blainville, 1835
568 Ichthyosauria indet.
569 (Figure 10)

570

571 MNHNL BM360 from the Rumelange locality contains one partial neural spine and two
572 proximal parts of bicipital ribs. The neural spine bears a dorsoventrally-oriented median
573 ridge on its anterior surface. The taxonomic values of these features is low and we assign
574 this material to Ichthyosauria indet.

575

576 PARVIPELVIA Motani, 1999
577 Parvipelvia indet.
578 (Figure 10)

579

580 MNHNL BM758 is an ichthyosaurian caudal centrum from the Rumelange locality. Its
581 shape is rounded, with a slightly flattened ventral surface, as in Parvipelvia (Merriam,
582 1908; Maisch & Matzke, 2000). Its height/length ratio is 100mm/38mm=2.63.

583

584 MNHNL BM725 from the Rumelange locality contains a cervical and a caudal
585 ichthyosaurian centra. The cervical centrum has a height/length ratio of
586 83mm/38mm=2.18 while the same ratio for the caudal centrum is 56mm/26mm=2.15.
587 The centra are rounded in shape and exhibit small and rounded apophyses, as in
588 parvipelvians and unlike in more primitive forms (Merriam, 1908; Huene, 1922).

589

590 MNHNL BM461 is an ichthyosaurian cervical centrum from the Rumelange locality. Its
591 height/length ratio is 72mm/30mm=2.4. The centrum is rounded in shape and exhibits
592 small and rounded apophyses, as in parvipelvians and unlike in more primitive forms
593 (Merriam, 1908; Huene, 1922).

594

595 MNHNL BM392 from the Rumelange locality contains five caudal centra and one
596 sclerotic plate of a single ichthyosaurian specimen. The height/length ratios of the first
597 four caudal centra are as follows: 72mm/30mm=2.4, 77mm/31mm=2.48,

598 72mm/29mm=2.48, 75mm/32mm=2.34. Their rounded shape indicates they belong to
599 Parvipelvia (Merriam, 1908; Huene, 1922; Maisch & Matzke, 2000). The sclerotic plate
600 bears radiating striations on its lateral surface and its internal and external edges are
601 crenulated, as is usually the case in ichthyosaurians (Andrews, 1910a; McGowan, 1973;
602 Fischer et al., 2014a). The external third of the plate is deflected, facing dorsally.
603

604 OPHTHALMOSAURIDAE Baur, 1897

605 Ophthalmosauridae indet.

606 (Figures 10, 11)

607

608 M NHNL BM780–BM781 is a nearly complete right ichthyosaurian surangular (M NHNL
609 BM780), associated with seven teeth (bearing the collection number M NHNL BM781),
610 originating from the Rumelange locality. The surangular is straight in all planes (only its
611 posterior quarter is slightly deflected medially) and bears a thickened dorsal margin,
612 giving it a tear-shaped cross-section. A lateral, anteroposteriorly elongated and
613 posteriorly-deepening concavity is present on the lateral surface ('surangular fossa'). This
614 depressed area terminates anteriorly to the level of the coronoid process. The coronoid
615 process is prominent and bears a rugose texture. A small, anteroposteriorly oriented ridge
616 is present directly posteromedially to the coronoid process and is likely part of the muscle
617 attachment. The dorsal margin of the surangular forms a concave plateau posteriorly to
618 the coronoid process. A prominent M.a.m.e process is present directly posteromedially
619 to the plateau; this process points dorsomedially. The angular facet extends anteriorly up
620 to the level of the coronoid process and covers the ventral half of the surangular
621 posteriorly. This indicates the presence of an extensive angular, which is an

622 ophthalmosaurid synapomorphy (Motani, 1999a). The teeth are peculiar in forming
623 externally-visible plicidentine, texturing the acellular cementum ring and the root by very
624 deep apicobasal grooves, as in *Ichthyosaurus*, *Protoichthyosaurus*, and
625 *Temnodontosaurus* (McGowan, 1973; Godefroit, 1993a; Vincent et al., 2014; Brusatte et
626 al., 2015; Lomax, Porro & Larkin, 2019). The root cross-section is oval as in all
627 parvipelvians bar most platypterygiines (Fischer et al., 2012; Fischer, 2016) and the base
628 of the enamel is easy to discern, which is usually the case in ophthalmosaurids (Fischer et
629 al., 2016). Accordingly, we refer this specimen as Ophthalmosauridae indet.

630

631 M NHNL BM779 is a fragmentary partially articulated ichthyosaurian, containing parts of
632 the skull, axial skeleton, and scapular girdle, originating from the Rumelange locality. It
633 belongs to Ophthalmosauridae, having a humerus with plate-like dorsal trochanter, a
634 massive deltopectoral crest, a humeral facet for an anterior accessory element, a
635 posterodistally deflected ulnar facet and a conspicuous acromial process on the scapula
636 (Motani, 1999a; Fischer et al., 2012, 2013; Moon, 2017; Zverkov & Efimov, 2019). It
637 resembles *Arthropterygius* in having an anteroposteriorly-short parietal symphysis, a

638 deep ulnar facet on the humerus, a small humeral facet for an anterior accessory epipodial
639 elements (Maxwell, 2010; Fernández & Maxwell, 2012; Zverkov & Prilepskaya, 2019).
640 However, the presence of several autapomorphies on the cranial and appendicular
641 elements indicates that this specimen constitutes a novel taxon, for which a dedicated
642 manuscript is in preparation.

643

644

645 PLESIOSAURIA De Blainville, 1835

646 CRYPTOCLIDIDAE Williston, 1925

647 Cryptoclididae indet.

648 (Figure 10)

649

650 M NHNL BM782 is a left plesiosaurian propodial from the Rumelange locality. The
651 propodial is straight in the proximodistal direction and slightly curves dorsally. The shaft
652 is fairly elongated. The dorsal tuberosity is wide and weakly demarcated from the
653 capitulum; they form together an evenly rounded proximal surface. A median dorsal boss
654 surrounded by muscle scars is present close to the distal end of the dorsal tuberosity. This
655 boss is median (along the axis of the shaft) and thus differs from the anteriorly placed
656 boss seen in the humeri of rhomaleosaurids and microcleidids (Owen, 1865; Smith &
657 Benson, 2014). The propodial markedly expands distally, forming a preaxial flange and a
658 longer postaxial flange. The preaxial flange forms a dorsoventrally-narrow triangular
659 facet distally. The postaxial flange forms a semioval posterior facet that forms a marked
660 angle (ca. 45°) with the ulnar/fibular. Both the radial/tibial and the anteroposteriorly- and
661 dorsoventrally-shorter ulnar/fibular facets are slightly convex. However, a distal ridge is
662 absent, unlike in *Colymbosaurus* (Benson & Bowdler, 2014; Arkhangelsky et al., 2019).
663 The shape of the propodial with its fairly slender shaft and large preaxial and postaxial
664 flanges indicate cryptocleidid affinities (Andrews, 1910a; Mehl, 1912; Brown, 1981;
665 O'Keefe & Wahl, 2003; Knutsen, Druckenmiller & Hurum, 2012c; Roberts et al., 2017,
666 2020); it resembles the humerus of humerus of *Murænosaurus leedsi* (Seeley, 1874b;
667 Andrews, 1910a) and the femur of *Tricleidus seeleyi* (Owen, 1865).

668

669

670 Results

671

672 We visualized the size and shape of Toarcian and Bajocian ichthyosaur centra (Figure
673 12). These data reveal a decrease in overall size and a much smaller range of shapes from
674 the late Toarcian to the Bajocian. However, a difference in sampling intensity between
675 these two assemblages is evident. We randomly sampled our late Toarcian dataset to
676 select 3 cervicals and 6 caudals 10000 times, yielding comparable datasets. We then
677 compared the range of height-length ratios between these datasets. The Bajocian height-

678 length ratios range is smaller than Toarcian height-length ratios range about 99.85% of
679 the times and two times smaller about 98.56% of the times. The data at hand thus
680 unambiguously suggest a reduction of the disparity of ichthyosaurian centrum shapes
681 across the Early–Middle Jurassic transition, i.e. going from an early neichthyosaurian-
682 dominated to an ophthalmosaurid-dominated assemblage.

683

684

685 Discussion

686

687 The Early–Middle Jurassic transition in marine reptiles, a state of the art

688 The very well-sampled (Benson et al., 2010) Lower Jurassic marine ecosystems of
689 western Europe housed a vast menagerie of neichthyosaurians, thalattosuchians, and
690 plesiosaurians (Owen, 1860; Huene, 1922, 1931; McGowan, 1974, 1979; Benton &
691 Taylor, 1984; Maisch & Matzke, 2000; O’Keefe, 2004; Großmann, 2007; Maisch, 2008,
692 2010; Benson, Evans & Druckenmiller, 2012; Martin et al., 2012; Bardet et al., 2014;
693 Lomax & Massare, 2016; Johnson, Young & Brusatte, 2020). Evidence from body size,
694 craniodental shape, and swimming capabilities suggests that these taxa occupied several
695 niches within shallow marine ecosystems (Hauff, 1953; Massare, 1987, 1997; Böttcher,
696 1989; Godefroit, 1994a; McGowan, 1996; Buchholz, 2001; Buchy, 2010; Fischer,
697 Guiomar & Godefroit, 2011; Martin et al., 2012; Dick, Schweigert & Maxwell, 2016;
698 Maxwell & Cortés, 2020). Such a diversity of forms made the Early Jurassic marine
699 ecosystems of western Europe an iconic representation of Mesozoic marine life (Taylor,
700 1997).

701

702 Even though a drastic reduction of apparent diversity is expected following the
703 lagerstätten effect of the early Toarcian localities (Benson et al., 2010), the Middle
704 Jurassic assemblages, when well sampled (i.e. not before the Callovian), are markedly
705 distinct from their Toarcian counterparts. For ichthyosaurians, the dominant and
706 diversified early neichthyosaurians (leptonectids, *Suevoleviathan*, *Hauffiopteryx*, and
707 *Temnodontosaurus*) are gone; the disappearance of non-thunnosaurian ichthyosaurians
708 marks the end of large (>6m) ichthyosaurian top predators, at least up until the Aptian,
709 when some derived platypterygiines will presumably fill similar roles (Fischer et al.,
710 2014b, 2016; Bardet, Fischer & Machalski, 2016; Fischer, 2016). Ichthyosaurians will
711 never re-evolve carinated teeth like those seen in *Temnodontosaurus platyodon* and
712 *Temnodontosaurus trigonodon* (Conybeare, 1822; Godefroit, 1993a), nor the sawfish-like
713 morphology of Leptonectidae (Swinton, 1930; Huene, 1951; McGowan, 1986, 2003;
714 Lomax, 2016). From the Callovian onwards, ichthyosaur assemblages overwhelmingly
715 consist of ophthalmosaurids (but see Fischer et al. (2013)), up to the final extinction of
716 ichthyosaurians (Fischer et al., 2016). While diverse ophthalmosaurid assemblages are
717 known by the Kimmeridgian (Fernández, 1997; Druckenmiller et al., 2012; Arkhangelsky

718 & Zverkov, 2014; Zverkov et al., 2015; Paparella et al., 2017; Moon & Kirton, 2018a;
719 Delsett et al., 2019; Zverkov & Efimov, 2019; Campos, Fernández & Herrera, 2020;
720 Zverkov & Jacobs, 2020), only the closely related *Ophthalmosaurus* and *Baptanodon* are
721 known in the Callovian of Europe and North America, respectively (Seeley, 1874c;
722 Marsh, 1895; Gilmore, 1902, 1906; Knight, 1903; Andrews, 1910b; Massare et al., 2006;
723 Moon & Kirton, 2018b).

