

A trace fossil made by a walking crayfish or crayfish-like arthropod from the Lower Jurassic Moenave Formation of southwestern Utah, USA

Makae Rose¹, Jerald D Harris^{Corresp., 2}, Andrew RC Milner³

¹ Department of Biology, Dixie State University, St. George, UT, United States

² Department of Physical Sciences, Dixie State University, St. George, UT, United States

³ St. George Dinosaur Discovery Site at Johnson Farm, St. George, UT, United States

Corresponding Author: Jerald D Harris
Email address: jharris@dixie.edu

New invertebrate trace fossils from the Lower Jurassic Moenave Formation at the St. George Dinosaur Discovery Site at Johnson Farm (SGDS) continue to expand the ichnofauna at the site. A previously unstudied arthropod locomotory trace, SGDS 1290, comprises two widely spaced, thick, gently undulating paramedial impressions flanked externally by small, tapered to elongate tracks with a staggered to alternating arrangement. The specimen is not a variant of any existing ichnospecies, but bears a striking resemblance to modern, experimentally generated crayfish walking traces, suggesting a crayfish or crayfish-like maker for the fossil. Because of its uniqueness, we place it in a new ichnospecies, *Siskemia eurypyge*. It is the first fossil crayfish or crayfish-like locomotion trace ever recorded.

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2 **arthropod from the Lower Jurassic Moenave Formation of**
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5 Makae Rose¹, Jerald D. Harris², and Andrew R.C. Milner³

6 ¹ Department of Biology, Dixie State University, St. George, UT, 84770, USA

7 ² Department of Physical Sciences, Dixie State University, St. George, UT, 84770, USA

8 ³ St. George Dinosaur Discovery Site at Johnson Farm, St. George, UT, 84790, USA

9

10 Corresponding Author:

11 Jerald D. Harris

12 225 South 700 East, St. George, UT, 84770, USA

13 Email address: jharris@dixie.edu

14

15 Abstract

16 New invertebrate trace fossils from the Lower Jurassic Moenave Formation at the St.
17 George Dinosaur Discovery Site at Johnson Farm (SGDS) continue to expand the ichnofauna at
18 the site. A previously unstudied arthropod locomotory trace, SGDS 1290, comprises two widely
19 spaced, thick, gently undulating paramedial impressions flanked externally by small, tapered to
20 elongate tracks with a staggered to alternating arrangement. The specimen is not a variant of any
21 existing ichnospecies, but bears a striking resemblance to modern, experimentally generated
22 crayfish walking traces, suggesting a crayfish or crayfish-like maker for the fossil. Because of its
23 uniqueness, we place it in a new ichnospecies, *Siskemia eurypyge*. It is the first fossil crayfish or
24 crayfish-like locomotion trace ever recorded.

25

26 Introduction

27 Paleoichnology, the study of ichnofossils (fossil tracks and traces), contributes a
28 substantial body of paleobiological information to the understanding of extinct organisms. This
29 is because ichnofossils are direct results of ancient animal behavior (Osgood, 1975) that could
30 only otherwise be inferred from body fossils. Furthermore, the global commonness of
31 ichnofossils compared to body fossils means that the ichnological record often can preserve
32 evidence of the presence of organisms not otherwise or poorly represented in the body fossil
33 record (Osgood, 1975), especially of invertebrates that lack hard parts and therefore fossilize
34 only under exceptional conditions. Except for conchostracans (*sensu* Kozur & Weems, 2010) and
35 ostracods, which have biomineralized carapaces, arthropods, when compared to their
36 evolutionary diversity, are among the less commonly preserved body-fossil components of
37 terrestrial (including freshwater) paleoecosystems except in various Konservat-Lagerstätten
38 (fossiliferous sites of exceptional preservational quality) (e.g., Charbonnier et al., 2010; Luque et
39 al., 2019; Selden & Nudds, 2012; Smith, 2012). Yet from the mid-Paleozoic through the
40 Cenozoic, arthropods—especially insects and arachnids—were certainly the most populous and
41 diverse metazoan components of most terrestrial ecosystems (Labandeira & Beall, 1990), and
42 their paleoecological importance cannot be underestimated.

43 Arthropod ichnofossils can be more common and abundant than arthropod body fossils,
44 and may indicate the presences of various arthropod taxa in terrestrial paleoecosystems for which
45 body fossils may be entirely absent. Burrows (domichnia) made by arthropods comprise one
46 such body of evidence. For example, several burrow ichnotaxa in eolian sandstones have been
47 attributed to arthropods (Ekdale, Bromley & Loope, 2007). Some *Macanopsis*, *Psilonichnus*, and
48 *Skolithos* burrows may have been made by spiders (Uchman, Vrenozi & Muceku, 2018); other
49 *Psilonichnus* have been attributed to crabs (Curran, Savarese & Glumac, 2016). *Camborygma*
50 burrows, attributed to crayfish (Hasiotis & Mitchell, 1993), are the primary body of evidence for
51 crayfish in the fossil record. Perhaps more familiarly, walking tracks (repichnia) of arthropods
52 have an extensive geological history, spanning from the Cambrian (and possibly latest
53 Precambrian [Chen et al., 2018]) through the Holocene (e.g., Eiseman & Charney, 2010). They

54 constitute some of the earliest evidence of metazoan life venturing onto land (reviewed in Minter
55 et al., 2016a; Minter et al., 2016b) and are known from virtually every paleoenvironment, from
56 near shore and shallow marine environments (e.g., Collette, Hagadorn & Lacelle, 2010;
57 MacNaughton et al., 2002; Pirrie, Feldmann & Buatois, 2004; Shillito & Davies, 2018; Trewin &
58 McNamara, 1994) and, terrestrially, from proglacial systems (e.g., Anderson, 1981; Lima,
59 Minter & Netto, 2017; Lima et al., 2015; Uchman, Kazakauskas & Gaigalas, 2009; Walter,
60 1985) to desert ergs (e.g., Gilmore, 1927; Good & Ekdale, 2014; Sadler, 1993).

61 The St. George Dinosaur Discovery Site at Johnson Farm (SGDS) in St. George,
62 Washington County, Utah (Fig. 1) has been called a Konzentrat-Ichnolagerstätte (Hunt & Lucas,
63 2006a) because of its exceptional concentration of well-preserved ichnofossils from a broad
64 spectrum of terrestrial organisms. The site preserves a detailed “snapshot” of an earliest Jurassic
65 ecosystem from on- and offshore portions of a freshwater, lacustrine paleoenvironment. The
66 “Dinosaur Discovery” part of the name of the site indicates the concentration of dinosaur tracks
67 at this locality (Milner, Lockley & Johnson, 2006; Milner, Lockley & Kirkland, 2006; Milner et
68 al., 2009), but tracks of other vertebrates (Lockley, Kirkland & Milner, 2004; Milner, Lockley &
69 Johnson, 2006) and a moderately diverse invertebrate ichnofauna (Lucas et al., 2006a) are also
70 preserved. Burrows pertaining to *Helminthoidichnites*, *Palaeophycus*, and *Skolithos* are abundant
71 at the site, but locomotory trackways made by arthropods, referred to cf. *Bifurculapes*,
72 *Diplichnites*, and cf. *Kouphichnium*, are also present. Ichnospecies of *Bifurculapes* have been
73 variably attributed to insects, possibly beetles, and to crustaceans (Getty, 2016, 2018; Hitchcock,
74 1858, 1865); ichnospecies of *Diplichnites* have been attributed to myriapods (Briggs, Rolfe &
75 Brannan, 1979; Davis, Minter & Braddy, 2007; Pollard, Selden & Watts, 2008; Shillito &
76 Davies, 2018), notostracans (Lucas et al., 2006a; Minter et al., 2007), and other arthropods
77 (Melchor & Cardonatto, 2014; Minter et al., 2007); and ichnospecies of *Kouphichnium* have
78 been attributed to limulids (Caster, 1944; King, Stimson & Lucas, 2019; Lomax & Racay, 2012).

79 A previously unstudied SGDS specimen, SGDS 1290, is an arthropod locomotory trace
80 that differs markedly from any other SGDS specimen, indicating the presence of a heretofore
81 unrecognized component of the SGDS ichnofauna. SGDS 1290 is an arthropod locomotory trace
82 because it includes distinct footprints in a discernible cycle, but lacks any features of vertebrate
83 tracks, such as distinct toes (*sensu* Seilacher, 2007). The trace thus resembles numerous other
84 fossil traces attributed to arthropods, as well as those generated experimentally. The fossil was
85 discovered and collected March 11, 2010 by SGDS volunteer Jon Cross.

86

87 Geological Setting

88 Most of the fossils preserved at the SGDS, including the ichnofossil described below,
89 occur in the Whitmore Point Member of the Moenave Formation (Kirkland & Milner, 2006;
90 Kirkland et al., 2014), which conformably overlies the Dinosaur Canyon Member of the
91 Moenave Formation and disconformably underlies the Springdale Sandstone Member, which
92 itself has been assigned as both the lowest member of the Kayenta Formation (Lucas & Tanner,
93 2006) and the uppermost member of the Moenave Formation (Steiner, 2014a). The richest source

94 of the ichnofossils at the SGDS, again including the trace described below, occur within a fine-
95 grained sandstone near the base of the Whitmore Point Member initially called the “Main Track
96 Layer” (Kirkland & Milner, 2006; Milner, Lockley & Johnson, 2006; Milner, Lockley &
97 Kirkland, 2006) and, later and more formally, the Johnson Farm Sandstone Bed (unit 40 of
98 Kirkland et al., 2014) (Fig. 2). SGDS 1290 comes from the uppermost strata of this unit,
99 specifically one of several thinly bedded, apparently conformable, fine-grained-sandstone “Top
100 Surface Tracksite” horizons (*sensu* Kirkland et al., 2014; Milner, Lockley & Johnson, 2006;
101 Milner, Lockley & Kirkland, 2006). The Whitmore Point Member preserves sediments deposited
102 in and around the large, freshwater Lake Whitmore (formerly called Lake Dixie) (Kirkland &
103 Milner, 2006; Kirkland et al., 2014); at the SGDS itself, the Johnson Farm Sandstone Bed
104 preserves ichnofossils and sedimentary structures made in both subaerial and subaqueous
105 conditions, indicating a shoreline paleoenvironment (Milner, Lockley & Kirkland, 2006).
106 Ichnologically, invertebrate ichnofossils in this paleoenvironment pertain to the *Scoyenia*
107 ichnofacies (Buatois & Mángano, 2004; Lucas et al., 2006a), while the associated vertebrate
108 ichnofauna pertains to the *Eubrontes* ichnocoenosis of the *Grallator* ichnofacies (Hunt & Lucas,
109 2006b; Hunt & Lucas, 2006c).

110 Age determinations for the Whitmore Point Member have varied. The unit typically has
111 been placed entirely within the Hettangian (earliest Jurassic) largely on biostratigraphic grounds
112 (see discussions in Kirkland et al., 2014; Milner et al., 2012; Parrish et al., 2019; and Tanner &
113 Lucas, 2009), but was also determined to straddle the Triassic–Jurassic boundary (201.3 ± 0.2
114 Ma) on magnetostratigraphic grounds (Donohoo-Hurley, Geissman & Lucas, 2010), in which
115 system the Johnson Farm Sandstone Bed would be Rhaetian (latest Triassic) in age. However,
116 Steiner (2014b) recovered Hettangian paleomagnetic sequences from the Whitmore Point
117 Member, and Suarez et al. (2017) calibrated the magnetostratigraphic data of Donohoo-Hurley,
118 Geissman & Lucas (2010) with high-precision U–Pb dates to re-situate the Triassic–Jurassic
119 boundary stratigraphically farther down in the Dinosaur Canyon Member of the Moenave
120 Formation, also making the Whitmore Point Member entirely earliest Jurassic in age. The
121 Johnson Farm Sandstone Bed and its fossils therefore are Hettangian in age, approximately 200
122 million years old.