724
725 This leaves a fairly long interval (late Toarcian-Callovian) for which the ichthyosaur
726 fossil record is very scarce and geographically dispersed: Sander & Bucher (1993)
727 reported a large cf. *Stenopterygius* from the late Toarcian of southern France; Maisch &
728 Matzke (2000: p72) mention the presence of *Temnodontosaurus* in the late Toarcian of
729 southern Germany; Vincent et al. (2013) reported *Temnodontosaurus* sp., and
730 *Stenopterygius*-like forms in the late Toarcian of southern France and a coracoid bearing
731 many similarities with MNHN DOU353 from the early Aalenian of southern France;
732 finally, Brusatte et al. (2015) described *Dearcmahara shawcrossi* from the late Toarcian-
733 Bajocian of Scotland. The only diagnosable ichthyosaur remains unambiguously known
734 from Aalenian deposits is the holotype of *Stenopterygius aalensis* from Germany
735 (Maxwell, Fernández & Schoch, 2012) (see also Arnaud et al. (1976) and Maxwell et al.
736 (2012) for additional indeterminate specimens). Fragmentary ophthalmosaurid specimens
737 have been reported in the Aalenian-Bajocian boundary of Argentina (Fernández, 2003)
738 and the early Bajocian of Argentina (Fernández, 1999; Gasparini et al., 2007) and Canada
739 (Druckenmiller & Maxwell, 2014). It can be inferred from these data that a replacement
740 of early neioichthyosaurs by ophthalmosaurids took place across the Early–Middle
741 Jurassic transition, but the generalized absence of Aalenian-Bathonian ichthyosaur fossils
742 worldwide and Early Jurassic fossils in South America make it difficult to identify a
743 precise turnover.

744
745 While less abundant than ichthyosaurians in early Toarcian Lagerstätten (Hauff, 1953),
746 plesiosaurians already had also evolved a vast array of morphotypes by the Toarcian
747 (O’Keefe, 2001a,b; Benson, Evans & Druckenmiller, 2012; Smith & Araújo, 2017),
748 presumably filling as many ecological niches: long-necked and small headed microclidids
749 (Owen, 1865; Bardet, Godefroit & Sciau, 1999), gigantic, apex predatory
750 rhomaelosaurids (Taylor, 1992; Cruickshank, 1994; Smith & Dyke, 2008; Smith &
751 Vincent, 2010; Smith & Benson, 2014), and small-sized, moderately long-necked, and
752 long-snout early pliosaurids (Benson et al., 2011; Ketchum & Benson, 2011a; Vincent,
753 2011; Fischer et al., 2017). Microcleidids supposedly go extinct after the Toarcian; their
754 last definite record in the latest Toarcian of France (Sciau, Crochet & Mattei, 1990;
755 Bardet, Godefroit & Sciau, 1999). However, an Aalenian specimen from France regarded
756 as an indeterminate elasmosaurid by Vincent et al. (2007) possesses microcleidid features
757 (8-shaped cervical rib facet, lateral ridge on cervical centra) and might possibly be

758 regarded as a member of that clade. Nevertheless, the long-necked morphotype is then
759 colonised by cryptoclydids by the Callovian-Oxfordian (Brown, 1981).

760

761 Only a couple of rhomaleosaurid lineages will survive up to the Middle/Late Jurassic, in
762 high latitudes (Gasparini, 1997; Sato & Wu, 2008; Benson, Zverkov & Arkhangelsky,
763 2015). Thalassophonean pliosaurids concomitantly radiated and became the main marine
764 apex predators from the Middle Jurassic to the early Late Cretaceous (Noe, 1999, 2001;
765 Ketchum & Benson, 2011a; Benson et al., 2013; Benson & Druckenmiller, 2014; Fischer
766 et al., 2017; Zverkov et al., 2018). A narrative where thalassophoneans filled (either
767 competitively or passively) the niches previously occupied by *Temnodontosaurus* and
768 derived rhomaleosaurids is sensible, but these animals have clearly distinct body plans
769 and probably hunted differently (Massare, 1988; Buchholtz, 2001; O'Keefe, 2001b,
770 2002). In any case, there is evidence for the continuous presence of gigantic apex
771 predatory plesiosaurians in Europe and elsewhere: *Simolestes keilenei* from the upper
772 Bajocian of eastern France (described by Godefroit (1994b), and currently under revision
773 by S. Sachs), an indeterminate fragmentary jaw from the lower Bajocian of Switzerland
774 (Sachs, Klug & Kear, 2019) and 'pliosaroid' tooth from the Aalenian-Bajocian of
775 Australia (Long & Cruickshank, 1998; Kear, 2012). Whether these specimens are
776 rhomaleosaurids or thalassophoneans obviously yields different outcomes as they might
777 push the origin of apex predatory pliosaurids to the Bajocian, but these fossils suggest
778 that the niche of very large apex predators was not vacated for a long period, if at all.
779

780 Recent evidence revealed a series of basal metriorhynchoid crocodyliforms in the late
781 Toarcian–Aalenian interval (Wilberg, 2015; Ösi et al., 2018; Aiglstorfer, Havlik &
782 Herrera, 2020) suggesting an intense diversification across the Early–Middle Jurassic
783 transition (Aiglstorfer, Havlik & Herrera, 2020). The effect of this radiation is evident by
784 the late Middle and Late Jurassic, where abundant thalattosuchian taxa are known,
785 occupying several ecological niches (Gasparini, Pol & Spalletti, 2006; Young & de
786 Andrade, 2009; Young et al., 2012, 2014b; Herrera, Gasparini & Fernández, 2013; Foffa
787 et al., 2017, 2018). Teleosauroids do not appear much affected by this transition, but this
788 might be due to the problematic taxonomy of the group (Johnson, Young & Brusatte,
789 2020).

790

791 **An abrupt turnover**

792

793 In light of the short review above, it appears clear that fossiliferous successions in a
794 geographically-restricted areas are crucial to decipher both the tempo and the severity of
795 the marine reptile turnover at the early–Middle Jurassic transition. The late Toarcian–
796 Bajocian successions of Luxembourg and Belgium we analysed above yields novel
797 biostratigraphic information that precise the turnover dynamics of ichthyosaurians and

798 plesiosaurians during the Early–Middle Jurassic transition (the crocodyliform record,
799 recently reviewed by Johnson et al. (2018), is scarce and does not provide relevant
800 information regarding this turnover). In a global, broad-brush view of marine reptile
801 macroevolution, the presence of rhomaleosaurids, microcleidids, and non-thunnosaurian
802 ichthyosaurians in the late Toarcian of Luxembourg (Table 2) was expected, as these
803 clades are already known to survive after the early Toarcian Lagerstätten (Bardet,
804 Godefroit & Sciau, 1999; Vincent et al., 2013; Benson, Zverkov & Arkhangelsky, 2015).
805 With the exception of possible remains from the Aalenian of France (Vincent, Bardet &
806 Morel, 2007; see discussion above), MNHN DOU978 and the holotype of *Microcleidus*
807 *tournemirensis* (Sciau, Crochet & Mattei, 1990; Bardet, Godefroit & Sciau, 1999) are the
808 youngest-known microcleidid to date and indicate that the clade extended at least up to
809 the very end of the Early Jurassic. Recording microcleidids, large rhomaleosaurids,
810 *Temnodontosaurus*-like forms, and more derived ichthyosaurians within the same basin
811 that formed the early Toarcian Lagerstätten yields important palaeobiogeographic
812 implications. These fossils indicate that the main marine reptile clades of the Early
813 Jurassic remained abundant and dominant in mid-latitude, epicontinental seas up to the
814 end of the Toarcian at least. Although much less prolific than the underlying strata, the
815 upper Toarcian fossil record suggest that no major turnover took place within the
816 Toarcian as the structuration of the upper tier of marine ecosystems remained intact. The
817 taxonomic diversity crash that follows the early Toarcian appears to be mainly a
818 Lagerstätten effect, confirming previous suspicions (Benson et al., 2010; Maxwell &
819 Vincent, 2015).

820
821 In stark contrast with the sense of continuity displayed by the upper Toarcian
822 occurrences, the Bajocian fossil record of Luxembourg suddenly records the dominant
823 clades of the Late Jurassic: cryptoclidids, ophthalmosaurids, and probable pliosaurids,
824 with no evidence – so far – for the presence of more ancient clades. The specimen
825 MNHN BM782 is oldest-known cryptoclidid and MNHN BM770 and MNHN
826 BM780–BM781 are the oldest ophthalmosaurids known, after a single partial forefin
827 from the Aalenian-Bajocian boundary of Argentina (Fernández, 2003) and fragmentary
828 basicranium from an equivalent of the Sauzei Zone in Canada (Druckenmiller &
829 Maxwell, 2014), which directly underlies the Humphriesianum Zone. The new
830 occurrences we report here indicate that ophthalmosaurids rapidly dispersed, being
831 almost simultaneously recorded in Canada, Argentina, and Luxembourg by the Early
832 Bajocian.

833
834 The Belgo-Luxembourgian marine reptile record provides a clearer picture of the marine
835 reptile turnover occurring at the Early–Middle Jurassic transition, indicating that this
836 replacement is restricted to the sole Aalenian stage instead of possibly spanning the entire
837 middle Toarcian–Bathonian interval. This reduces the uncertainty on the timing and

duration of this turnover to 4 instead of 14 million years and packs a series of extinctions and diversifications within a short period of abrupt climate cooling and changes in oceanic currents (Korte et al., 2015). This Aalenian shift from a Toarcian Warm Mode to Aalenian-Bajocian Cool Mode (Korte et al., 2015) also appears associated with a marked faunal disruption in belemnites, with random (i.e. non-morphologically-selective) extinctions leading to a distinct drop in belemnite biodiversity at least in the northwestern Peri-Tethys Ocean (Dera, Toumoulin & de Baets, 2016; Neige, Weis & Fara, 2020). This major disruption in the evolutionary history of Jurassic belemnites ended at the Aalenian-Bajocian boundary and resulted in a radiation of the suborder Belemnopseina that partially replaced the previously dominant Belemnitina in the Western Tethys (Weis, Mariotti & Riegraf, 2012; Weis, Sadki & Mariotti, 2017), furthermore entailing a distinct Boreal vs. Tethyan belemnite provincialism (Doyle, 1987; Mariotti, Santantonio & Weis, 2007; Weis & Mariotti, 2007; Mariotti et al., 2012; Dzyuba et al., 2019). At the present state of knowledge, we can only speculate about whether faunal changes in belemnites and marine reptiles were independently impacted by the same factors, or even causally connected, since belemnites are an essential component of the food spectrum in some marine reptiles (Massare, 1987; Böttcher, 1989; Dick, Schweigert & Maxwell, 2016), their faunal change could have triggered a disruption of the trophic chain up to the giant top predators.