123

124 **Materials & Methods**

125 Ichnological terminology for arthropod locomotory traces used herein follows Minter,
126 Braddy & Davis (2007) and Genise (2017). Minter, Braddy & Davis (2007) defined “tracks” as
127 discrete marks made by locomotory appendages, “impressions” as continuous traces made by
128 another portion of the anatomy of a trace maker, and “imprints” as discontinuous such traces;
129 they also provided terms for trackway arrangement and measurements. Genise (2017) outlined
130 various descriptive terms for individual track morphologies. Measurements of SGDS 1290 (Fig.
131 3C, Table 1) were taken using digital calipers. The measurements were: track length and width,
132 internal and external widths between paramedial impressions, distances between tracks and
133 adjacent paramedial impressions, and widths of left and right paramedial impressions.

134 Measurements pertaining to the paramedial impressions were taken adjacent to individual
135 tracks/track sets.

136 The electronic version of this article in Portable Document Format (PDF) will represent a
137 published work according to the International Commission on Zoological Nomenclature (ICZN),
138 and hence the new names contained in the electronic version are effectively published under that
139 Code from the electronic edition alone. This published work and the nomenclatural acts it
140 contains have been registered in ZooBank, the online registration system for the ICZN. The
141 ZooBank LSIDs (Life Science Identifiers) can be resolved and the associated information viewed
142 through any standard web browser by appending the LSID to the prefix <http://zoobank.org/>. The
143 LSID for this publication is: urn:lsid:zoobank.org:pub:D78963CE-11C8-4447-8E26-
144 BBCCF0E37143. The LSID for the herein described *Siskemia eurypyge* isp. nov. is:
145 urn:lsid:zoobank.org:act:769B0815-8991-4F0E-B32C-99C87A9D293B. The online version of
146 this work is archived and available from the following digital repositories: PeerJ, PubMed
147 Central and CLOCKSS.

148

149 **Description of SGDS 1290**

150 SGDS 1290, a natural cast (convex hyporelief), consists of two parallel, undulating,
151 paramedial ridges flanked externally by eight sets of small tracks that range in shape from ovoid
152 to tapered (teardrop-shaped) to elongate (Fig. 3A, B). Tracks within each set are closely
153 appressed; track sets are, however, spaced well apart from one another. The track sets are
154 oriented perpendicular to the trackway axis, though tapered and elongate individual tracks within
155 each set have long axes that parallel or are oblique to the trackway axis. Track sets average 4.5
156 mm long anteroposteriorly and 5.4 mm wide mediolaterally (Table 1). Left (L) and right (R) sets
157 of tracks are arranged in an alternating pattern. Based on Fairchild & Hasiotis (2011), the
158 tapering ends of the tapered tracks are presumed to be anterior reflectures, indicating the
159 direction of movement. Most tracks have long axes oriented parallel to the trackway axis; a few
160 (such as in sets L1, L4, and R2) are oblique to the axis. Track R1 is markedly elongate rather
161 than tapered, but also parallel to the trackway axis. Track sets L1, L3, and R4 consist of three
162 distinct but appressed tracks; sets L2, L4, and R2 consist of pairs of appressed tracks, and R1 and
163 R3 appear to consist of single tracks, although the possibility that each comprises multiple,
164 conjoined tracks cannot be ruled out.

165 The paramedial impressions typically are thick mediolaterally, though they vary and taper
166 briefly to nothingness in a few places (being more continuous than repeating, we consider them
167 “impressions” and not “imprints”). The impressions follow gently undulating (non-linear and
168 low amplitude) pathways. They span approximately 7.5 cm along the slab of rock. Overprints of
169 short segments of the paramedial impressions that are not accompanied by tracks are visible
170 behind the main trace segment on a slightly higher stratum. The impressions vary in width along
171 their lengths, ranging from 0–3.1 mm (mean = 2.0 mm) for the left impressions and 0–3.0 mm
172 (mean = 1.6 mm) for the right (Table 1). The width of the trace from left impression to right
173 impression averages 9.4 mm when measured from the lateral (external) edges and 5.6 mm when

174 measured between the medial (internal) edges (Table 1). The distances between the impressions
175 thus are greater than the distances between the impressions and their flanking tracks (mean = 1.6
176 mm); the ratio of the distance between a paramedial impression and its flanking track to the
177 distance between the medial edges of the paramedial impressions ranges from 0-0.78 (mean =
178 0.31; see Supplemental Material), so on average, the paramedial impressions are roughly three
179 times farther apart from each other than either is from its flanking tracks. The impressions taper
180 slightly in cross sectional view: they are wider at their bases and narrower at their rounded
181 apices.

182

183 **Comparisons to Arthropod Repichnial Ichnotaxa**

184 **Arthropod repichnia lacking medial or paramedial impressions**

185 Several arthropod locomotory (walking) ichnotaxa are readily distinguished from SGDS
186 1290 by (usually) lacking medial or paramedial impressions, but are worth comparing to SGDS
187 1290 to ascertain whether or not it might be a morphological variant of such ichnotaxa. These
188 ichnotaxa are *Acanthichnus* (Hitchcock, 1858), *Asaphoidichnus* (Miller, 1880), *Bifurculapes*
189 (Hitchcock, 1858), *Coenobichnus* (Walker, Holland & Gardiner, 2003), *Copeza* (Hitchcock,
190 1858), *Danstairia congesta* (Smith, 1909), *Diplichnites* (Dawson, 1873), *Eisenachichnus* (Kozur,
191 1981), *Foersterichnus* (Pirrie, Feldmann & Buatois, 2004), *Hamipes* (Hitchcock, 1858),
192 *Laterigradus* (de Carvalho et al., 2016), *Lithographus* (Hitchcock, 1858), *Hornburgichnium*
193 (Kozur, 1989), *Maculichna* (Anderson, 1975a), *Merostomichnites* (Packard, 1900),
194 *Mirandaichnium* (Aceñolaza, 1978), *Octopodichnus* (Gilmore, 1927), *Petalichnus* (Miller,
195 1880), *Pterichnus* (Hitchcock, 1865), *Tasmanadia* (Chapman, 1929), and *Umfolozia* (Savage,
196 1971). Most of these ichnotaxa further differ from SGDS 1290 in the shapes and configurations
197 of their tracks. The comparisons below specify track-making taxa only when one or more have
198 been proposed for the ichnotaxon.

199 *Acanthichnus* tracks (Fig. 4A), attributed to a chelicerate such as a solifugid, are
200 oppositely arranged, short, elongate impressions in two (or four, per Dalman & Lucas, 2015)
201 parallel rows; tracks either are parallel to or angle slightly away from the trackway axis (Dalman
202 & Lucas, 2015; Hitchcock, 1858). This morphology and organization are both unlike those of
203 SGDS 1290.

204 *Asaphoidichnus* tracks (Fig. 4B), attributed to trilobites, are elongate to crescentic,
205 possess 3–4 crescentic branches at one end, and are oriented oblique to the trackway axis (Miller,
206 1880). They are far more complex in structure than the tracks of SGDS 1290.

207 *Bifurculapes* (Fig. 4C), attributed to an insect, possibly a beetle (Getty, 2016), comprises
208 adjacent pairs (rarely triplets) of slightly staggered, elongate, straight to crescentic tracks that lie
209 parallel or slightly oblique to the trackway axis, unlike the tracks of SGDS 1290. Tracks in each
210 pair sometimes converge toward one end in *Bifurculapes*. This ichnotaxon typically does not
211 possess paramedial impressions, but a specimen described by Getty (2016: fig. 1) possess two
212 such traces, albeit faintly, that lie closer to the tracks than to the trackway axis, as in SGDS 1290.
213 These impressions are far less pronounced than their associated tracks, unlike those of SGDS

214 1290. Getty (2020) ascertained that *Bifurculapes* traces were made subaqueously and may have
215 been made by a terrestrial insect that would have left different tracks subaerially.

216 *Coenobichnus* tracks (Fig. 4D), attributed to a hermit crab, are thick, crescentic to
217 ellipsoidal, roughly parallel and closely appressed to the trackway axis, and asymmetrical, with
218 the left tracks larger than the right tracks (Walker, Holland & Gardiner, 2003), all of which
219 differentiate this ichnotaxon from SGDS 1290.

220 *Copeza* (possibly a variant and synonym of *Lithographus* [Lull, 1953; Rainforth, 2005];
221 Fig. 4E) consists of triplets of roughly oppositely arranged, linear, elongate tracks in which the
222 anteriormost lies roughly perpendicular to the trackway axis while the posteriormost pairs lie
223 parallel or oblique to the trackway axis (Lull, 1953). This rare ichnotaxon is thus unlike SGDS
224 1290.

225 *Danstairia congesta* (Fig. 4F) comprises circular to crescentic tracks in closely appressed
226 sets of up to six that are oriented oblique to the trackway axis; tracks often overlap to form V-
227 shaped structures (Walker, 1985), unlike in SGDS 1290.

228 *Diplichnites* (possibly including *Acripes* per Häntzschel [1975] and Hammersburg,
229 Hasiotis & Robison [2018]; also see below) tracks span a range of morphologies. *D. aenigma*
230 (Fig. 4G), ostensibly the ichnospecies lectotype except that no specimen was designated as such
231 (Stimson et al., 2018), typically comprises elongate, closely packed tracks in parallel rows on
232 either side of the trackway axis; the tracks lie perpendicular to the trackway axis (Dawson,
233 1873). *D. gouldi* Type A (Fig. 4H) comprises parallel rows of closely spaced, oppositely
234 arranged, simple, roughly circular to oblong to comma-shaped or irregular tracks with varying
235 orientations to the trackway axis; *D. gouldi* Type B comprises closely spaced, elongate
236 impressions oriented perpendicular, or nearly so, to the trackway axis, matching the general
237 description of *D. aenigma*; *D. gouldi* Type C is similar to Type B, but the tracks are oriented
238 oblique (~ 45 °) to the trackway axis (Trewin & McNamara, 1994). *D. cuithensis*, attributed to
239 large arthropleurid myriapods, is similar to both *D. aenigma* and *D. gouldi* Type B, but is very
240 large and has widely spaced rows of tracks (Briggs, Rolfe & Brannan, 1979). *D. binatus* tracks
241 often occur as closely appressed pairs of imprints (Webby, 1983). *D. triassicus* tracks are small
242 and circular to ovoid rather than elongate, and frequently paired on either side of the trackway
243 axis (e.g., Pollard, Selden & Watts, 2008); this ichnospecies has been alternately suggested to be
244 a synonym of *D. gouldi* (e.g., Lucas et al., 2006b) or to pertain to *Acripes* (e.g., Machalski &
245 Machalska, 1994; Pollard, 1985). *D. metzi* possesses a midline impression that is sometimes
246 interrupted by connections between tracks in the closely appressed rows (Fillmore et al., 2017).
247 The ichnogenus is in need of thorough review (Smith et al., 2003), but in all cases, the tracks are
248 of different morphologies and arrangements than those of SGDS 1290.

249 *Eisenachichnus* tracks (Fig. 4I) are elongate, paired, and oblique to the trackway axis
250 (rarely perpendicular), but the patterns of the pairs on either side of the trackway axis normally
251 are asymmetrical: on one side, the paired tracks lie end to end, while on the other, they are more
252 adjacent (Kozur, 1981). This morphology and arrangement are unlike those of SGDS 1290.

253 *Foersterichnus* (Fig. 4J), attributed to a crab, consists of widely spaced, paired rows of
254 elongate tracks in sets of 3–4; rows are parallel to oblique to the trackway axis, and tracks in
255 each set frequently overlap (Pirrie, Feldmann & Buatois, 2004). The wide spacing, clustering of
256 tracks, and orientations of the tracks are unlike those of SGDS 1290.

257 *Hamipes* tracks (Fig. 4K) consist of closely spaced, paired, elongate to crescentic
258 impressions oriented parallel to the trackway axis; the outer tracks are longer than their
259 accompanying inner tracks, and the tracks are staggered or alternately arranged (Getty, 2018;
260 Hitchcock, 1858). Track morphology readily differentiates *Hamipes* from SGDS 1290.