857

858

859

860 Conclusions

861 A generalised turnover affected marine reptile communities across the Early–Middle
862 Jurassic transition. However, the extremely poor fossil record of the middle Toarcian–
863 Bathonian interval leaves a ca. 14 million years window of the uncertainty for this
864 important event. Our thorough analysis of the marine reptile record of the late Toarcian to
865 Bajocian successions in a confined palaeogeographic setting (the Belgo-Luxembourgian
866 sub-basin) indicates that:

867

- 868 (i) the structuration of the upper tier of marine ecosystems remained unchanged
869 up to the very end of the Early Jurassic, with the presence of large
870 rhomaleosaurid plesiosaurians, microcroclidid plesiosaurians, as well as
871 *Temnodontosaurus*-like and baracromian ichthyosaurians in the late Toarcian.
- 872 (ii) the dominant clades of the Late Jurassic, cryptoclidid plesiosaurians and
873 ophthalmosaurid ichthyosaurians, arose and dispersed earlier than expected,
874 being recorded in the early Bajocian of Luxembourg.
- 875 (iii) the within-System marine reptile turnover of the Early–Middle Jurassic
876 transition is more abrupt than previously supposed, being restricted to the sole
877 Aalenian stage, i.e. 4 million years.

878

879 **Acknowledgements**

880

881 We thank the curatorial staff of the IRSNB for their help. We also thank the main
882 contributors of the Paleobiology Database data we used: Prof. R.B.J. Benson, Dr. M.T.
883 Carrano, Dr. J.P. Tennant, Dr. H.P. Street, Prof. R.J. Butler, Prof. M. Clapham and Prof.
884 M. Uhen. Finally, we warmly thank Dr. M. Trotta for her help in generating geographic
885 maps. Reviewers+Editor.

886

887

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- 1505
- 1506

1507

1508 **Figure Captions**

1509

1510 **Figure 1. Number of collections of marine reptiles (Ichthyosauria, Plesiosauria,
1511 Thalattosuchia, Pleurosauria, Angolachelonia) per stage.** Data extracted from the
1512 paleobiology database on the 26th March 2020 (see Acknowledgements for the main
1513 contributors of these data). Silhouettes originate from phylopic: *Dakosaurus* by Dmitry
1514 Bogdanov & T Michael Keesey; *Meyerasaurus*, *Eurhinosaurus*, *Temnodontosaurus*,
1515 *Plesiopterys*, and *Ophthalmosaurus* by Gareth Monger; *Rhomaleosaurus* and
1516 *Stenopterygius* by Scott Hartmann; *Peloneustes* by Nobu Tamura & T Michael Keesey;
1517 *Albertonectes* by Frank Denota.

1518

1519 **Figure 2. Map of the fossiliferous localities investigated.** Produced using data from
1520 Esri, HERE, Garmin, and OpenStreetMap.

1521

1522 **Figure 3. Thoursense Zone fauna, late Toarcian, Luxembourg.** Selected anatomy of
1523 specimen M NHNL TM212. (A–B) left coracoid in ventral (A) and anteroventral (B)
1524 views. (C, E, F) left humerus in anterodorsal (C), distal (E), and proximal (F) views. (D)
1525 left forefin in ventral view.

1526

1527 **Figure 4. Pseudoradiosa Zone fauna, late Toarcian, Luxembourg.** (A–E)
1528 rhomaleosaurid left humerus M NHNL DOU307 in ventral (A), anterior (B), dorsal (C),
1529 proximal (D), and distal (E) views. (F–H) rhomaleosaurid humerus M NHNL KA109 in
1530 dorsal/ventral (F), proximal (H), and distal (H) views.

1531

1532 **Figure 5. Ichthyosaurians of the late Toarcian of Belgium and Luxembourg.** (A–B)
1533 large, *Temnodontosaurus*-like caudal centrum IRSNB Vert-00000-00803 in anterior (A)
1534 and lateral (B) views. (C–D) probable thunnosaurian caudal centrum IRSNB Vert-00000-
1535 00808 in anterior (C) and dorsal (D) views. (E–F) ichthyosaurian distal tooth IRSNB
1536 Vert-06462-0003 in labial (E) and mesial (F) views. (G–H) probable non-
1537 ophthalmosaurid left angular IRSNB Vert-06462-0002 in lateral (G) and medial (H)
1538 views. (I–K) ichthyosaurian right quadrate IRSNB Vert-00000-00802 in lateral (I),
1539 medial (J), and posterior (K) views.

1540

1541 **Figure 6. Ichthyosaurians of the Aalensis Zone, late Toarcian, Luxembourg.** (A–C)
1542 left coracoid of a non-baracromian parvipelvian M NHNL DOU353 in ventral (A), lateral
1543 (B), and posterior (C) views. (D–E) large parvipelvian caudal centrum M NHNL DOU378
1544 in dorsal (D), anterior (E), and cross-sectional (F) views. (G) small parvipelvian caudal
1545 centrum M NHNL DOU998 in anteroventral view. (H–I) parvipelvian dorsal centrum

1546 M NHNL DOU944 in anterior (H), and lateral (I) views. (J) parvipelvian dorsal centrum
1547 M NHNL DOU352 in anterior view.

1548

1549 **Figure 7. Plesiosaurian centra of the Aalensis Zone, late Toarcian, Luxembourg.** (A–
1550 B) rhomaleosaurid pectoral vertebra M NHNL DOU954 in anterior (A) and ventral (B)
1551 views. (C–G) cf. *Microcleidus* cervical vertebra in anterior (C), ventral (D), dorsal (E),
1552 lateral (F), and oblique (G) views. (H–J) rhomaleosaurid pectoral vertebra M NHNL
1553 DOU723 in anterior (H), lateral (I), and posterior (J) views. (K–N) plesiosaurian sacral
1554 vertebra M NHNL DOU724 in dorsal (K), anterior (L), posterolateral (M), and ventral (N)
1555 views. (O–Q) rhomaleosaurid pectoral vertebra M NHNL DOU722 in anterior (H), lateral
1556 (I), and posterior (J) views.

1557

1558

1559 **Figure 8. Plesiosaurians of the Aalensis Zone, late Toarcian, Luxembourg.** (A)
1560 plesiosaurian tooth crown M NHNL DOU906 in ?mesial view. (B) rhomaleosaurid right
1561 humerus M NHNL DOU558 in dorsal view. (C–F) rhomaleosaurid right humerus
1562 M NHNL DOU324a in proximal (C), dorsal (D), posterior (E), and distal (F) views. (G)
1563 juvenile rhomaleosaurid propodial M NHNL DOU324b.

1564

1565 **Figure 9. Late Aalenian fauna, Luxembourg.** (A–D) non-cryptoclidid plesiosaroid
1566 cervical centrum M NHNL BU157 in anterior (A), dorsal (B), ventral (C), and oblique (D)
1567 views. (E–G) plesiosaur juvenile caudal centrum in lateral (E), anterior (F), and
1568 anteroventral (G) views. (H–I) fragmentary ichthyosaurian rostrum.

1569

1570 **Figure 10. Humphresianum Zone fauna, early Bajocian, Luxembourg.** (A–F)
1571 cryptoclidid propodial M NHNL BM782 in anterior (A), proximal (B), dorsal (C), ventral
1572 (D), posterior (E), and distal (F) views. (G–H) parvipelvian M NHNL BM392 centra (G)
1573 and sclerotic element in lateral view (H). (I–O) ophthalmosaurid surangular and teeth
1574 M NHNL BM780_781: anterior tooth in labial (I) and basal (J) views; mid-rostrum tooth
1575 in labial (K) and basal (L) views; posterior tooth in labial view (M); right angular in
1576 lateral (N) and medial (O) views.

1577

1578 **Figure 11. The new Humphresianum Zone ophthalmosaurid, early Bajocian,**
1579 **Luxembourg.** Selected anatomy of specimen M NHNL BM779. (A–C) right exoccipital
1580 in anterolateral (A), posteromedial (B), and posterior (C) views. (D–E) suraoccipital in
1581 posterior (D) and ventral (E) views. (F–H) left scapula in anterior (F), medial (G), and
1582 lateral (H) views. (I–K) right quadrate in lateral (I), anterior (J), and condylar (K) views.
1583 (L) right parietal in dorsal view. (M–N) right humerus in posterior (M) and dorsal (N)
1584 views. (O–P) right ulna in posterior (O) and dorsal (P) views.

1585

1586 **Figure 12. Ichthyosaurian centra proportions.** (A) Height over length; the grey lines
1587 indicate a ratio of 2 and 3. (B) Height-length ratio, emphasizing the reduction of shape
1588 disparity from the late Toarcian to the Bajocian. The silhouette originate from phylopic:
1589 *Stenopterygius* by Scott Hartmann.

Figure 1

Number of collections of marine reptiles (Ichthyosauria, Plesiosauria, Thalattosuchia, Pleurosauria, Angolachelonia) per stage.