261 *Laterigradus* tracks (Fig. 4L), attributed to sideways-walking crabs, comprises
262 asymmetrical trackways consisting of sets of up to four tracks (de Carvalho et al., 2016).
263 Individual tracks vary widely in shape, ranging from elongate to tapered to comma-shaped to
264 roughly circular. Track sets fall within a narrow trackway width and exhibit different stride
265 lengths along the course of a trackway. While some individual track shapes resemble those of
266 SGDS 1290, the overall arrangement and layout of the tracks is distinct.

267 *Lithographus* (including *Permichnium sensu* Minter & Braddy, 2009) tracks (Fig. 4M),
268 which match those made by pterygote insects, especially extant cockroaches (Davis, Minter &
269 Braddy, 2007), comprise trios (or pairs, in the case of the *Permichnium* variant) of elongate to
270 comma-shaped, rather than circular or tapering, tracks that are arranged at different angles to one
271 another, some of which are oriented perpendicular to the trackway axis, and others of which are
272 oblique to the trackway axis (Guthörl, 1934; Hitchcock, 1858; Minter & Braddy, 2009). These
273 track morphologies and arrangements are distinct from those of SGDS 1290. *Hornburgichnium*
274 reportedly is similar to *Permichnium*, but has three tracks on either side of the midline instead of
275 two, and at least one of each set is oriented parallel to the trackway axis (Kozur, 1989); it may
276 also be a variant of *Lithographus* (Lucas et al., 2005a). Trackways of *Lithographus* can transition
277 into trackways that Hitchcock (1858) called *Hexapodichnus* (Davis, Minter & Braddy, 2007;
278 Minter, Braddy & Davis, 2007), so the latter may be considered a behavioral and/or substrate-
279 consistency variant of the former, and also unlike SGDS 1290.

280 *Maculichna* (including *Guandacolichnus* and *Paganzichnus* of Pazos [2000] per Buatois
281 & Mángano [2003]) tracks (Fig. 4N) comprise pairs (sometimes more) of small, circular to
282 slightly elongate tracks arranged in closely appressed, slightly staggered rows. Pairings are
283 oriented virtually parallel to the trackway axis (Anderson, 1975a); occasionally, short segments
284 of linear, singular medial or closely spaced, paired paramedial imprints are also preserved that
285 can be offset to one side of the trackway axis (Anderson, 1975a: fig. 8b, d, e). The pairing of
286 *Maculichna* tracks differs from that of SGDS 1290. Aceñolaza & Buatois (1991, 1993) and
287 Archer & Maples (1984) described *Maculichna* traces that exhibit the pairing of classic
288 *Maculichna* from South Africa, but in which tracks are more ellipsoidal to shaped like slightly
289 inflated isosceles triangles; the long axes of the triangles are oriented close to perpendicular to
290 the trackway axis. Pazos (2000) recognized this morphology as the separate ichnotaxon
291 *Paganzichnus*. This morphology is also unlike that of SGDS 1290.

292 The ichnospecies holotype of *Merostomichnites narragansettensis* (Fig. 4O) consists of
293 parallel rows of roughly oppositely arranged circular to elongate to comma-shaped tracks whose
294 long axes are perpendicular to the trackway axis (Packard, 1900). *Merostomichnites beecheri*
295 tracks are circular and connected across the trackway axis by curvilinear, shallow, M-shaped
296 imprints, creating a sort of segmented midline impression (Packard, 1900: fig. 4). These track
297 and trace morphologies do not match those of SGDS 1290.

298 *Mirandaichnium* (Fig. 4P) consists of two rows of elongate, linear tracks that terminate
299 laterally in small, circular impressions. Tracks are oriented perpendicular or oblique to the
300 trackway axis, oppositely situated, and often grouped into series of eight (Aceñolaza & Buatois,
301 1993; Buatois et al., 1998), unlike those of SGDS 1290.

302 *Octopodichnus* (Fig. 4Q) ichnospecies, attributed to arachnids, have different
303 morphologies. *O. didactylus* tracks are circular to crescentic to bifurcate or trifurcate oriented
304 parallel to the trackway axis in alternating, arcuate sets of four (Sadler, 1993). *O. minor* tracks
305 have a similar organization, but the tracks are more amorphous; *O. raymondi* tracks consist of
306 clusters of four circular to crescentic marks arranged in checkmark-like patterns (Sadler, 1993).
307 These track morphologies and distinctive arrangements are substantially unlike those of SGDS
308 1290.

309 *Petalichnus* (Fig. 4R), attributed to trilobites (Braddy & Almond, 1999), comprises sets
310 of 2–3 elongate to crescentic tracks oriented perpendicular to the trackway axis (Miller, 1880).
311 Anderson (1975b) and Braddy & Almond (1999) diagnosed *Petalichnus* tracks as sometimes
312 bifurcate and occurring in series of 9–12; they further noted that the ichnotaxon needs review.
313 Both track morphology and organization are unlike those of SGDS 1290.

314 *Pterichnus* tracks (Fig. 4S), attributed to isopods, frequently are segmented and are more
315 linear and elongate than any in SGDS 1290. Tracks (or series of segments) are oriented oblique
316 (Types 1 and 2 of Gaillard et al., 2005) or parallel (Types 3 and 4 of Gaillard et al., 2005) to the
317 trackway axis, and approximately symmetrically arranged in two parallel rows (Gaillard et al.,
318 2005; Hitchcock, 1865) that are somewhat closer together than are those of the morphologically
319 similar *Diplichnites*. Types 3 and 4 of Gaillard et al. (2005) morphologically grade into
320 *Diplopodichnus* (Uchman et al., 2011). Hammersburg, Hasiotis & Robison (2018) suggested that
321 *Pterichnus* comprises undertracks of *Lithographus* and is thus a junior synonym of that
322 ichnotaxon. In any case, *Pterichnus* tracks are readily distinguished from SGDS 1290.

323 *Tasmanadia* traces (Fig. 4T) consist of two rows of closely packed, elongate, linear tracks
324 oriented generally perpendicular to the trackway axis; occasionally, tracks overlap at one end,
325 creating narrow, V-shaped structures (Chapman, 1929; Glaessner, 1957). Morphologically, this
326 ichnotaxon resembles *Diplichnites gouldi* Type B and *Umfolozia* (but lacks the organization of
327 the latter), and differs from SGDS 1290 for the same reasons as those ichnotaxa.

328 *Umfolozia* (Fig. 4U), attributed to syncarid and peracarid crustaceans (Lima, Minter &
329 Netto, 2017; Savage, 1971), consists of parallel rows of irregularly shaped to crescentic tracks
330 oriented perpendicular or oblique to the trackway axis that follow a unique repeating pattern
331 (Anderson, 1981; Savage, 1971) unlike anything discernible in SGDS 1290. Aceñolaza &

332 Buatois (1993) noted morphological similarities between *Mirandaichnium*, *Tasmanadia*, and
333 *Umfolozia* and postulated similar track makers.

334 In summary, SGDS 1290 is not a variant of any of these ichnotaxa.

335

336 **Arthropod repichnia possessing one medial impression**

337 Several other arthropod locomotory (walking) ichnotaxa are also readily distinguished
338 from SGDS 1290 by possessing singular medial, rather than paired paramedial, impressions;
339 again, comparison is warranted to ascertain whether or not SGDS 1290 is a morphological
340 variant of such ichnotaxa. These ichnotaxa are *Kouphichnium* (Caster, 1938; Nopcsa, 1923),
341 “*Merostomichnites* isp.” (Hanken & Stormer, 1975), *Oniscoidichnus* (Brady, 1947, 1949),
342 *Palmichnium* (Richter, 1954), *Paleohelcura* (Gilmore, 1926), *Protichnites* (Owen, 1852),
343 *Robledoichnus* (Kozur & Lemone, 1995), *Shalemichnus* (Kozur & Lemone, 1995), *Stiallia*
344 (Smith, 1909), and *Stiaria* (Smith, 1909). As with traces lacking any medial impressions, these
345 ichnotaxa further differ from SGDS 1290 in track morphology. As before, possible track makers
346 are specified only when available.

347 *Kouphichnium* traces (Fig. 4V) are attributed to limulids and occur in a variety of
348 configurations. Tracks in clear *Kouphichnium* walking traces that possess singular medial
349 impressions (e.g., many *K. lithographicum*, but not, for example, *K. minkinensis* [King, Stimson
350 & Lucas, 2019; q.v. Gaillard, 2011a; Shu et al., 2018]) typically occur in sets of up to five in
351 rows oriented oblique to the medial impression and trackway axis; individual tracks range from
352 circular and ellipsoidal to elongate, and can split into anywhere from 2–5 branches at their ends
353 (Caster, 1938, 1944; King, Stimson & Lucas, 2019; Shu et al., 2018). Well-preserved
354 *Kouphichnium* tracks are dissimilar to those of SGDS 1290.

355 Traces referred to as “*Merostomichnites*” (Fig. 4W) and attributed to the eurypterid
356 *Mixopterus* by Hanken & Stormer (1975) consist of three elongate and crescentic tracks in
357 oblique rows on either side of an intermittent medial impression; the tracks increase in size
358 laterally, and some split into two or more branches on one end. In any of these cases, however,
359 the tracks are substantially more complex than those of SGDS 1290.

360 *Oniscoidichnus* tracks (Fig. 4X) are elongate to crescentic, oriented perpendicular or
361 oblique to the trackway axis, closely packed and closely appressed to the single midline
362 impression (Brady, 1947; Davies, Sansom & Turner, 2006). In all these details, *Oniscoidichnus*
363 traces differ markedly from SGDS 1290.

364 Ichnospecies of *Palmichnium* (Fig. 4Y), also attributed to eurypterids, vary in
365 morphology. Generally, they comprise complex sets of tracks lying lateral to a medial
366 impression that can be either continuous or discontinuous. Tracks range in shape from elongate
367 to crescentic to ovoid to chevron shaped, and they generally parallel the trackway axis. The
368 tracks occur in oblique rows in sets of up to four; in some traces, the lateralmost tracks are
369 elongate and curved, while the more medial tracks are linear and oriented parallel to the
370 trackway axis (Braddy & Milner, 1998; Minter & Braddy, 2009; Poschmann & Braddy, 2010;
371 Richter, 1954). Tracks are both more numerous and differently shaped than those of SGDS 1290.

372 *Paleohelcura* (including *Mesichnium* per Braddy [1995] and *Triavestigia* per Kozur,
373 Löffler & Sittig [1994]; possibly a junior synonym of *Stiaria*; Fig. 4Z) traces, attributed to
374 scorpions (Brady, 1947; Davis, Minter & Braddy, 2007), comprise small, circular tracks in sets
375 of three in either rows, triangular arrangements, or checkmark-like patterns that lie external and
376 oblique to the medial impression (Gilmore, 1926; Lagnaoui et al., 2015; Sadler, 1993). This
377 distinctive layout is unlike that of SGDS 1290. Peixoto et al. (2020) attributed traces lacking a
378 medial impression and comprising closely appressed pairs or triplets of mostly elliptical tracks
379 from the Upper Jurassic or Lower Cretaceous of Brazil to a new ichnospecies of *Paleohelcura*
380 and attributed them to a pterygote insect track maker. Tracks in this ichnospecies are arranged in
381 rows oriented oblique to the trackway axis, and track sets in this ichnospecies lie close to the
382 midline. This morphology is also unlike that of SGDS 1290.

383 *Protichnites* traces (Fig. 4AA) comprise thick, often segmented medial impressions
384 (sometimes absent except on trackway turns) flanked by oppositely arranged, subcircular to
385 ellipsoidal to irregularly shaped tracks with varying orientations to the trackway axis (Burton-
386 Kelly & Erickson, 2010; Collette, Gass & Hagadorn, 2012; Hagadorn & Seilacher, 2009). They
387 differ substantially from the tracks of SGDS 1290.

388 *Robledoichnus* tracks (Fig. 4BB), attributed to flying insects, resemble tracks of
389 *Eisenachichnus* but possess a discontinuous, faint medial trace consisting entirely of periodic, V-
390 shaped marks flanked by asymmetrical pairs of tracks. On one side, the tracks are short, tapered,
391 and oriented oblique to the trackway axis; on the other side, the tracks are longer and crescentic,
392 oriented closer to perpendicular to the trackway axis (Kozur & Lemone, 1995). Lucas et al.
393 (2005b) considered *Robledoichnus* a probable junior synonym of *Paleohelcura* or *Stiaria*, and
394 the ichnotaxon differs from SGDS 1290 for similar reasons as those ichnotaxa, in addition to the
395 asymmetry.