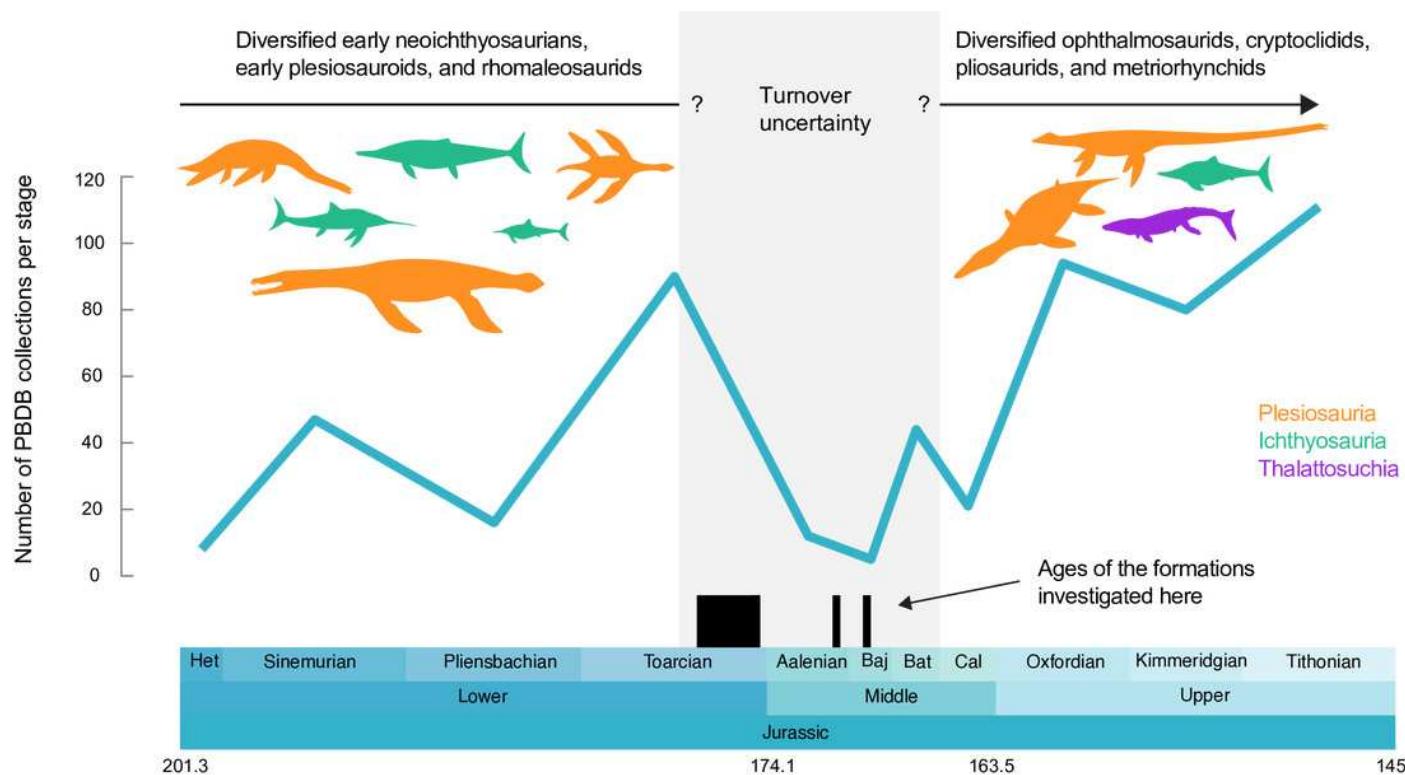


Figure 2

Map of the fossiliferous localities investigated

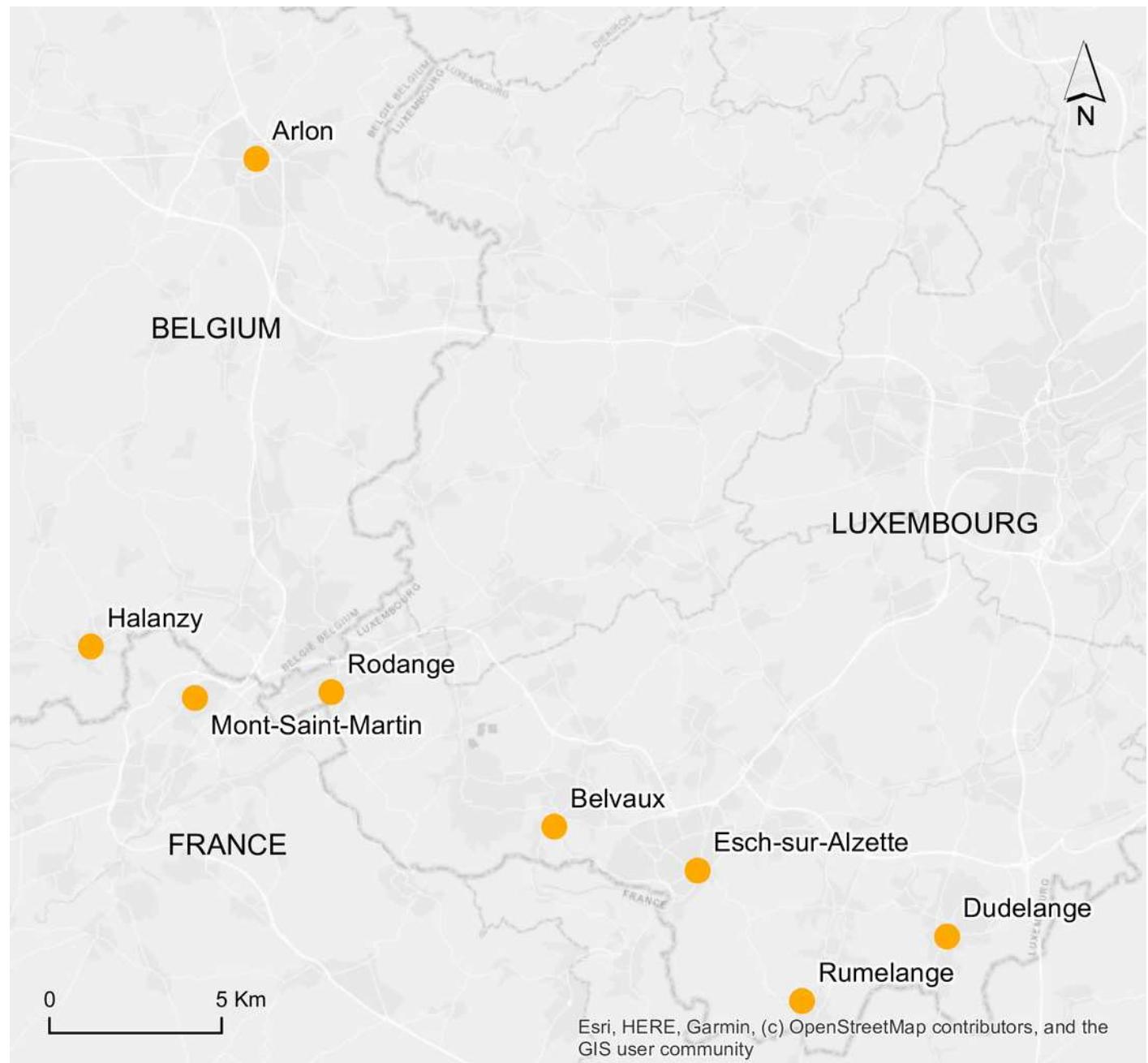
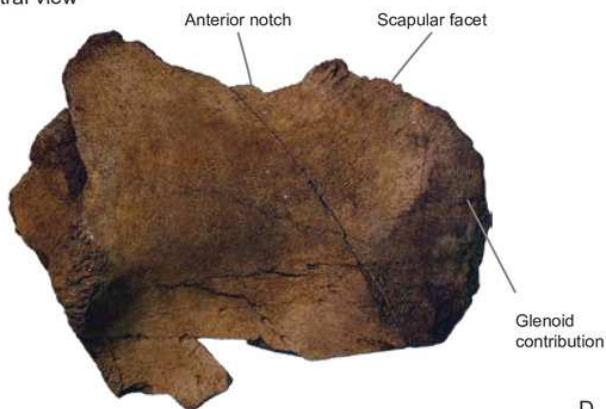


Figure 3

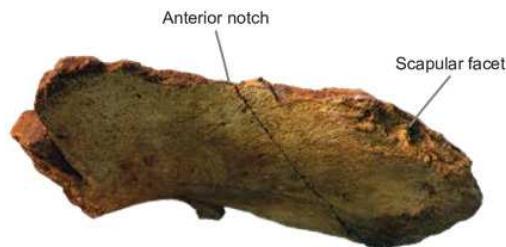
Thoursense Zone fauna, late Toarcian, Luxembourg.

Thouarsense Zone, late Toarcian

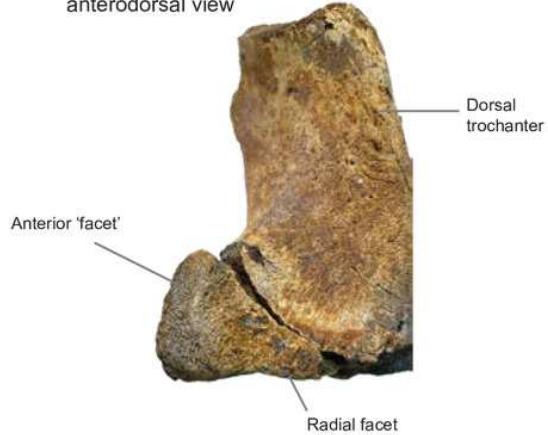
A Left coracoid, ventral view



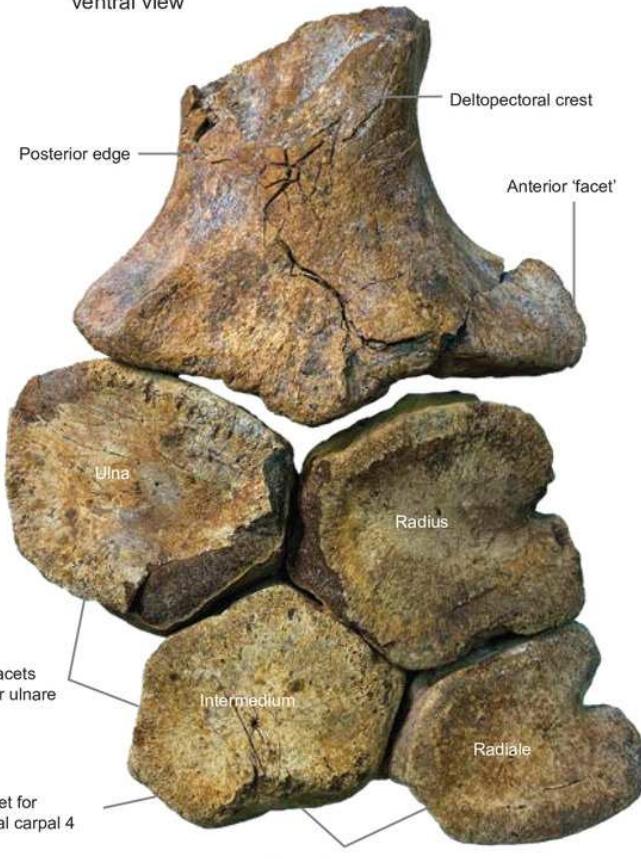
B Left coracoid, anteroventral view



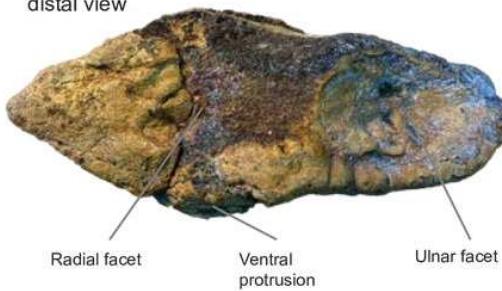
C Left humerus, anterodorsal view



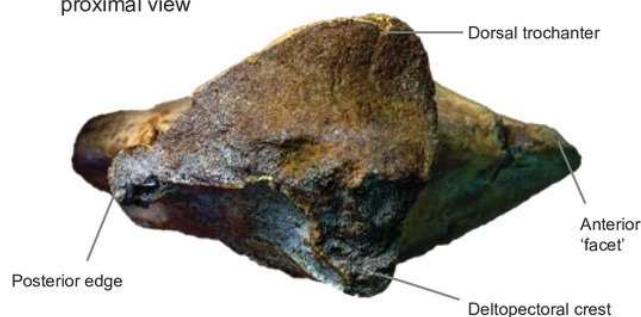
D Left forefin, ventral view



E Left humerus, distal view



F Left humerus, proximal view



50 mm

Figure 4

Pseudoradiosa Zone fauna, late Toarcian, Luxembourg.

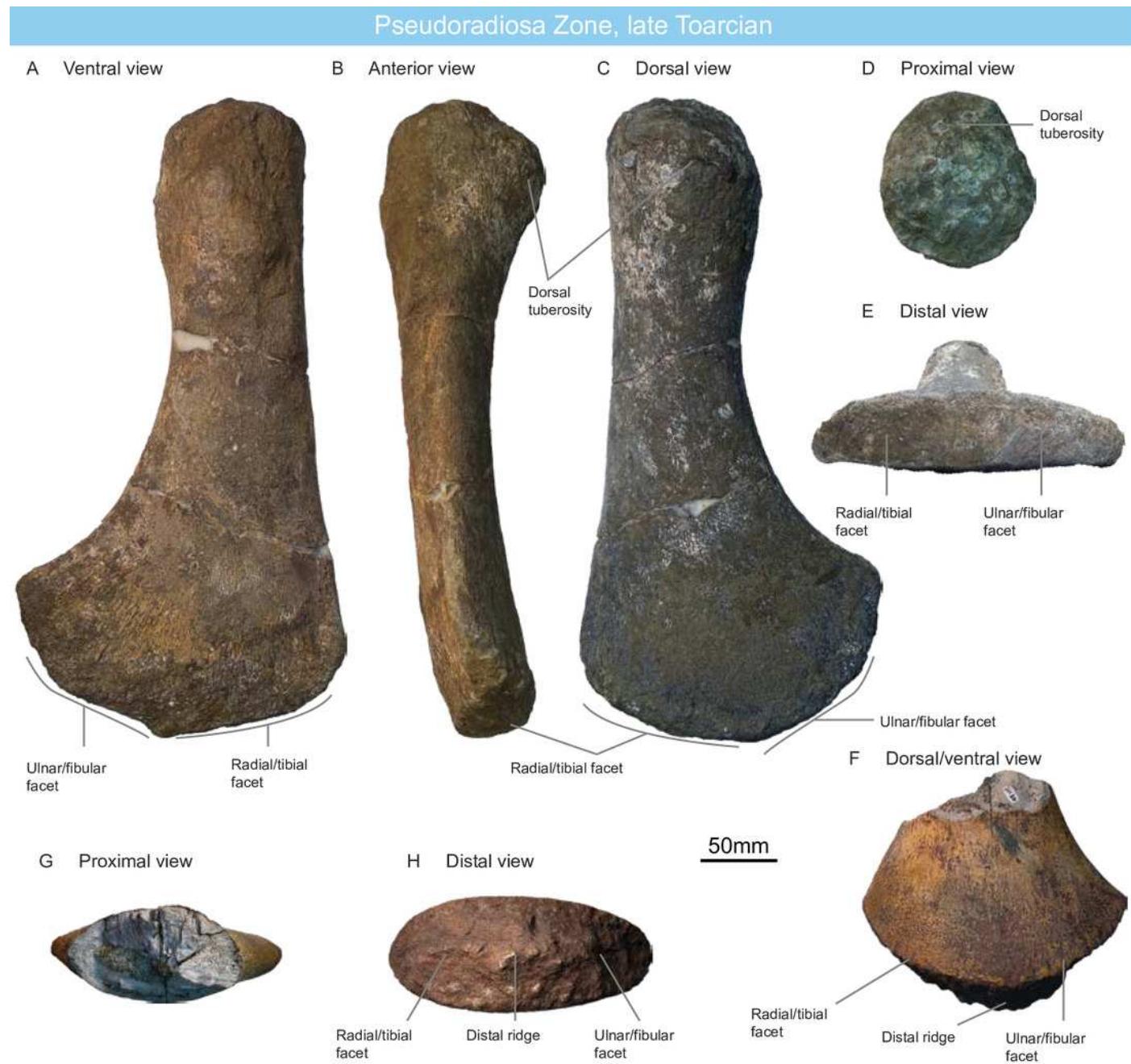


Figure 5

Ichthyosaurians of the Aalensis Zone, late Toarcian of Luxembourg and Belgium.

Dispansum-Pseudoradiosa-Aalensis Zone, late Toarcian

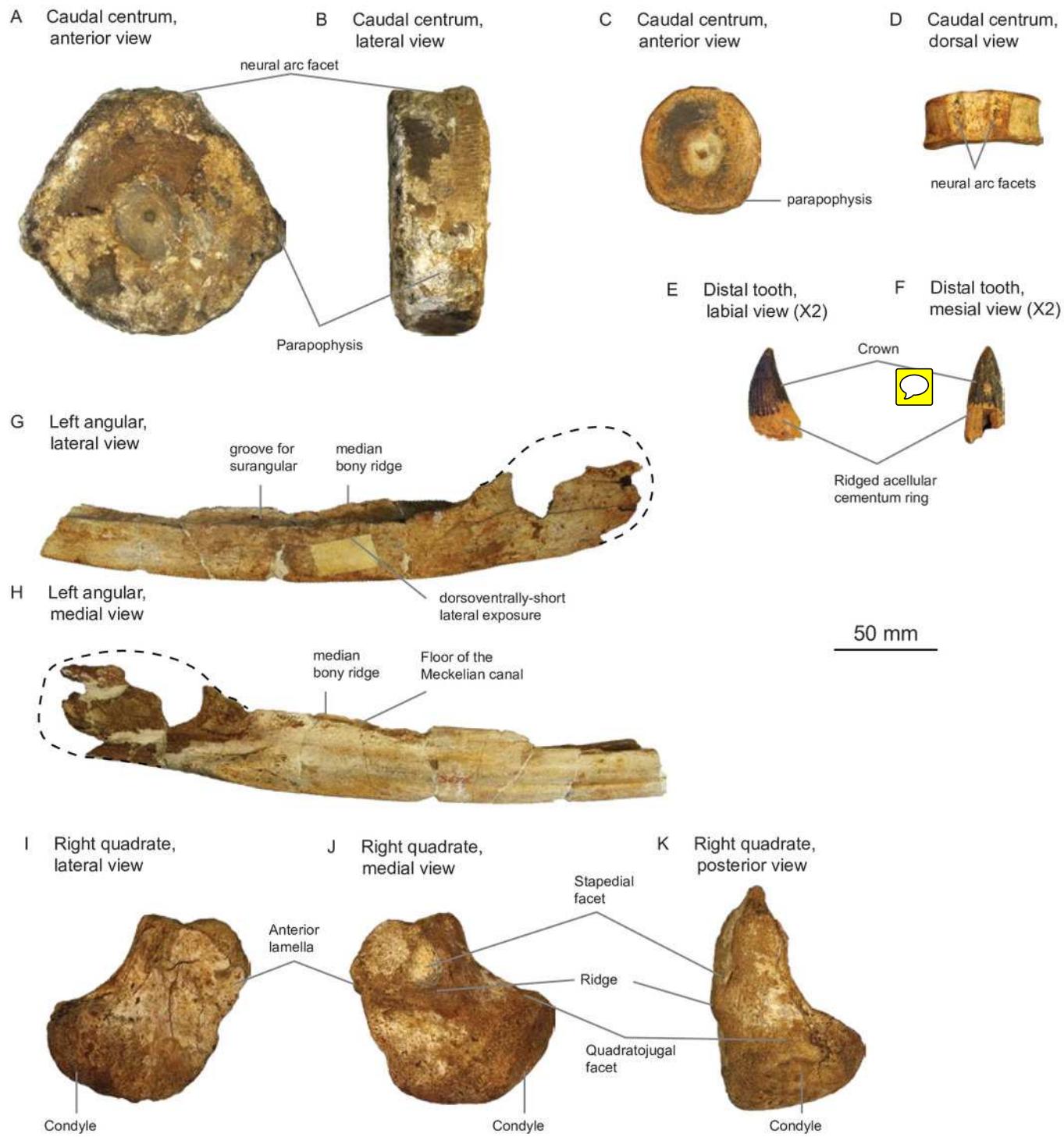


Figure 6

Ichthyosaurians of the Aalensis Zone, late Toarcian, Luxembourg.

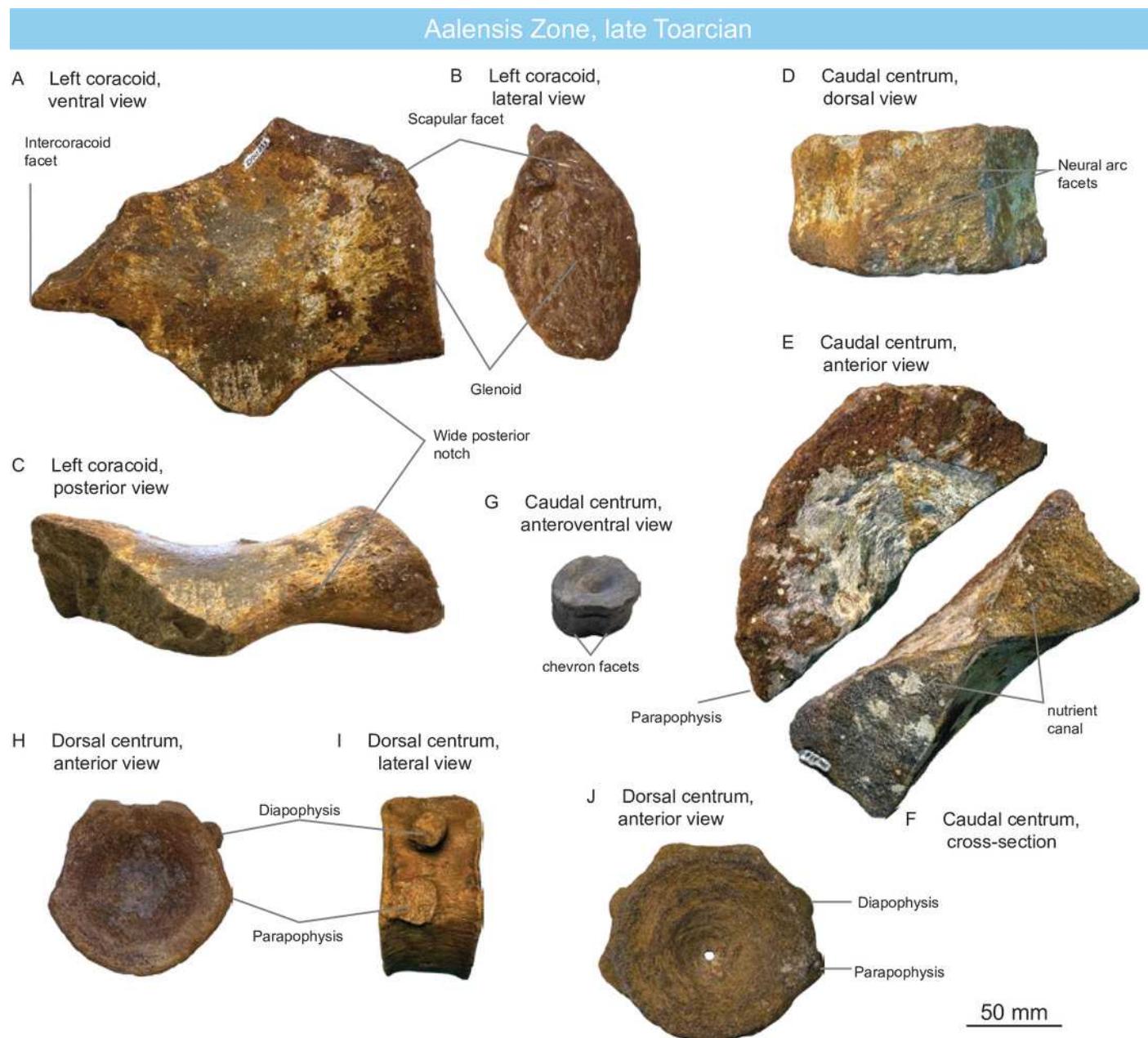


Figure 7

Plesiosaurian centra of the Aalensis Zone, late Toarcian, Luxembourg.