396 *Shalemichnus* traces (Fig. 4CC), for which only half a trackway is known, consist of a
397 straight medial impression punctuated at intervals by V-shaped marks. This impression is
398 flanked by sets of three tapered tracks in straight rows oriented perpendicular to the trackway
399 axis; individual tracks have their long axes parallel to the trackway axis (Kozur & Lemone,
400 1995). Minter & Braddy (2009) considered *Shalemichnus* a junior synonym of *Stiaria*. The
401 tracks of *Shalemichnus* bear some similarity to those of SGDS 1290, but the paramedial
402 impressions of SGDS 1290 lack the V-shaped markings of the medial impression of
403 *Shalemichnus*.

404 *Stiallia* traces (Fig. 4DD) consist of paired rows of long, linear impressions parallel or
405 slightly oblique to the trackway axis and that frequently overlap. *Stiallia pilosa* lacks any medial
406 or paramedial impressions, but *Stiallia* (*Carrickia* of Smith [1909]) *berriana* possesses a medial
407 row of crescentic to chevron-shaped marks (Smith, 1909; Walker, 1985). Pollard (1995)
408 suggested that *Stiallia* could be an arthropod swimming, rather than a walking, trace, though it
409 also resembles traces made by bristletail insects walking in highly saturated mud (Getty et al.,
410 2013: fig. 6F, G). *Stiallia* tracks are markedly unlike those of SGDS 1290.

411 *Stiaria* tracks (including some ichnospecies of *Danstairia* of Smith [1909]; Fig. 4EE)),
412 attributed to scorpionids (Braddy, 2003; Lucas, Lerner & Voigt, 2013) and monuran insects
413 (Genise, 2017; Kopaska-Merkel & Buta, 2013), are oppositely situated groups of 2–4 generally
414 circular to tapered tracks in a linear to crescentic arrangement lying roughly perpendicular to the
415 trackway axis (Walker, 1985). In some specimens and ichnospecies of *Stiaria*, the singular
416 medial impression actually varies in position, meandering from medial to almost lateral to their
417 tracks (Fillmore, Lucas & Simpson, 2012: fig. 26d; Walker, 1985: fig. 5b, c). In some
418 Mississippian specimens from Pennsylvania, the medial impression is flanked by thin,
419 discontinuous, but closely appressed paramedial imprints (Fillmore, Lucas & Simpson, 2012: fig.
420 26d–g). Track arrangement alone differentiates *Stiaria* from SGDS 1290. Genise (2017) asserted
421 that *Stiaria* should be considered a junior synonym of *Siskemia* (the latter has page priority over
422 the former).

423 As with locomotory traces lacking medial impressions, SGDS 1290 is not a variant of
424 any of these ichnotaxa.

425

426 **Arthropod repichnia possessing three or more medial and paramedial impressions**

427 *Mitchellichnus* (Fig. 4FF), attributed to archaeognathan insects (Getty et al., 2013), is
428 distinguished from SGDS 1290 by possessing three medial impressions (Walker, 1985).
429 *Mitchellichnus* tracks are complex, comprising two distinct types and arrangements. An inner
430 set, lying close to the medial impressions, consists of apparently elongate tracks in sets of up to
431 six that lie parallel to slightly oblique to the trackway axis; an outer set consists of larger,
432 amorphous impressions (Walker, 1985). Tracks are thus more numerous in *Mitchellichnus* than
433 in SGDS 1290, and the tracks differ in arrangement and morphology. Like *Stiaria*, Genise (2017)
434 asserted that *Mitchellichnus* should be considered a junior synonym of *Siskemia*.

435 *Keircalia* (Fig. 4GG) is distinguished from SGDS 1290 by possessing four medial
436 impressions (Smith, 1909; Walker, 1985). *Keircalia* tracks are crescentic to irregularly shaped,
437 generally are oriented perpendicular to the trackway axis, and have no discernible arrangement
438 (Walker, 1985). Both track morphology and organization are unlike those of SGDS 1290.

439

440 **Arthropod repichnia possessing paired paramedial impressions**

441 A few ichnotaxa, as well as some experimentally produced tracks of extant arthropods,
442 resemble SGDS 1290 by possessing paired paramedial impressions in at least some specimens.
443 Such ichnotaxa are *Danstairia vagusa* (Smith, 1909), *Glaciichnium* (Walter, 1985),
444 *Warvichnium* (Walter, 1985), and *Siskemia* (Smith, 1909); similar extant traces include those
445 made by notostracans (Trusheim, 1931) and crayfish (Fairchild & Hasiotis, 2011).

446 *Danstairia vagusa* (Fig. 4HH) possesses intermittent, thin, linear imprints that do not
447 always parallel their accompanying trackways. Tracks are circular to triangular, generally have
448 their long axes perpendicular to the trackway axis, and lack any coherent layout (Walker, 1985),
449 unlike those of SGDS 1290. *D. vagusa* somewhat resembles *Keircalia* traces, but its tracks are
450 spaced more widely apart.

451 *Glaciichnium* traces (Fig. 4II), which resemble traces made by isopods (Gibbard &
452 Stuart, 1974; Lima, Minter & Netto, 2017; Uchman, Kazakauskas & Gaigalas, 2009; Uchman et
453 al., 2011), comprise 1–3 elongate, linear tracks (“bars” that are divided into segments [Uchman,
454 Kazakauskas & Gaigalas, 2009]) that lie oblique to the trackway axis and are staggered on either
455 side of that axis, unlike the tracks of SGDS 1290; their linear, serial but discontinuous
456 paramedial imprints are widely spaced, consistently abutting the medial ends of the tracks
457 (Walter, 1985), farther apart than those of SGDS 1290. Walter (1985) and Lima et al. (2015)
458 described the paramedial imprints in Brazilian specimens as comprising successive pairs of C-
459 shaped imprints rather than strictly linear structures, further unlike SGDS 1290. Some
460 *Glaciichnium* traces also possess a medial imprint as well (Uchman, Kazakauskas & Gaigalas,
461 2009; Walter, 1985).

462 *Warvichnium* traces (Fig. 4JJ) are complex, comprising pairs to multiple sets of linear,
463 discontinuous medial and paramedial imprints flanked by varying numbers of linear to crescentic
464 tracks in two or more sets: an inner set, close to the medial imprints, that are oriented slightly
465 oblique to the trackway axis, and an outer set oriented closer to perpendicular to the trackway
466 axis (Walter, 1985), quite unlike SGDS 1290. Getty (2020) noted similarities between
467 *Warvichnium* and subaqueous *Bifurculapes*.

468 Among described arthropod repichnia, SGDS 1290 architecturally most closely
469 resembles ichnospecies of *Siskemia* by possessing discreet, compact (not linear) tracks and track
470 sets flanking paired paramedial impressions. Three ichnospecies of *Siskemia* are presently
471 recognized (Walker, 1985):

- 472 • *S. bipediculus* (Fig. 4KK) comprises small, circular tracks in closely appressed pairs
473 (occasionally trios) in rows oriented perpendicular or slightly oblique to the trackway axis;
474 the pairs are spaced apart at approximately regular intervals and evenly distant from the
475 uniformly straight and narrow paramedial impressions (Walker, 1985). The paramedial
476 impressions lie close to the midline axis, well away from their adjacent tracks (the average
477 ratio of the distance between a paramedial impression and its flanking track to the distance
478 between the medial edges of the paramedial impressions is 1.34; see Supplemental
479 Material).
- 480 • *S. elegans* (Fig. 4LL) has similarly shaped tracks in sets of up to four tracks; the sets
481 similarly lie well away from the likewise straight, narrow, and closely appressed
482 paramedial impressions (the average ratio of the distance between a paramedial
483 impression and its flanking track to the distance between the medial edges of the
484 paramedial impressions is 1.75; see Supplemental Material). *S. bipediculus* and *S. elegans*
485 differ primarily in the orientations of their track rows to the trackway axis and the
486 continuities and thicknesses of their paramedial impressions (Walker, 1985), though these
487 could be behavioral and/or substrate-driven variants.
- 488 • *Siskemia latavia* (Fig. 4MM; called “*lata-via*” by Smith [1909] and Walker [1985], but the
489 ICZN does not permit hyphens in genus or species names) tracks comprise tapered or
490 ovoid tracks arranged in roughly triangular sets of three. Most individual tracks have their

491 long axes oriented oblique to the trackway axis; track sets have varying orientations to the
492 trackway axis. Tracks in individual *S. latavia* sets usually are spaced well apart from each
493 other; rarely do two individual tracks in a set contact one another. *S. latavia* tracks lack the
494 regular spacing and arrangements of those of *S. elegans* and *S. bipediculus*, and occasional
495 individual tracks lie close to the paramedial impressions, farther medially than in either of
496 the other two *Siskemia* ichnospecies (the average ratio of the distance between a
497 paramedial impression and its flanking track to the distance between the medial edges of
498 the paramedial impressions is 0.80; see Supplemental Material). Each paramedial
499 impression of *S. latavia* is slightly wider than those of the other two *Siskemia* ichnospecies
500 (probably a function of the larger overall size of specimens attributed to this
501 ichnospecies), but retain the close appression to the trackway midline and the uniform
502 straightness.

503 The tracks of SGDS 1290 vary more in morphology than those of any known *Siskemia*
504 ichnospecies, but grossly share their layout. Tracks in all three *Siskemia* ichnospecies have a
505 staggered distribution, similar to, but less pronounced than, that of SGDS 1290. SGDS 1290
506 differs most markedly from any of the three *Siskemia* ichnospecies in the morphology and
507 positions of the paramedial impressions: in SGDS 1290, the impressions vary in thickness along
508 their lengths and undulate, in contrast to the thin, straight impressions of all three *Siskemia*
509 ichnospecies. Additionally, the impressions in SGDS 1290 lie farther apart than those of the
510 three *Siskemia* ichnospecies. In fact, all of the ichnospecies of *Siskemia* erected by Smith (1909),
511 as well as both specimens later attributed to this ichnogenus (e.g., Getty et al., 2013; McNamara,
512 2014; Pollard, Steel & Undersrud, 1982) and *Siskemia*-like traces made by extant,
513 archaeognathan insects (Getty et al., 2013), have such thin, linear, closely appressed paramedial
514 impressions (sometimes offset toward one side of the trackway). The only time when
515 archaeognathan traces approach the paramedial impression spacing of SGDS 1290 is when both
516 abdominal styli (laterally) and gonostyli (medially) of the trace makers register impressions and
517 imprints in wet mud, producing two sets of paramedial traces (Getty et al., 2013: fig. 6K, L), but
518 even then the linear, lateralmost paramedial impressions still do not resemble the thick,
519 undulating impressions of SGDS 1290. Simultaneously, in such wet mud, archaeognathan tracks
520 are elongate and oriented oblique to the trackway axis, unlike those in SGDS 1290. In total,
521 SGDS 1290 does not fit neatly into any known *Siskemia* ichnospecies and does not seem to be an
522 archaeognathan insect trace.