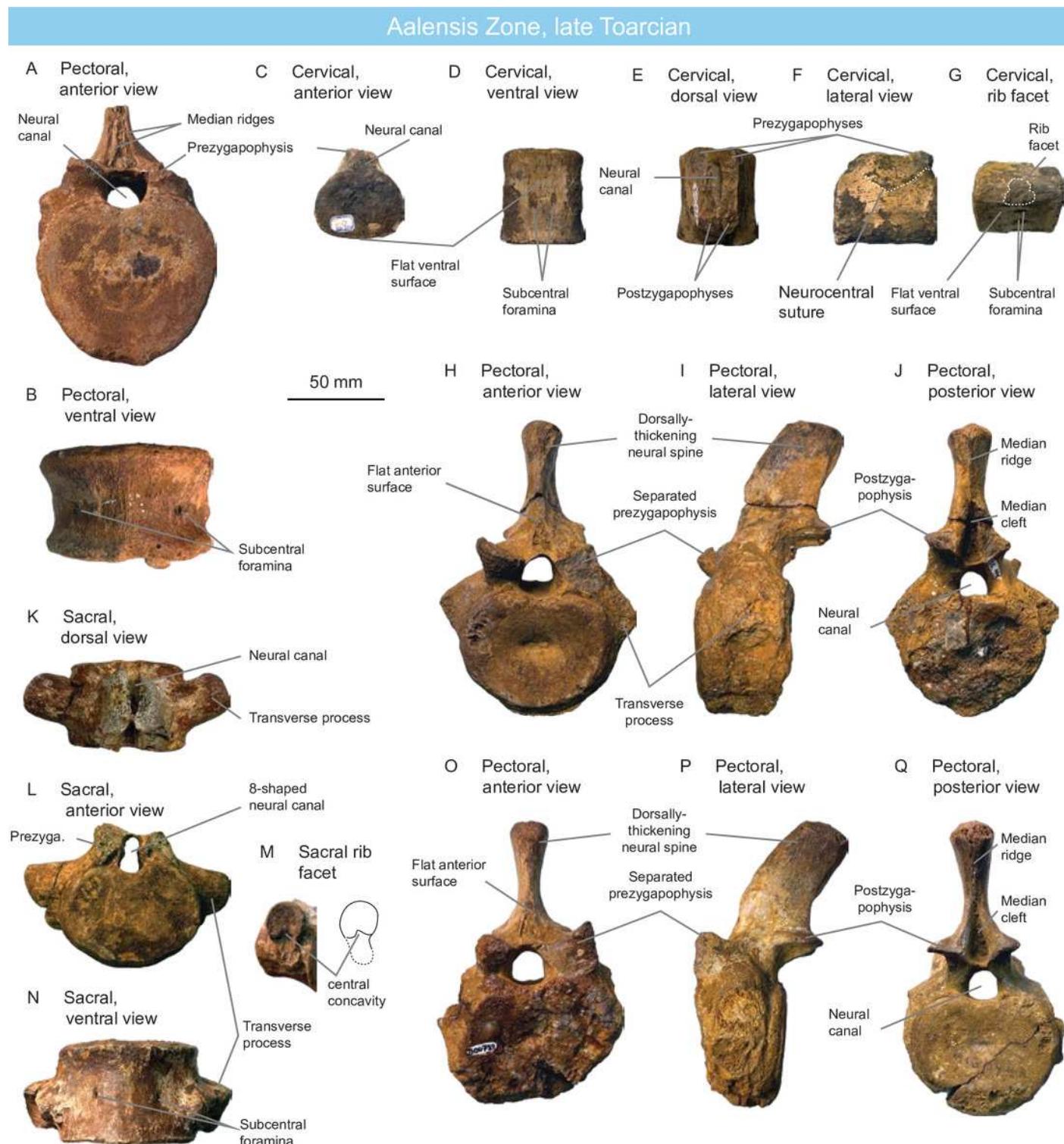


Figure 8

Plesiosaurians of the Aalensis Zone, late Toarcian, Luxembourg.

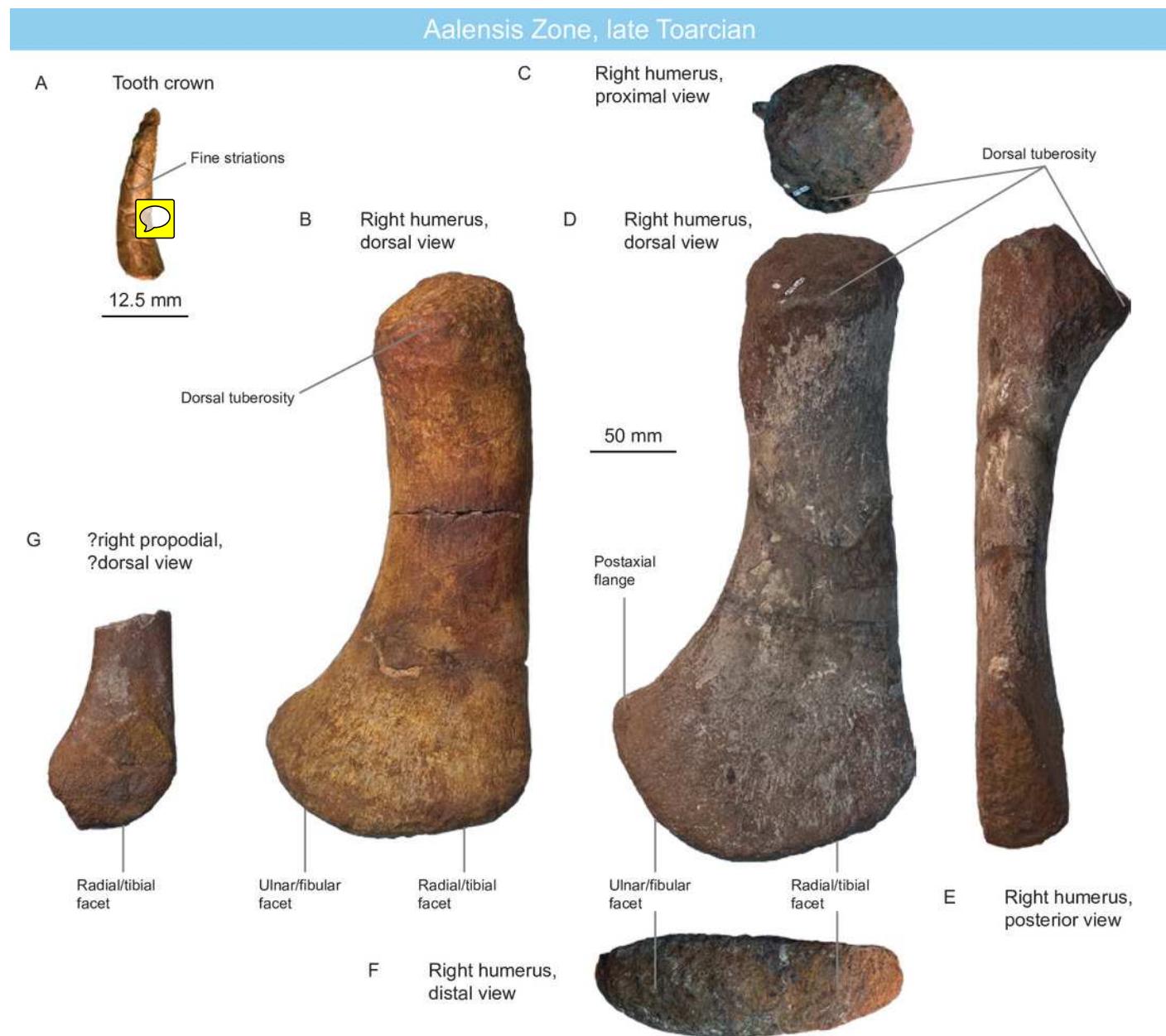


Figure 9

Late Aalenian fauna, Luxembourg.

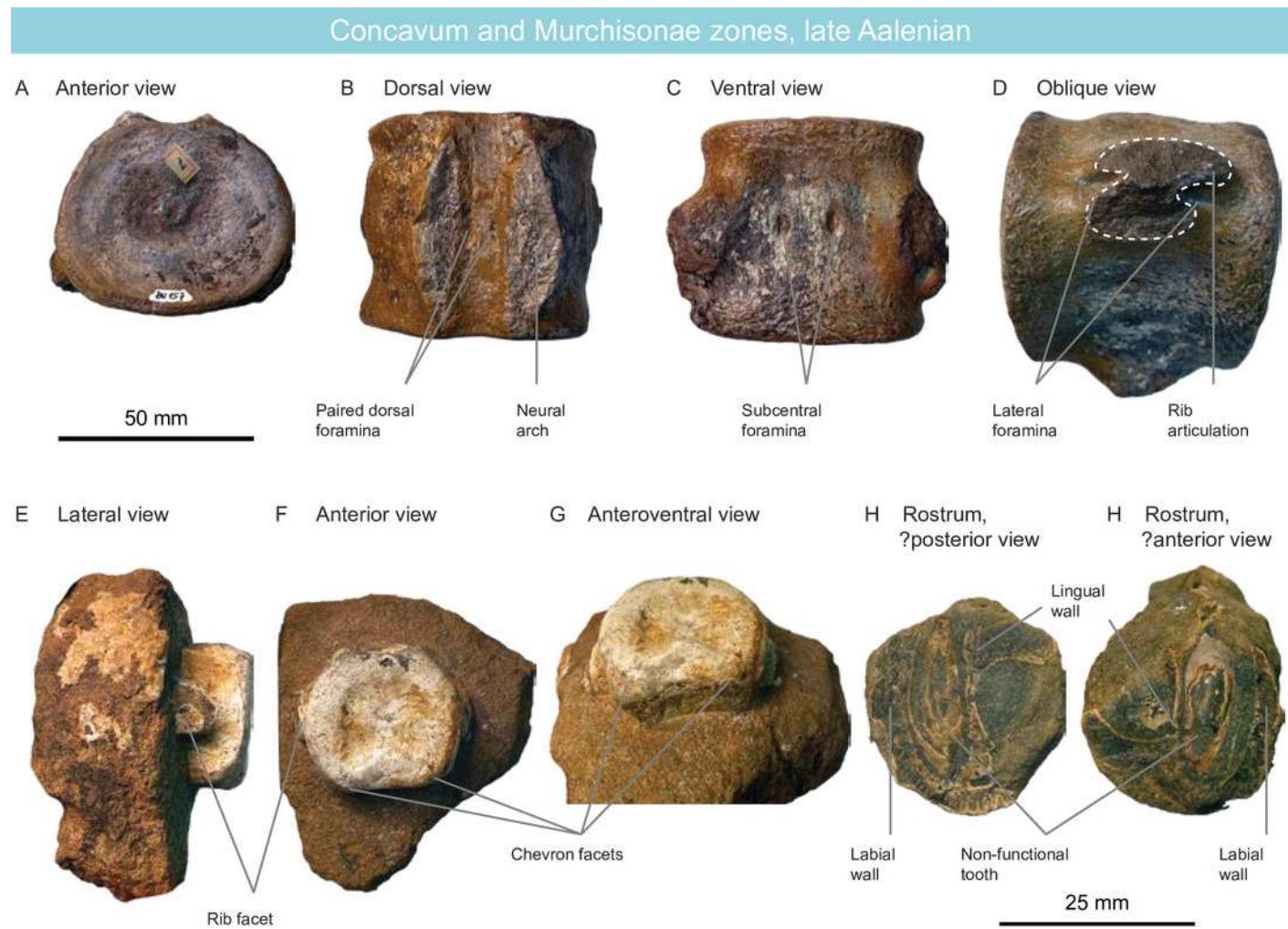


Figure 10

Humphresianum Zone fauna, early Bajocian, Luxembourg.