523 Among traces made by extant arthropods, SGDS 1290 bears similarities to traces made
524 by both notostracans and crayfish. Interpretive drawings of experimental traces made by
525 notostracans figured by Trusheim (1931) depict elongate, crescentic, or tapered tracks oriented
526 perpendicular to paramedial impressions; the tracks are arranged oppositely, unlike those of
527 SGDS 1290. Additionally, the thin, linear paramedial impressions figured by Trusheim (1931)
528 lie so far from the trackway axis that they often contact their accompanying tracks, a
529 phenomenon that only occurs in SGDS 1290 near L4 and R4, where the lateral margins of its
530 undulating paramedial impressions meander particularly far laterally. Tasch (1969) noted,

531 though, that the drawings presented by Trusheim (1931) were misleading compared to his own
532 experimentally generated notostracan traces. However, he described the morphologies of his
533 notostracan tracks only as “minute en echelon stripes” (Tasch, 1969: 327), which does not
534 adequately specify how they differed from those of Trusheim (1931); track details are impossible
535 to discern in his lone photographic figure (Tasch, 1969: pl. 1.2). Gand et al. (2008) also
536 conducted neoichnological experiments with notostracans, recovering locomotory traces that
537 were less orderly than those illustrated by Trusheim (1931) (Gand et al., 2008: figs. 16.1, 16.2,
538 17.1). Their extant notostracan tracks comprised multiple tracks with rather chaotic distributions
539 lateral to their paramedial impressions, unlike the regular distribution seen in SGDS 1290. Gand
540 et al. (2008) found their notostracan traces to fall within the “etho-morphotype” of *Acripes*, as
541 exemplified by their novel ichnospecies *A. multiformis* from the Permian of France. (Linck
542 [1943] and Pollard [1985] also referred *Acripes* [*Merostomichnites* of Linck, 1943] tracks to
543 notostracans, but not based on neoichnological experiments.) *A. multiformis* traces, unlike classic
544 *Acripes* (Matthew, 1910), possess paramedial imprints, albeit inconsistently. Hammersburg,
545 Hasiotis & Robison (2018), Häntzschel (1975), Miller (1996), and Pemberton, MacEachern &
546 Gingras (2007) all supported classic *Acripes* as a junior synonym of *Diplichnites*; the issue of
547 synonymy is beyond the scope of this paper, but we note at least that the tracks in fossils that
548 Gand et al. (2008) called *A. multiformis* differ from SGDS 1290 in the same ways that
549 *Diplichnites* tracks do (see above). Lastly, Knecht et al. (2009: figs. 5, 6) also illustrated traces
550 made by extant notostracans (Fig. 4NN), which are “tidier” than those of Gand et al. (2008) and
551 resemble those of classic *Acripes* and *Diplichnites*, albeit with discontinuous paramedial and
552 curvilinear lateral (external) imprints. The tracks in these traces comprise irregular, ellipsoidal
553 sets oriented oblique to the trackway axis, unlike those of SGDS 1290. In total, SGDS 1290 is
554 unlikely to be a notostracan trace.

555 Fairchild & Hasiotis (2011) conducted neoichnological experiments with crayfish to
556 examine their locomotory traces. These traces varied in morphology depending on substrate
557 conditions (sediment grain size and saturation) and slope; in general, when clearest, they consist
558 of sets of 1–4 circular, tapering, ellipsoidal, or elongate tracks, occasionally of different sizes,
559 that are oriented parallel to the trackway axis and that lie lateral to a pair of undulating, variably
560 thick paramedial impressions that lie closer to their flanking tracks than to each other (Figs. 3D,
561 4OO). Morphologically, the tracks and impressions match those of SGDS 1290, although the
562 tracks made by extant crayfish often are larger than those of SGDS 1290 when produced in dry
563 substrate (Fairchild & Hasiotis, 2011: fig. 9). Track sets in extant crayfish traces have complex
564 arrangements: when comprised of multiple tracks, they frequently are in rows oriented
565 perpendicular to the trackway axis, but sometimes rows are oblique to the trackway axis. When
566 fewer tracks are preserved, sets can appear to lie in single, parallel rows on either side of the
567 paramedial impressions. Track positions can be opposite to staggered to alternating, also as in
568 SGDS 1290. In both track and paramedial impression morphology, as well as in overall trace
569 architecture, SGDS 1290 strongly resembles crayfish traces made in damp silt and clay
570 (Fairchild & Hasiotis, 2011: fig. 2e, f), dry and saturated, very fine-grained sand (Fig. 3D;

571 Fairchild & Hasiotis, 2011: fig. 4a, b, e, f), dry and damp, fine-grained sand (Fairchild &
572 Hasiotis, 2011: fig. 5a–d), and saturated medium sand (Fairchild & Hasiotis, 2011: fig. 6e, f).
573 SGDS 1290 is preserved in, and was presumably registered in, a fine-grained sandstone,
574 lithologically matching one set of experimental conditions in Fairchild & Hasiotis (2011).
575 However, SGDS 1290 is not as detailed as many of the experimentally generated crayfish traces
576 in comparable sediments. This could indicate one or more things: the fossil could be a slight
577 overtrack (*sensu* Bertling et al., 2006) rather than a direct natural cast; the trace maker may have
578 been partly buoyant; and/or trace-maker behavior and sediment consistency combined such that
579 the lighter limbs did not impress as deeply as the heavier tail.

580

581 Discussion

582 Trace maker

583 The stronger resemblance of SGDS 1290 to experimental crayfish locomotion traces than
584 to any known ichnotaxon, or other documented extant arthropod trace, implies a crayfish or
585 crayfish-like maker for the fossil. Whether SGDS 1290 had a crayfish-like or an actual crayfish
586 maker depends on whether the term “crayfish” is used to refer to members of a monophyletic
587 clade (Parastacidae + (Astacidae + (Cambaridae + Cricoidoscelosidae)); Karasawa, Schweitzer
588 & Feldmann, 2013) of freshwater (and brackish water if *Protastacus* is included, *sensu*
589 Karasawa, Schweitzer & Feldmann [2013]) lobsters, or, more broadly, to any freshwater, lobster-
590 like crustacean regardless of phylogenetic position, which presumes that more than one crayfish-
591 like lineage colonized terrestrial environments in the past. Here we use the term in the
592 monophyletic sense: true crayfish comprise Astacida (*sensu* Karasawa, Schweitzer & Feldmann,
593 2013; Schram & Dixon, 2004; Shen, Braband & Scholtz, 2015). Whether or not the maker of
594 SGDS 1290 was a true crayfish is unclear: the oldest undisputed crayfish body fossils are from
595 the Early Cretaceous (Garassino, 1997; Martin et al., 2008; Shen, 2003; Taylor, Schram & Shen,
596 1999), although unnamed, Late Jurassic specimens from western Colorado also have been called
597 crayfish (Hasiotis, Kirkland & Callison, 1998). Despite this, a Triassic or earlier origin for true
598 crayfish has been hypothesized frequently (Breinholt, Pérez-Losada & Crandall, 2009; Crandall
599 & Buhay, 2008; Porter, Pérez-Losada & Crandall, 2005; Schram, 2001; Schram & Dixon, 2004;
600 Wolfe et al., 2019) and possibly substantiated by fossil burrows referred to crayfish from the
601 Early Permian (Hembree & Swaninger, 2018) and Late Permian–Early Triassic (Baucon et al.,
602 2014; Hasiotis & Mitchell, 1993).

603 Several Late Triassic body-fossil specimens also have been reported as crayfish (Hasiotis,
604 1995; Hasiotis & Mitchell, 1993; Miller & Ash, 1988; Olsen & Huber, 1997; Santucci &
605 Kirkland, 2010), but the identities of these specimens as true astacidans has not been established.
606 Miller & Ash (1988) placed a Late Triassic specimen from Petrified Forest National Park,
607 Arizona in *Enoploclytia*, which is an erymid lobster, not an astacidan, genus. That generic
608 placement subsequently has been contested (Amati, Feldmann & Zonneveld, 2004; Schweitzer et
609 al., 2010; Urreta, 1989), so the specimen needs detailed restudy, but if it pertains to Erymididae
610 rather than Astacida, then it indicates that a lineage of erymid lobsters colonized terrestrial

611 environments, possibly before true (monophyletic) crayfish. Some older analyses (reviewed in
612 Rode & Babcock, 2003) postulated crayfish origins within Erymidae, but more recent
613 phylogenetic analyses (Devillez, Charbonnier & Barreil, 2019; Karasawa, Schweitzer &
614 Feldmann, 2013; Rode & Babcock, 2003; Schram & Dixon, 2004; Stern & Crandall, 2015) have
615 recovered (a frequently paraphyletic) Erymidae with members at varying distances from
616 Astacida. If those hypothesized phylogenetic relationships are correct, then no erymid can, in a
617 monophyletic sense, be considered a crayfish, even if it was a freshwater taxon. But regardless of
618 semantics or phylogenetic relationships, crayfish and erymid morphological similarities suggest
619 that their locomotory traces might be indistinguishable, making a definitive attribution of SGDS
620 1290 impossible.

621 A crayfish or crayfish-like trace maker for SGDS 1290 is tenable both chronologically
622 and ecologically. As mentioned above, multiple crayfish-like morphotypes have been found in
623 the Upper Triassic Chinle Formation of Arizona and Utah (Hasiotis, 1995; Miller & Ash, 1988;
624 Santucci & Kirkland, 2010). The Moenave Formation overlies the Chinle Formation in
625 southwestern Utah, so crayfish or crayfish-like decapods plausibly could have been present in
626 and around freshwater Lake Whitmore both geographically and stratigraphically. As-yet-
627 undescribed, crayfish or crayfish-like body fossils also have been recovered from lacustrine
628 sediments of the uppermost Triassic Chatham Group (Newark Supergroup) in North Carolina
629 (Olsen & Huber, 1997), attesting to how widespread such arthropods were in terrestrial
630 environments in North America even prior to the Jurassic.

631

632 **Ichnotaxonomy**

633 To date, no fossil arthropod locomotory ichnotaxon has been attributed definitively to a
634 crayfish or crayfish-like trace maker. Several such ichnotaxa have been attributed, for various
635 reasons, to crustaceans (e.g., Braddy, 2003; Gand et al., 2008; Lima, Minter & Netto, 2017;
636 Pirrie, Feldmann & Buatois, 2004; Savage, 1971; Walker, Holland & Gardiner, 2003);
637 additionally, some purported crustacean tracks have not been assigned to particular ichnotaxa
638 (e.g., Imaizumi, 1967; Karasawa, Okamura & Naruse, 1990; Matsuoka et al., 1993), including
639 mortichnial trackways leading to *Eryma*, *Eryon*, and *Mecochirus* lobster body fossils from the
640 marine, Upper Jurassic lithographic limestones of Germany (Glaessner, 1969: fig. 243A;
641 Seilacher, 2008: fig. 2; Viohl, 1998: fig. 6). None of these German taxa are crayfish, though
642 morphological similarities of *Eryma* and *Mecochirus* to crayfish might mean that they would
643 have produced indistinguishable locomotory ichnofossils during normal, forward locomotion.
644 None of their traces have been granted detailed description or ichnotaxonomic assignment.

645 Only three locomotory ichnotaxa have been attributed specifically to crayfish. First,
646 Heidtke (1990) erected *Pollichianum repichnum* for Early Permian ichnofossils from Germany
647 that he attributed to the “crawfish” (in the English abstract; “Krebses” in the German abstract)
648 *Uronectes fimbriatus*, also from the Early Permian of Germany. However, *Uronectes* has long
649 been recognized as a syncarid (Brooks, 1962; Calman, 1934; Perrier et al., 2006), not an
650 astacidan, or even a decapod, so the term appearing in the English abstract likely is a simple

651 translation error. Furthermore, however, *P. repichnum* is not differentiable from the resting trace
652 (cubichnion) *P. cubichnum* (O'Brien, Braddy & Radley, 2009) and therefore is a junior synonym
653 and is not a locomotory trace. In any case, *Pollichianum* is morphologically quite unlike both
654 SGDS 1290 and experimentally generated crayfish traces (Fairchild & Hasiotis, 2011). Second,
655 Bolliger & Gubler (1997) hypothesized that their novel, early Miocene ichnospecies *Hamipes*
656 *molassicus* was made by a buoyed (presumably swimming) crayfish. Getty (2018) referred these
657 specimens to *Conopsoides*; later, Getty & Burnett (2019) suggested that at least some of the
658 specimens may pertain to *Acanthichnus*, and they differ from SGDS 1290 for the same reasons
659 outlined above for *Acanthichnus*. Third, de Gibert et al. (2000) attributed Early Cretaceous,
660 Spanish specimens that they assigned to *Hamipes didactylus* to crayfish. Getty (2018) attributed
661 these tracks to *Bifurculapes* and maintained a crustacean track maker for *H. didactylus*, but was
662 not more specific. However, neither *Bifurculapes* nor *Hamipes* resemble experimentally
663 generated crayfish traces (Fairchild & Hasiotis, 2011), or any of the mortichnial decapod traces,
664 and thus are unlikely to have been made by a crayfish-like decapod, at least while walking.
665 Lastly, we also note that unnamed trackways attributed to crayfish from the Upper Triassic
666 Chinle Formation of Utah were mentioned, but not described, by Hasiotis (1991); Fairchild &
667 Hasiotis (2011) did not note whether or not these were similar to their experimentally generated
668 traces. Additionally, an unnamed "crayfish locomotion trace" was figured, but not described, by
669 Rainforth & Lockley (1996: fig. 1b); it does not resemble either SGDS 1290 or experimentally
670 generated crayfish traces (Fairchild & Hasiotis, 2011).

671 As detailed above, SGDS 1290 does not fit neatly into any existing ichnospecies of
672 *Siskemia*. Whether or not to place it in a novel ichnospecies, or even ichnogenus, is, therefore, an
673 open question. Bertling et al. (2006), Gaillard (2011b), and Minter, Braddy & Davis (2007)
674 provided solid criteria for the erection of new ichnotaxa, the latter particularly for arthropods.
675 One criterion is that a new ichnotaxon ideally should be represented by a substantial number of
676 specimens that demonstrate behavioral and substrate-based morphological variation; this
677 prevents erecting several ichnotaxa for minor, readily explained variations in trace morphology.
678 SGDS 1290, as a singular specimen, certainly does not meet that criterion, but Minter, Braddy &
679 Davis (2007) also allowed that truly unique morphologies exhibited by singular specimens can
680 support an ichnotaxon. In terms of uniqueness, another criterion is whether or not a new
681 morphotype falls onto a continuum, established or hypothetical, of morphologies within an
682 established ichnotaxon. SGDS 1290 is closest morphologically to ichnospecies of *Siskemia*, but
683 has several distinctions from any established ichnospecies therein, particularly the thick and
684 undulating paramedial impressions and the wider spacing between the paramedial impressions
685 and consequent closer appression of the paramedial impressions to the tracks: the average ratios
686 of the distance between a paramedial impression and its flanking track to the distance between
687 the medial edges of the paramedial impressions are 1.34 for *S. bipediculus*, 1.75 for *S. elegans*,
688 and 0.80 for *S. latavia* compared to 0.31 for SGDS 1290 (see Supplemental Material). No
689 published specimen of *Siskemia* demonstrates the features of SGDS 1290; nor do *Siskemia*-like
690 traces made by archaeognathan insects in experimental conditions (Getty et al., 2013). Thus,

691 SGDS 1290 does not appear to fall within the established *Siskemia* continuum. The greater
692 prominence (depth) of the paramedial impressions of SGDS 1290 than their associated tracks
693 suggests either a trace maker with heavier tail elements than the gonostyli of an archaeognathan
694 insect or an archaeognathan trace maker with unusually large styli adopting an unusual posture
695 (possibly partly buoyant), flexing its caudal region downward to create deep styli impressions
696 but not deep track impressions. We consider the latter unlikely; thus, SGDS 1290 does not
697 appear to fall within a hypothetical *Siskemia* continuum, either. However, SGDS 1290 falls
698 within the continuum of trace morphologies made by extant crayfish in experimental conditions
699 (Fairchild & Hasiotis, 2011). No philosophical basis has been established for the recognition of
700 novel ichnotaxa on the basis of comparison with traces made by extant organisms; only by
701 comparison with fossil ichnotaxa because extant traces cannot be the basis for an ichnotaxon
702 (Bertling et al., 2006; International Commission on Zoological Nomenclature, 1999).

703 SGDS 1290 clearly is morphologically distinctive. Lacking a sufficient number of
704 specimens with which to determine ranges of morphological variation, however, erecting a new
705 ichnogenus for it clearly is unwarranted. Yet we feel that its unique morphology warrants
706 ichnotaxonomic distinction. Given its distant similarity to *Siskemia* ichnospecies, we therefore
707 herein place it in a new ichnospecies of that ichnogenus.

708

709 **Systematic Ichnology**

710

711 **Ichnofamily** Protichnidae Uchman, Gaździcki & Błazejowski 2018

712

713 **Ichnogenus** *Siskemia* Smith 1909

714

715 **Type Ichnospecies** *Siskemia elegans* Smith 1909

716 **Diagnosis.** Trace consisting of parallel rows of grouped tracks on either side of two parallel,
717 paramedial impressions. Each group of tracks consists of up to four imprints arranged in series,
718 transversely or obliquely to the midline of the trackway (following Walker, 1985). Walker
719 (1985) further specified that *Siskemia* was diagnosed by paramedial impressions with maximum
720 widths of 0.5 mm, but following Bertling et al. (2006), size is not a suitable ichnotaxobase.

721

722 **Ichnospecies** *Siskemia eurypyge* isp. nov.

723 Figure 3A, B

724

725 **Diagnosis.** Two parallel, undulating, paramedial impressions flanked externally by closely
726 appressed sets of 1–3 small, ovoid to tapered to elongate tracks; tapered and elongate tracks have
727 long axes parallel or oblique to the trackway axis. Track sets are oriented perpendicular to the
728 trackway axis. Left and right tracks are arranged in a staggered to alternating pattern. Paramedial
729 impressions are mediolaterally thick, but discontinuous, tapering out of existence briefly in some
730 places. Impressions are gently undulating (low amplitude). The paramedial impressions lie far

731 from the trackway axis, generally closer to (and sometimes in contact with) the tracks than to the
732 midline axis or each other.

733 **Holotype.** Natural cast specimen SGDS 1290, St. George Dinosaur Discovery Site, St. George,
734 Utah, USA.

735 **Type locality.** “Bug Crossing Quarry,” SGDS Loc. 87, St. George Dinosaur Discovery Site, St.
736 George, Washington County, Utah, USA (Fig. 1).

737 **Stratigraphy.** “Top Surface” of Johnson Farm Sandstone Bed (unit 40 of Kirkland et al., 2014),
738 Whitmore Point Member, Moenave Formation (Fig. 2). Hettangian, Lower Jurassic.

739 **Derivation of name.** From the Greek ευρυς (eurys), meaning “broad” or “wide,” and πυγή
740 (pyge), meaning “rump.” The ichnospecies name refers to the wider spacing between the
741 paramedial impressions, made by the rear end of the trace maker, than those of other *Siskemia*
742 ichnospecies.

743

744 Conclusions

745 SGDS 1290, from the Lower Jurassic (Hettangian) Whitmore Point Member of the
746 Moenave Formation, consists of two paramedial impressions that are flanked by staggered to
747 alternating sets of tapered or elongate tracks. The traces closely resemble those made by extant
748 crayfish (Fairchild & Hasiotis, 2011) and are similar in components to traces placed in the
749 ichnogenus *Siskemia* (Smith 1909; Walker, 1985). In previously recognized *Siskemia*
750 ichnospecies, the paramedial impressions are thin, relatively straight, and closely appressed to
751 the trackway axis. But in SGDS 1290, paramedial impressions have the opposite morphology
752 and arrangement: they are thick and lie closer to their tracks than the medial axis of the trackway.
753 SGDS 1290 paramedial impressions also undulate, which is not seen in any previously known
754 *Siskemia* ichnospecies. Thus, we erect a new ichnospecies, *Siskemia eurypyge*, to house SGDS
755 1290 and as-yet undiscovered traces with this morphology and arrangement.

756 Placing SGDS 1290 in *Siskemia* extends the known range of the ichnogenus into the
757 Early Mesozoic. All other reported occurrences of the ichnogenus are Paleozoic in age: Early
758 Silurian (McNamara, 2014; Trewin & McNamara, 1994), Late Silurian (Davies, Sansom &
759 Turner, 2006), Early Devonian (Pollard, Steel & Undersrud, 1982; Pollard & Walker, 1984;
760 Smith, 1909; Walker, 1985), and Pennsylvanian (Getty et al., 2013). However, age should not be
761 a factor in ichnotaxonomy (Bertling et al., 2006). At least some Paleozoic *Siskemia* traces likely
762 were made by archaeognathan insects (Getty et al., 2013), which are extant and for which body
763 fossils are known as early as the Devonian. Based on their similarity to traces made by extant
764 crayfish (Fairchild and Hasiotis, 2011), *S. eurypyge* likely was made by a crayfish or crayfish-
765 like crustacean, for which body fossils are known as early as the Late Triassic and which also are
766 extant. Thus, *Siskemia* ispp. traces would be expected to occur from Early Silurian to Recent, but
767 thus far have not been documented except for the occurrences noted above.

768 SGDS 1290 expands the ichnological record of crayfish and crayfish-like animals to
769 include repichnia in addition to domichnia. Fossil burrows (*Camborygma* ispp.), usually
770 attributed to crayfish, are well known at some sites and in some formations (e.g., Hasiotis, 1995;

771 Hasiotis & Bown, 1996; Hasiotis & Honey, 1995; Hasiotis & Mitchell, 1993; Hasiotis, Kirkland
772 & Callison, 1998; see Schram & Dixon [2004] concerning pre-Cretaceous examples), attesting to
773 the presences—and, in some places, abundances—of crayfish and/or crayfish-like taxa in
774 Mesozoic–Cenozoic freshwater paleoecosystems. Yet locomotion traces made by these
775 burrowers oddly have never before been documented as ichnofossils, possibly because they
776 infrequently venture far from their burrows in substrates suitable for registering locomotory
777 traces, as with modern crayfish (Martin, 2013). SGDS 1290 is the first documented locomotory
778 ichnofossil made by a freshwater crayfish or crayfish-like organism, as well as the first fossil
779 evidence of such a taxon in the Lower Jurassic Moenave Formation and indeed the Early Jurassic
780 of the southwestern US. The absence of *Camborygma* burrows in the Moenave Formation that
781 would have been made by the SGDS 1290 trace maker is puzzling, and may be a consequence of
782 a lack of recognition; alternatively, the producer of SGDS 1290 was not a burrower.

783

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792

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1361 **FIGURE CAPTIONS**

1362

1363 **Figure 1: Location of the St. George Dinosaur Discovery Site at Johnson Farm.** (A) Map
 1364 showing the location of the St. George Dinosaur Discovery Site in St. George, Utah. (B)
 1365 The museum at the St. George Dinosaur Discovery Site. Photograph by Jerald D. Harris.

1366 **Figure 2: Stratigraphic section at and immediately around the St. George Dinosaur**
 1367 **Discovery Site (SGDS) in St. George, Utah.** Arthropod locomotory trackway SGDS
 1368 1290 comes from the “Top Surface Tracksite” layers of the Johnson Farm Sandstone Bed
 1369 (red arrows).

1370 **Figure 3: Arthropod locomotory ichnofossil SGDS 1290.** (A) Photograph of specimen; scale
 1371 in mm. (B) Schematic diagram of specimen. Arrow indicates direction of movement. L =
 1372 left track; R = right track; Ov = overtracks. Numbers indicate position of tracks in
 1373 sequence from posterior to anterior. Photograph by Andrew R.C. Milner. (C) Schematic
 1374 diagram of specimen, showing examples of how measurements were taken. IEW =
 1375 external width between paramedial impressions; IIW = internal width between
 1376 paramedial impressions; IW = paramedial impression width; TL = track length; TTI =
 1377 distance between track and adjacent paramedial impression; TW = track width. (D)
 1378 Photograph of extant crayfish walking trace made in saturated, very fine sand for
 1379 comparison to fossil in (A); scale in cm. Arrow indicates direction of travel; photograph
 1380 rotated to have the same orientation as (A). Modified from Fairchild & Hasiotis (2011:
 1381 fig. 4E); reproduced with permission from SEPM.