Humphriesianum Zone, early Bajocian

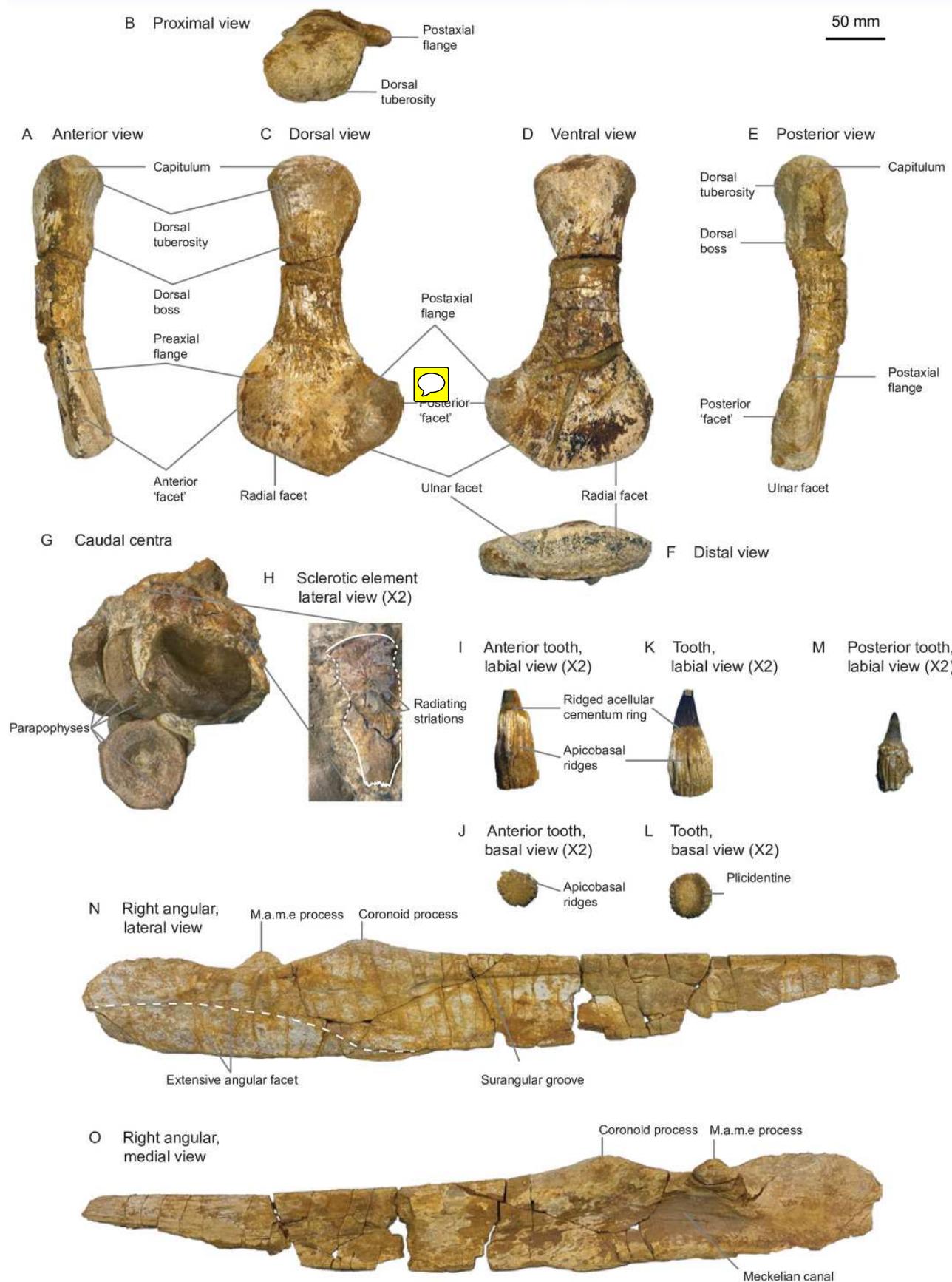


Figure 11

The new Humphresianum Zone ophthalmosaurid, early Bajocian, Luxembourg.

Humphriesianum Zone, early Bajocian

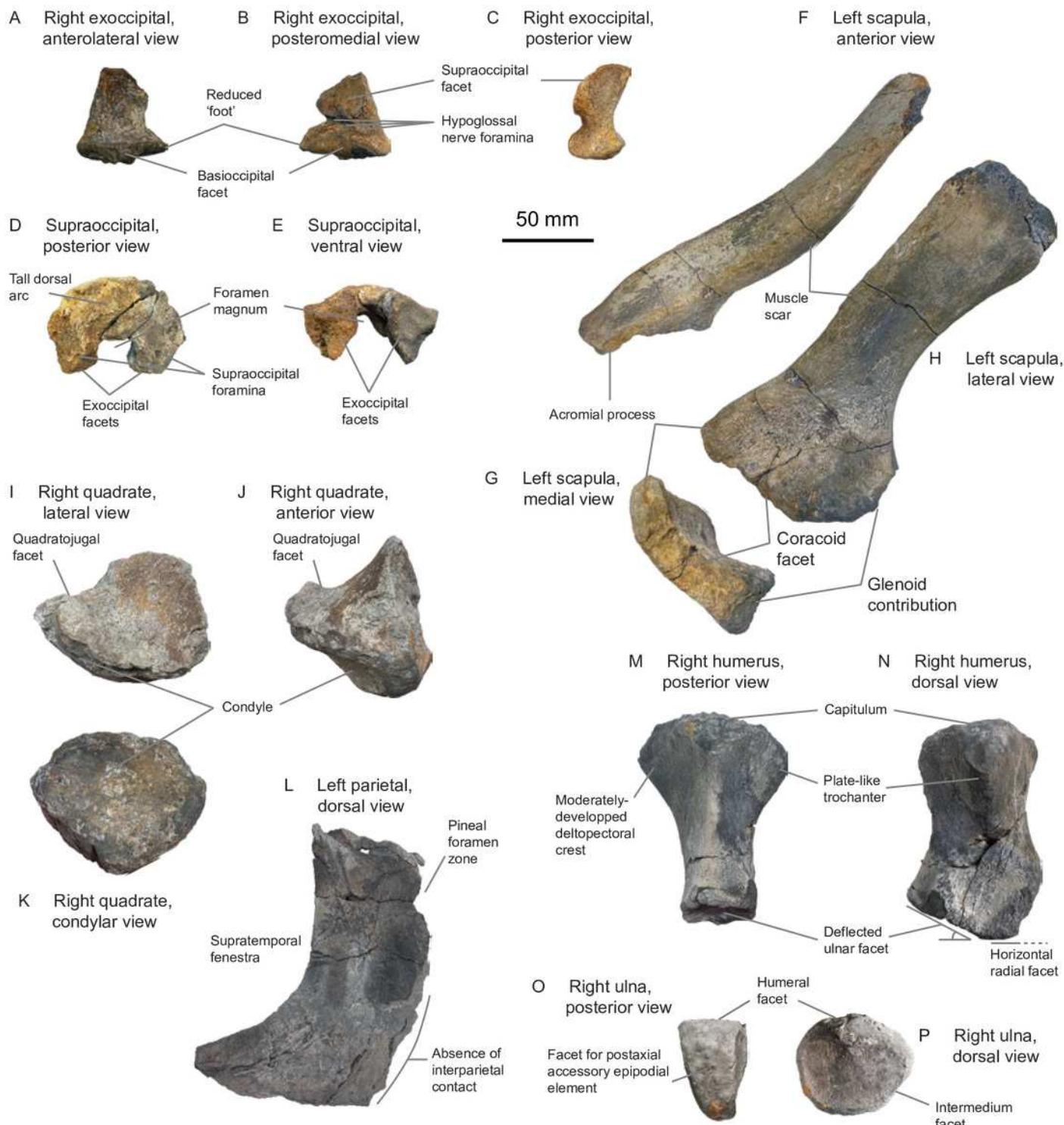


Figure 12

Ichthyosaurian centra proportions.

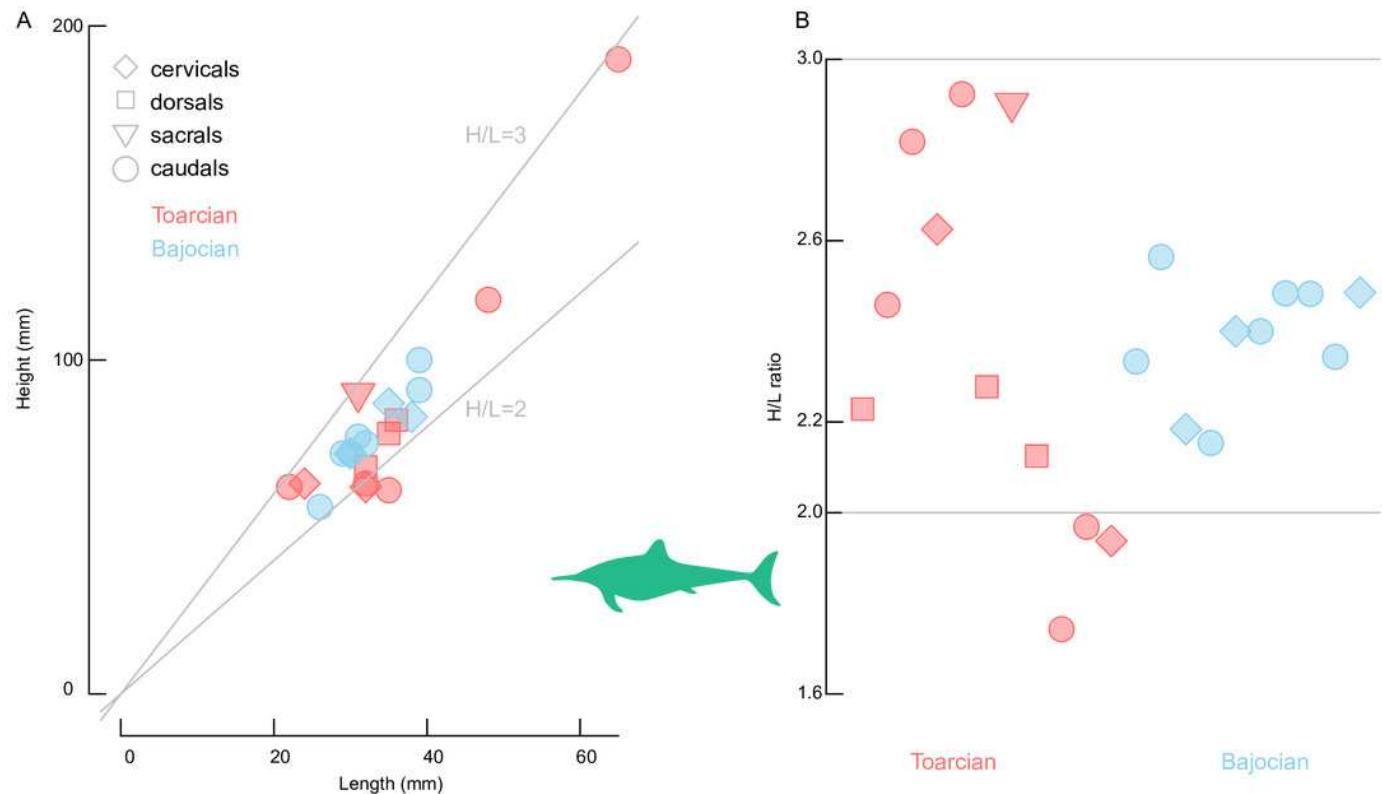


Table 1(on next page)

Centra dimensions (in millimeters).