1382 **Figure 4: Schematic diagrams of locomotory (presumably walking) ichnofossils attributed**
 1383 **to arthropods (A–MM) and extant walking traces made by arthropods (NN–OO)**
 1384 **compared to SGDS 1290 (PP).** Diagrams not to scale. (A) *Acanthichnus cursorius*
 1385 (traced from Hitchcock, 1858). (B) *Asaphoidichnus trifidus* (traced from Miller, 1880).
 1386 (C) *Bifurculapes laqueatus* (traced from Getty, 2016). (D) *Coenobichnus currani* (traced
 1387 from Walker, Holland & Gardiner, 2003). (E) *Copeza triremis* (traced from Hitchcock,
 1388 1858). (F) *Danstairia congesta* (traced from Walker, 1985). (G) *Diplichnites aenigma*
 1389 (traced from Dawson, 1873). (H) *Diplichnites gouldi* Type A (traced from Trewin and
 1390 McNamara, 1994). (I) *Eisenachichnus inaequalis* (traced from Kozur, 1981). (J)
 1391 *Foersterichnus rossensis* (traced from Pirrie, Feldmann & Buatois, 2004). (K) *Hamipes*
 1392 *didactylus* (traced from Getty, 2018). (L) *Laterigradus lusitanicus* (traced from de
 1393 Carvalho et al., 2016). (M) *Lithographus hieroglyphicus* (traced from Hitchcock, 1858).
 1394 (N) *Maculichna varia* (traced from Anderson, 1975a). (O) *Merostomichnites*
 1395 *narrangansettensis* (traced from Packard, 1900). (P) *Mirandaichnium famatinense* (traced
 1396 from Aceñolaza, 1978). (Q) *Octopodichnus didactylus* (traced from Gilmore, 1927). (R)
 1397 *Petalichnus multipartatus* (Miller, 1880). (S) *Pterichnus centipes* (traced from Hitchcock,
 1398 1865). (T) *Tasmanadia twelvetreesi* (traced from Glaessner, 1957). (U) *Umfolozia*
 1399 *sinuosa* (traced from Anderson, 1981). (V) *Kouphichnium lithographicum* (traced from
 1400 Gaillard, 2011). (W) “*Merostomichnites isp.*” (traced from Hanken & Stormer, 1975). (X)

1401 *Oniscoidichnus filiciformis* (traced from Brady, 1947). (Y) *Palmichnium antarcticum*
1402 (traced from Braddy & Milner, 1998). (Z) *Paleohelcura tridactyla* (traced from Gilmore,
1403 1926). (AA) *Protichnites septemnotatus* (traced from Owen, 1852). (BB) *Robledoichnus*
1404 *lucasi* (traced from Kozur & Lemone, 1995). (CC) *Shalemichnus sittigi*, half of trackway
1405 (traced from Kozur & Lemone, 1995). (DD) *Stiallia berriana* (traced from Smith, 1909).
1406 (EE) *Stiaria quadripedia* (traced from Walker, 1985). (FF) *Mitchellichnus ferrydenensis*
1407 (traced from Walker, 1985). (GG) *Keircalia multipedia* (traced from Walker, 1985).
1408 (HH) *Danstairia vagusa* (traced from Walker, 1985). (II) *Glaciichnium liebegastensis*
1409 (traced from Walter, 1985). (JJ) *Warvichnium ulbrichi* (traced from Walter, 1985). (KK)
1410 *Siskemia bipediculus* (traced from Walker, 1985). (LL) *Siskemia elegans* (traced from
1411 Walker, 1985). (MM) *Siskemia latavia* (traced from Walker, 1985). (NN) Extant
1412 notostracan traces (traced from Knecht et al., 2009). (OO) Extant crayfish trace in
1413 saturated, very fine sand (traced from Fairchild & Hasiotis, 2011). (PP) SGDS 1290.
1414

1415 **TABLE CAPTIONS**

1416

1417 **Table S1: Measurements (in mm) of arthropod locomotory ichnofossil SGDS 1290.**

1418 Measurements in parentheses are approximated based on faint portions of paramedial
1419 impressions. See Fig. 3C for a depiction of the measurements taken.

Figure 1

Location of the St. George Dinosaur Discovery Site at Johnson Farm.

(A) Map showing the location of the St. George Dinosaur Discovery Site in St. George, Utah.

(B) The museum at the St. George Dinosaur Discovery Site. Photograph by Jerald D. Harris.

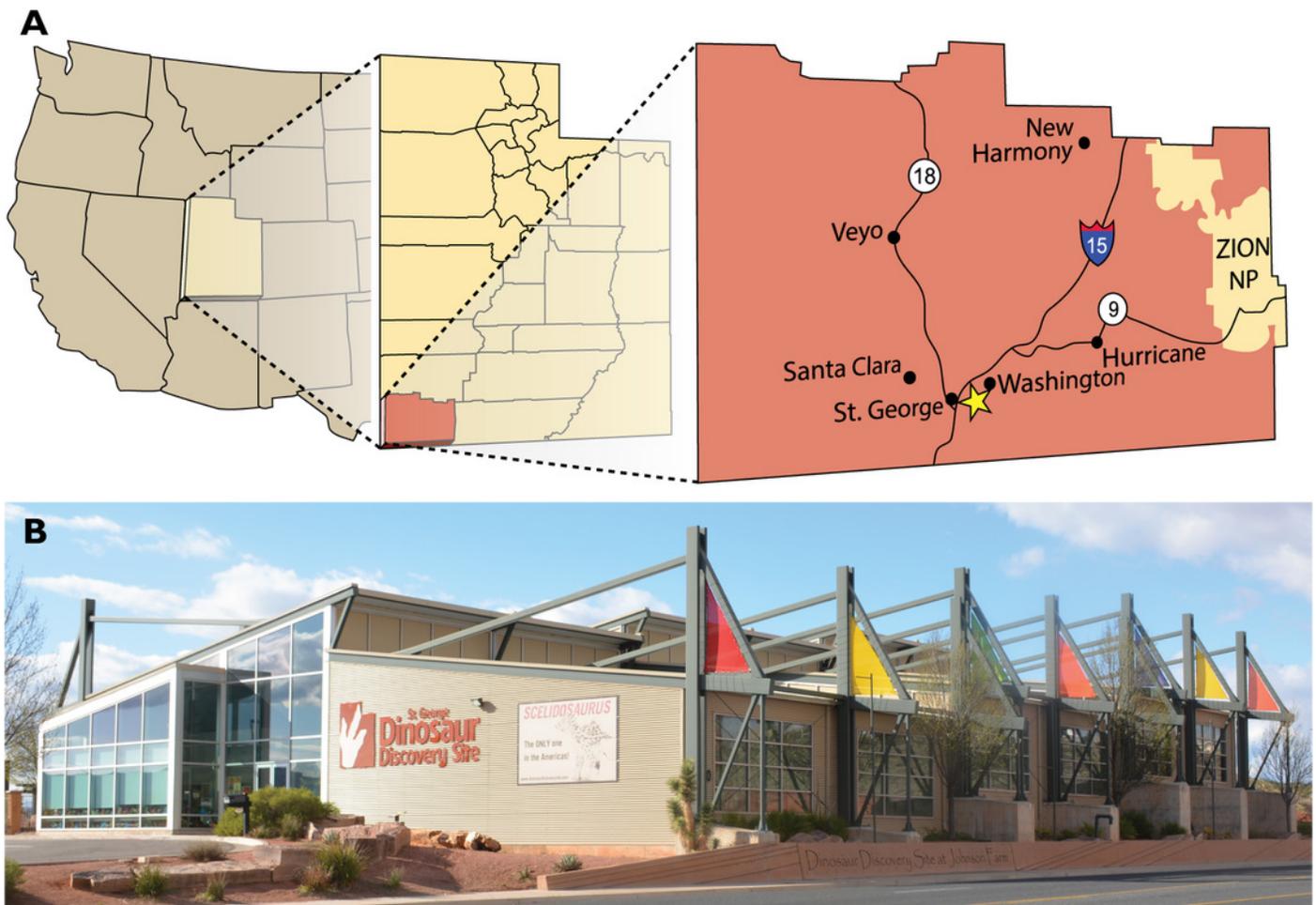


Figure 2

Stratigraphic section at and immediately around the St. George Dinosaur Discovery Site (SGDS) in St. George, Utah.

Arthropod locomotory trackway SGDS 1290 comes from the “Top Surface Tracksite” layers of the Johnson Farm Sandstone Bed (red arrows).

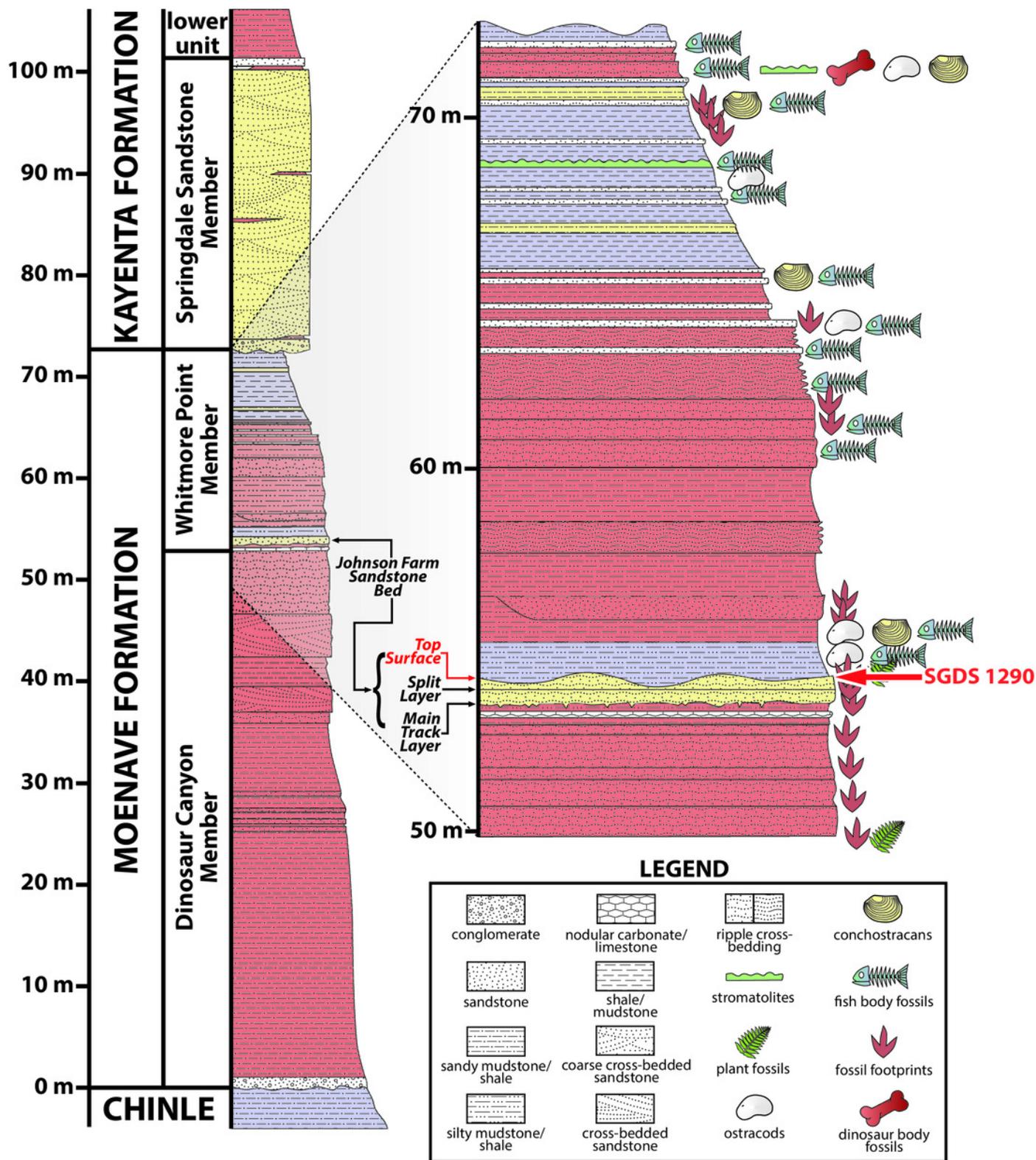


Figure 3

Arthropod locomotory ichnofossil SGDS 1290.