1 **Table 1.** Centra measurements in millimeters.

Specimen	Clade	Age	Axial_zone	Height	Length	Width
ULgPA35961	Ichthyosauria	Toarcian	dorsal	78	35	86
IRSNB_Vert-00000-00803	Ichthyosauria	Toarcian	caudal	118	48	118
IRSNB_Vert-00000-00808	Ichthyosauria	Toarcian	caudal	62	22	60
IRSNB_Vert-06462-0005	Ichthyosauria	Toarcian	cervical	63	24	64
IRSNB_Vert-11312-00007	Plesiosauria	Toarcian	dorsal	67	44	81
MNHNL_DOU378	Ichthyosauria	Toarcian	caudal	190	65	NA
MNHNL_DOU978	Plesiosauria	Toarcian	cervical	41	55	53
MNHNL_DOU944	Ichthyosauria	Toarcian	dorsal	82	36	NA
MNHNL_DOU998	Ichthyosauria	Toarcian	sacral	90	31	NA
MNHNL_DOU722	Plesiosauria	Toarcian	dorsal	50	37	60
MNHNL_DOU723	Plesiosauria	Toarcian	dorsal	55	30	63
MNHNL_DOU954	Plesiosauria	Toarcian	pectoral	63	42	73
MNHNL_DOU724	Plesiosauria	Toarcian	sacral	39	40	57
MNHNL_DOU369a	Ichthyosauria	Toarcian	dorsal	68	32	NA
MNHNL_DOU369b	Ichthyosauria	Toarcian	caudal	61	35	NA
MNHNL_DOU369c	Ichthyosauria	Toarcian	caudal	63	32	NA
MNHNL_HU384	Plesiosauria	Aalenian	caudal	32	NA	NA
MNHNL_BU157	Plesiosauria	Aalenian	cervical	67	62	71
TMNHNL_MP212	Ichthyosauria	Toarcian	cervical	62	32	56
MNHNL_BM779	Ichthyosauria	Bajocian	caudal	91	39	NA
MNHNL_BM758	Ichthyosauria	Bajocian	caudal	100	39	NA
MNHNL_BM725a	Ichthyosauria	Bajocian	cervical	83	38	NA
MNHNL_BM725b	Ichthyosauria	Bajocian	caudal	56	26	NA
MNHNL_BM461	Ichthyosauria	Bajocian	cervical	72	30	NA
MNHNL_BM392a	Ichthyosauria	Bajocian	caudal	72	30	NA
MNHNL_BM392b	Ichthyosauria	Bajocian	caudal	77	31	NA
MNHNL_BM392c	Ichthyosauria	Bajocian	caudal	72	29	NA
MNHNL_BM392d	Ichthyosauria	Bajocian	caudal	75	32	NA
MNHNL_BM779	Ichthyosauria	Bajocian	cervical	87	35	90

Table 2(on next page)

List of the material and taxonomic identifications.

1 **Table 2. List of the material and taxonomic identifications.**

Specimen	Locality	Stage	Zone	Element	Clade	Identification	Note
ULgPA359 61	Arlon region	Toarcian	Dispansum to Aalensis	Dorsal centrum	Ichthyosau- ria	Parvipelvia indet.	
IRSNB Vert- 06455- 0001	Mont-Saint- Martin (France)	Toarcian	Dispansum to Aalensis	Centrum	Ichthyosau- ria	Ichthyosauria indet.	
IRSNB Vert- 00000- 00801	Halanzы (Belgium)	Toarcian	Dispansum to Aalensis	Centrum	Ichthyosau- ria	Ichthyosauria indet.	
IRSNB Vert- 06659- 0001	NA (Luxem- bourg)	Toarcian	Dispansum to Aalensis	Centrum	Ichthyosau- ria	Ichthyosauria indet.	
IRSNB Vert- 06462- 0003	Halanzы (Belgium)	Toarcian	Dispansum to Aalensis	Tooth	Ichthyosau- ria	Ichthyosauria indet.	
IRSNB Vert- 00000- 00802	NA	Toarcian	Dispansum to Aalensis	Right quadrate	Ichthyosau- ria	Ichthyosauria indet.	
IRSNB Vert- 00000- 00800	Halanzы (Belgium)	Toarcian	Dispansum to Aalensis	Left surangular	Ichthyosau- ria	Parvipelvia indet.	Probable non- ophthalmosaurid
IRSNB Vert- 06462- 0002	Halanzы (Belgium)	Toarcian	Dispansum to Aalensis	Left angular	Ichthyosau- ria	Parvipelvia indet.	Probable non- ophthalmosaurid
IRSNB Vert- 00000- 00803	NA (Luxem- bourg)	Toarcian	Dispansum to Aalensis	Caudal centrum	Ichthyosau- ria	Parvipelvia indet.	<i>Temnodontosau- rus</i> -like
IRSNB Vert- 00000- 00804	NA (Luxem- bourg)	Toarcian	Dispansum to Aalensis	Dorsal centrum	Ichthyosau- ria	Parvipelvia indet.	<i>Temnodontosau- rus</i> -like
IRSNB Vert- 00000- 00808	NA	Toarcian	Dispansum to Aalensis	Caudal centrum	Ichthyosau- ria	Parvipelvia indet.	Possible thunnosaurian
IRSNB Vert- 06462- 0005	Halanzы (Belgium)	Toarcian	Dispansum to Aalensis	Cervical centrum	Ichthyosau- ria	Parvipelvia indet.	

IRSNB Vert- 11312- 00007	Halanzyn (Belgium)	Toarcian	Dispansum to Aalensis	Dorsal centrum	Plesiosauria	Plesiosauria indet.	
MNHNL TM212	Dudelange	Toarcian	Thouarsense	Partial skeleton	Ichthyosau- ria	Baracromia indet.	Non- ophthalmosaurid
MNHNL DOU9 44	NA (Luxembourg)	Toarcian	Aalensis	Dorsal centrum	Ichthyosau- ria	Parvipelvia indet.	
MNHNL DOU9 98	NA (Luxembourg)	Toarcian	Aalensis	Caudal centrum	Ichthyosau- ria	Parvipelvia indet.	
MNHNL DOU3 69	NA (Luxembourg)	Toarcian	Aalensis	Centra	Ichthyosau- ria	Parvipelvia indet.	Probable non- thunnosaurian
MNHNL DOU3 53	NA (Luxembourg)	Toarcian	Aalensis	Coracoid	Ichthyosau- ria	Parvipelvia indet.	Non- baracromian
MNHNL DOU3 78	NA (Luxembourg)	Toarcian	Aalensis	Posterior dorsal to anterior caudal centrum	Ichthyosau- ria	Parvipelvia indet.	Probable non- thunnosaurian
MNHNL HU242	Rumelange	Aalenian	Concavum	Partial rostrum	Ichthyosau- ria	Ichthyosauria indet.	
MNHNL BM360	Rumelange	Bajocian	Humphriesian um	Partial neural spine and ribs	Ichthyosau- ria	Ichthyosauria indet.	
MNHNL BM758	Rumelange	Bajocian	Humphriesian um	Caudal centrum	Ichthyosau- ria	Parvipelvia indet.	
MNHNL BM725	Rumelange	Bajocian	Humphriesian um	Cervical and caudal centra	Ichthyosau- ria	Parvipelvia indet.	
MNHNL BM461	Rumelange	Bajocian	Humphriesian um	Cervical centrum	Ichthyosau- ria	Parvipelvia indet.	
MNHNL BM392	Rumelange	Bajocian	Humphriesian um	Five caudal centrum and one sclerotic plate	Ichthyosau- ria	Parvipelvia indet.	
MNHNL BM780 _781	Rumelange	Bajocian	Humphriesian um	Surangular, teeth	Ichthyosau- ria	Ophthalmosauri- dae indet.	
MNHNL BM779	Rumelange	Bajocian	Humphriesian um	Partial skeleton	Ichthyosau- ria	Ophthalmosauri- dae indet.	
MNHNL DOU3 07	Rodange	Toarcian	Pseudoradiosa	Humerus	Plesiosauria	Rhomaleosaurid ae indet.	

MNHNL KA109	Esch-sur- Alzette	Toarcia n	Pseudoradiosa	Humerus	Plesiosauri a	Rhomaleosaurid ae indet.	
MNHNL DOU9 06	Esch-sur- Alzette	Toarcia n	Aalensis	Tooth crown	Plesiosauri a	Plesiosauria indet.	
MNHNL DOU7 24	Probably Esch- sur- Alzette	Toarcia n	Aalensis	Sacral vertebra	Plesiosauri a	Plesiosauria indet.	
MNHNL DOU7 22	Esch-sur- Alzette	Toarcia n	Aalensis	Pectoral vertebra	Plesiosauri a	Rhomaleosaurid ae indet.	
MNHNL DOU7 23	Esch-sur- Alzette	Toarcia n	Aalensis	Pectoral vertebra	Plesiosauri a	Rhomaleosaurid ae indet.	
MNHNL DOU9 54	NA (Luxem bourg)	Toarcia n	Aalensis	Posterior pectoral/ant ior dorsal	Plesiosauri a	Rhomaleosaurid ae indet.	
MNHNL DOU5 58	Belvaux	Toarcia n	Aalensis	Humerus	Plesiosauri a	Rhomaleosaurid ae indet.	
MNHNL DOU3 24A	Esch-sur- Alzette	Toarcia n	Aalensis	Humerus	Plesiosauri a	Rhomaleosaurid ae indet.	
MNHNL DOU3 24B	Esch-sur- Alzette	Toarcia n	Aalensis	Humerus	Plesiosauri a	Rhomaleosaurid ae indet.	Juvenile
MNHNL DOU9 78	NA (Luxem bourg)	Toarcia n	Aalensis	Cervical centrum	Plesiosauri a	cf. Microcleidus	
MNHNL HU384	Esch-sur- Alzette	Aalenia n	Murchisonae	Caudal centrum	Plesiosauri a	Plesiosauria indet.	Juvenile
MNHNL BU157	Rumelange	Aalenia n	Concavum or Murchisonae	Cervical vertebra	Plesiosauri a	Plesiosauroidea indet.	Non-cryptoclidid
MNHNL BM782	Rumelange	Bajocia n	Humphriesian um	Propodial	Plesiosauri a	Cryptoclididae indet.	