(A) Photograph of specimen; scale in mm. (B) Schematic diagram of specimen. Arrow indicates direction of movement. L = left track; R = right track; Ov = overtracks. Numbers indicate position of tracks in sequence from posterior to anterior. Photograph by Andrew R.C. Milner. (C) Schematic diagram of specimen, showing examples of how measurements were taken. IEW = external width between paramedial impressions; IIW = internal width between paramedial impressions; IW = paramedial impression width; TL = track length; TTI = distance between track and adjacent paramedial impression; TW = track width. (D) Photograph of extant crayfish walking trace made in saturated, very fine sand for comparison to fossil in (A); scale in cm. Arrow indicates direction of travel; photograph rotated to have the same orientation as (A). Modified from Fairchild & Hasiotis (2011: fig. 4E); reproduced with permission from SEPM.

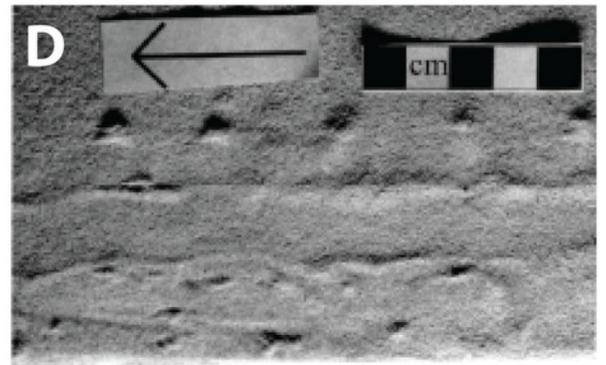
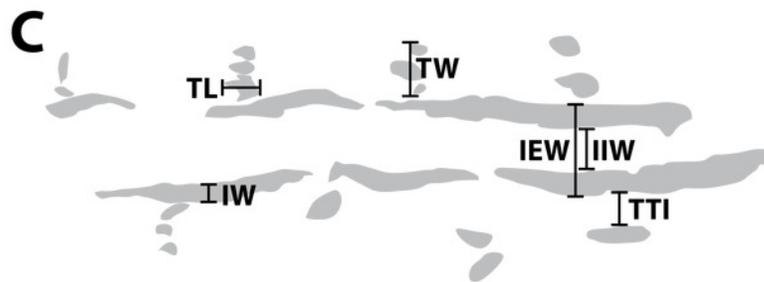
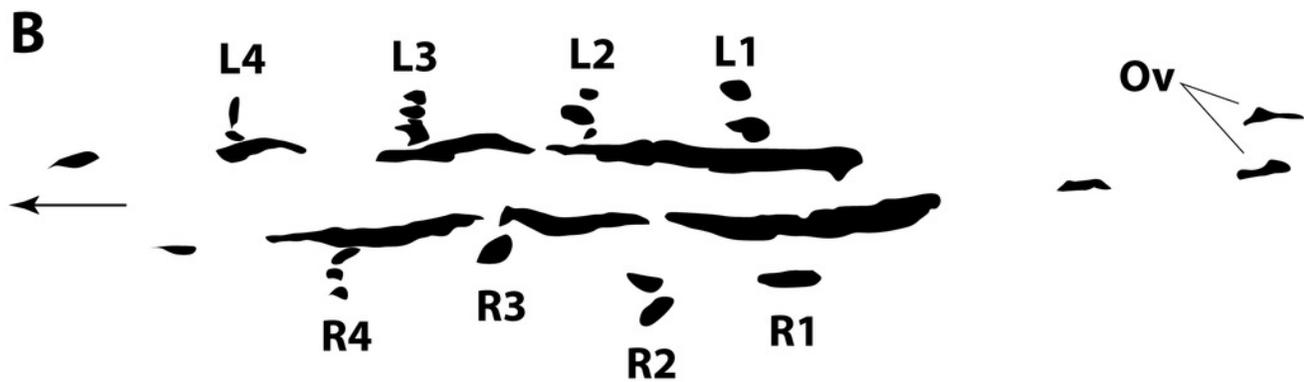


Figure 4

Schematic diagrams of locomotory (presumably walking) ichnofossils attributed to arthropods (A–MM) and extant walking traces made by arthropods (NN–OO) compared to SGDS 1290 (PP).

Diagrams not to scale. (A) *Acanthichnus cursorius* (traced from Hitchcock, 1858). (B) *Asaphoidichnus trifidus* (traced from Miller, 1880). (C) *Bifurculapes laqueatus* (traced from Getty, 2016). (D) *Coenobichnus currani* (traced from Walker, Holland & Gardiner, 2003). (E) *Copeza triremis* (traced from Hitchcock, 1858). (F) *Danstairia congesta* (traced from Walker, 1985). (G) *Diplichnites aenigma* (traced from Dawson, 1873). (H) *Diplichnites gouldi* Type A (traced from Trewin and McNamara, 1994). (I) *Eisenachichnus inaequalis* (traced from Kozur, 1981). (J) *Foersterichnus rossensis* (traced from Pirrie, Feldmann & Buatois, 2004). (K) *Hamipes didactylus* (traced from Getty, 2018). (L) *Laterigradus lusitanicus* (traced from de Carvalho et al., 2016). (M) *Lithographus hieroglyphicus* (traced from Hitchcock, 1858). (N) *Maculichna varia* (traced from Anderson, 1975a). (O) *Merostomichnites narrangansettensis* (traced from Packard, 1900). (P) *Mirandaichnium famatinense* (traced from Aceñolaza, 1978). (Q) *Octopodichnus didactylus* (traced from Gilmore, 1927). (R) *Petalichnus multipartatus* (Miller, 1880). (S) *Pterichnus centipes* (traced from Hitchcock, 1865). (T) *Tasmanadia twelvetreesi* (traced from Glaessner, 1957). (U) *Umfolozia sinuosa* (traced from Anderson, 1981). (V) *Kouphichnium lithographicum* (traced from Gaillard, 2011). (W) “*Merostomichnites* isp.” (traced from Hanken & Stormer, 1975). (X) *Oniscoidichnus filiciformis* (traced from Brady, 1947). (Y) *Palmichnium antarcticum* (traced from Braddy & Milner, 1998). (Z) *Paleohelcura tridactyla* (traced from Gilmore, 1926). (AA) *Protichnites septemnotatus* (traced from Owen, 1852). (BB) *Robledoichnus lucasi* (traced from Kozur & Lemone, 1995). (CC) *Shalemichnus sittigi*, half of trackway (traced from Kozur & Lemone, 1995). (DD) *Stiallia berriana* (traced from Smith, 1909). (EE) *Stiaria quadripedia* (traced from Walker, 1985). (FF)

Mitchellichnus ferrydenensis (traced from Walker, 1985). (GG) *Keircalia multipedia* (traced from Walker, 1985). (HH) *Danstairia vagusa* (traced from Walker, 1985). (II) *Glaciichnium liebegastensis* (traced from Walter, 1985). (JJ) *Warvichnium ulbrichi* (traced from Walter, 1985). (KK) *Siskemia bipediculus* (traced from Walker, 1985). (LL) *Siskemia elegans* (traced from Walker, 1985). (MM) *Siskemia latavia* (traced from Walker, 1985). (NN) Extant notostracan traces (traced from Knecht et al., 2009). (OO) Extant crayfish trace in saturated, very fine sand (traced from Fairchild & Hasiotis, 2011). (PP) SGDS 1290.

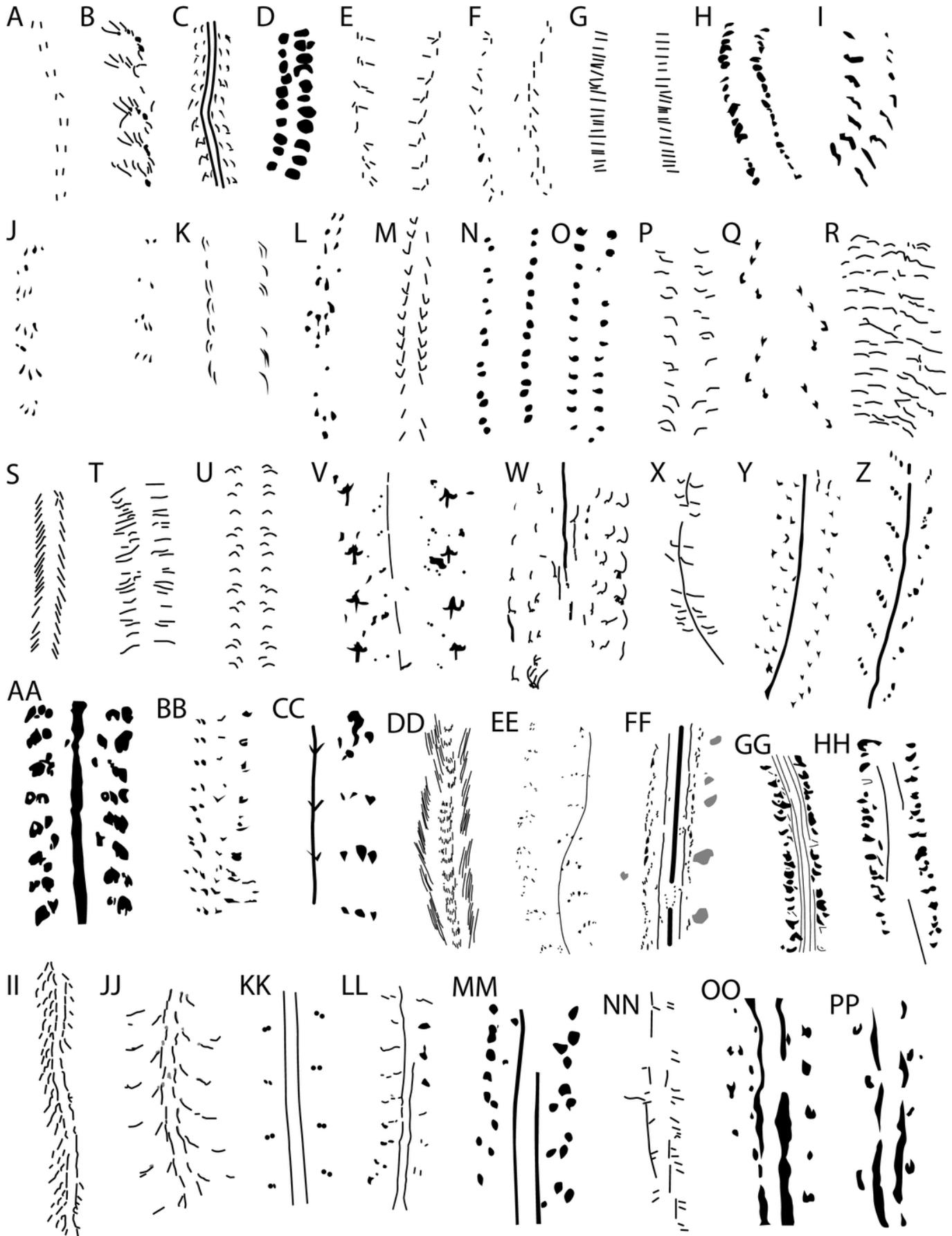


Table 1 (on next page)

Measurements (in mm) of arthropod locomotory trace fossil SGDS 1290.

Measurements in parentheses are approximated based on faint portions of paramedial impressions.

1 **Table 1: Measurements (in mm) of arthropod locomotory trace fossil SGDS 1290.**

2 Measurements in parentheses are approximated based on faint portions of paramedial
3 impressions.

4

Track Position	Length	Width	Impression Internal Width	Impression External Width	Track to Impression Distance	Left Impression Width	Right Impression Width
L1	4.3	6.4	4.7	9.1	1.8	3.1	2.5
L2	4.2	5.7	5.9	9.5	1.5	2.6	2.2
L3	5.2	5.8	5.8	9.7	0.0	1.6	1.6
L4	3.2	5.6	(8.1)	(11.2)	0.0	1.9	0.0
R1	6.8	2.5	4.2	9.2	3.0	3.0	2.4
R2	4.6	5.3	5.4	8.8	4.2	2.7	0.0
R3	4.5	6.4	6.5	9.2	2.4	1.3	1.3
R4	3.0	5.8	6.9	10.6	0.0	0.0	3.0

5

6